

Bidirectional Reflectance Distribution Function (BRDF) of Deep Convective Clouds (DCC) Derived from PARASOL Measurements and Compared to Radiative Transfer Computation and Model

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**CALCON 2014** 



## **DCC BRDF**:

## **Space Measurements and Modeling**

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## Summary

- Introduction
- The PARASOL instrument
- Methodology to derive the observed BRDF
- BRDF model derived from PARASOL measurements
- Modeling of DCC
- Analysis on the principal plane comparison Measurements // Computations
- Sensitivity analysis
- Conclusion



### Introduction

#### DCC are used for many years for calibration purposes

- not only trending, but also cross-calibration, interband, field-of-view
- GSICS reference calibration methods
- that's my powerful white diffuser (see Fougnie et al., IEEE TGARS, 2009)

#### PARASOL is an instrument allowing a bidirectional characterization of any targets of the Earth-surface system

- because DCC were used for PARASOL calibration purposes, an archive of CC observations was build and is available (8-years)
- it can be used to characterized BRDF of DCC

#### Model and radiative transfer computations are also possible

- also developed for calibration purposes
- can be used to derive BRDF and compare with measurements

## **The PARASOL Instrument**





### **Geometrical sampling**





### The Mean DCC

- Database = full archive 2005-2012
- Assume a « mean DCC » :

BRDF depends on :

- \* Cloud particle type
- \* Cloud optical thickness
- \* Cloud structure

hypothesis = all selected pixels always observe the same « mean DCC »

- All viewing directions are covered
  - 1 pixel = 16 views per track
  - N pixels = Nx16 views per track
  - 16 tracks per cycle
- •All solar angles are covered :
  - along the year : from 15 to 45°
  - orbital drift after 2009 = access to large angles, up to 60°



### Selection of DCC – The perfect storm

### How suitable DCC are selected for PARASOL (not TIR bands) :

#### • Operational procedure : every month, acquisitions are collected over

- oceanic sites in Guinée et Maldives
- $\rho$ nua>0.7, neighborhood (5x5) < , 400<Papp< 50hPa
- "nadir/zenith" geometries :  $\theta$ s<30° et  $\theta$ v<40° (avoiding shadow)
  - This nadir/zenith geometry is for calibration purposes -> here extended to all available angles





Criteria	Characteristics	Properties
reflectance in band 865 > 0.7	intensive scattering inside the cloud	dense scattered cloud
Inter-tropical sites	favorable areas for the convection mechanism	convective cloud
oceanic site	negligible surface contribution	predominance of the cloud
apparent pressure < 400HPa	top of the cloud > 11km low molecular and aerosol impact	very high cloud
cloud size > 70x70 km²	spatially large cloud structure	large cloud
rms for band 865 < 3% over 30x30km²	homogeneity minimization of structure effect	homogenous cloud
solar angle < 30° viewing angle < 40°	"nadir/zenith" viewing low bidirectional and shadow effects	reference geometry

### Methodology – construction of a mean DCC

- A BRDF structure is initialized with
  - ◆2° x2° bins for VZA and RAA
  - 10° range for SZA : [20°;30°], [30°;40°], [40°;50°], [50;60°] (to few for [0°;20°])
- All Measured reflectance from the archive (all geometries, all dates) is stored in its corresponding (VZA, SZA, RAA) box
  - ◆BRDF assumed to be symetrical to principal plane, i.e. 0° <RAA<180°</p>
- For each box and each wavelength, are computed :
  - mean TOA reflectance
  - standard-deviation
  - number of pixel per bin
- Visualization : polar plot









#### ■ 30° < SZA < 40°









## **Modeling DCC**

### Hu et al. model (2004) : used by GSICS for cross-calibration over DCC

- based on broadband CERES data collection
- extrapolated to narrowband using a non linear regression neural network
- paramaters are taken from hualb.data file (reference GSICS)

### Lafrance et al. model (2002) : used by CNES for calibration over DCC

- based on Discrete Ordinate Code narrowband
- 16 layers, droplet in 1-6km, ice in 6-15km, molecular scattering + aerosol background
  - DCC are very white for VIS (most of contributors are minor)
  - cloud particle type (CPT) and cloud optical thickness (COT) are the dominant parameters
- CPT
  - Historical CPT : plate and hexogonal crystals based on Macke et al. (1996)
  - Revised CPT : <u>RHM</u> (Rough Hexagonal Monocrystals) derived by Labonnote et al. and based on Hess et al. reference (1998)





### Analysis on the principal plane



- Excellent consistency between RHM and Hu
- Good consistency with measurements for VZA>45°
- Some differences for large angles





### Analysis on the principal plane



- Excellent consistency between RHM and Hu
- Good consistency with measurements for VZA<40°</li>
- Some differences for large angles





### **Sensitivity analysis**

(from Fougnie and Bach, 2009 and Lafrance et al., 2002)





### **Sensitivity analysis**

- Cross-calibration strategy :
- at a given date & geolocation // COT and CPT are given by the actual geophysics
- exact simultaneity, i.e. solar and viewing geometries, is not always possible
- Behavior versus VZA for SZA=25° and 45° and for 16 COT
  - computed at 670nm for CPT=RHM (principal plane)





### Sensitivity analysis – The Crescent BRDF

#### Optimisation for Cross-calibration

- A crescent (RAA//VZA) could to be preferred



- moderate variation with both SZA and VZA
- toward the backscattering half-space and/or large viewing angles
- doesn't simply correspond to the backscattering direction



·50

140

Meas. Hu



### Conclusion

A first DCC BRDF characterization based on the mean TOA observation by PARASOL was derived

Comparison with reference model (Hu) and computations show :

- A very good consistency between measurements and computation
- This is a statistical mean observation, BDRF may depend on various parameters
- Here 490/670/865 are presented
  - 565, 763, 765, 910, 1020 available but sensitive gaseous absorption, straylight for 443
- Main difference between measurements & modeling
  - MODEL : Assumption of a flat homogeneous level for cloud
  - MEASU : roughness of the DCC, shadowing effects
  - SZA variation can be correlated to seasonal change of microphysics (CPT)
- Recommandation for cross-calibration purposes is to select the most isotropic part of the BRDF
  - avoid geometries around the specular direction, ideally prefer « the crescent » (large angles on the backscattering half space)

### References

# Thank you !

#### ■ Macke et al., *J. Atmos. Sci.*, 1996

Single Scattering Properties of Atmospheric Ice Crystals

#### ■ Hess et al., J. Quant. Remote-Sensing. Tech., 1998

Scattering Matrice of Imperfect Hexagonal Ice Crystals

#### ■ C.-Labonnote et al., *J. Geophys. Res.*, 2001

Polarized Light Scattering by Inhomogeneous Hexagonal Monocrystals. Validation with POLDER.

#### ■ Lafrance et al., *IEEE TGARS*, 2002

Interband Calibration Over Clouds for POLDER Space Sensor

#### Hu et al., *IEEE TGARS*, 2004

Application of DCC Albedo Observation to satellite-based Study of the Terrestrial Atmosphere

#### ■ Fougnie et al., *IEEE TGARS*, 2009

Monitoring of Radiometric Sensitivity Changes of Space Sensors Using DCC – Operational Application to PARASOL

#### ■ Fougnie et al., CALCON, 2011

Adding an On-board Diffuser in Front of the PARASOL Instrument

#### ■ Fougnie et al., CALCON, 2013

Definition, Adjustment and Validation of a Physical Model to Describe the PARASOL Radiometric Tending



## **Backup Slides**

# **c**nes



### How lambertian are DCC – Spectral Ratio

