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P. W. Lommen

D. C. Wilkin

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1975/76 PROGRESS REPORT

**PROGRAMMING PHASE OF
WATER RESPONSE ECOSYSTEM MODEL:
II. ABIOTIC SUBMODELS**

P. W. Lommen* and K. A. Marshall
Utah State University
(*now at HDR Ecosciences, Santa Barbara, California)

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Ecology Center, Utah State University, Logan, Utah 84322

This report describes a portion of the Desert Biome Water Response Ecosystem Model. Five Research Memoranda comprise the full description: Introduction and support programs (RM 76-36); Abiotic submodels (RM 76-37); Animal submodel (RM 76-38); Perennial plant, nitrogen and decomposition submodels (RM 76-39); and Annual plant submodel (RM 76-40). The objectives of the Water Response Model, information on the arrangement of material distributed among the five Research Memoranda and descriptions of program MAIN and support programs F1, F3 and FTAVE are contained in Research Memorandum 76-36, **Programming phase of water response ecosystem model: I. Introduction and support programs**. The relationships between various sections of the model, their interactions and location in the report series are summarized in Table 1 of RM 76-36.

INTRODUCTION

This research memorandum describes submodels PSWG, HEAT and WATER (and their support programs) of the Desert Biome Water Response Model. These submodels provide environmental (or driving) variables to the plant, animal and soil-microbe submodels. PSWG (for Pseudo Stochastic Weather Generator) provides meteorological variables (air temperature, precipitation, plus several

others). HEAT calculates soil temperature profile, given soil heat flow characteristics and air temperature. WATER determines soil water potential profile given initial conditions, soil moisture flow characteristics and several water inputs and outputs (precipitation, evaporation, transpiration, runoff). Transpiration, through its effect on soil water potential, is the only biological effect on any of the ten or so environmental variables determined in these submodels.

A. PSWG
(including support programs EVAP, RINT,
IPROB, RNOR)

K. A. Marshall and P. W. Lommen

GENERAL DESCRIPTION OF PSWG SUBMODEL

PSWG is a submodel which generates driving variables for the Water Response Model. (The name PSWG is generated from Pseudo Stochastic Weather Generator).

These variables are maximum, minimum and mean time-step air temperature at 2 m; precipitation and its intensity for the time-step; potential evaporation for the time-step; and mean values for the time-step of relative humidity, wind speed, fraction of possible sunlight and photoperiod.

For any run, one of two versions is used depending on data availability. Version I is used for debugging, tuning and validation runs. For debugging runs, all driving variables are stochastically generated. For tuning and

validation runs where driving variables must correspond to the weather of a given year or years, daily temperature and precipitation records for the site are used. The remaining driving variables are chosen stochastically because daily records for the site are either not available at all or are not complete enough. Variables are stochastically chosen with the aid of random numbers and parameters generated from six years' data of the variable in question obtained from the United States Weather Bureau at the nearest complete weather station to our site. For Curlew Valley simulation, data from Pocatello, Idaho, were used. For Rock Valley simulations, Las Vegas, Nevada, data were used. Refer to Table A-1 for definitions of variables for both versions of PSWG. Definitions will also be found in the complete program listings at the end of this section.

Version II of PSWG is used for five-year (or longer) simulations. It generates nothing stochastically. Instead, it reads temperature, precipitation, wind speed, fraction of possible sunlight and relative humidity data from a weather data file. (For five-year Curlew Valley runs, data were used from the U.S. Weather Bureau station at Pocatello.)

Table A-1. Variable dictionary for PSWG, Versions I and II

AMT(6)	AMT(I) IS AMOUNT OF PRECIPITATION IN MM IN EVENT CLASS I.
DECL	ANGLE CALCULATED IN PHOTOPERIOD CALCULATION.
EVAP	SUBROUTINE WHICH CALCULATES POTENTIAL EVAPORATION.
FACTOR	PRECIPITATION FACTOR BY WHICH EACH EVENT IS MULTIPLIED.
HUM(10,13)	INTEGRATED PROBABILITY DISTRIBUTION OF RELATIVE HUMIDITY CLASS BY PERIOD OF YEAR.
IPROB	FUNCTION WHICH DETERMINES DEPENDENT VARIABLE GIVEN INTEGRATED PROBABILITY DISTRIBUTION, BY RANDOMLY CHOOSING INDEPENDENT VARIABLE IN RANGE 0-1.
ISEED	SEED FROM WHICH RNOR GENERATES NEXT RANDOM NUMBER.
IYEST	EQUALS 1 IF THERE WAS NO PRECIPITATION YESTERDAY, IF EQUALS 2 THEN THERE WAS PRECIPITATION YESTERDAY AND PROBABILITY OF PRECIPITATION TODAY IS INCREASED.
J	CLASS INDEX, USED IN PRECIP, WIND, SUN, HUM GENERATION.
JDAY	JULIAN DATE AT BEGINNING OF TIMESTEP.
LAT	LATITUDE OF SITE.
MDAY	CURRENT JULIAN DAY (INCREMENTED IN PSWG FROM JDAY TO JDAY+PMDT).
MTIME	INDEX OF 4 WEEK PERIOD OF YEAR IN WHICH MDAY FALLS.
NNN	COUNTER TO DETERMINE WHEN TO START PRINTING ON NEXT PAGE.
PMDT	TIMESTEP LENGTH, DAYS.
PREC(6,13)	INTEGRATED PROBABILITY DISTRIBUTION OF PRECIPITATION CLASS BY PERIOD OF YEAR.
PRECIP	PRECIPITATION FOR CURRENT DAY, MM.
PMJDAT	JULIAN DATE AT BEGINNING OF TIMESTEP. VALUE DETERMINED IN MAIN PROGRAM.
PWFAC	FACTOR FOR ADJUSTING AVERAGE WIND SPEED IF SITE IS KNOWN TO HAVE DIFFERENT AVERAGE WIND SPEED.
RAIN(2,13)	ARRAY HOLDING VALUES OF (1. = PROBABILITY OF PRECIP. TODAY) GIVEN PERIOD OF YEAR AND WHETHER OR NOT WE HAD PRECIP. YESTERDAY.

Table A-1, continued

RCHECK(20)	VARIABLE USED TO READ AND WRITE COMMENTS IN WITH INITIAL DATA.
READ	IF *.TRUE, THEN READ TEMPERATURE PRECIPITATION DATA FROM SEPARATE DATA FILE.
RINT	SUBROUTINE WHICH CALCULATES PRECIPITATION INTENSITY.
RNOR	RANDOM NORMAL NUMBER GENERATOR (68% OF VALUES LIE BETWEEN -1 AND +1).
SHUM	RUNNING SUM OF UP TO PMDT DAYS' RELATIVE HUMIDITY.
SMAX	RUNNING SUM OF UP TO PMDT DAYS' MAXIMUM TEMPERATURES, DEGREES CELCIUS.
SMIN	RUNNING SUM OF UP TO PMDT DAYS' MINIMUM TEMPERATURES, DEGREES CELCIUS.
SPREC	RUNNING SUM OF UP TO PMDT DAYS' PRECIPITATIONEVENTS,MM.
SSUN	RUNNING SUM OF UP TO PMDT DAYS' PER CENT POSSIBLE SUNLIGHT
SUN(10,13)	INTEGRATED PROBABILITY DISTRIBUTION OF PER CENT POSSIBLE SUNLIGHT CLASS BY PERIOD OF YEAR,
SWIND	RUNNING SUM OF UP TO PMDT DAYS' AVERAGE WIND SPEEDS, KM/HR.
TMAX	MAXIMUM TEMPERATURE FOR CURRENT DAY, DEGREES CELCIUS.
TMIN	MINIMUM TEMPERATURE FOR CURRENT DAY, DEGREES CELCIUS.
TRAIN(20)	ARRAY WHICH HOLDS DAILY PRECIP VALUES FOR TIMESTEP (USED BY RAIN INTENSITY ROUTINE).
WIND(7,13)	INTEGRATED PROBABILITY DISTRIBUTION OF WIND SPEED CLASS BY PERIOD OF YEAR,
XMAX(1),XMAX(2)	PARAMETERS FOR SINE WAVE FIT TO SEVERAL YEARS' DAILY MAXIMUM TEMPERATURES, TAKEN AT NEAREST WEATHER BUREAU SITE.
XMAX(3)	STANDARD DEVIATION OF DATA ABOUT SINE WAVE FIT USING XMAX(1) AND XMAX(2).
XMIN(1),XMIN(2)	PARAMETERS FOR LINEAR REGRESSION BETWEEN TMAX AND TMIN.
XMIN(3)	STANDARD DEVIATION OF TMIN.
ZAIRT	AVERAGE DAILY AIR TEMPERATURE FOR TIMESTEP, DEGREES CELCIUS, DRIVING VARIABLE USED BY OTHER SUBMODELS.
ZESUM	SUM FROM OCTOBER 1 OF ZEVAP, MM.
ZEVAP	POTENTIAL EVAPORATION, MM/TIMESTEP, DRIVING VARIABLE USED BY OTHER SUBMODELS.
ZPHPD	PHOTOPERIOD OF FIRST DAY IN TIMESTEP, HOURS, DRIVING VARIABLE USED BY OTHER SUBMODELS.
ZRAIN	TOTAL PRECIPITATION FOR TIMESTEP, MM, DRIVING VARIABLE USED BY OTHER SUBMODELS.
ZRH	AVERAGE DAILY RELATIVE HUMIDITY, DECIMAL FRACTION FROM 0-1, DRIVING VARIABLE USED BY OTHER SUBMODELS.
ZRINT	PRECIPITATION INTENSITY, MM/HR, DRIVING VARIABLE USED BY OTHER SUBMODELS.
ZRSUM	SUM FROM OCTOBER 1 OF ZRAIN, MM.
ZSUN	AVERAGE DAILY PER CENT POSSIBLE SUNLIGHT FOR TIMESTEP, DECIMAL FRACTION FROM 0-1, DRIVING VARIABLE USED BY OTHER SUBMODELS.
ZTMAX	AVERAGE DAILY MAXIMUM TEMPERATURE FOR TIMESTEP, DEGREES CELCIUS, DRIVING VARIABLE USED BY OTHER SUBMODELS.
ZTMIN	AVERAGE DAILY MINIMUM TEMPERATURE FOR TIMESTEP, DEGREES CELCIUS, DRIVING VARIABLE USED BY OTHER SUBMODELS.
ZWIND	AVERAGE DAILY WIND SPEED FOR TIMESTEP, DRIVING VARIABLE USED BY OTHER SUBMODELS.

PROGRAM DESCRIPTION

PSWG, VERSION I

Only the important segments of the FORTRAN code are shown and described. Sequence numbers are shown to aid in reference to the full code listing which follows the Program Description. All comment cards, specification statements and bookkeeping sections have also been deleted. Definitions of variable names may be found in Table A-1, which also appears at the beginning of the program listing.

```

DO 3 I=1,PMDT

MTIME=MDAY/28+1
IF (MTIME .GT. 13) MTIME=13

IF(.NOT.READ)GOTO100
020 READ(6,25,END=500)TMAX,TMIN,PRECIP
025 FORMAT(7X,2F6.1,36X,F6.1)
GOTO30

500 READ=,FALSE,
GOTO100

IF(TMAX.LT.TMIN)GOTO40
IF((TMAX.LT.=30.)OR.(TMAX.GT.45.))GOTO40
IF((TMIN.LT.=30.)OR.(TMIN.GT.45.))GOTO40
IF(((PRECIP.LT.50.)AND.(PRECIP.GE.0.0)).OR.(PRECIP.GT.5000.))
* GOTO60

040 WRITE(6,50)PMJDAT,TMAX,TMIN,PRECIP
050 FORMAT(' ',I5,3F12.5)
STOP

```

PSG1 183

Beginning of main loop in program. Go through loop once for each day of time-step in order to get daily weather information for appropriate average or summed value.

PSG1 184
PSG1 185

Determine MTIME, index of four-week period of year in which MDAY, current Julian day, falls.

PSG1 189
PSG1 190
PSG1 191
PSG1 192

If flag is set (i.e., if READ is .TRUE.) then read temperature and precipitation data from separate weather file. Then check if data are reasonable.

PSG1 194
PSG1 197

If control reaches here then all data have been read and remainder of temperatures and precipitation amounts will be generated stochastically.

PSG1 202
PSG1 203
PSG1 204
PSG1 205
PSG1 206

Check data just read and see if they are reasonable.

PSG1 210
PSG1 211
PSG1 212

We reach here only if a datum was judged unreasonable. Write offending card and stop execution.

```

060 CONTINUE
  SMAX=SMAX+TMAX
  SMIN=SMIN+TMIN
  SPREC=SPREC+PRECIP
  TRAIN(I)=PRECIP
  GOTO2

```

```

PSG1 215
PSG1 216
PSG1 217
PSG1 218
PSG1 220
PSG1 221

```

Keep running sums over PMDT of temperature and precipitation information. Also load appropriate place in TRAIN array for rain intensity calculation. Go on to wind speed section.

```

100 CONTINUE
  TMAX = XMAX(1) + XMAX(2)*SIN(.017214*MOD((PMJDAT+344.),365.))
  C = 1.570796)
  TMIN = TMAX + XMAX(3)*RNOR(ISEED)
  TMIN=XMIN(1)+XMIN(2)*TMAX
  TMIN=TMIN+XMIN(3)*RNOR(ISEED)
  IF (TMIN .GT. TMAX) GO TO 1
  SMAX=SMAX+TMAX
  SMIN=SMIN+TMIN

```

```

PSG1 225
PSG1 234
PSG1 235
PSG1 237
PSG1 238
PSG1 239
PSG1 240
PSG1 241
PSG1 242

```

Control reaches here if we wish to stochastically generate the current day's temperatures. First, calculate the mean maximum temperature for the day. Next, generate the day's minimum temperature from the maximum. Next, make sure the minimum is less than or equal to the maximum. Finally, add these temperatures to the sums of temperatures for previous days of the time-step.

```

IYEST=IPROB(RAIN(IYEST,MTIME),1)
IF(IYEST .NE. 2) TRAIN(I)=0.0
IF (IYEST .NE. 2) GO TO 2
J=IPROB(PREC(1,MTIME),5)
SPREC=SPREC+AMT(J)
TRAIN(I)=AMT(J)

```

```

PSG1 246
PSG1 247
PSG1 248
PSG1 249
PSG1 250
PSG1 251

```

Stochastically generate the day's precipitation. First determine IYEST for current day which depends on previous day's value. (If previous value was 2 rather than 1, then probability of IYEST = 2 today is considerably increased.) If IYEST = 1, then there is no precipitation today and control goes to wind speed section. If IYEST = 2, then there is precipitation today and so the amount is then determined. Calculate J, the size class of the event, an integer from 1 to 6 which depends on time of year. Once size class is determined, the amount corresponding to this class is added to SPREC, and loaded in TRAIN(I).

```

002 J=IPROB(WIND(1,MTIME),7)
  S=WIND*SWIND+(J-1)*8.05

```

```

PSG1 254
PSG1 255

```

Generate today's wind speed. First find J, the size class. Today's wind speed is then $(J-1) \cdot 8.05$ km/hr (8.05 km/hr = 5.00 mi/hr). Add this value to the time-step running total of wind speeds.

```

J=IPROB(SUN(1,MTIME),10)
SSUN=SSUN+(J-1)*0.1
J=IPROB(HUM(1,MTIME),10)
SHUM=SHUM+(J-1)*0.1

```

```

PSG1 258
PSG1 259
PSG1 262
PSG1 263

```

Determine fraction of possible sunlight and relative humidity similarly to wind speed. Possible values taken on will be 0, 0.1, 0.2, . . . , 1.0.


```
003 MDAY=MDAY+1
```

```
PSG1 264
```

The main loop of PSWG is ended by incrementing the Julian date.

```
X*1./PMOT
ZTHAX=SMAX*X
ZTMIN=SMIN*X
ZWIND=SWIND*X
```

```
PSG1 269
PSG1 270
PSG1 271
PSG1 272
```

Determine time-step averages for maximum and minimum air temperatures and for wind speed.

```
ZWIND = ZWIND * PWFAC
```

```
PSG1 276
```

If site is known to have different average wind speed than measuring station, multiply wind speed just determined by this factor.

```
ZSUN=SSUN*X
ZRH=SHUM*X
ZRAIN=SPREC
ZAIRT=(ZTHAX+ZTMIN)/2.
```

```
PSG1 279
PSG1 280
PSG1 281
PSG1 282
```

ZSUN, ZRH and ZAIRT are averages for the time-step. ZRAIN is the total time-step precipitation.

```
DECL=ARCSIN(SIN(6.2831853*23.45/360.))*
C COS(MOD((PMJDAT+193.),365.)*6.2831853/365.)
X=(SIN(LAT*6.2831853/360.)*TAN(DECL)+SIN(.875*6.2831853/360.))/
C COS(LAT*6.2831853/360.)
IF (X .GT. 1.) X=1.
IF (X .LT. -1.) X=-1.
ZPHPD=12. + .1333333*ARCSIN(X)*360./6.2831853
```

```
PSG1 285
PSG1 286
PSG1 287
PSG1 288
PSG1 289
PSG1 290
PSG1 291
```

Calculate ZPHPD, photoperiod, in hours, of first day of time-step.

```
CALL RINT(ZRINT, TRAIN, ZRAIN, PMOT, PMJDAT, ISEED)
CALL EVAP
```

```
PSG1 295
PSG1 300
```

Determine rain intensity and potential evaporation.

```
ZRAIN=FACTOR*ZRAIN
```

```
PSG1 306
```

Change time-step precipitation if FACTOR is not equal to 1. By changing FACTOR from one run to another, keeping all else the same, we determine the response of the entire model to changes in precipitation, a primary objective of the Water Response Model, as its name implies.

```
ENTRY ZINIT
```

```
PSG1 333
```

Entry point for initialization purposes. Called from MAIN once. Variables peculiar to PSWG are then read and written, with comment cards at the beginning, middle and end.

```
CALL EVINIT
CALL RIINIT
```

PSG1 418
PSG1 419

Call initialization sections of subroutines EVAP and RINT.

PSWG, VERSION II

Version II is much more simple than Version I.

```
DO 3 I=1,PMDT
```

PSG2 120

Begin main loop. Go through loop once for each day of time-step.

```
READ(6,30,END=500)TMAX,TMIN,PRECIP,WIND,PPS,RH
30 FORMAT(10X,2F5.0,F5.2,F5.1,2F5.2)
```

PSG2 122
PSG2 123

Read today's weather data.

```
TMAX=(TMAX-32.)*0.55555
TMIN=(TMIN-32.)*0.55555
PRECIP=PRECIP*25.4
WIND=WIND*1.60934
```

PSG2 133
PSG2 134
PSG2 137
PSG2 140

Convert temperature, precipitation and wind speed to metric units.

```
IF(TMIN.GT.TMAX)GOTO40
IF((TMAX.GT.42.)OR.(TMIN.LT.-43.))GOTO40
IF((PRECIP.GT.40.)OR.(PRECIP.LT.0.0))GOTO40
IF((WIND.GT.60.)OR.(WIND.LT.0.0))GOTO40
IF((PPS.GT.1.0)OR.(PPS.LT.0.0))GOTO40
IF((RH.LE.1.00).AND.(RH.GE.0.0))GOTO60
40 WRITE(6,50)TMAX,TMIN,PRECIP,WIND,PPS,RH,PMJDAT
50 FORMAT(' ',7F10.3)
STOP
```

PSG2 144
PSG2 145
PSG2 146
PSG2 147
PSG2 148
PSG2 149
PSG2 152
PSG2 153
PSG2 154

Check if values just read are reasonable. If not, print offending card and stop. If reasonable, go on to summing section.

```
SMAX=SMAX+TMAX
SMTN=SMIN+TMIN
SPREC=SPREC+PRECIP
SWIND=SWIND+WIND
SSUN=SSUN+PPS
SHUM=SHUM+RH
```

PSG2 158
PSG2 159
PSG2 160
PSG2 161
PSG2 162
PSG2 163

Keep running sums of values for appropriate averages to be calculated below.

```
TRAIN(I)=PRECIP
```

PSG2 166

Load TRAIN(I) for use in rain intensity subroutine.

3 CONTINUE

PG2 168

End of main loop.

From here on, Version II is the same as Version I, with one exception: the entry point of Version II is much simpler than that of Version I because all the parameters for stochastically generating variables are not read in.

SUBROUTINE EVAP

EVAP calculates potential evaporation from soil surface in millimeters, during a time-step. The Blaney-Criddle method is used and closely follows the approach of Griffin et al. 1974.

Refer to Table A-2 for definitions of variables used in this program. This table is also found in the complete program listing at the end of this section.

ZEVAP=F1(ZAIRT, ARRAY1, NPTS)

EVAP 056

Here, ZEVAP is multiplying factor (≤ 1) depending on air temperature, which has been empirically determined.

Table A-2. Variable dictionary for EVAP

A, B, C, Z	TEMPORARY VARIABLES USED IN DAYLIGHT CALCULATION.
ARRAY(5,2)	PAIRS OF DATA POINTS FOR INTERPOLATION BY F1 IN CALCULATING BLANEY CRIDDLE FACTOR.
CVVCOV	FRACTION OF SOIL SURFACE COVERED BY ANNUAL AND PERENNIAL VEGETATION.
DALITE(365)	ARRAY INDEXED BY JULIAN DATE, GIVES NUMBER OF MINUTES OF DAYLIGHT ON THAT JULIAN DATE.
DLTFR(365)	ARRAY INDEXED BY JULIAN DATE AND IS FRACTION OF YEAR'S DAYLIGHT WHICH FALLS ON THAT JULIAN DATE.
F1	FUNCTION WHICH LINEARLY INTERPOLATES BETWEEN PAIRS OF DATA POINTS TO DETERMINE DEPENDENT VARIABLE.
LAT	LATITUDE OF SITE, DEGREES.
MM	FLAG USED ONLY IN CASE PHJDAT IS EVER ZERO.
NPTS	NUMBER OF PAIRS OF DATA POINTS ACTUALLY USED IN ARRAY.
PMOT	LENGTH OF TIMESTEP, DAYS.
PHJDAT	JULIAN DATE OF FIRST DAY OF TIMESTEP.
RCHECK(20)	ARRAY USED TO READ AND WRITE COMMENTS IN WITH INITIAL DATA.
TOTMIN	MINUTES OF DAYLIGHT IN YEAR.
ZAIRT	AVERAGE TIMESTEP AIR TEMPERATURE, DEGREES CELCIUS.
ZEVAP	POTENTIAL EVAPORATION FOR TIMESTEP, MM, DRIVING VARIABLE USED BY OTHER SUBMODELS.

```

IF(PMJDAT .GT. 0) GO TO 670
PMJDAT=1
MM=1
670 CONTINUE

```

```

EVAP 057
EVAP 058
EVAP 059
EVAP 060

```

PMJDAT is used as a subscript in the next line. These lines simply set it equal to 1 in case it should ever be 0.

```

ZEVAP=ZEVAP*((1.8*ZAIR7 + 32.)*DLTFR(PMJDAT)*(1.-CVVCOV)*25.4*PMDT EVAP 063

```

This, in essence, is subroutine EVAP. The potential evaporation equals the empirical factor times the temperature in degrees Fahrenheit, times the fraction of the year's daylight occurring today, times the fraction of uncovered soil surface, times 25.4 to change from inches to millimeters, times PMDT to change from one day to the value for the entire time-step.

```

CVVCOV=0,0

```

```

EVAP 064

```

Reset CVVCOV. It is calculated each time-step by one or both vegetation submodels.

```

IF(MM .EQ. 0) GO TO 680
PMJDAT=0
MM=0
680 CONTINUE

```

```

EVAP 065
EVAP 066
EVAP 067
EVAP 068

```

Set PMJDAT back to 0 if it was changed to 1 in lines 57-60 above.

```

ENTRY EVINIT

```

```

EVAP 073

```

Entry point called once from PSWG to read initial data and do daylight calculations which need to be done only once.

```

DO 650 I=1, 365
A=730, -.274*LAT+.00793*(LAT**2)
B=34, 2, .78*LAT+.1*(LAT**2)
C=1
Z=2, *3, 1416*((C+285,)/365, )
DALITE(I)=A+B*SIN(Z)
TOTMIN=DALITE(I)+TOTMIN
650 CONTINUE

```

```

EVAP 108
EVAP 109
EVAP 110
EVAP 111
EVAP 112
EVAP 114
EVAP 115
EVAP 116

```

In this loop calculate the length of daylight in minutes for each day of the year at latitude LAT. Also calculate TOTMIN, the total minutes of daylight in a year.

```

DO 660 I=1, 365
DLTFR(I)=DALITE(I)/TOTMIN
660 CONTINUE

```

```

EVAP 119
EVAP 120
EVAP 121

```

Calculate DLTFR, the fraction of daylight which occurs each day of the year.

SUBROUTINE RINT

This program calculates ZRINT, precipitation intensity in mm/hr. ZRINT is used by WATER in determining rate of infiltration into soil. Intensity is important mainly if the rate of water arriving at the soil surface is greater than the infiltration rate. Under these conditions, runoff or runoff could occur.

Refer to Table A-3 for definitions of variables used in this program. This table is also repeated in the complete program listing at the end of this section.

DO 550 I=1,PMDT

RINT 070

Make one pass through this loop for each day of time-step.

Table A-3. Variable dictionary for RINT

A, B	PARAMETERS IN HIGH INTENSITY CALCULATION.
ARRAY(10)	HOLDS UP TO 5 PAIRS OF DATA POINTS FOR INTERPOLATION BY FUNCTION F1 IN DETERMINING NORMAL INTENSITY.
BA	INTERMEDIATE VARIABLE IN HIGH INTENSITY CALCULATION.
B1	INTERMEDIATE VARIABLE IN HIGH INTENSITY CALCULATION.
F1	FUNCTION WHICH LINEARLY INTERPOLATES BETWEEN NPTS PAIRS OF DATA POINTS.
ISEED	INTEGER FROM WHICH RANDOM NUMBERS ARE GENERATED.
N	INDEX OF TIME OF YEAR. IN MIDDLE OF YEAR N#1. OTHER TIMES OF YEAR N#2.
NPTS	NUMBER OF PAIRS OF DATA POINTS IN NORMAL INTENSITY CALCULATION.
PDAY1,PDAY2	BEGINNING AND ENDING JULIAN DATES OF CENTRAL SEGMENT OF YEAR.
PMDT	LENGTH OF TIMESTEP, DAYS.
PMJDAT	JULIAN DATE OF BEGINNING OF TIMESTEP.
PTH	PRECIP AMOUNT ABOVE WHICH HIGH INTENSITY CAN OCCUR
PA(2)	COMPARED WITH R TO SEE IF HIGH INTENSITY ACTUALLY OCCURRED.
R	TEMPORARY VARIABLE. SET EQUAL TO RANDOM NUMBER
RCHECK(20)	ARRAY USED FOR READING AND WRITING COMMENTS IN WITH INITIALIZATION DATA.
TR	TEMPORARY VARIABLE.
TRAIN(20)	DAILY PRECIPITATION AMOUNTS, MM. ARRAY INDEXED BY DAY OF TIMESTEP
X	TODAY'S RAIN
Y	TODAY'S PRECIPITATION INTENSITY, MM/HR.
YSUM	SUM OVER TIMESTEP OF DAILY PRECIP INTENSITY TIMES PRECIP AMOUNT.
ZRAIN	TIMESTEP PRECIPITATION, MM. DRIVING VARIABLE USED BY WATER SUBMODEL.
ZRINT	RAIN INTENSITY, MM/HR. DRIVING VARIABLE USED BY WATER SUBMODEL.

```

X=TRAIN(I)

```

RINT 074

Get today's precipitation amount, x, from appropriate location in TRAIN, which was loaded in PSWG.

```

IF(X .GT. 1.E=5) GO TO 100
GO TO 550

```

RINT 077
RINT 079

If $x \geq 1.0 \times 10^{-5}$, this is considered precipitation so an intensity must be determined. Otherwise, go to end of loop.

```

100 IF(X .LT. PTH) GO TO 500

```

RINT 082

If there is precipitation today but the amount is less than threshold value PTH, go to statement 500 and calculate normal intensity for this time of year.

```

R=RANDOM(ISEED)

```

RINT 085

If we reach here, a high intensity is possible. Choose a random number uniformly distributed in interval 0 to 1.

```

N=2
IF((PMJDAT .GE. PDAY1) .AND. (PMJDAT .LE. PDAY2)) N=1

```

RINT 091
RINT 092

Determine which of two segments of year we're in. In central segment (summer, roughly) $N = 1$.

```

IF(R .GT. PA(N)) GO TO 500

```

RINT 095

If random number chosen above is greater than a threshold PA, which varies with segment of year, then go to section calculating normal intensity for this time of year.

```

R=R/PA(N)

```

RINT 097

If we reach here, a high intensity will result for today's precipitation. Generate a new random number.

```

150 TR=BA*(1.=R)
IF(TR .LE. 1.E=10) TR=1.E=10
Y*=B1*ALOG(TR)
IF(Y .GT. 25.) Y=25.
GO TO 520

```

RINT 099
RINT 100
RINT 102
RINT 105
RINT 106

Calculate intensity from exponential distribution. TR is not allowed to be smaller than 10^{-10} in order to keep ALOG function manageable. Also, intensity is kept less than 25 mm/hr (a value it would almost never reach anyway). Go on to YSUM calculation.

Lines 99 and 102 require several lines of algebra to derive. From "Local Climatological Data" sheets obtained from

the U.S. Weather Bureau for the station of interest (Pocatello, Idaho, for Curlew Valley runs), a histogram was constructed of

(frequency of precipitation events per .01 in/hr interval
where intensity ≥ 0.1 in/hr)

vs.

(intensity, in/hr).

Data over a period of several years were used. This histogram was fit with the expression:

$$\text{frequency} = ae^{-bx}, \quad (\text{A-1})$$

where x is the intensity, in/hr.

Then, if intensity > 0.1 in/hr, probability for intensity to be between x and $x + dx$ is

$$P(x)dx = (ae^{-bx}) (dx/.01). \quad (\text{A-2})$$

If intensity > 0.1 in/hr, the probability for it to be between .1 and ∞ is 1, i.e.:

$$1 = \int_{0.1}^{\infty} (ae^{-bx}) (dx/0.01) = [a/(.01b)]e^{-.1b}. \quad (\text{A-3})$$

If we let S be the probability that 0.1 in/hr $<$ intensity $\leq x^*$, then

$$S = \int_{0.1}^{x^*} (ae^{-bx}) (dx/0.01). \quad (\text{A-4})$$

If the integral is carried out,

$$S = 1 - [a/(0.01b)]e^{-bx^*}. \quad (\text{A-5})$$

Now, if we let $S = R$, where R is a random number chosen uniformly over the range from 0 to 1, and solve Equation A-5 for x , we get

$$x^* = - (1/b) \ln [(0.01b/a) (1 - R)]. \quad (\text{A-6})$$

The values of x^* will then have a distribution given by ae^{-bx} , as required.

The transformation of Equation A-5 into FORTRAN code is straightforward. Parameters a and b become A and B . The factor $0.01b/a$ becomes $0.01 * B/A = BA$; $1/b$ becomes $1./B \Rightarrow 25.4/B = BI$, so that intensities come out in mm/hr, not in/hr. Thus, $(0.01b/a) (1-R)$ becomes $BA*(1.-R) = TR$; x^* becomes Y ; and $Y = -BI*ALOG (TR)$.

500 Y=F1(PHJDAT, ARRAY, NPT8)

RINT 110

This is the normal intensity calculation. This line is missed only if intensity is high (i.e., this line is seldom missed). Function F1 interpolates between values found in ARRAY, dependent on time of year.

```
520 YSUM=X*Y + YSUM
```

```
RINT 113
```

Continue running sum over time-step of daily intensity times amount.

```
IF(ZRAIN .LE. 1.E-6) GO TO 570
ZRINT = YSUM/ZRAIN
GO TO 580
570 ZRAIN=0.0
ZRINT=0.0
580 IF((ZRINT*24.*PMDT) .LT. ZRAIN) ZRINT=ZRAIN/(24.*PMDT)
```

```
RINT 118
RINT 119
RINT 120
RINT 121
RINT 122
RINT 123
```

Determine average rain intensity, making sure of two things first: 1) that we don't divide by zero, and 2) that the precipitation amount we'd get if precipitation fell at rate ZRINT for the entire time-step is at least as large as the total time-step precipitation already determined in PSWG.

```
ENTRY RIINIT
```

```
RINT 129
```

Entry point for reading and writing initial data and comments, and for calculations which need to be done only once.

```
BA=0.01*B/A
BI=25.4/B
```

```
RINT 176
RINT 177
```

Calculate BA and BI which are used in high intensity calculation (lines 99 and 102).

FUNCTION IPROB (A,N)

IPROB (A,N) determines an index from 1 to N+1 given an integrated probability distribution A and a uniformly chosen random number X, $0 \leq X \leq 1$. A is an array with N values such that $A(1) < A(2) < A(3) < \dots < A(N) \leq 1$. If $A(1) > X$, then IPROB = 1. If $A(2) > X > A(1)$, then IPROB = 2, etc. If $X > A(N)$, then IPROB = N+1. Array A can be set up in many ways. The index determined can be for precipitation size class, wind speed class, relative humidity (RH) class, etc. For example, we have RH classes set up so that if IPROB = 1, RH = 0, if IPROB = 2, RH = 0.1, . . . , if IPROB = 11, RH = 1.0.

```
X=RANDOM(ISEED)
```

```
IPROB 04
```

Choose random number between 0 and 1.

```
DO 1 I=1,N
IF (A(I) .GT. X) GO TO 2
001 CONTINUE
```

```
IPROB 05
IPROB 06
IPROB 07
```

This DO loop takes each value of A, starting with the smallest, and checks it against the value of X, until a value is found greater than X. The loop is then exited.

```
002 IPROB=I
```

```
IPROB 08
```

Set value of function equal to current value of I. Usually this means $A(I-1) < X < A(I)$. If DO loop was exited normally, i.e., $X >$ all values of A, then $I = N + 1$.

COMPLETE PROGRAM LISTING

SUBROUTINE PSWG, VERSION I

```

SUBROUTINE PSWG                                PSG1 001
C THE PURPOSE OF THIS SUBROUTINE IS TO PROVIDE DRIVING VARIABLES FOR . PSG1 002
C USE BY SUBMODELS OF THE WATER RESPONSE MODEL. PSG1 003
C THIS VERSION OF PSWG USED IN TUNING AND VALIDATION RUNS. PSG1 004
C THIS VERSION OF PSWG CAN EITHER PSEUDO STOCHASTICALLY GENERATE ALL PSG1 005
C WEATHER VARIABLES OR IT CAN READ TEMP AND PRECIP DATA AND GENERATE PSG1 006
C THE REST. PSG1 007
C WRITTEN BY KIM MARSHALL PSG1 008
C DESERT BIOME PSG1 009
C UTAH STATE UNIVERSITY UMC 52 PSG1 010
C LOGAN, UTAH 84322 PSG1 011
C AUGUST 1976 PSG1 012
C PSG1 013
C PSG1 014
C PSG1 015
C PSG1 016
C PSG1 017
C PSG1 018
CCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCC PSG1 019
C VARIABLE DICTIONARY C PSG1 020
C AMT(6) AMT(I) IS AMOUNT OF PRECIPITATION IN MM IN EVENT C PSG1 021
C CLASS I. C PSG1 022
C DECL ANGLE CALCULATED IN PHOTOPERIOD CALCULATION. C PSG1 023
C EVAP SUBROUTINE WHICH CALCULATES POTENTIAL EVAPORATION. C PSG1 024
C FACTOR PRECIPITATION FACTOR BY WHICH EACH EVENT IS C PSG1 025
C MULTIPLIED. C PSG1 026
C HUM(10,13) INTEGRATED PROBABILITY DISTRIBUTION OF RELATIVE C PSG1 027
C HUMIDITY CLASS BY PERIOD OF YEAR. C PSG1 028
C IPROB FUNCTION WHICH DETERMINES DEPENDENT VARIABLE GIVEN C PSG1 029
C INTEGRATED PROBABILITY DISTRIBUTION, BY RANDOMLY C PSG1 030
C CHOOSING INDEPENDENT VARIABLE IN RANGE 0=1. C PSG1 031
C ISEED SEED FROM WHICH RNOR GENERATES NEXT RANDOM NUMBER. C PSG1 032
C IYEST EQUALS 1 IF THERE WAS NO PRECIPITATION YESTERDAY. IF C PSG1 033
C EQUALS 2 THEN THERE WAS PRECIPITATION YESTERDAY AND C PSG1 034
C PROBABILITY OF PRECIPITATION TODAY IS INCREASED. C PSG1 035
C J CLASS INDEX, USED IN PRECIP, WIND, SUN, HUM GENERATION. C PSG1 036
C JDAY JULIAN DATE AT BEGINNING OF TIMESTEP. C PSG1 037
C LAT LATITUDE OF SITE. C PSG1 038
C MDAY CURRENT JULIAN DAY (INCREMENTED IN PSWG FROM C PSG1 039
C JDAY TO JDAY+PHDT). C PSG1 040
C MTIME INDEX OF 4 WEEK PERIOD OF YEAR IN WHICH MDAY FALLS. C PSG1 041
C NNN COUNTER TO DETERMINE WHEN TO START PRINTING ON C PSG1 042
C NEXT PAGE. C PSG1 043
C PHDT TIMESTEP LENGTH, DAYS. C PSG1 044
C PREC(6,13) INTEGRATED PROBABILITY DISTRIBUTION OF PRECIPITATION C PSG1 045
C CLASS BY PERIOD OF YEAR. C PSG1 046
C PRECIP PRECIPITATION FOR CURRENT DAY, MM. C PSG1 047
C PMJDAT JULIAN DATE AT BEGINNING OF TIMESTEP. VALUE C PSG1 048
C DETERMINED IN MAIN PROGRAM. C PSG1 049
C PWFAC FACTOR FOR ADJUSTING AVERAGE WIND SPEED IF SITE IS C PSG1 050
C KNOWN TO HAVE DIFFERENT AVERAGE WIND SPEED. C PSG1 051
C RAIN(2,13) ARRAY HOLDING VALUES OF (1, = PROBABILITY OF PRECIP. C PSG1 052
C TODAY) GIVEN PERIOD OF YEAR AND WHETHER OR NOT WE HAD C PSG1 053
C PRECIP. YESTERDAY. C PSG1 054
C RCHECK(20) VARIABLE USED TO READ AND WRITE COMMENTS IN WITH C PSG1 055
C INITIAL DATA. C PSG1 056
C READ IF =.TRUE, THEN READ TEMPERATURE PRECIPITATION DATA C PSG1 057
C FROM SEPARATE DATA FILE. C PSG1 058
C RINT SUBROUTINE WHICH CALCULATES PRECIPITATION INTENSITY. C PSG1 059
C RNOR RANDOM NORMAL NUMBER GENERATOR (60% OF VALUES LIE C PSG1 060
C PSG1 061
C PSG1 062
C PSG1 063
C PSG1 064
C PSG1 065
C PSG1 066
C PSG1 067
C PSG1 068
C PSG1 069
C PSG1 070
C PSG1 071
C PSG1 072
C PSG1 073
C PSG1 074
C PSG1 075
C PSG1 076
C PSG1 077
C PSG1 078
C PSG1 079
C PSG1 080
C PSG1 081
C PSG1 082
C PSG1 083
C PSG1 084
C PSG1 085
C PSG1 086

```

C	BETWEEN -1 AND +1).	C PSG1 087
C		C PSG1 088
C	SHUM	C PSG1 089
C	RUNNING SUM OF UP TO PMDT DAYS' RELATIVE HUMIDITY.	C PSG1 090
C	SMAX	C PSG1 091
C	RUNNING SUM OF UP TO PMDT DAYS' MAXIMUM TEMPERATURES,	C PSG1 092
C	DEGREES CELCIUS.	C PSG1 093
C	SMIN	C PSG1 094
C	RUNNING SUM OF UP TO PMDT DAYS' MINIMUM TEMPERATURES,	C PSG1 095
C	DEGREES CELCIUS.	C PSG1 096
C	SPREC	C PSG1 097
C	RUNNING SUM OF UP TO PMDT DAYS' PRECIPITATIONEVENTS,MM.	C PSG1 098
C	SSUN	C PSG1 099
C	RUNNING SUM OF UP TO PMDT DAYS' PER CENT POSSIBLE	C PSG1 100
C	SUNLIGHT	C PSG1 101
C	SUN(10,13)	C PSG1 102
C	INTEGRATED PROBABILITY DISTRIBUTION OF PER CENT	C PSG1 103
C	POSSIBLE SUNLIGHT CLASS BY PERIOD OF YEAR,	C PSG1 104
C	SWIND	C PSG1 105
C	RUNNING SUM OF UP TO PMDT DAYS' AVERAGE WIND SPEEDS,	C PSG1 106
C	KM/HR.	C PSG1 107
C	TMAX	C PSG1 108
C	MAXIMUM TEMPERATURE FOR CURRENT DAY, DEGREES CELCIUS.	C PSG1 109
C	TMIN	C PSG1 110
C	MINIMUM TEMPERATURE FOR CURRENT DAY, DEGREES CELCIUS.	C PSG1 111
C	TRAIN(20)	C PSG1 112
C	ARRAY WHICH HOLDS DAILY PRECIP VALUES FOR TIMESTEP	C PSG1 113
C	(USED BY RAIN INTENSITY ROUTINE).	C PSG1 114
C	WIND(7,13)	C PSG1 115
C	INTEGRATED PROBABILITY DISTRIBUTION OF WIND SPEED CLASS	C PSG1 116
C	BY PERIOD OF YEAR,	C PSG1 117
C	XMAX(1),XMAX(2)	C PSG1 118
C	PARAMETERS FOR SINE WAVE FIT TO SEVERAL YEARS'	C PSG1 119
C	DAILY MAXIMUM TEMPERATURES, TAKEN AT NEAREST WEATHER	C PSG1 120
C	BUREAU SITE.	C PSG1 121
C	XMAX(3)	C PSG1 122
C	STANDARD DEVIATION OF DATA ABOUT SINE WAVE FIT USING	C PSG1 123
C	XMAX(1) AND XMAX(2).	C PSG1 124
C	XMIN(1),XMIN(2)	C PSG1 125
C	PARAMETERS FOR LINEAR REGRESSION BETWEEN	C PSG1 126
C	TMAX AND TMIN.	C PSG1 127
C	XMIN(3)	C PSG1 128
C	STANDARD DEVIATION OF TMIN.	C PSG1 129
C	ZAIRT	C PSG1 130
C	AVERAGE DAILY AIR TEMPERATURE FOR TIMESTEP, DEGREES	C PSG1 131
C	CELCIUS, DRIVING VARIABLE USED BY OTHER SUBMODELS.	C PSG1 132
C	ZESUM	C PSG1 133
C	SUM FROM OCTOBER 1 OF ZEVAP, MM.	C PSG1 134
C	ZEVAP	C PSG1 135
C	POTENTIAL EVAPORATION, MM/TIMESTEP.	C PSG1 136
C	DRIVING VARIABLE USED BY OTHER SUBMODELS.	C PSG1 137
C	ZPHPD	C PSG1 138
C	PHOTOPERIOD OF FIRST DAY IN TIMESTEP, HOURS. DRIVING	C PSG1 139
C	VARIABLE USED BY OTHER SUBMODELS.	C PSG1 140
C	ZRAIN	C PSG1 141
C	TOTAL PRECIPITATION FOR TIMESTEP, MM. DRIVING VARIABLE	C PSG1 142
C	USED BY OTHER SUBMODELS.	C PSG1 143
C	ZRH	C PSG1 144
C	AVERAGE DAILY RELATIVE HUMIDITY, DECIMAL FRACTION FROM	C PSG1 145
C	0=1. DRIVING VARIABLE USED BY OTHER SUBMODELS.	C PSG1 146
C	ZRINT	C PSG1 147
C	PRECIPITATION INTENSITY, MM/HR. DRIVING VARIABLE USED	C PSG1 148
C	BY OTHER SUBMODELS.	C PSG1 149
C	ZRSUM	C PSG1 150
C	SUM FROM OCTOBER 1 OF ZRAIN, MM.	C PSG1 151
C	ZSUN	C PSG1 152
C	AVERAGE DAILY PER CENT POSSIBLE SUNLIGHT FOR TIMESTEP,	C PSG1 153
C	DECIMAL FRACTION FROM 0=1. DRIVING VARIABLE USED	C PSG1 154
C	BY OTHER SUBMODELS.	C PSG1 155
C	ZTMAX	C PSG1 156
C	AVERAGE DAILY MAXIMUM TEMPERATURE FOR TIMESTEP,	C PSG1 157
C	DEGREES CELCIUS, DRIVING VARIABLE USED BY OTHER	C PSG1 158
C	SUBMODELS.	C PSG1 159
C	ZTMIN	C PSG1 160
C	AVERAGE DAILY MINIMUM TEMPERATURE FOR TIMESTEP,	C PSG1 161
C	DEGREES CELCIUS. DRIVING VARIABLE USED BY OTHER	C PSG1 162
C	SUBMODELS.	C PSG1 163
C	ZWIND	C PSG1 164
C	AVERAGE DAILY WIND SPEED FOR TIMESTEP. DRIVING	C PSG1 165
C	VARIABLE USED BY OTHER SUBMODELS.	C PSG1 166
C	CC	C PSG1 167
C		C PSG1 168
C		C PSG1 169
C		C PSG1 170
C		C PSG1 171
C	SHUM=0.	C PSG1 172
C	SMIN=0.	C PSG1 173
C	SPREC=0.	C PSG1 174
C	SSUN=0.	C PSG1 175
C	SWIND=0.	C PSG1 176
C	SHUM=0.	C PSG1 176

```

      JDAY=PHJDAY
      MDAY=JDAY
C
C
C
C
C MAIN LOOP
      DO 3 I=1,PMDY
      MTIME=HDAY/28+I
      IF (MTIME.GT. 13) MTIME=13
C
C
C READ IN TEMP AND PRECIP DATA IF FLAG SET
      IF (.NOT.READ)GOTO100
020 READ(8,25,END=500)TMAX,YMIN,PRECIP
025 FORMAT(7X,2F6.1,3GX,F6.1)
      GOTO30
C
C 500 READ=.FALSE.
C IF WE REACH HERE THEN ADDITIONAL TEMPS AND PRECIP WILL BE
C SUPPLIED BY THE STOCHASTIC WEATHER GENERATOR.
      GOTO100
C
C 030 CONTINUE
C
C NOW CHECK DATA FOR BEING REASONABLE
      IF (TMAX.LT.YMIN)GOTO40
      IF ((TMAX.LT.=30.) .OR. (TMAX.GT.45.))GOTO40
      IF ((YMIN.LT.=30.) .OR. (YMIN.GT.45.))GOTO40
      IF ((PRECIP.LT.50.) .AND. (PRECIP.GE.0.0)) .OR. (PRECIP.GT.5000.)
      * GOTO60
C SOMETIMES ZERO PRECIP IN DSCODES IS LISTED AS A RIDICULOUSLY LARGE NO
C
C SOMETHING WRONG = WRITE OFFENDING CARD AND STOP
040 WRITE(6,30)PHJDAY,TMAX,YMIN,PRECIP
050 FORMAT(' ',15,3F12.5)
      STOP
C
C
C 060 CONTINUE
      SMAX=SMAX+TMAX
      SMIN=SMIN+YMIN
      SPREC=SPREC+PRECIP
C LOAD PRECIP INFORMATION FOR ZRINT CALCULATION
      TRAIN(I)=PRECIP
      GOTO2
C
C
C
C 100 CONTINUE
C TEMPERATURE SECTION
C TMAX IS A QUADRATIC FIT TO THE YEAR'S TEMPERATURES
C RNOR IS 'RANDOM NORMAL' NUMBER GENERATOR ( ,68 OF VALUES BETWEEN
C =1 AND +1)
C XMAX(3) IS THE STANDARD DEVIATION ABOUT THIS FIT.
C THE FIRST TMAX BELOW IS A SINE WAVE FIT TO THE YEAR'S MEAN MAX TEMP
C .017214 = 2 * PI / 365.
C 1.570796 = PI / 2.
      TMAX = XMAX(1) + XMAX(2)*SIN(.017214*MOD((PHJDAY+344.),365.))
      C = 1.570796)
C THE FOLLOWING TMAX IS THE DAY'S MAXIMUM TEMPERATURE.
001 TMAX = TMAX + XMAX(3)*RNOR(1SEED)
      YMIN=XMIN(1)+XMIN(2)*TMAX
      YMIN=YMIN+XMIN(3)*RNOR(1SEED)
      IF (YMIN.GT. YMAX) GO TO 1
      SMAX=SMAX+TMAX
      SMIN=SMIN+YMIN
C
C
C PRECIPITATION SECTION
C IYEST HAS SOMETHING TO DO WITH WHETHER OR NOT WE HAD RAIN YESTERDAY.
      IYEST=IPROB(RAIN(IYEST,MTIME),1)
      IF (IYEST.NE. 2) TRAIN(I)=0.0
      IF (IYEST.NE. 2) GO TO 2
      J=IPROB(PREC(1,MTIME),5)
      SPREC=SPREC+AMT(J)
      TRAIN(I)=AMT(J)
C
C
C WIND SECTION
002 J=IPROB(WIND(1,MTIME),7)
      SWIND=SWIND+(J-1)*8.05
C
C PERCENT POSSIBLE SUNLIGHT SECTION
      J=IPROB(SUN(1,MTIME),10)
      SSUN=SSUN+(J-1)*0.1
C
C RELATIVE HUMIDITY SECTION
      J=IPROB(HUM(1,MTIME),10)
      SHUM=SHUM+(J-1)*0.1
003 MDAY=MDAY+1
C END MAIN LOOP
C

```

```

PSG1 177
PSG1 178
PSG1 179
PSG1 180
PSG1 181
PSG1 182
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PSG1 266

```

```

C
C
C      X#1./PMDT
C      ZTMAX#SMAX*X
C      ZTMIN#SMIN*X
C      ZWIND#SWIND*X
C
C      HERE'S A FACTOR FOR ADJUSTING AVERAGE WIND SPEED IF SITE IS KNOWN TO
C      BE DIFFERENT FROM MEASURING STATION
C      ZWIND = ZWIND * PWFAC
C      ZWIND IS IN KM/HR
C
C      ZSUN#SSUN*X
C      ZRH#SHUM*X
C      ZRAIN#SPREC
C      ZAIRT#(ZTMAX+ZTMIN)/2.
C
C      THIS IS WILKIN'S PHOTOPERIOD CALCULATION
C      DECL#AR SIN( SIN(6.2831853*23.45/360.)*
C      COS(MOD((PMJDAT+193.),365.)*6.2831853/365.))
C      X#( SIN(LAT*6.2831853/360.)*TAN(DECL)+SIN(.875*6.2831853/360.))/
C      COS(LAT*6.2831853/360.)
C      IF (X.GT. 1.) X#1.
C      IF (X.LT. -1.) X#-1.
C      ZPHPD#12. + .1333333*AR SIN(X)*360./6.2831853
C
C
C      DETERMINE RAIN INTENSITY
C      CALL RINT(ZRINT, ZRAIN, ZRH, ZPMPD, ZRAIN, PMDT, PMJDAT, ISEED)
C      ZRINT IS RETURNED IN MM/HR
C
C
C      DETERMINE ZEVAP
C      CALL EVAP
C
C      ZEVAP IS IN MM/TIMESTEP
C
C
C      THIS IS THE WEATHER MODIFICATION
C      ZRAIN#FACTOR*ZRAIN
C
C
C      SUM SOME VARIABLES FROM OCT 1 = = WATER YEAR
C      IF(PMJDAT.LT.270)GOTO900
C      IF(ABS(PMJDAT-270) .GE. PMDT) GO TO 900
C      ZRSUM#0.0
C      ZESUM#0.0
C      900 ZRSUM#ZRSUM+ZRAIN
C      ZESUM#ZESUM+ZEVAP
C
C
C      IF(NNN .LT. 58) GO TO 910
C      WRITE(4,930)
C      NNN#0
C      910 WRITE(4,940)PMJDAT,ZTMAX,ZTMIN,ZRH,ZPMPD,ZRAIN,ZRINT,ZRSUM,ZSUN
C      *,ZEVAP,ZESUM,ZWIND
C      NNN#NNN+1
C      930 FORMAT('1', 'PMJDAT      ZTMAX      ZTMIN      ZRH      ZPMPD      ZRAIN
C      *N      ZRINT      ZRSUM      ZSUN      ZEVAP      ZESUM      ZWIND
C      *D', /)
C      940 FORMAT('1', 4X, I3, 2(4X,F6.2),5X, F4.2, 3(5X, F5.2),
C      * 2(5X, F7.2), 5X, F5.2, 5X, F7.2, 5X, F5.2)
C
C      RETURN
C
C
C      ENTRY ZINIT
C
C      COMMON TIME, TSTART, TEND, DT, DTPR, DTPL,
C      1CAAR(6,8), CADEBT(6,8), CADETH, CAUMST, CDDL(12,4), CHD(10), CHDX(10),
C      2CHOXX(10), CNFIX, CN8DF, CN8DS, CN8OMF, CN8OMS, CN8UP, CVDETH(12,4),
C      3CVLTPR(11,4), CVPHEN(6,8), CVPMS(6), CVRDBT(6,10), CVT8PR(6,8), CVVCOV,
C      4CWINF, CWPSI(10), PMDT, PMDTPL, PMDTPR, PMPGPS, PMJDAT, PMN, PMNCDH, PMNSP,
C      5PHX(10), XAA(4), XAAVWT(4,8), XAAWT(4), XAFT8(4), XAFTWT(4), XANUMB(4,8),
C      6XABA(4), XABAWT(4), XAYNG(4), XAYWT(4), XH8OLT(10), XNMN, XN8OMF, XN8OMS,
C      7XVFG(6,5), XVLITR(12,4), XVPLNT(6,8), XVTOTL, XHTHTA(10), XH8TND, ZAIRT,
C      8ZEVAP, ZESUM, ZPMPD, ZRAIN, ZRSUM, ZRH, ZRINT, ZRSUM, ZSUN, ZTMAX, ZTMIN,
C      9ZWIND
C
C      LOGICAL READ
C
C      REAL LAT
C
C      INTEGER PMJDAT
C
C      COMMON /SUNNY/ IYEST, ISEED
C
C      DIMENSION XMAX(3), XMIN(3), RAIN(2,13), PREC(6,13), AMT(6),
C      * WIND(7,13), SUN(10,13), HUM(10,13)
C      DIMENSION TRAIN(20)
C      DIMENSION RCHECK(20)

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C	WRITE(4,930)	P8G1 358
	NNN#0	P8G1 359
C	READ(5,710) RCHECK	P8G1 360
	WRITE(6,720) RCHECK	P8G1 361
	READ(5,710) RCHECK	P8G1 362
	WRITE(6,720) RCHECK	P8G1 363
C	READ(5,/) ISEED	P8G1 364
	WRITE(6,/) ISEED	P8G1 365
C	READ(5,/) LAT	P8G1 366
	WRITE(6,/) LAT	P8G1 367
C	READ(5,/) XMAX,XMIN	P8G1 368
	WRITE(6,/) XMAX	P8G1 369
	WRITE(6,/) XMIN	P8G1 370
C	READ(5,/) RAIN	P8G1 371
	WRITE(6,/) RAIN	P8G1 372
C	READ(5,/) PREC	P8G1 373
	WRITE(6,/) PREC	P8G1 374
C	READ(5,/) AMT	P8G1 375
	WRITE(6,/) AMT	P8G1 376
C	READ(5,/) WIND	P8G1 377
	WRITE(6,/) WIND	P8G1 378
C	READ(5,710) RCHECK	P8G1 379
	WRITE(6,720) RCHECK	P8G1 380
C	READ(5,/) SUN	P8G1 381
	WRITE(6,/) SUN	P8G1 382
C	READ(5,/) HUM	P8G1 383
	WRITE(6,/) HUM	P8G1 384
C	READ(5,/) PWFAC	P8G1 385
	WRITE(6,/) PWFAC	P8G1 386
C	READ(5,/) READ	P8G1 387
	WRITE(6,/) READ	P8G1 388
C	READ(5,/) FACTOR	P8G1 389
	WRITE(6,/) FACTOR	P8G1 390
C	READ(5,710) RCHECK	P8G1 391
	WRITE(6,720) RCHECK	P8G1 392
C	710 FORMAT(20A4)	P8G1 393
	720 FORMAT(' ', 20A4)	P8G1 394
C	IYEST#1	P8G1 395
C	CALL INITIALIZATION SECTION OF SUBROUTINES EVAP AND RINT	P8G1 396
	CALL EVINIT	P8G1 397
	CALL RIINIT	P8G1 398
C	RETURN	P8G1 399
C	END	P8G1 400
		P8G1 401
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		P8G1 422
		P8G1 423

SUBROUTINE PSWG, VERSION II

	SUBROUTINE PSWG	P8G2 001
C		P8G2 002
C		P8G2 003
C	THIS VERSION OF PSWG USED IN 5-YEAR RUNS.	P8G2 004
C	WEATHER VARIABLES ARE READ FROM WEATHER DATA FILE.	P8G2 005
C		P8G2 006
C		P8G2 007
C		P8G2 008
C	CC	P8G2 009
C	VARIABLE DICTIONARY	P8G2 010
C		P8G2 011
C		P8G2 012
C		P8G2 013
C	DECL ANGLE CALCULATED IN PHOTOPERIOD CALCULATION.	P8G2 014
C		P8G2 015

C	EVAP	SUBROUTINE WHICH CALCULATES POTENTIAL EVAPORATION.	C	P8G2	016
C			C	P8G2	017
C	FACTOR	PRECIPITATION FACTOR BY WHICH EACH EVENT IS	C	P8G2	018
C		MULTIPLIED.	C	P8G2	019
C			C	P8G2	020
C	ISEED	SEED FROM WHICH RNOR GENERATES NEXT RANDOM NUMBER.	C	P8G2	021
C			C	P8G2	022
C	JDAY	JULIAN DATE AT BEGINNING OF TIMESTEP.	C	P8G2	023
C			C	P8G2	024
C	LAT	LATITUDE OF SITE.	C	P8G2	025
C			C	P8G2	026
C	NNN	COUNTER TO DETERMINE WHEN TO START PRINTING ON	C	P8G2	027
C		NEXT PAGE.	C	P8G2	028
C			C	P8G2	029
C	PMDY	TIMESTEP LENGTH, DAYS.	C	P8G2	030
C			C	P8G2	031
C	PRECIP	PRECIPITATION FOR CURRENT DAY, MM.	C	P8G2	032
C			C	P8G2	033
C	PHJDAY	JULIAN DATE AT BEGINNING OF TIMESTEP. VALUE	C	P8G2	034
C		DETERMINED IN MAIN PROGRAM.	C	P8G2	035
C			C	P8G2	036
C	PHFAC	FACTOR FOR ADJUSTING AVERAGE WIND SPEED IF SITE IS	C	P8G2	037
C		KNOWN TO HAVE DIFFERENT AVERAGE WIND SPEED.	C	P8G2	038
C			C	P8G2	039
C	RCHECK(20)	VARIABLE USED TO READ AND WRITE COMMENTS IN WITH	C	P8G2	040
C		INITIAL DATA.	C	P8G2	041
C			C	P8G2	042
C	RINT	SUBROUTINE WHICH CALCULATES PRECIPITATION INTENSITY.	C	P8G2	043
C			C	P8G2	044
C	SHUM	RUNNING SUM OF UP TO PMDY DAYS' RELATIVE HUMIDITY.	C	P8G2	045
C			C	P8G2	046
C	SHAX	RUNNING SUM OF UP TO PMDY DAYS' MAXIMUM TEMPERATURES,	C	P8G2	047
C		DEGREES CELCIUS.	C	P8G2	048
C			C	P8G2	049
C	SHIN	RUNNING SUM OF UP TO PMDY DAYS' MINIMUM TEMPERATURES,	C	P8G2	050
C		DEGREES CELCIUS.	C	P8G2	051
C			C	P8G2	052
C	SPREC	RUNNING SUM OF UP TO PMDY DAYS' PRECIPITATIONEVENTS,MM.	C	P8G2	053
C			C	P8G2	054
C	SSUN	RUNNING SUM OF UP TO PMDY DAYS' PER CENT POSSIBLE	C	P8G2	055
C		SUNLIGHT	C	P8G2	056
C			C	P8G2	057
C	SWIND	RUNNING SUM OF UP TO PMDY DAYS' AVERAGE WIND SPEEDS.	C	P8G2	058
C		KM/HR.	C	P8G2	059
C			C	P8G2	060
C	TMAX	MAXIMUM TEMPERATURE FOR CURRENT DAY, DEGREES CELCIUS.	C	P8G2	061
C			C	P8G2	062
C	TMIN	MINIMUM TEMPERATURE FOR CURRENT DAY, DEGREES CELCIUS.	C	P8G2	063
C			C	P8G2	064
C	TRAIN(20)	ARRAY WHICH HOLDS DAILY PRECIP VALUES FOR TIMESTEP	C	P8G2	065
C		(USED BY RAIN INTENSITY ROUTINE).	C	P8G2	066
C			C	P8G2	067
C	ZAIRT	AVERAGE DAILY AIR TEMPERATURE FOR TIMESTEP, DEGREES	C	P8G2	068
C		CELCIUS. DRIVING VARIABLE USED BY OTHER SUBMODELS.	C	P8G2	069
C			C	P8G2	070
C	ZESUM	SUM FROM OCTOBER 1 OF ZEVAP, MM.	C	P8G2	071
C			C	P8G2	072
C	ZEVAP	POTENTIAL EVAPORATION, MM/TIMESTEP.	C	P8G2	073
C		DRIVING VARIABLE USED BY OTHER SUBMODELS.	C	P8G2	074
C			C	P8G2	075
C	ZPHPD	PHOTOPERIOD OF FIRST DAY IN TIMESTEP, HOURS. DRIVING	C	P8G2	076
C		VARIABLE USED BY OTHER SUBMODELS.	C	P8G2	077
C			C	P8G2	078
C	ZRAIN	TOTAL PRECIPITATION FOR TIMESTEP, MM. DRIVING VARIABLE	C	P8G2	079
C		USED BY OTHER SUBMODELS.	C	P8G2	080
C			C	P8G2	081
C	ZRH	AVERAGE DAILY RELATIVE HUMIDITY, DECIMAL FRACTION FROM	C	P8G2	082
C		0-1. DRIVING VARIABLE USED BY OTHER SUBMODELS.	C	P8G2	083
C			C	P8G2	084
C	ZRINT	PRECIPITATION INTENSITY, MM/HR. DRIVING VARIABLE USED	C	P8G2	085
C		BY OTHER SUBMODELS.	C	P8G2	086
C			C	P8G2	087
C	ZRSUM	SUM FROM OCTOBER 1 OF ZRAIN, MM.	C	P8G2	088
C			C	P8G2	089
C	ZSUN	AVERAGE DAILY PER CENT POSSIBLE SUNLIGHT FOR TIMESTEP,	C	P8G2	090
C		DECIMAL FRACTION FROM 0-1. DRIVING VARIABLE USED	C	P8G2	091
C		BY OTHER SUBMODELS.	C	P8G2	092
C			C	P8G2	093
C	ZTMAX	AVERAGE DAILY MAXIMUM TEMPERATURE FOR TIMESTEP,	C	P8G2	094
C		DEGREES CELCIUS. DRIVING VARIABLE USED BY OTHER	C	P8G2	095
C		SUBMODELS.	C	P8G2	096
C			C	P8G2	097
C	ZTMIN	AVERAGE DAILY MINIMUM TEMPERATURE FOR TIMESTEP,	C	P8G2	098
C		DEGREES CELCIUS. DRIVING VARIABLE USED BY OTHER	C	P8G2	099
C		SUBMODELS.	C	P8G2	100
C			C	P8G2	101
C	ZWIND	AVERAGE DAILY WIND SPEED FOR TIMESTEP. DRIVING	C	P8G2	102
C		VARIABLE USED BY OTHER SUBMODELS.	C	P8G2	103
C			C	P8G2	104
C			C	P8G2	105

```

C
C
C   SMAX=0.
C   SMIN=0.
C   SPREC=0.
C   SWIND=0.
C   SSUN=0.
C   SHUM=0.
C   JDAY=PMJDAY
C   MDAY=JDAY
C
C
C
C   MAIN LOOP
C   DO 3 I=1,PHDT
C   READ TODAY'S DATA
C   READ(6,30,END=500)TMAX,TMIN,PRECIP,WIND,PPS,RH
C   30 FORMAT(10X,2F5.0,F5.2,F5.1,2F5.2)
C
C   GOTO501
C   500 CONTINUE
C   STOP
C   501 CONTINUE
C
C   CONVERT TO METRIC UNITS
C
C   TEMPS FROM F TO C
C   TMAX=(TMAX-32.)*0.555555
C   TMIN=(TMIN-32.)*0.555555
C
C   PRECIP FROM INCHES TO MM
C   PRECIP=PRECIP*25.4
C
C   AVERAGE WIND SPEED FROM MILES / HR TO KM/HR
C   WIND=WIND*1.60934
C
C
C   CHECK IF VALUES ARE SENSIBLE
C   IF(TMIN.GT.TMAX)GOTO40
C   IF((TMAX.GT.42.).OR.(TMIN.LT.=43.))GOTO40
C   IF((PRECIP.GT.40.).OR.(PRECIP.LT.0.0))GOTO40
C   IF((WIND.GT.60.).OR.(WIND.LT.0.0))GOTO40
C   IF((PPS.GT.1.0).OR.(PPS.LT.0.0))GOTO40
C   IF((RH.LE.1.00).AND.(RH.GE.0.0))GOTO60
C
C   IF WE REACH HERE SOMETHING IS WRONG. WRITE OFFENDING CARD AND STOP.
C   40 WRITE(6,50)TMAX,TMIN,PRECIP,WIND,PPS,RH,PMJDAY
C   50 FORMAT(' ',7F10.3)
C   STOP
C
C
C   60 CONTINUE
C   SMAX=SMAX+TMAX
C   SMIN=SMIN+TMIN
C   SPREC=SPREC+PRECIP
C   SWIND=SWIND+WIND
C   SSUN=SSUN+PPS
C   SHUM=SHUM+RH
C
C   LOAD TRAIN FOR USE IN SUBROUTINE RINT
C   TRAIN(I)=PRECIP
C
C
C   3 CONTINUE
C   END MAIN LOOP
C
C
C
C
C   X=1./PHDT
C
C
C   ZTMAX=SMAX*X
C   ZTMIN=SMIN*X
C   ZWIND=SWIND*X
C
C   HERE IS A FACTOR FOR ADJUSTING AVERAGE WIND SPEED IF SITE IS KNOWN TO
C   BE DIFFERENT FROM MEASURING STATION
C   ZWIND = ZWIND * PWFAC
C   ZSUN=SSUN*X
C   ZRH=SHUM*X
C   ZRAIN=SPREC
C   ZAIRY=(ZTMAX+ZTMIN)/2.
C
C
C   THIS IS WILKIN'S PHOTOPERIOD CALCULATION
C   DECL=ARCSIN(STN(6.2831853*23.45/360.))*
C   COS(MOD((PMJDAY+193.),365.)*6.2831853/365.))
C   X=(SIN(LAT*6.2831853/360.)*TAN(DECL)+8IN(.875*6.2831853/360.))/
C   COS(LAT*6.2831853/360.)
C   IF (X .GT. 1.) X=1.
C   IF (X .LT. -1.) X=-1.
C   ZPHPP=12. + .1333333*ARCSIN(X)*360./6.2831853
C
C

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P8G2 195

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C DETERMINE ZRINT                                P8G2 196
  CALL RINT(ZRINT, TRAIN, ZRAIN, PMDT, PMJDAT, ISEED) P8G2 197
C                                                P8G2 198
C DETERMINE ZEVAP                                P8G2 199
  CALL EVAP                                       P8G2 200
C                                                P8G2 201
C                                                P8G2 202
C                                                P8G2 203
C                                                P8G2 204
C HERE IS THE RAINFALL MODIFICATION              P8G2 205
  ZRAIN=FACTOR*SPREC                             P8G2 206
C                                                P8G2 207
C                                                P8G2 208
C                                                P8G2 209
C SUMMING SECTION                               P8G2 210
C DO FOLLOWING SUMS FROM OCTOBER 1 - - WATER YEAR P8G2 211
  IF(PMJDAT.LT.270)GOTO900                       P8G2 212
  IF(ABS(PMJDAT-270) .GE. PMDT) GO TO 900         P8G2 213
  ZRSUM=0.0                                       P8G2 214
  ZESUM=0.0                                       P8G2 215
900 ZRSUM=ZRSUM+ZRAIN                             P8G2 216
  ZESUM=ZESUM+ZEVAP                              P8G2 217
  IF(NNN .LT. 56) GO TO 910                     P8G2 218
  WRITE(4,930)                                    P8G2 219
  NNN=0                                           P8G2 220
910 WRITE(4,940)PMJDAT,ZTMAX,ZTMIN,ZRH,ZPHPD,ZRAIN,ZRINT,ZRSUM,ZSUN P8G2 221
  *,ZEVAP,ZESUM,ZWIND                            P8G2 222
  NNN=NNN+1                                       P8G2 223
930 FORMAT('11', 'PMJDAT  ZTMAX  ZTMIN  ZRH  ZPHPD  ZRAIN P8G2 224
  *N  ZRINT  ZRSUM  ZSUN  ZEVAP  ZESUM  ZWINP8G2 225
  *D', /)                                         P8G2 226
940 FORMAT('1', 4X, I3, 2(4X,F6.2),5X, F4.2, 3(5X, F5.2), P8G2 227
  * 2(5X, F7.2), 5X, F5.2, 5X, F7.2, 5X, F5.2) P8G2 228
C                                                P8G2 229
C                                                P8G2 230
C RETURN                                         P8G2 231
C                                                P8G2 232
C ENTRY ZINIT                                    P8G2 233
C                                                P8G2 234
C                                                P8G2 235
  COMMON TIME,TSTART,TEND,DT,DTPR,DTPL,         P8G2 236
  1CAAR(6,8),CADEST(6,8),CAETH,CAUWST,CDDL(12,4),CHD(10),CHDX(10), P8G2 237
  2CHDX(10),CNFIX,CNSDF,CNSDS,CNBOMF,CNSOMS,CNSUP,CVDETH(12,4), P8G2 238
  3CVLTPR(11,4),CVPHEN(6,8),CVPHS(6),CVRDST(6,10),CVTSPR(6,8),CVVCOV, P8G2 239
  4CWINF,CWPST(10),PMDT,PHDTPL,PHDTPR,PMFGPS,PMJDAT,PMN,PMNCOM,PMNSP, P8G2 240
  5PWK(10),YAA(4),XAAVNT(2,8),XAAWT(4),XAFTS(4),XAFWT(4),XANUMS(4,8), P8G2 241
  6XABA(4),XABANT(4),XAYNG(4),XAYNT(4),XHSOLT(10),XNMN,XNSOMF,XNSOMS, P8G2 242
  7XVFG(6,5),XVLITR(12,4),XVPLNT(6,8),XVTOTL,XWHTA(10),XWSTNO,ZATRY, P8G2 243
  8ZEVAP,ZESUM,ZPHPD,ZRAIN,ZRSUM,ZRH,ZRINT,ZRISUM,ZSUN,ZTMAX,ZTMIN, P8G2 244
  9ZWIND                                           P8G2 245
C                                                P8G2 246
C REAL LAT                                       P8G2 247
C                                                P8G2 248
C COMMON /BUNNY/ IVEBT,ISEED                   P8G2 249
C                                                P8G2 250
C DIMENSION TRAIN(20)                          P8G2 251
  DIMENSION RCHECK(20)                          P8G2 252
C                                                P8G2 253
C                                                P8G2 254
C WRITE(4,930)                                  P8G2 255
  NNN=0                                           P8G2 256
C                                                P8G2 257
C                                                P8G2 258
C READ(5,710) RCHECK                            P8G2 259
  WRITE(6,720) RCHECK                           P8G2 260
C                                                P8G2 261
C READ(5,/) ISEED                               P8G2 262
  WRITE(6,/) ISEED                              P8G2 263
C                                                P8G2 264
C READ(5,/) LAT                                 P8G2 265
  WRITE(6,/) LAT                                P8G2 266
C                                                P8G2 267
C READ(5,/) PWFAC                              P8G2 268
  WRITE(6,/) PWFAC                             P8G2 269
C                                                P8G2 270
C READ(5,/)FACTOR                              P8G2 271
  WRITE(6,/) FACTOR                            P8G2 272
C                                                P8G2 273
C READ(5,710) RCHECK                            P8G2 274
  WRITE(6,720) RCHECK                         P8G2 275
C                                                P8G2 276
710 FORMAT(20A4)                                P8G2 277
720 FORMAT('1', 20A4)                          P8G2 278
C                                                P8G2 279
C CALL INITIALIZATION SECTIONS OF EVAP AND RINT P8G2 280
  CALL EVINIT                                    P8G2 281
  CALL RINIT                                     P8G2 282
C                                                P8G2 283
C                                                P8G2 284
C RETURN                                         P8G2 285
  END                                           P8G2 286

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SUBROUTINE EVAP

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SUBROUTINE EVAP
C EVAP 001
C EVAP 002
C EVAP 003
C EVAP CALCULATES POTENTIAL EVAPORATION FROM SOIL SURFACE DURING EVAP 004
C TIMESTEP, IN MM, EVAP 005
C THIS VERSION OF EVAP USES BLANEY CRIDDLE METHOD, EVAP 006
C IT CLOSELY FOLLOWS CALCULATION DONE BY GRIFFIN AND HANKS IN THEIR EVAP 007
C D.B. RESEARCH MEMORANDUM, EVAP 008
C EVAP 009
C EVAP 010
CCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCC EVAP 011
C EVAP 012
C VARIABLE DICTIONARY FOR EVAP
C EVAP 013
C EVAP 014
C EVAP 015
C A, B, C, Z TEMPORARY VARIABLES USED IN DAYLIGHT CALCULATION, EVAP 016
C EVAP 017
C ARRAY(5,2) PAIRS OF DATA POINTS FOR INTERPOLATION BY F1 IN EVAP 018
C CALCULATING BLANEY CRIDDLE FACTOR, EVAP 019
C EVAP 020
C CVVCOV FRACTION OF SOIL SURFACE COVERED BY ANNUAL AND EVAP 021
C PERENNIAL VEGETATION, EVAP 022
C EVAP 023
C DALITE(365) ARRAY INDEXED BY JULIAN DATE , GIVES NUMBER OF EVAP 024
C MINUTES OF DAYLIGHT ON THAT JULIAN DATE, EVAP 025
C EVAP 026
C DLTFR(365) ARRAY INDEXED BY JULIAN DATE AND IS FRACTION OF YEAR'S EVAP 027
C DAYLIGHT WHICH FALLS ON THAT JULIAN DATE, EVAP 028
C EVAP 029
C F1 FUNCTION WHICH LINEARLY INTERPOLATES BETWEEN PAIRS OF EVAP 030
C DATA POINTS TO DETERMINE DEPENDENT VARIABLE, EVAP 031
C EVAP 032
C LAT LATITUDE OF SITE, DEGREES, EVAP 033
C EVAP 034
C MM FLAG USED ONLY IN CASE PHJDAT IS EVER ZERO, EVAP 035
C EVAP 036
C NPTS NUMBER OF PAIRS OF DATA POINTS ACTUALLY USED IN ARRAY, EVAP 037
C EVAP 038
C PMDT LENGTH OF TIMESTEP, DAYS, EVAP 039
C EVAP 040
C PHJDAT JULIAN DATE OF FIRST DAY OF TIMESTEP, EVAP 041
C EVAP 042
C RCHECK(20) ARRAY USED TO READ AND WRITE COMMENTS IN WITH EVAP 043
C INITIAL DATA, EVAP 044
C EVAP 045
C TOTMIN MINUTES OF DAYLIGHT IN YEAR, EVAP 046
C EVAP 047
C ZAIRT AVERAGE TIMESTEP AIR TEMPERATURE, DEGREES CELCIUS, EVAP 048
C EVAP 049
C ZEVAP POTENTIAL EVAPORATION FOR TIMESTEP, MM, DRIVING EVAP 050
C VARIABLE USED BY OTHER SUBMODELS, EVAP 051
C EVAP 052
CCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCC EVAP 053
C EVAP 054
C ZEVAP=F1(ZAIRT, ARRAY1, NPTS)
C EVAP 055
C IF(PHJDAT .GT. 0) GO TO 670
C EVAP 056
C PHJDAT=1
C EVAP 057
C MM=1
C EVAP 058
C 670 CONTINUE
C EVAP 059
C IF ZAIRT IN NEXT EQUATION LOOKS LIKE IT'S BEING CHANGED TO DEGREES
C EVAP 060
C F, THAT'S BECAUSE IT IS,
C EVAP 061
C ZEVAP=ZEVAP*(1.8*ZAIRT + 32.)*DLTFR(PHJDAT)*(1.-CVVCOV)*25.4*PMDT
C EVAP 062
C CVVCOV=0.0
C EVAP 063
C IF(MM .EQ. 0) GO TO 680
C EVAP 064
C PHJDAT=0
C EVAP 065
C MM=0
C EVAP 066
C 680 CONTINUE
C EVAP 067
C EVAP 068
C EVAP 069
C EVAP 070
C RETURN
C EVAP 071
C EVAP 072
C ENTRY EVINIT
C EVAP 073
C EVAP 074
C COMMON TIME, TSTART, TEND, DT, DTPR, DTPL,
C EVAP 075
C 1CAAR(6,8), CADEBT(6,8), CADETH, CAUWST, COOL(12,4), CHD(10), CHDX(10),
C EVAP 076
C 2CHDXX(10), CNFIX, CNBDF, CNBDS, CNSOMF, CNSOM9, CNSUP, CVDETH(12,4),
C EVAP 077
C 3CVLTFR(11,4), CVPHEN(6,8), CVPHS(6), CVRDST(6,10), CVTSPR(6,8), CVVCOV,
C EVAP 078
C 4CWINP, CWPST(10), PMDT, PMDTPL, PMDTPR, PMFGPS, PHJDAT, PHN, PHNCOM, PMNSP,
C EVAP 079
C 5PHK(10), XAA(4), XAAVHT(4,8), XAAMT(4), XAFYS(4), XAFNY(4), XANUMB(4,8),
C EVAP 080
C 6XASA(4), XASAMT(4), XAYNG(4), XAYNT(4), XHSOLT(10), XNMN, XNSOMF, XNSOMS,
C EVAP 081
C 7XVFG(6,5), XLITR(12,4), XVPLNT(6,8), XVTOTL, XWHTA(10), XWBTND, ZAIRT,
C EVAP 082
C 8ZEVAP, ZESUM, ZPHPD, ZRAIN, ZR8UM, ZRH, ZRINT, ZRISUM, ZSUN, ZTMAX, ZTMIN,
C EVAP 083
C 9ZWIND
C EVAP 084
C
C DIMENSION ARRAY1(5,2), DALITE(365), DLTFR(365)
C EVAP 085
C DIMENSION RCHECK(20)
C EVAP 086
C EVAP 087
C EVAP 088

```

```

      INTEGER PMJDAT
C
      REAL LAT
C
C
      READ(5,710) RCHECK
      WRITE(6,720) RCHECK
C
      READ(5,/) LAT
      WRITE(6,/) LAT
      READ(5,/) ARRAY1
      WRITE(6,/) ARRAY1
      READ(5,/) NPTS
      WRITE(6,/) NPTS
C
      710 FORMAT(20A4)
      720 FORMAT(' ', 20A4)
C
C FROM GRIFFIN, HANKS BIOME REPORT
      DO 650 I=1, 365
      A=730,=.274*LAT+.00793*(LAT**2)
      B=34,2=.78*LAT+.1*(LAT**2)
      C=1
      Z=2,*.3,1416*((C+285,)/365.)
C LENGTH OF DAY IN MINUTES
      DALITE(I)=A+B*SIN(Z)
      TOTMIN=DALITE(I)+TOTMIN
      650 CONTINUE
C
C FRACTION OF DAYLIGHT FOR EACH DAY
      DO 660 I=1, 365
      DLTR(I)=DALITE(I)/TOTMIN
      660 CONTINUE
C
      WRITE(6,681) DALITE
      WRITE(6,681) TOTMIN
      681 FORMAT(' ', 10F12,5)
C
      READ(5,710) RCHECK
      WRITE(6,720) RCHECK
C
      RETURN
      END
      EVAP 089
      EVAP 090
      EVAP 091
      EVAP 092
      EVAP 093
      EVAP 094
      EVAP 095
      EVAP 096
      EVAP 097
      EVAP 098
      EVAP 099
      EVAP 100
      EVAP 101
      EVAP 102
      EVAP 103
      EVAP 104
      EVAP 105
      EVAP 106
      EVAP 107
      EVAP 108
      EVAP 109
      EVAP 110
      EVAP 111
      EVAP 112
      EVAP 113
      EVAP 114
      EVAP 115
      EVAP 116
      EVAP 117
      EVAP 118
      EVAP 119
      EVAP 120
      EVAP 121
      EVAP 122
      EVAP 123
      EVAP 124
      EVAP 125
      EVAP 126
      EVAP 127
      EVAP 128
      EVAP 129
      EVAP 130
      EVAP 131

```

SUBROUTINE RINT

```

      SUBROUTINE RINT(ZRINT, TRAIN, ZRAIN, PMDT, PMJDAT, ISEED)
C
C RAIN INTENSITY SUBROUTINE 4=76 PAUL LOMMEN
C RINT CALCULATES PRECIPITATION INTENSITY IN MM/HR. USED BY WATER.
C
CCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCC
C
C VARIABLE DICTIONARY FOR RINT
C
C A, B PARAMETERS IN HIGH INTENSITY CALCULATION.
C
C ARRAY(10) HOLDS UP TO 5 PAIRS OF DATA POINTS FOR INTERPOLATION
C BY FUNCTION F1 IN DETERMINING NORMAL INTENSITY.
C
C BA INTERMEDIATE VARIABLE IN HIGH INTENSITY CALCULATION.
C
C B1 INTERMEDIATE VARIABLE IN HIGH INTENSITY CALCULATION.
C
C F1 FUNCTION WHICH LINEARLY INTERPOLATES BETWEEN NPTS PAIRS
C OF DATA POINTS.
C
C ISEED INTEGER FROM WHICH RANDOM NUMBERS ARE GENERATED.
C
C N INDEX OF TIME OF YEAR. IN MIDDLE OF YEAR N=1.
C OTHER TIMES OF YEAR N=2.
C
C NPTS NUMBER OF PAIRS OF DATA POINTS IN NORMAL INTENSITY
C CALCULATION.
C
C PDAY1,PDAY2 BEGINNING AND ENDING JULIAN DATES OF CENTRAL
C SEGMENT OF YEAR.
C
C PMDT LENGTH OF TIMESTEP, DAYS.
C
C PMJDAT JULIAN DATE OF BEGINNING OF TIMESTEP.
C
C PTH PRECIP AMOUNT ABOVE WHICH HIGH INTENSITY CAN OCCUR
C
      RINT 001
      RINT 002
      RINT 003
      RINT 004
      RINT 005
      RINT 006
      RINT 007
      RINT 008
      RINT 009
      RINT 010
      RINT 011
      RINT 012
      RINT 013
      RINT 014
      RINT 015
      RINT 016
      RINT 017
      RINT 018
      RINT 019
      RINT 020
      RINT 021
      RINT 022
      RINT 023
      RINT 024
      RINT 025
      RINT 026
      RINT 027
      RINT 028
      RINT 029
      RINT 030
      RINT 031
      RINT 032
      RINT 033
      RINT 034
      RINT 035
      RINT 036
      RINT 037

```

```

C PA(2) COMPARED WITH R TO SEE IF HIGH INTENSITY ACTUALLY C RINT 038
C OCCURRED. C RINT 039
C R TEMPORARY VARIABLE. SET EQUAL TO RANDOM NUMBER C RINT 040
C RINT 041
C RCHECK(20) ARRAY USED FOR READING AND WRITING COMMENTS IN WITH C RINT 042
C INITIALIZATION DATA. C RINT 043
C TR TEMPORARY VARIABLE. C RINT 044
C RINT 045
C TRAIN(20) DAILY PRECIPITATION AMOUNTS, MM. ARRAY INDEXED BY C RINT 046
C DAY OF TIMESTEP C RINT 047
C X TODAY'S RAIN C RINT 048
C Y TODAY'S PRECIPITATION INTENSITY, MM/HR. C RINT 049
C YSUM SUM OVER TIMESTEP OF DAILY PRECIP INTENSITY TIMES C RINT 050
C PRECIP AMOUNT. C RINT 051
C ZRAIN TIMESTEP PRECIPITATION, MM. DRIVING VARIABLE USED BY C RINT 052
C WATER SUBMODEL. C RINT 053
C ZRINT RAIN INTENSITY, MM/HR. DRIVING VARIABLE USED BY C RINT 054
C WATER SUBMODEL. C RINT 055
C CCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCC C RINT 056
C YSUM=0.0 C RINT 057
C RINT 058
C DO 550 I=1,PHDT C RINT 059
C RINT 060
C X IS TODAY'S PRECIP C RINT 061
C Y IS TODAY'S PRECIP INTENSITY C RINT 062
C X=TRAIN(I) C RINT 063
C DO WE HAVE ANY RAIN TODAY? C RINT 064
C IF(X .GT. 1.E+5) GO TO 100 C RINT 065
C GO TO 550 C RINT 066
C IF X LARGE ENOUGH ALLOW FOR HIGH INTENSITIES C RINT 067
C 100 IF(X .LT. PTH) GO TO 500 C RINT 068
C WILL NEED A RANDOM NUMBER TO DETERMINE IF WE HAVE HIGH INTENSITY C RINT 069
C R=RANDOM(1SEED) C RINT 070
C DIVIDE YEAR INTO TWO SEGMENTS, ALLOWING DIFFERENT HIGH INTENSITY C RINT 071
C PROBABILITIES IN EACH C RINT 072
C IN CENTRAL SEGMENT N=1 C RINT 073
C N=2 C RINT 074
C IF((PMJDAT .GE. PDAY1) .AND. (PMJDAT .LE. PDAY2)) N=1 C RINT 075
C N NOW TELLS US WHICH SEGMENT OF YEAR WE'RE IN C RINT 076
C IF(R .GT. PA(N)) GO TO 500 C RINT 077
C GENERATE NEW (SEMI) RANDOM NUMBER C RINT 078
C R=R/PA(N) C RINT 079
C 150 TR=BA*(1.=R) C RINT 080
C IF(TR .LE. 1.E-10) TR=1.E-10 C RINT 081
C Y=B1*ALOG(TR) C RINT 082
C LIMIT Y TO AN INCH AN HOUR C RINT 083
C IF(Y .GT. 25.) Y=25. C RINT 084
C GO TO 520 C RINT 085
C NORMAL INTENSITY CALCULATION C RINT 086
C 500 Y=F1(PMJDAT, ARRAY, NPT8) C RINT 087
C C RINT 088
C 520 YSUM=X*Y + YSUM C RINT 089
C 550 CONTINUE C RINT 090
C IF(ZRAIN .LE. 1.E-6) GO TO 570 C RINT 091
C ZRINT = YSUM/ZRAIN C RINT 092
C GO TO 580 C RINT 093
C 570 ZRAIN=0.0 C RINT 094
C ZRINT=0.0 C RINT 095
C 580 IF((ZRINT*24.*PHDT) .LT. ZRAIN) ZRINT=ZRAIN/(24.*PHDT) C RINT 096
C C RINT 097
C RETURN C RINT 098
C RINT 099
C RINT 100
C RINT 101
C RINT 102
C RINT 103
C RINT 104
C RINT 105
C RINT 106
C RINT 107
C RINT 108
C RINT 109
C RINT 110
C RINT 111
C RINT 112
C RINT 113
C RINT 114
C RINT 115
C RINT 116
C RINT 117
C RINT 118
C RINT 119
C RINT 120
C RINT 121
C RINT 122
C RINT 123
C RINT 124
C RINT 125
C RINT 126
C RINT 127

```

C		RINT 128
	ENTRY RINIT	RINT 129
C		RINT 130
	DIMENSION PA(2), ARRAY(10), RCHECK(20)	RINT 131
	DIMENSION TRAIN(20)	RINT 132
C		RINT 133
	INTEGER PMDT	RINT 134
C		RINT 135
	READ(5,600) RCHECK	RINT 136
	WRITE(6,610) RCHECK	RINT 137
	WRITE(6,630)	RINT 138
C		RINT 139
	READ(5,600) RCHECK	RINT 140
	WRITE(6,610) RCHECK	RINT 141
	READ(5,/) PTH	RINT 142
	WRITE(6,620) PTH	RINT 143
	WRITE(6,630)	RINT 144
C		RINT 145
	READ(5,600) RCHECK	RINT 146
	WRITE(6,610) RCHECK	RINT 147
	READ(5,/) PDAY1, PDAY2	RINT 148
	WRITE(6,620) PDAY1, PDAY2	RINT 149
	WRITE(6,630)	RINT 150
C		RINT 151
	READ(5,600) RCHECK	RINT 152
	WRITE(6,610) RCHECK	RINT 153
	READ(5,/) ARRAY	RINT 154
	WRITE(6,620) ARRAY	RINT 155
	WRITE(6,630)	RINT 156
C		RINT 157
	READ(5,600) RCHECK	RINT 158
	WRITE(6,610) RCHECK	RINT 159
	READ(5,/) NPT8	RINT 160
	WRITE(6,620) NPT8	RINT 161
	WRITE(6,630)	RINT 162
C		RINT 163
	READ(5,600) RCHECK	RINT 164
	WRITE(6,610) RCHECK	RINT 165
	READ(5,/) PA	RINT 166
	WRITE(6,620) PA	RINT 167
	WRITE(6,630)	RINT 168
C		RINT 169
	READ(5,600) RCHECK	RINT 170
	WRITE(6,610) RCHECK	RINT 171
	READ(5,/) A, B	RINT 172
	WRITE(6,620) A, B	RINT 173
	WRITE(6,630)	RINT 174
C		RINT 175
	BA=0.01*B/A	RINT 176
	B1=25.4/B	RINT 177
	WRITE(6,*) BA, B1	RINT 178
C		RINT 179
	READ(5,600) RCHECK	RINT 180
	WRITE(6,610) RCHECK	RINT 181
	WRITE(6,630)	RINT 182
	WRITE(6,630)	RINT 183
	WRITE(6,630)	RINT 184
C		RINT 185
C		RINT 186
	600 FORMAT(20A4)	RINT 187
	610 FORMAT(' ', 20A4)	RINT 188
	620 FORMAT(10F12.5)	RINT 189
	630 FORMAT(' ')	RINT 190
C		RINT 191
C		RINT 192
	RETURN	RINT 193
	END	RINT 194

FUNCTION IPROB

	FUNCTION IPROB(A,N)	IPROB 01
	DIMENSION A(1)	IPROB 02
	COMMON /SUNNY/ IYEST,ISEED	IPROB 03
	X=RANDOM(ISEED)	IPROB 04
	DO 1 I=1,N	IPROB 05
	IF (A(I) .GT. X) GO TO 2	IPROB 06
001	CONTINUE	IPROB 07
002	IPROB=I	IPROB 08
	RETURN	IPROB 09
	END	IPROB 10

FUNCTION RNOR

```

FUNCTION RNOR(IR)
C THIS FUNCTION COURTESY DR. REX HURST, DEPT. COMPUTER SCIENCE AND
C APPLIED STATISTICS, UTAH STATE UNIVERSITY, LOGAN, UTAH 84322
$SET OWN
DATA I/0/
$RESET OWN
IF (I .GT. 0) GO TO 30
010 X=2,*RANDOM(IR)=1.
Y=2,*RANDOM(IR)=1.
S=X*X+Y*Y
IF (S .GE. 1.) GO TO 10
S=SQRT(-2,*ALOG(S)/S)
RNOR=X*S
$SET OWN
GO2=Y*S
$RESET OWN
I=1
GO TO 40
030 RNOR=GO2
I=0
040 RETURN
END
RNOR 01
RNOR 02
RNOR 03
RNOR 04
RNOR 05
RNOR 06
RNOR 07
RNOR 08
RNOR 09
RNOR 10
RNOR 11
RNOR 12
RNOR 13
RNOR 14
RNOR 15
RNOR 16
RNOR 17
RNOR 18
RNOR 19
RNOR 20
RNOR 21
RNOR 22

```

LITERATURE CITED

GRIFFIN, R. A., R. J. HANKS, and S. CHILDS. 1974. Model for estimating water, salt, and temperature distribution in the soil profile. US/IBP Desert Biome Res. Memo. 74-61. Utah State Univ., Logan. 12 pp.

B. HEAT

P. W. Lommen

GENERAL DESCRIPTION OF THE HEAT SUBMODEL

This submodel determines average temperature during a time-step for each soil layer. Soil temperature information is needed by VEG in determining root respiration rates, by DCMP in determining decomposition rates and by WATER in checking if the soil is frozen. Temperature determinations more frequent than once each time-step (generally 4 days) would be more accurate but, given the overall model objectives, not likely to be more useful.

The approach used is very similar to that of Hanks et al. (1971). The soil surface temperature is taken to be the mean daily air temperature at 2 m for the time-step (determined in the weather submodel). The temperature at the bottom of the profile (60-cm depth was used for the Curlew Valley simulation) is set equal to the mean air temperature for the previous 30 days (adapted from the approach of Beckman et al. 1973). Once these boundary conditions are established the temperature profile is determined by the solution of Equation B-1:

$$\frac{\partial T}{\partial t} = \frac{\partial}{\partial z} \left(\sigma \frac{\partial T}{\partial z} \right), \quad (\text{B-1})$$

where

- T = the soil temperature at a given point at a given instant of time, °C;
- t = time, day;
- z = depth in soil, cm;
- σ = soil thermal diffusivity (the ratio of thermal conductivity, $\text{cal}\cdot\text{cm}^{-1}\cdot\text{day}^{-1}\cdot\text{C}^{-1}$, to specific heat, $\text{cal}\cdot\text{cm}^{-3}\cdot\text{C}^{-1}$), cm^2/day .

This differential equation is converted to a difference equation for computation. A set of nodes is introduced in the soil profile (at Curlew Valley the depths of these eight nodes are 0, 3, 10, 20, 30, 40, 50 and 60 cm) and the Crank-Nicholson iteration scheme (Richtmyer 1957) used to generate $N-2$ equations with $N-2$ unknowns. N is the number of soil nodes and the $N-2$ unknowns are the temperatures at all except top and bottom nodes. A tri-diagonal matrix method of solution (Richtmyer 1957) is then applied to the $N-2$ equations.

DISCUSSION

Thermal diffusivity is assumed to be independent of soil water content. Because of the boundary conditions and the relatively long time-step used (4 days for Curlew Valley), soil temperatures turn out to be fairly insensitive to the values used for thermal diffusivity. Values for thermal conductivity and specific heat were taken from Hanks et al. (1971).

CURLEW VALLEY IMPLEMENTATION

Thermal conductivity is equal to $86.4 \text{ cal}\cdot\text{cm}^{-1}\cdot\text{day}^{-1}\cdot\text{C}^{-1}$ throughout profile (Griffin et al. 1974). Specific heat, taken from the same source, is $0.3 \text{ cal}\cdot\text{cm}^{-3}\cdot\text{C}^{-1}$ throughout profile. The set of eight soil nodes at depths of 0, 3, 10, 20, 30, 40, 50 and 60 cm was chosen with a number of factors in mind: 1) almost all roots occur in the first 60 cm; 2) variations occur very slowly at 60 cm; 3) caliche layer depth is about 60 cm (the water submodel uses these same nodes; 4) 10-cm increments are convenient to work with.

PROGRAM DESCRIPTION

Only the important segments of the code in the program listing are shown and described. Sequence numbers are shown to aid in reference to full code listing which follows the program description. All comment cards, specification statements and bookkeeping sections have been left out. Almost all initialization has been deleted also. Definitions of variable names may be found in Table B-1, which also appears at the beginning of the program listing.

Table B-1. Variable dictionary for HEAT

CHD(I)	DEPTH TO SOIL NODE, CM.
CHDX(I)	THICKNESS OF LAYER AROUND NODE, COUNTING FROM 2ND NODE, CM.
CHDXX(I)	DISTANCE BETWEEN ADJACENT NODES, CHDXX(I) = CHD(I+1) - CHD(I), CM.
DBGFLG	IF=.TRUE. WRITE DEBUGGING INFORMATION THIS TIME STEP
DEBUG1	NUMBER OF TIME STEPS BETWEEN DEBUGGING OUTPUTS
DT	=PMDT, LENGTH OF TIME STEP, DAY.
FTAVE(PMDT,N)	FUNCTION TO AVERAGE PREVIOUS N DAYS' AIR TEMPERATURES BY AVERAGING APPROPRIATE NUMBER OF PREVIOUS TIME STEP AVERAGE TEMPERATURES.
PHCV(I)	SPECIFIC HEAT OF SOIL LAYER, (REMEMBER, 1ST LAYER IS CENTERED ON 2ND NODE.) CAL CM-3 C-1
PHK(I)	THERMAL CONDUCTIVITY OF SOIL BETWEEN NODES, (1ST VALUE IS FOR REGION BETWEEN NODES 1 AND 2), CAL CM-1 DAY-1 C-1.
PMDT	LENGTH OF TIME STEP, DAY.
PMJDAT	JULIAN DATE OF BEGINNING OF TIMESTEP.
PMN	NUMBER OF SOIL NODES.
THA(I)	THA(I) IS THE NEGATIVE OF THE COEFFICIENT OF THU(I+1) IN THE ITH EQUATION IN THE SET OF EQUATIONS TO SOLVE FOR THE THU ARRAY.
THB(I)	THB(I) IS THE COEFFICIENT OF THU(I) IN THE ITH EQUATION IN THE SET OF EQUATIONS TO SOLVE FOR THE THU ARRAY.
THC(I)	THC(I) IS THE NEGATIVE OF THE COEFFICIENT OF THU(I-1) IN THE ITH EQUATION IN THE SET OF EQUATIONS TO SOLVE FOR THE THU ARRAY.
THCCC(20)	AN ARRAY USED TO READ AND WRITE COMMENTS.
THD(I)	THD(I) IS THE CONSTANT, RIGHT HAND SIDE OF THE ITH EQUATION IN THE SET OF EQUATIONS TO SOLVE FOR THE THU ARRAY.
THM	AN INTEGER EQUAL TO THE LARGEST WHOLE NUMBER OF TIMESTEPS IN 30 DAYS.
THMR	A REAL VARIABLE EQUAL TO THM.
THTA(I)	TEMPERATURE OF SOIL NODES AT BEGINNING OF TIMESTEP, DEGREES C.
THTB(I)	TEMPERATURE OF SOIL NODES AT END OF TIMESTEP, DEGREES C.
THTHA(I)	TEMPORARY ARRAY USED IN CALCULATING MATRIX ELEMENTS IN CALL TO TRI-DIAGONAL MATRIX SUBROUTINE.
THTHB(I)	ANOTHER TEMPORARILY USED ARRAY FOR CALCULATING MATRIX ELEMENTS IN CALL TO TRI-DIAGONAL MATRIX SUBROUTINE.
THU(I)	THIS ARRAY CONTAINS THE TEMPERATURES OF THE INNER NODES (2 THROUGH PMN-1) AT THE END OF PMDT. THESE TEMPERATURES ARE THE SOLUTIONS OBTAINED BY THE TRI-DIAGONAL MATRIX SUBROUTINE.
TIME	TIME AT START OF TIMESTEP. EQUALS TSTART PLUS (NUMBER OF TIMESTEPS) * PMDT
TSTART	JULIAN DATE OF BEGINNING OF SIMULATION
XHSDLY(I)	AVERAGE TEMPERATURE OF SOIL NODE DURING PMDT, DEGREES C.
ZAIRT	AVERAGE AIR TEMPERATURE FOR PRESENT TIMESTEP, DETERMINED BY PSWG, DEGREES C.
ZHAIRT(3)	HOLDS THM PRESENT AND PAST TIME STEP AVERAGE AIR TEMPERATURES. ZHAIRT(1) HOLDS PRESENT TIME STEP TEMP. ZHAIRT(2) HOLDS PREVIOUS TIME STEP TEMP, ETC.

```
17 THTB(1) = ZAIRT
```

```
HEAT 195
```

Assume surface temperature equals average air temperature for time-step.

```

NTHM=THM+1
18 ZHAIRT(NTHM)=ZHAIRT(NTHM-1)
NTHM=NTHM-1
IF(NTHM .GE. ?) GO TO 18
ZHAIRT(1)=ZAIRT

```

```

HEAT 200
HEAT 201
HEAT 202
HEAT 203
HEAT 204

```

Variable NTHM is essentially a DO loop index which decreases by one each time (Burroughs FORTRAN doesn't accept negative DO loop increments). This loop shifts temperature in ZHAIRT, dropping the oldest time-step temperature from ZHAIRT (THM + 1) and loading present temperature in ZHAIRT (1).

```
THTB(PMN)=FTAVE(PMDT,30)
```

```
HEAT 208
```

Set the temperature of the bottom node equal to the average temperature of the previous 30 days. Accomplish this using function FTAVE.

```

28 DO 30 I= 1, PMN-2
30 THD(I) = THTA(I) + THTA(I) + THTA(I+1) +
  * (THTB(I) - THTA(I) - THTA(I+1)) + THTA(I+2) + THTA(I+1)
  THD(1) = THD(1) + THTB(1) + THTA(1)
  THD(PMN-2) = THD(PMN-2) + THTB(PMN) + THTA(PMN-1)

```

```

HEAT 223
HEAT 224
HEAT 225
HEAT 228
HEAT 229

```

Calculate the right-hand sides of the PMN—2 difference equations. For details see "Derivation of Difference Equations."

```
CALL TDM(THA, THB, TNC, THD, THU, PMN-2)
```

```
HEAT 239
```

TDM is the subroutine which solves PMN—2 equations for PMN—2 unknowns (temperatures at nodes 2 through PMN—1). The equations are of the form:

$$-C_i U_{i-1} + B_i U_i - A_i U_{i+1} = D.$$

The U 's (or THU's) are the temperatures determined by TDM and returned to HEAT. For more information see the detailed program description of subroutine TDM.

```

DO 32 I=1,PMN-2
32 THTB(I+1) = THU(I)

```

```

HEAT 246
HEAT 247

```

Load THU array, just determined by TDM, into the appropriate places in the THTB array.

```

DO 34 I=1,PMN
34 XHSOLT(I) = (THTA(I) + THTB(I)) / 2.

```

```

HEAT 250
HEAT 251

```

Make XHSOLT the average temperature at the soil nodes during PMDT.

ENTRY HINIT

HEAT 274

HEAT INITIALIZATION SECTION

Initial data are read in and written here. Also, calculations which need to be done only once each run are done here.

READ(5,80)THCC
WRITE(6,81)
WRITE(6,82)THCC

HEAT 281
HEAT 282
HEAT 283

First, a comment card is read (THCC is used only to read and write comments). This card is intended to be a general comment about HEAT, e.g., INITIALIZATION DATA FOR HEAT. Then come four spaces and this comment is written.

READ(5,80)THCC
WRITE(6,82)THCC
READ(5,83)CHD
WRITE(6,84)CHD

HEAT 285
HEAT 286
HEAT 287
HEAT 288

Another comment card is read and written, saying something about CHD (the depths of the nodes). The last characters on the card are the dimensions of the variable. Then values of the variable CHD, dimensioned 10, are read in and written.

An analogous procedure is then followed for reading in CHDX, PHCV, PHK, XHSOLT and ZHAIRT.

READ(5,80)THCC
WRITE(6,82)THCC

HEAT 317
HEAT 318

Finally, a comment is read and written, which says END READING INITIAL VALUES FOR HEAT.

DO 110 I=1,PMN-1
110 CHDX(I)=CHD(I+1)-CHD(I)

HEAT 328
HEAT 329

This array is simply the distances between adjacent soil nodes.

THM=30. / PMDT

HEAT 338

THM is an integer equal to the largest whole number of time-steps in 30 days.

DO 140 I=1,PMN-1
THTA(I)=PHK(I) / CHDX(I)
140 THTB(I)=2.*PHCV(I)*CHDX(I) / PMDT
DO 160 I=1,PMN-2
THA(I) = THTA(I+1)
THB(I) = THTB(I) + THTA(I) + THTA(I+1)
160 THC(I)=THTA(I)
THA(PMN-2)=0.
THC(I) = 0.

HEAT 342
HEAT 343
HEAT 344
HEAT 347
HEAT 348
HEAT 349
HEAT 350
HEAT 351
HEAT 352

These calculations load arrays THA, THB and THC for use when the tridiagonal matrix subroutine (TDM) is called. Since there is no water dependence on specific heat or thermal conductivity, these arrays do not change during the run and can be calculated just once. For details see "Derivation of Difference Equations."

COMPLETE PROGRAM LISTING

```

$SET SEPARATE                          HEAT 001
$SET DIM                                 HEAT 002
C                                         HEAT 003
C                                         HEAT 004
SUBROUTINE HEAT                          HEAT 005
C                                         HEAT 006
C SOIL HEAT PROFILE SUBMODEL OF DESERT BIOME WATER RESPONSE MODEL HEAT 007
C MARCH 1976.   WRITTEN BY:   PAUL W. LOMMEN HEAT 008
C                                       ECOLOGY CENTER, UMC 52 HEAT 009
C                                       UTAH STATE UNIVERSITY HEAT 010
C                                       LOGAN, UTAH , 84322 HEAT 011
C                                         HEAT 012
C                                         HEAT 013
C VVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVV HEAT 014
C VVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVV HEAT 015
C                                         HEAT 016
VARIABLE DICTIONARY FOR HEAT            HEAT 017
C                                         HEAT 018
C                                         HEAT 019
C CHD(10)    DEPTH TO SOIL NODE, CM. HEAT 020
C                                         HEAT 021
C CHDX(10)   THICKNESS OF LAYER AROUND NODE, COUNTING FROM 2ND HEAT 022
C            NODE, CM. HEAT 023
C CHDXX(10)  DISTANCE BETWEEN ADJACENT NODES, CHDXX(I) = CHD(I+1) HEAT 024
C            - CHD(I), CM. HEAT 026
C DBGFLG     IF=.TRUE. WRITE DEBUGGING INFORMATION THIS TIME STEP HEAT 027
C DEBGI      NUMBER OF TIME STEPS BETWEEN DEBUGGING OUTPUTS HEAT 029
C DT         =PMDT, LENGTH OF TIME STEP, DAY. HEAT 030
C FTAVE(PMDT,N) FUNCTION TO AVERAGE PREVIOUS N DAYS' AIR TEMPERATURES HEAT 031
C            BY AVERAGING APPROPRIATE NUMBER OF PREVIOUS TIME HEAT 032
C            STEP AVERAGE TEMPERATURES. HEAT 033
C PHCV(10)   SPECIFIC HEAT OF SOIL LAYER, (REMEMBER, 1ST LAYER HEAT 034
C            IS CENTERED ON 2ND NODE.) CAL CM-3 C-1 HEAT 035
C PHK(10)    THERMAL CONDUCTIVITY OF SOIL BETWEEN NODES, HEAT 036
C            (1ST VALUE IS FOR REGION BETWEEN NODES 1 AND 2), HEAT 037
C            CAL CM-1 DAY-1 C-1. HEAT 038
C PMDT       LENGTH OF TIME STEP, DAY. HEAT 039
C PHJDAT     JULIAN DATE OF BEGINNING OF TIMESTEP. HEAT 040
C PMN        NUMBER OF SOIL NODES. HEAT 041
C THA(I)     THA(I) IS THE NEGATIVE OF THE COEFFICIENT OF HEAT 042
C            THU(I+1) IN THE ITH EQUATION IN THE SET OF HEAT 043
C            EQUATIONS TO SOLVE FOR THE THU ARRAY. HEAT 044
C THB(10)    THB(I) IS THE COEFFICIENT OF THU(I) IN THE ITH HEAT 045
C            EQUATION IN THE SET OF EQUATIONS TO SOLVE FOR THE HEAT 046
C            THU ARRAY. HEAT 047
C THC(10)    THC(I) IS THE NEGATIVE OF THE COEFFICIENT OF HEAT 048
C            THU(I-1) IN THE ITH EQUATION IN THE SET OF HEAT 049
C            EQUATIONS TO SOLVE FOR THE THU ARRAY. HEAT 050
C THCCC(20)  AN ARRAY USED TO READ AND WRITE COMMENTS. HEAT 051
C THD(10)    THD(I) IS THE CONSTANT, RIGHT HAND SIDE OF THE ITH HEAT 052
C            EQUATION IN THE SET OF EQUATIONS TO SOLVE FOR THE HEAT 053
C            THU ARRAY. HEAT 054
C THM        AN INTEGER EQUAL TO THE LARGEST WHOLE NUMBER OF HEAT 055
C            TIMESTEPS IN 30 DAYS. HEAT 056
C THMR       A REAL VARIABLE EQUAL TO THM. HEAT 057
C THTA(10)   TEMPERATURE OF SOIL NODES AT BEGINNING OF TIMESTEP, HEAT 058
C            DEGREES C. HEAT 059
C THTB(10)   TEMPERATURE OF SOIL NODES AT END OF TIMESTEP, HEAT 060
C            DEGREES C. HEAT 061
C THTHA(10)  TEMPORARY ARRAY USED IN CALCULATING MATRIX HEAT 062
C            ELEMENTS IN CALL TO TRI-DIAGONAL MATRIX SUBROUTINE. HEAT 063
C THTHB(10)  ANOTHER TEMPORARILY USED ARRAY FOR CALCULATING HEAT 064
C            MATRIX ELEMENTS IN CALL TO TRI-DIAGONAL MATRIX HEAT 065
C            SUBROUTINE. HEAT 066
C THU(10)    THIS ARRAY CONTAINS THE TEMPERATURES OF THE INNER HEAT 067
C            NODES (2 THROUGH PMN-1) AT THE END OF PMDT. THESE HEAT 068
C            TEMPERATURES ARE THE SOLUTIONS OBTAINED BY THE HEAT 069
C            TRI-DIAGONAL MATRIX SUBROUTINE. HEAT 070
C            HEAT 071
C            HEAT 072
C            HEAT 073
C            HEAT 074
C            HEAT 075
C            HEAT 076
C            HEAT 077
C            HEAT 078
C            HEAT 079
C            HEAT 080
C            HEAT 081
C            HEAT 082
C            HEAT 083
C            HEAT 084
C            HEAT 085
C            HEAT 086
C            HEAT 087
C            HEAT 088
C            HEAT 089
C            HEAT 090

```


DERIVATION OF DIFFERENCE EQUATIONS

As mentioned in the general description of HEAT, the temperature profile is determined by the solution of Equation B-1:

$$\frac{\partial T}{\partial t} = \frac{\partial}{\partial z} \left(\sigma \frac{\partial T}{\partial z} \right). \quad (\text{B-1})$$

where

- T = the soil temperature at a given point at a given instant of time, °C;
- t = time, day;
- z = depth in soil, cm;
- σ = soil thermal diffusivity (the ratio of thermal conductivity, $\text{cal}\cdot\text{cm}^{-1}\cdot\text{day}^{-1}\cdot\text{C}^{-1}$, to specific heat, $\text{cal}\cdot\text{cm}^{-3}\cdot\text{C}^{-1}$), cm^2/day .

The difference equations are derived as follows:

for a thin layer of soil during a time interval Δt ,

$$(\text{heat in}) - (\text{heat out}) = (\text{heat stored}). \quad (\text{B-2})$$

In particular, for layer i during Δt ,

$$(\text{heat in})_i = - \left(K \frac{\partial T}{\partial z} \right) \text{ at top of layer } i, \quad (\text{B-3})$$

where

counting of layers starts at the soil surface and goes down; heat flowing down is positive:

(heat in) $_i$ in Equation B-3 has units of $\text{cal}\cdot\text{cm}^{-2}\cdot\text{day}^{-1}$;

- K = soil thermal conductivity, $\text{cal}\cdot\text{cm}^{-1}\cdot\text{day}^{-1}\cdot\text{C}^{-1}$;
- T = temperature of layer i , °C;
- z = depth in soil, cm.

Flesh out Equation B-3,

$$(\text{heat in})_i = (-K_{i-1/2}) \left\{ \left[\frac{T_i - T_{i-1}}{\Delta z_{i-1/2}} + \frac{T_i^* - T_{i-1}^*}{\Delta z_{i-1/2}} \right] \left(\frac{1}{2} \right) \right\}, \quad (\text{B-4})$$

where

- T_i = temperature of layer i at beginning of Δt , assumed located at the center of a layer, called a node. The first node is at the soil surface at 0-cm depth. The second is at 3-cm depth (at Curlew Valley) and is at the center of layer 1, which extends from 1- to 5-cm depth. Node 3 at Curlew is at 10-cm depth and is at the center of layer 2, which extends from 5- to 15-cm depth, etc. The bottom layer is centered on the next to the bottom node. There are PMN nodes

and, hence, PMN—2 layers. These distinctions between nodes, layers, regions between nodes, etc., appear confusing but must be made in using the difference equations;

- T_{i-1} = temperature of layer $i-1$ at beginning of Δt ;
- T_i^* = temperature of layer i at end of Δt ;
- T_{i-1}^* = temperature of layer $i-1$ at end of Δt ;
- $K_{i-1/2}$ = conductivity in region between centers of layers $i-1$ and i (i.e., between nodes i and $i+1$);
- $\Delta z_{i-1/2}$ = distance from center of layer $i-1$ to center of layer i (i.e., from node i to $i+1$).

The quantity in braces is the temperature gradient determined by the Crank-Nicholson scheme (Richtmyer 1957) and is the average of the temperature gradients at the beginning and end of Δt .

By similar reasoning:

$$\begin{aligned} (\text{heat out})_i &= - \left(K \frac{\partial T}{\partial z} \right) \text{ at bottom of layer } i, \\ &= - (K_{i+1/2}) \left\{ \left[\frac{T_{i+1} - T_i}{\Delta z_{i+1/2}} + \frac{T_{i+1}^* - T_i^*}{\Delta z_{i+1/2}} \right] \left(\frac{1}{2} \right) \right\}. \end{aligned} \quad (\text{B-5})$$

The heat stored in $\text{cal}\cdot\text{cm}^{-2}\cdot\text{day}^{-1}$

$$= \frac{(Cv_i)(\Delta x_i)}{\Delta t} (T_i^* - T_i), \quad (\text{B-6})$$

where

- Cv_i = specific heat of layer i , $\text{cal}\cdot\text{cm}^{-3}\cdot\text{C}^{-1}$;
- Δx_i = thickness of layer i , cm;
- Δt = length of time-step, day.

By substituting Equations B-4, B-5 and B-6 in Equation B-2 we get:

$$\begin{aligned} & \frac{-K_{i-1/2}}{2} \left[\frac{T_i - T_{i-1} + T_i^* - T_{i-1}^*}{\Delta z_{i-1/2}} \right] \\ & + \frac{K_{i+1/2}}{2} \left[\frac{T_{i+1} - T_i + T_{i+1}^* - T_i^*}{\Delta z_{i+1/2}} \right] \\ & = \frac{Cv_i \Delta x_i}{\Delta t} (T_i^* - T_i). \end{aligned} \quad (\text{B-7})$$

Multiply through by two and collect T^* terms on the left. After two lines of algebra we get:

$$-C_i T_{i-1}^* + B_i T_i^* - A_i T_{i+1}^* = D_i, \quad (\text{B-8})$$

where

$$C_i = \frac{K_{i-1/2}}{\Delta z_{i-1/2}}, \quad (\text{B-9})$$

$$B_i = \frac{K_{i-1/2}}{\Delta z_{i-1/2}} + \frac{K_{i+1/2}}{\Delta z_{i+1/2}} + \frac{2Cv_i \Delta x_i}{\Delta t}, \quad (\text{B-10})$$

$$A_i = \frac{K_{i+1/2}}{\Delta z_{i+1/2}}, \quad (\text{B-11})$$

$$D_i = T_{i-1} \left(\frac{K_{i-1/2}}{\Delta z_{i-1/2}} \right) + T_i \left[\frac{2Cv_i \Delta x_i}{\Delta t} - \frac{K_{i-1/2}}{\Delta z_{i-1/2}} - \frac{K_{i+1/2}}{\Delta z_{i+1/2}} \right] + T_{i+1} \left(\frac{K_{i+1/2}}{\Delta z_{i+1/2}} \right). \quad (\text{B-12})$$

For top and bottom layers we must make small changes in Equations B-8 to 12 because boundary condition must be included. The temperature at the surface at the end of the time-step is set equal to the average air temperature during Δt :

$$T_s^* = \text{ZAIRT}. \quad (\text{B-13})$$

The temperature at the bottom node T_B^* (there are PMN nodes) at the end of Δt is set equal to the average air temperature during the previous 30 days.

For the top layer then, (heat in)₁ contains only one unknown temperature, not two as shown in Equation B-4.

That is

$$(\text{heat in})_1 = - (K_{1/2}) \left[\frac{T_1 - T_s}{\Delta z_{1/2}} - \frac{T_1^* - T_s^*}{\Delta z_{1/2}} \right] \left(\frac{1}{2} \right). \quad (\text{B-14})$$

In Equation B-14, only T_1^* , the temperature of layer 1 (i.e., node 2) is unknown. Equations B-5 and 6 for the top layer with $i=1$ are unchanged. Substituting Equations B-14, 5 and 6 in Equation B-2 now yields by the same procedure as before:

$$B_1 T_1^* - A_1 T_2^* = D_1 \quad (\text{B-15})$$

where

$$B_1 = \frac{K_{1/2}}{\Delta z_{1/2}} + \frac{K_{3/2}}{\Delta z_{3/2}} + \frac{2Cv_1 \Delta x_1}{\Delta t}, \quad (\text{B-16})$$

$$A_1 = \frac{K_{3/2}}{\Delta z_{3/2}}, \quad (\text{B-17})$$

$$D_1 = T_s \frac{K_{1/2}}{\Delta z_{1/2}} + T_1 \left[\frac{2Cv_1 \Delta x_1}{\Delta t} - \frac{K_{1/2}}{\Delta z_{1/2}} - \frac{K_{3/2}}{\Delta z_{3/2}} \right] + T_2 \frac{K_{3/2}}{\Delta z_{3/2}} + T_s^* \frac{K_{1/2}}{\Delta z_{1/2}}. \quad (\text{B-18})$$

For the bottom layer, (heat out)_{bottom} contains only one unknown temperature. After several lines of algebra we then get (say the bottom layer is layer number M):

$$-C_m T_{m-1}^* + B_m T_m^* = D_m, \quad (\text{B-19})$$

where

$$C_m = \frac{K_{m-1/2}}{\Delta z_{m-1/2}}, \quad (\text{B-20})$$

$$B_m = \frac{K_{m-1/2}}{\Delta z_{m-1/2}} + \frac{K_{m+1/2}}{\Delta z_{m+1/2}} + \frac{2Cv_m \Delta x_m}{\Delta t}, \quad (\text{B-21})$$

$$D_m = T_{m-1} \left(\frac{K_{m-1/2}}{\Delta z_{m-1/2}} \right) + T_m \left[\frac{2Cv_m \Delta x_m}{\Delta t} - \frac{K_{m-1/2}}{\Delta z_{m-1/2}} - \frac{K_{m+1/2}}{\Delta z_{m+1/2}} \right] + T_B \left(\frac{K_{m+1/2}}{\Delta z_{m+1/2}} \right) + T_B^* \left(\frac{K_{m+1/2}}{\Delta z_{m+1/2}} \right). \quad (\text{B-22})$$

Now summarize these equations for A , B , C and D .

$$A_i = \frac{K_{i+1/2}}{\Delta z_{i+1/2}} \quad \text{for } i = 1 \text{ to PMN} - 3, \\ = 0 \quad \text{for } i = \text{PMN} - 2 \text{ (bottom layer)}, \quad (\text{B-23})$$

$$B_i = \frac{K_{i-1/2}}{\Delta z_{i-1/2}} + \frac{K_{i+1/2}}{\Delta z_{i+1/2}} + \frac{2Cv_i \Delta x_i}{\Delta t} \\ \text{for } i = 1 \text{ to PMN} - 2, \quad (\text{B-24})$$

$$C_i = 0 \quad \text{for } i = 1, \\ = \frac{K_{i-1/2}}{\Delta z_{i-1/2}} \quad \text{for } i = 2 \text{ to PMN} - 2, \quad (\text{B-25})$$

$$\begin{aligned}
D_i &= T_s \frac{K_{1/2}}{\Delta z_{1/2}} + T_1 \left[\frac{2C_{v1} \Delta x_1}{\Delta t} - \frac{K_{1/2}}{\Delta z_{1/2}} - \frac{K_{3/2}}{\Delta z_{3/2}} \right] \\
&\quad + T_2 \frac{K_{3/2}}{\Delta z_{3/2}} + T_s^* \frac{K_{1/2}}{\Delta z_{1/2}} \\
&\quad \text{for } i = 1, \\
&= T_{i-1} \frac{K_{i-1/2}}{\Delta z_{i-1/2}} + T_i \left[\frac{2C_{vi} \Delta x_i}{\Delta t} - \frac{K_{i-1/2}}{\Delta z_{i-1/2}} \right. \\
&\quad \left. - \frac{K_{i+1/2}}{\Delta z_{i+1/2}} \right] + T_{i+1} \frac{K_{i+1/2}}{\Delta z_{i+1/2}} \\
&\quad \text{for } i = 2 \text{ to } \text{PMN} - 3, \\
&= T_{i-1} \frac{K_{i-1/2}}{\Delta z_{i-1/2}} + T_i \left[\frac{2C_{vi} \Delta x_i}{\Delta t} - \frac{K_{i-1/2}}{\Delta z_{i-1/2}} \right. \\
&\quad \left. - \frac{K_{i+1/2}}{\Delta z_{i+1/2}} \right] + T_B \frac{K_{i+1/2}}{\Delta z_{i+1/2}} + T_B^* \frac{K_{i+1/2}}{\Delta z_{i+1/2}} \\
&\quad \text{for } i = \text{PMN} - 2. \tag{B-26}
\end{aligned}$$

It can easily be seen in Equations B-23, 24 and 25 that there are no temperatures involved. Also, the conductivity, K , and specific heat, C_v , are not dependent on water content or temperature. (A water content dependence can be included if it is felt such detail is warranted. See Hanks et al. 1971). Thus, the A 's, B 's and C 's need to be calculated only once. The D 's must be calculated each time-step.

Transform Equations B-23, 24, 25 and 26 into computer code:

- a. Thermal conductivity $K_{1/2}$ becomes PHK(1)
 $K_{3/2}$ becomes PHK(2)
 \vdots
 $K_{\text{PMN}-2+1/2}$ becomes PHK(PMN-1)
- b. Specific heat C_{vi} becomes PHCV(I) for I = 1 to PMN-2
- c. Distance from node to node $\Delta z_{1/2}$ becomes CHDXX(1)
 \vdots
 $\Delta z_{\text{PMN}-2+1/2}$ becomes CHDXX(PMN-1)
- d. Thickness of layer Δx_i becomes CHDX(I) for I = 1 to PMN-2

- e. Temperatures at beginning of time-step

T_i becomes THTA(I+1)
for I = 1 to PMN-2
 T_s becomes THTA(1)
 T_B becomes THTA(PMN)

- f. Temperatures at end of time-step

T_s^* becomes THTB(1)
 T_B^* becomes THTB(PMN)

- g. Time-step length

Δt becomes PMDT

- h. Tridiagonal matrix elements

A_i, B_i, C_i, D_i become THA(I), THB(I), THC(I), THD(I) for I = 1, PMN-2.

Now, if we define temporary variables THTHA(I) = PHK(I)/CHDXX(I) and THTHB(I) = 2. * PHCV(I) * CHDX(I)/PMDT, then Equation B-23 becomes

$$\begin{aligned}
\text{THA(I)} &= \text{THTHA(I+1)} & \text{for I = 1 to PMN - 3,} \\
&= 0 & \text{for I = PMN - 2,}
\end{aligned}$$

Equation B-24 becomes

$$\begin{aligned}
\text{THC(I)} &= 0 & \text{for I = 1,} \\
&= \text{THTHA(I)} & \text{for I = 2 to PMN - 2,}
\end{aligned}$$

Equation B-25 becomes

$$\begin{aligned}
\text{THB(I)} &= \text{THTHA(I)} + \text{THTHA(I+1)} + \text{THTHB(I)} \\
&\text{for I = 1 to PMN - 2,}
\end{aligned}$$

Equation B-26 becomes

$$\begin{aligned}
\text{THD(I)} &= \text{THTA(1)} * \text{THTHA(1)} + \text{THTA(2)} * (\text{THTHB(1)} - \text{THTHA(1)} \\
&\quad - \text{THTHA(2)}) + \text{THTA(3)} * \text{THTHA(2)} + \text{THTB(1)} * \text{THTHA(1)} \\
&\quad \text{for I = 1,} \\
&= \text{THTA(I)} * \text{THTHA(I)} + \text{THTA(I+1)} * (\text{THTHB(I)} - \text{THTHA(I)} \\
&\quad - \text{THTHA(I+1)}) + \text{THTA(I+2)} * \text{THTHA(I+1)} \\
&\quad \text{for I = 2 to PMN - 3,} \\
&= \text{THTA(PMN-2)} * \text{THTHA(PMN-2)} + \text{THTA(PMN-1)} \\
&\quad * (\text{THTHB(PMN-2)} - \text{THTHA(PMN-2)} - \text{THTHA(PMN-1)}) \\
&\quad + \text{THTA(PMN)} * \text{THTHA(PMN-1)} + \text{THTB(PMN)} * \text{THTHA(PMN-1)} \\
&\quad \text{for I = PMN - 2.}
\end{aligned}$$

As mentioned earlier, THA, THB and THC have no water content or temperature dependence and hence are calculated only once, in the initialization section. THD is dependent on temperature and hence is calculated each time-step.

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C. WATER
(including WBAL and WTIME)
P. W. Lommen

**GENERAL DESCRIPTION OF THE
WATER SUBMODEL**

In the Water Response Model, soil water potential is needed by the plant submodels in order to predict growth. This is provided by submodel WATER. Soil water potential also helps determine rates in nitrogen and decomposition submodels. Inputs required each time-step are precipitation, evaporation and transpiration. Soil parameters needed once are: 1) relation between soil water potential and soil water content (theta); 2) relation between hydraulic conductivity and theta; 3) wet and dry limits to soil water potential; 4) conductivity of caliche layer; 5) soil layer thicknesses and locations; 6) average runoff:runoff ratio per day; 7) largest permissible value of change in theta during water time-step (which is generally much smaller than overall model time-step); 8) initial values of theta and depth of water standing on surface.

Various papers (Hanks and Bowers 1962; Hanks et al. 1969; Nimah and Hanks 1973; Griffin et al. 1974) have provided ideas which have contributed to this effort.

A set of nodes is introduced in the soil profile (the same nodes as for submodel HEAT) as reference points for soil layers. Figure C-1 is a schematic diagram of submodel WATER. The soil water potential in soil layer i is the solution of Equation C-1:

$$C_i \left[\frac{\partial H}{\partial t} \right] = \left[\frac{\partial}{\partial z} \left(K \frac{\partial H}{\partial z} \right) \right] + A_i \quad (C-1)$$

A_i is the removal of water by roots in layer i (transpiration) and is proportional to water available in layer i and relative root amount in layer i .

The remainder of the equation is a standard diffusion equation: H is the soil water potential in layer i , bars; K is the soil hydraulic conductivity at layer i , $\text{cm}^2 \cdot \text{bar}^{-1} \cdot \text{day}^{-1}$; C_i is the specific water capacity of layer i , bar^{-1} ; Z is soil depth, centimeters; t is time, days.

The upper boundary condition on Equation C-1 is determined by the surface flux of water which depends on evaporation, precipitation and precipitation intensity, and standing water. The bottom boundary condition is determined by the caliche layer conductivity and water content, both of which are read in with initial parameters.

A set of difference equations is then generated. It is solved by a tridiagonal matrix method of Richtmyer (1957). For more details, see "Derivation of Difference Equations" below, and Detailed Program Description of subroutine TDM in section D of this research memorandum.

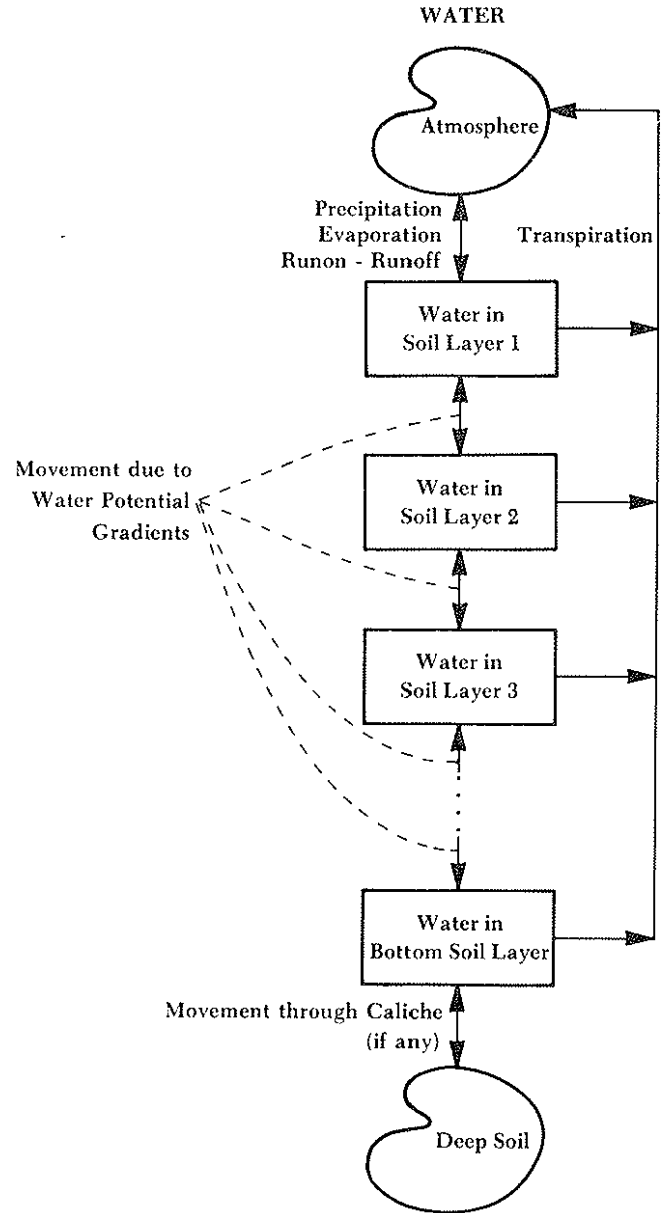


Figure C-1. Schematic diagram of submodel WATER.

PROGRAM DESCRIPTION

Only the important segments of the FORTRAN code are shown and described. Sequence numbers are shown to aid in reference to the full code listing that follows the program description. All comment cards, specification statements and bookkeeping sections have been deleted also. Definitions of variable names may be found in Table C-1, which also appears at the beginning of the program listing.

Before proceeding to the code, several additional comments should be made. First, volumetric water content of the soil, a dimensionless quantity (cm of water per cm of soil), is a very important and much-used variable throughout the program and will be called *theta* in the description. Similarly, soil water potential, measured in bars pressure, will be abbreviated SWP

Second, water subdivides PMDT, the overall model

time-step (typically 4 days), as needed. If water is moving rapidly into the soil profile during a precipitation event, e.g., the WATER time-step, TWDT will be cut down to 30 min (the minimum allowed). WATER tries to find the largest TWDT consistent with the requirement that theta does not change more than PWTLIM (typically .015) during TWDT. We want TWDT large enough to be efficient but small enough to be accurate. There is a fair amount of code concerned with determining the best value of TWDT.

Third, in making difference equations from the main differential equation, it has been necessary to define nodes, layers and regions between nodes for the soil profile. This can be confusing. Figure C-2 shows how this was done for Curlew Valley. There are PMN nodes or reference points, PMN-1 regions between nodes and PMN-2 layers. Also shown in Figure C-2 are the variables which are indexed by layers, nodes and regions.

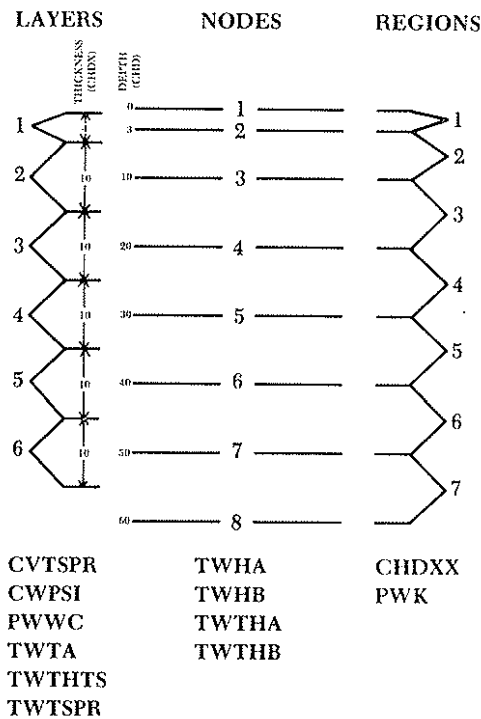


Figure C-2. Relations between layers, nodes and regions between nodes for the soil profile as used in Curlew Valley runs. The first layer is centered on the second node, etc. Also shown are layer width (CHDX) and node depth (CHD), both in centimeters. Finally, arrays indexed by layer are listed under column LAYERS, etc.

Table C-1. Variable dictionary for WATER

LIST OF VARIABLES FOR WATER	
CHD(10)	DEPTHS, CM, OF NODES BELOW SURFACE (A NODE IS A POINT WHERE A WATER POTENTIAL IS CALCULATED) STARTING AT SURFACE
CHDX(10)	THICKNESS OF REGION CENTERED ON NODE WHERE COUNTING STARTS AT FIRST NODE BELOW SURFACE AND GOES DOWN
CHDXX(10)	$CHDXX(I) = CHD(I+1) - CHD(I)$
CVTSPR(10)	MM OF WATER TAKEN FROM EACH LAYER DURING PMDT
CHINF	NET TOP SURFACE FLUX FOR PMDT, CM, DOES NOT INCLUDE TRANSPIRATION.
CHPSI(10)	SOIL WATER POTENTIAL, BARS, VALUE AT END OF PMDT
DBGFLG	IF=.TRUE. WRITE DEBUGGING INFORMATION THIS TIME STEP
DEBUG1	NUMBER OF TIME STEPS BETWEEN DEBUGGING OUTPUTS
FLGFLG	IF=.TRUE. WRITE DEBUGGING INFORMATION THIS TWDT
PHDT	TIME STEP DETERMINED IN MAIN, DAYS
PHN	NUMBER OF NODES, SAME AS NUMBER OF NODES IN HEAT, COUNTING STARTS AT SURFACE
PHDELW	INCREMENT IN VHC, TYPICALLY 0.01-0.02
PHW(60)	TABLE OF HYDRAULIC PRESSURE HEAD VERSUS THETA, BARS, USING SAME THETA SCALE AND SPACING AS PHKIN
PHLERY	ALLOWABLE LOW PRESSURE LIMIT FOR ANY LAYER
PHHRET	ALLOWABLE HIGH PRESSURE LIMIT FOR ANY LAYER
PHK(10)	HYDRAULIC CONDUCTIVITY, CM ² BAR ⁻¹ DAY ⁻¹ , PHK(I) IS CONDUCTIVITY AVERAGED OVER TWDT, IN REGION BETWEEN NODES I AND I+1, WHERE COUNTING STARTS AT SURFACE.
PHKCAL	CONDUCTIVITY OF CALICHE LAYER
PHKIN(60)	TABLE OF CONDUCTIVITY VERSUS THETA, CM ² BAR ⁻¹ DAY ⁻¹ , STARTING WITH VALUE AT THETA = 0.0, READ IN IN WINIT.
PHKSUM(60)	ARRAY USED TO CALCULATE AVERAGE K, (CM-BARS)/DAY
PHLLIM	LOWER LIMIT OF THETA, DIMENSIONLESS CALCULATED IN WINIT AND = SUM OF CONDUCTIVITY X DELTA H
PHM	NUMBER OF ENTRIES IN CONDUCTIVITY AND PRESURE TABLES
PHRRRF	RUNON TO RUNOFF RATIO PER DAY
PHW(60)	ARRAY CONTAINING UNIFORMLY DISTRIBUTED VALUES OF THETA
PHWTLIM	LARGEST CHANGE IN THETA ALLOWED FOR NODES 2 TO PHN-1 DURING ANY TIME STEP TWDT. TYPICAL VALUE .01 TO .02
PHWC(10)	WATER CAPACITY, BAR ⁻¹ , OF REGIONS SURROUNDING NODES, COUNTING STARTS AT FIRST NODE BELOW SURFACE
P	A VERY TEMPORARY VARIABLE - USED IN REFINED TWDT ESTIMATE, TRANSPIRATION SECTION, REDUCTION OF TWDT IF DELTA THETA TOO LARGE, AND AS RATIO OF WATER IN TO WATER STORED AS CALCULATED IN WBAL.
S	A VERY TEMPORARY VARIABLE - USED IN REFINED TWDT ESTIMATE, TRANSPIRATION SECTION, CONDUCTIVITY SECTION, AND REDUCTION OF TWDT IF DELTA THETA TOO LARGE.
T	A VERY TEMPORARY VARIABLE - USED IN REFINED TWDT ESTIMATE.
TDM	A SUBROUTINE USED BY WATER AND HEAT TO SOLVE A SET OF SIMULTANEOUS DIFFERENCE EQUATIONS, THE LEFT HAND SIDES OF WHICH FORM A TRI-DIAGONAL MATRIX.
TIME	CURRENT JULIAN DATE - BOOKKEEPING DONE IN MAIN PROGRAM.
TSTART	JULIAN DATE OF START OF RUN - INITIALIZED IN MAIN PROGRAM
TWAA	FRACTION OF PHDELW INTERVAL, USED IN INTERPOLATING
TWA(10)	PARAMETER A FOR TRI-DIAGONAL MATRIX ROUTINE (TDM ROUTINE)
TWAAA(10)	ARRAY CONTAINING VALUES OF TWTHA(I) FOR I=2,PHN-1

Table C-1, continued

TWB(10)	PARAMETER B FOR TOM ROUTINE
TWBE(10)	CONTAINS ORDERED VALUES OF TWAAA.
TW	A VERY TEMPORARY VARIABLE - USED IN SURFACE FLUX CALCULATION, TRANSPIRATION SECTION, AND CONDUCTIVITY SECTION.
TWC(10)	PARAMETER C FOR TOM ROUTINE
TWCCC	USED TO READ AND WRITE COMMENTS IN INPUT DATA
TWD(10)	PARAMETER D FOR TOM ROUTINE
TWDT	ACTUAL TIME STEP WATER IS USING
TWDTHP(10)	CHANGE IN THETA DURING LAST TWDT
TWDTP	LENGTH OF PREVIOUS WATER TIME STEP
TWDT0	HOLDS CURRENT VALUE OF TWDT AS IT IS BEING REDUCED BECAUSE DELTA THETA IS TOO LARGE.
TWEVAP	EVAPORATIVE DEMAND, MM, WATER TO EVAPORATE FROM SURFACE DURING REMAINDER OF PMDT
TWFRZN	EQUALS .TRUE. IF SOIL PROFILE FROZEN, EQUALS .FALSE. IF NOT FROZEN.
TWI	TEMPORARY VARIABLE USED TO SET UP PWT ARRAY.
TWHA(10)	PRESSURE HEAD AT BEGINNING OF TWDT, BARS
TWHB(10)	PRESSURE HEAD AT END OF TWDT, BARS
TWHHA	INTERPOLATED VALUES OF PWH
TWIN	EQUALS .TRUE. IF SURFACE PRESSURE IS WITHIN LIMITS
TWJA	INTEGER USED IN PICKING VALUES OF K, H OFF TABLES
TWJB	INTEGER USED IN PICKING VALUES OF K, H OFF TABLES
TWJM	INTEGER COUNTER USED TO LIMIT NUMBER OF TWDT HALVINGS
TWKS	INTERPOLATED VALUE OF PWKSUM
TWONE	EQUALS .TRUE. IF FIRST DELTA THETA APPROXIMATION HAS ALREADY BEEN DONE DURING CURRENT TWDT.
TWR	RAIN, MM, WATER TO INFILTRATE, LEAVE STANDING OR RUNOFF DURING REMAINDER OF PMDT.
TWRAIN	EQUALS .TRUE. IF RAIN OCCURS DURING TWDT
TWRP	EQUALS .TRUE. IF TWRAIN WAS TRUE LAST TWDT.
TWSF	SURFACE FLUX OF WATER OR VAPOR, CM
TWSFB	FLUX OF WATER MOVING INTO PROFILE DURING TWDT THROUGH CALICHE, CM.
TWSFBT	FLUX OF WATER MOVING INTO PROFILE DURING PMDT THROUGH CALICHE, CM.
TWSFC	CALCULATED SURFACE FLUX USING TWHB(1) AND PWK(1).
TWSFD	EQUALS TWSF-TWSFC
TWSTND	STANDING WATER, MM, AFTER RUNON-RUNOFF ASSUMPTIONS HAVE BEEN CONSIDERED.
TWSTRD	CHANGE DURING TIME STEP IN WATER STORED IN LAYERS OF PROFILE, CM. (SHOULD EQUAL TWIN)
TWTA(10)	WATER AVAILABLE IN LAYERS, DIMENSIONLESS
TWTHAA(10)	VOLUMETRIC WATER CONTENT AT BEGINNING OF PMDT
TWTHA(10)	VOLUMETRIC WATER CONTENT (VMC) AT BEGINNING OF TWDT
TWTHB(10)	VMC AT END OF TWDT
TWTHC	THETA AVERAGED OVER TWDT FOR WHATEVER LAYER WE'RE CALCULATING AT THE MOMENT

Table C-1, continued

TWTK	PROPORTIONALITY CONSTANT WHICH IS CALCULATED TO MAKE TRANSPIRATION DEMAND FROM A LAYER PROPORTIONAL TO WATER AVAILABLE AND RELATIVE ROOT AMOUNTS IN A LAYER.
TWHTS(10)	RUNNING SUM OVER TWDT'S OF TRANSPIRATION BY LAYER, CM
TWSPR(10)	TRANSPIRATION LOSS OF WATER FROM REGION SURROUNDING NODE, CM, (TOP AND BOTTOM NODES EXCLUDED) DURING TWDT
TWTTOT	TOTAL TIME THAT HAS ELAPSED IN PMDT. DOES NOT INCLUDE CURRENT TWDT, DAYS
TWTN(6)	TRANSPIRATIONAL DEMAND BY FUNCTIONAL GROUP, CM
TWU(10)	WATER POTENTIALS IN BARS OF SOIL LAYERS AS RETURNED FROM SUBROUTINE TDM.
TWWA(10)	USED IN SETTING UP CALL TO TDM
TWWB(10)	USED IN SETTING UP CALL TO TDM
TWWIN	FLUX OF WATER MOVING IN TO PROFILE THIS TIME STEP, EQUALS FLUX IN AT TOP + FLUX IN FROM BOTTOM - TRANSPIRATION, CM. (SHOULD EQUAL TWSTRD)
U	A VERY TEMPORARY VARIABLE-USED IN REFINED TWDT ESTIMATE
WBAL	SUBROUTINE WHICH CHECKS OVERALL ACCURACY OF CALCULATIONS DONE IN WATER - DETERMINES WATER INTO OR OUT OF PROFILE AND COMPARES IT TO CHANGE IN WATER STORED IN PROFILE.
WTIME	FUNCTION WHICH RETURNS ELAPSED CPU TIME IN SECONDS FROM START OF RUN.
XHSGLT(10)	SOIL TEMPERATURE PROFILE, C. INDEXED BY NODES, SO XHSOLT(2), E.G., IS TEMPERATURE OF FIRST LAYER.
XWSIND	STATE VARIABLE CALCULATED BY WATER EQUAL TO AMOUNT OF WATER STANDING ON SOIL SURFACE, MM.
XWHTA(10)	THETA OF LAYER AT END OF PMDT
Y	A VERY TEMPORARY VARIABLE - USED ONLY TO HOLD VALUE OF FUNCTION WTIME.
Z	A DUMMY USED AS AN ARGUMENT IN WTIME
ZEVAP	MM WATER POTENTIALLY EVAPORATED FROM SOIL SURFACE IN PMDT VALUE OBTAINED FROM SUBROUTINE PSWG
ZRAIN	RAINFALL, MM IN PMDT
ZRINT	RAINFALL INTENSITY, MM/HR

SUBROUTINE WATER

```
IF(XHSOLT(2) .LE. -1) TWFRZN=.TRUE.
```

```
WATR 290
```

If temperature of the top soil layer is less than or equal to -1 C, the entire soil profile is considered frozen.

```
DO 50 I=1,6
  DO50 J=1,PMN-2
50 TWTW(I)=TWTW(I)+1.E-5*CVTSR(I,J)
```

```
WATR 319
WATR 320
WATR 322
```

Determine TWTW(I), transpiration demand in centimeters for plant functional group I, for $I = 1$ to 6.

```
IF(.NOT. TWFRZN) GO TO 70
TWRAIN=.FALSE.
TWEVAP=0.0
TWRP=.FALSE.
GOTO130
```

```
WATR 334
WATR 335
WATR 336
WATR 337
WATR 338
```

If soil is frozen, set evaporation equal to zero, set flags indicating there is no precipitation this TWDT, nor was there any last TWDT, and go to section refining TWDT estimate.

```
70 IF((TWR+TWSTND).LE.4.)GOTO80
  TWRAIN = .TRUE.
  TWDT=.02083
  IF(TWRP)GOTO90
  GO TO 120
80 TWRAIN=.FALSE.
90 TWDT=TWDTP
```

```
WATR 339
WATR 340
WATR 341
WATR 342
WATR 343
WATR 344
WATR 345
```

If the amount of water remaining on the surface to be infiltrated is less than or equal to 4 mm, set TWRAIN = FALSE (not raining this TWDT), set TWDT = TWDTP (previous value) and go on to section refining TWDT estimate. If water amount is greater than 4 mm (no rain this TWDT), set TWDT equal to .02083 (= 1/48 day = 30 min = minimum value) and check if rain during last TWDT. If yes, set TWDT = TWDTP and go on to section refining TWDT estimate. If no, keep TWDT at minimum value and skip section refining TWDT estimate.

```
R=0.3*PWTLM
S=0.9*PWTLM
T=0.0
DO110 I=2,PMN-1
  U=ABS(TWDTHP(I))
  IF(U.LT.R)GOTO110
  IF(U.GT.S)GOTO120
  IF(U.GT.T)T=U
110 CONTINUE
  IF(T.LT.R)T=R
  IF(T.LE.0.0)GOTO120
  TWDT=.9*TWDT*(PWTLM/T)
```

```
WATR 352
WATR 353
WATR 354
WATR 355
WATR 356
WATR 357
WATR 358
WATR 359
WATR 360
WATR 361
WATR 362
WATR 363
```

The only circumstances under which we do not reach this point are if TWRAIN = .TRUE. and TWRP = .FALSE. (The most common example of this would be that the surface was dry last TWDT but rain has occurred since.) The point of this section is to increase TWDT as much as possible within the constraint that theta does not change more than PWTLM during TWDT. R, S and T are defined as shown. Each pass through DO loop (one pass per soil layer), U is set equal to magnitude of change in theta during previous TWDT. If U is ever found to be greater than $0.9 * PWTLM$, remainder of this section is skipped (TWDT

retains present value). At the completion of the loop plus the next statement, T is the larger of: 1) $0.3 * PWTLIM$ and 2) the largest value of U (maximum value $0.9 * PWTLIM$). After a precautionary check for zero value of T, TWDT is increased by the factor $PWTLIM/T$. The factor 0.9 is a precautionary measure so TWDT isn't increased too much because if TWDT is set too large here, a later check may reduce it.

```
120 IF(TWDT .GT. (PMDT-THTTOT)) TWDT=PMDT-THTTOT
```

WATR 365

Check the sum of TWDT's to be sure it isn't greater than PMDT.

```
130 CONTINUE
```

WATR 375

Control returns to this point if TWDT is reduced in the line below.

```
IF(TWFRZN) GO TO 150
```

WATR 380

If soil is frozen, skip potential surface flux calculation.

```
TW=AMEN1(TWR,(ZRINT*TWDT*24.))
```

WATR 391

The contribution to the surface flux from precipitation, TW, is the minimum of the following: 1) $ZRINT * TWDT * 24$ (which is the amount of precipitation in TWDT at rate ZRINT); and 2) TWR, the remaining precipitation to infiltrate into the soil. At the end of each TWDT, $ZRINT * TWDT * 24$ is subtracted from TWR (line 801). Thus, line 391 is necessary to ensure that the sum of $ZRINT * TWDT * 24$ does not become larger than ZRAIN, total precipitation for the time-step.

```
TWSF=TW+TWSTND-TWEVAP*TWDT/PMDT
140 TWSF=TWSF*0.1
```

WATR 392
WATR 395

Potential surface flux, in centimeters, equals TW from line 391 above, plus standing water, minus that portion of evaporation for PMDT which occurs this TWDT.

```
150 TWSF=0.0
```

WATR 397

We reach here only if soil is frozen (line 380). If frozen, then set surface flux equal to zero and go on.

```
R=TWDT/PMDT
CO1701=1,PMN-2
170 TWTSPR(I)=0.0
DO 210 J=1,6
  TWTK=0.0
  DO180I=1,PMN-2
    TWTA(I)=TWTHA(I+1)-PWLLIM
180 TWTK=TWTK+TWTA(I)*CVRDST(J,I)
  IF(TWTK .LE. 1.E-8) GO TO 190
  TWTK=R+TWTK(J)/TWTK
190 CONTINUE
  DO200I=1,PMN-2
200 TWTSPR(I)=TWTSPR(I)+TWTK+TWTA(I)*CVRDST(J,I)
210 CONTINUE
```

WATR 413
WATR 414
WATR 415
WATR 418
WATR 419
WATR 420
WATR 422
WATR 423
WATR 426
WATR 427
WATR 428
WATR 429
WATR 430
WATR 431

Determine $TWTSPR(I)$, centimeters, transpiration demand by layer. It is proportional to water available in the layer $TWTA(I)$, and the fraction of roots in the layer $CVRDST(J,I)$. Transpiration demand is assumed uniformly distributed over PMDT.

Thus, each TWDT, the demand is $R = TWDT/PMDT$ times the total demand. Line 418 begins the main DO loop

over plant functional groups (FG's). TWTK first becomes the sum over layers of water available in the layer times fraction of roots of FG J in the layer. Water available is dimensionless and equal to the value of theta at the beginning of TWDT, TWTHA(I+1), minus the lower limit of theta, PWLLIM. The indexing looks strange because TWTHA is indexed over nodes, and TWTA over layers (see Fig. C-2). Line 426 will be true if there is no FG represented by current value of J. (For Curlew Valley simulation, e.g., perennial FG's fill places 1, 2, 3 in XVPLNT, CVRDST, etc., and annual FG's fill places 5, 6, leaving place 4 vacant.) In line 427, TWTK becomes the proportionality factor needed. Lines 429, 430 are a DO loop over soil layers, where, in each pass through the loop, the transpirational demand due to roots of FG J in layer I is added to the demand from layer I. The algebra behind these lines of code is not obvious so more explanation is necessary. If one puts the equation in lines 423 and 427 into line 430 and writes it in normal algebraic form we get:

$$\Delta[\text{TWTSPR}(I)]_J = \frac{\text{TWDT}}{\text{PMDT}} [\text{TWTW}(J)] \frac{[\text{TWTA}(I)] [\text{CVRDST}(J,I)]}{\sum_I [\text{TWTA}(I)] [\text{CVRDST}(J,I)]}$$

where $\Delta[\text{TWTSPR}(I)]_J$ means the addition to $\text{TWTSPR}(I)$ from FG J. If we sum both sides over I (layers), we find that the total transpirational demand from FG J for this TWDT is $(\text{TWDT}/\text{PMDT})[\text{TWTW}(J)]$, which is just what we want since $\text{TWTW}(J)$ is the demand from FG J over PMDT.

```

TH=0.9
D0230I=1,PMN=2
S=TWTSPR(I)/CHDX(I)
IF(S.LE.(TWTA(I)+0.5))GOTO220
TWTSPR(I)=TWTA(I)+0.5*CHDX(I)
S=TWTSPR(I)
270 CONTINUE
IF(TW.LT.S)TW=S
230 CONTINUE

```

```

WATR 4 34
WATR 4 35
WATR 4 36
WATR 4 38
WATR 4 39
WATR 4 40
WATR 4 41
WATR 4 42
WATR 4 43

```

S = centimeters of transpiration from a layer \div thickness of layer = change in theta for layer from transpiration (line 436). Check all layers and set TW = largest value of S . If S ever represents more than half the water available from a layer, reduce transpiration and S to a value equal to half the water available. After this loop, TWTSPR is actual transpiration. A precaution should be noted here. It is possible for the plant submodel(s) to demand more transpiration than WATER is removing from the soil. For one or two TWDT's, this is not serious, but if it occurs over several consecutive PMDT's, the plants will be growing as if there is more soil water available than actually exists. This situation should, and can, be avoided by keying transpiration to soil water in such a manner that, as water is removed from the soil, transpiration becomes zero before theta reaches PWLLIM.

```

IF((TW.LE.PWLLIM).OR.(TWJM.GE.2).OR.(TWDT.LE.0.02083))GOTO240
TWJM=TWJM+1
TWDT=TWDT+0.9*PWLLIM/TW
IF(TWDT.LE.0.02083)TWDT=0.02083
GO TO 130
240 CONTINUE

```

```

WATR 4 45
WATR 4 46
WATR 4 47
WATR 4 48
WATR 4 49
WATR 4 50

```

TWDT will not be reduced if: 1) TW , calculated just above, is less than or equal to the largest allowable change in theta; 2) if it has already been reduced twice before; or 3)

if it is already at the minimum value of 30 min. If it is too large, we increment the counter, reduce TWDT and ensure that its value is not below minimum.

Overview of Conductivity and Specific Water Capacity Calculation

In order to accurately simulate water content and potential, it is important to determine hydraulic conductivity, PHK, as accurately as possible. Average PHK for a region between nodes (see Fig. C-2) over TWDT can be calculated if theta is known at the beginning and end of TWDT. But to determine theta at the end of TWDT, we must use average PWK during TWDT. Thus, an iteration scheme would appear to be in order. With a large subroutine such as this, however, it appears likely that it would be easy to set something up which would eat up considerable computer time. To avoid excessive computer-time use, no iterations are performed and, whenever possible, the change in theta for the current TWDT is estimated from the change in theta during the previous TWDT. "Whenever possible" turns out to be over 90% of the time.

```
IF(TWONE) GO TO 280
TWONE=.TRUE.
```

```
WATR 466
WATR 467
```

Use the flag TWONE to ensure going through this section only once each TWDT.

```
IF((TWRP .AND. TWRAIN) .OR. ((.NOT. TWRP) .AND. (.NOT. TWRAIN)))
1 GO TO 250
GO TO 280
250 CONTINUE
```

```
WATR 468
WATR 469
WATR 470
WATR 471
```

If there was rain last TWDT and it is raining now, or if no rain occurred last TWDT and there is no rain now, we can estimate the change in theta this TWDT from change last TWDT. Otherwise, go to the section calculating conductivity and specific water capacity.

```
DO270I=2,PMN
```

```
WATR 472
```

DO loop through nodes (all except top one).

```
S=TWDT/TWDTP
IF(S.GT.1.)S=SQRT(S)
TW=TMTHB(I)+S*TWDTHP(I)
IF(TW .GE. PWT(PMN)) GO TO 260
IF(TW .LT. PMLLIM) TW=PMLLIM
```

```
WATR 474
WATR 475
WATR 476
WATR 477
WATR 478
```

Tentative change in theta is equal to change last TWDT times S. S is the ratio of present and past time-steps if ratio ≤ 1 . If ratio > 1 , S equals the square root of the ratio as a precaution against overestimating change in theta. Finally, perform checks to ensure that theta is within limits.

```
TWTHB(I)=TW
```

WATR 479

Set theta for current node equal to tentative value.

```
IF(I.GT.2)GOTO270
TWJA=TWTHB(I) / PWDELW + 1.
TWAA=(TWTHB(I)-PWT(TWJA))/PWDELW
TWHB(I)=PWH(TWJA)+(PWH(TWJA+1)-PWH(TWJA))*TWAA
GO TO 270
```

WATR 480
WATR 483
WATR 485
WATR 486
WATR 487

If current node is below second node, skip these lines. Otherwise, calculate TWHB, SWP at end of TWDT, for first layer (second node) by interpolating between appropriate values in PWH table. It is used below in the calculation for actual surface flux (line 574).

```
DO 320I=2,PMN-1
```

WATR 494

DO loop over PMN—2 inner nodes. In this loop, calculate specific water capacity of each of the PMN—2 layers and hydraulic conductivity of the PMN—3 regions between the inner nodes.

```
TWTHC=(TWTHA(I)+TWTHB(I)) / 2.
IF(TWTHC.GT. PWT(PMN)) TWTHC=PWT(PMN)
IF(TWTHC.LT. PWH(I)) TWTHC=PWH(I)
```

WATR 495
WATR 496
WATR 497

Calculate average water content during TWDT and ensure that values lie between limits.

```
TWJA=TWTHC/PWDELW + 1.
IF(TWJA.GE. PMN) TWJA=PMN-1
TWAA=(TWTHC-PWT(TWJA))/PWDELW
TWKSA(I)=PWKSUM(TWJA)+(PWKSUM(TWJA+1)-PWKSUM(TWJA))*TWAA
TWHHA(I)=PWH(TWJA)+(PWH(TWJA+1)-PWH(TWJA))*TWAA
```

WATR 499
WATR 500
WATR 502
WATR 504
WATR 505

Using value of average water content just found, interpolate between appropriate values in PWKSUM and PWH tables, and call these values TWKSA and TWHHA.

```
PWKC(I-1) = PWDELW/(PWH(TWJA+1)-PWH(TWJA))
```

WATR 508

The specific water capacity is simply the inverse of the slope of the SWP vs. theta relationship.

```
IF(I.LE.2)GOTO310
```

WATR 509

Need TWKSA and TWHHA for two adjacent nodes to calculate PWK, hydraulic conductivity, for region between.

```
IF(TWJB.EQ. TWJA) GO TO 300
```

WATR 510

If average water content of previous node and current node are close (between same pairs of values in PWKSUM and PWH table), calculate PWK differently. This avoids any chance of a divide-by-zero (or nearly zero) situation.

PWK(I-1)=(TWKSA(I)-TWKSA(I-1))/(TWHHA(I)-TWHHA(I-1))

WATR 511

Calculate hydraulic conductivity. Hanks and Bowers (1962) call it a "round about" method in that diffusivity is calculated from measured conductivity values, and then average conductivity is determined from diffusivity. If one goes through their algebra, PWK is seen to be average conductivity over a region of soil, weighted with respect to SWP. PWKSUM, calculated in line 1013, is

$$PWKSUM(M) = \sum_{i=1}^M K_i (\Delta SWP)_i,$$

where

- i = index which goes through theta values in the conductivity and SWP tables ($i = 1 \Rightarrow$ theta = PWTLM; $i = 2 \Rightarrow$ theta = 2·PWTLM, . . . ; $i = m \Rightarrow$ theta = M·PWTLM);
- K_i = measured conductivity (read in as PWKIN) at theta = $i \cdot PWTLM$;
- $(\Delta SWP)_i$ = change in soil water potential (from PWH array) from theta = $i \cdot PWTLM$ to theta = $(i + 1) \cdot PWTLM$.

If theta at the top of the region corresponds to $M \cdot PWTLM$ and theta at the bottom corresponds to $N \cdot PWTLM$, then rewrite line 511:

$$\text{average conductivity} = \frac{\sum_{i=1}^M K_i (\Delta SWP)_i - \sum_{i=1}^N K_i (\Delta SWP)_i}{SWP_M - SWP_N}$$

Say $M > N$, then

$$\text{average conductivity} = \frac{\sum_{i=N}^M K_i (\Delta SWP)_i}{\sum_{i=N}^M (\Delta SWP)_i}$$

Thus, average conductivity for a region is an average over measured conductivity, weighted with respect to SWP, as asserted above.

To illustrate, do a numerical example with theta at the node at the top of the region = 0.25 and = 0.15 at the bottom. Use input data for Curlew Valley run.

theta	0.15	0.25
TWKSA	0.4047	0.9157
TWHHA	-8.090	-.8800
PWKIN	0.0198	0.4630

$$\begin{aligned} \text{weighted average conductivity} &= \frac{TWKSA_{.15} - TWKSA_{.25}}{TWHHA_{.15} - TWHHA_{.25}} \\ &= \frac{0.4047 - .9157}{-8.090 - (-.8800)} = 0.0709. \end{aligned}$$

$$\text{unweighted average conductivity} = \frac{0.01980 + 0.4630}{2} = 0.2414.$$

The weighted average value, which is the one used in the program, is considerably smaller than the unweighted average. This is consistent with the findings of Hanks and Bowers (1962) that the dry soil (lower conductivity) controls the flow of water more than the wetter soil.

```
300 FHK(I-1)=(PWKSUM(TWJA+1)-PWKSUM(TWJA)) / (PWH(TWJA+1)-PWH(TWJA))  MATR 513
```

If theta values at nodes outlining the current region are between the same pairs of values in PWKSUM and PWH tables, use this formula which is a special case of the one in line 511 and avoids the possibility of dividing by zero or a very small number.

```
310 THJB=TWJA
```

MATR 514

We need to know previous value of TWJA on the next pass through the loop.

```
320 CONTINUE
```

MATR 517

End of loop.

Overview of Top Boundary Condition Section

Potential top surface flux was calculated in line 392. Here, we check if we can meet potential by setting soil surface conditions to wettest conditions (if potential flux > 0, or into soil) or to driest (if potential flux < 0). If potential flux can be met, actual flux is set equal to potential flux and program goes on. Since no submodel, including WATER, uses soil surface conditions, the surface node is always either at its wettest or driest condition, and more precise values are not determined. If potential flux cannot be met, actual flux is calculated, and difference between actual and potential is determined. If this difference is positive, it ultimately becomes standing water (line 794). If negative, it eventually is lost.

```
IF(TMSF.LT.0.0)GOTO330
TWHB(1)=PWHWET
TWHC(1)=PWT(PWM)
GOTO340
320 CONTINUE
TWHB(1)=PWHDRY
TWHC(1)=PWLLIM
```

MATR 540
MATR 544
MATR 545
MATR 546
MATR 548
MATR 551
MATR 552

Check sign of surface flux and set surface conditions wettest if positive, driest if negative.

```
340 CONTINUE
C CALCULATE PWK(1)
TWHC=TWHB(1)
IF(TWHC.GT.PWT(PWM)) TWHC=PWT(PWM)
```

MATR 554
MATR 555
MATR 556
MATR 557

```

      IF(TWTHC .LT. PWLLIM) TWTHC=PWLLIM           WATR 558
C  PWT(TWJA) AND PWT(TWJA+1) BRACKET TWTHC       WATR 559
      TWJA=TWTHC/PWDELW + 1.                       WATR 560
      IF(TWJA .GE. PMN) TWJA=PMN-1                WATR 561
C  CHECK IF WATER CONTENTS OF FIRST TWO NODES ARE VERY CLOSE WATR 562
      IF((TWKSA(2) .GE. PWKSUM(TWJA)) .AND. (TWKSA(2) .LE. PWKSUM(TWJA+1)
1 ))) GO TO 350                                     WATR 563
C  THAA IS FOR INTERPOLATING                       WATR 564
      THAA=(TWTHC-PWT(TWJA))/PWDELW                WATR 565
      TWKSA(1)=PWKSUM(TWJA)+(PWKSUM(TWJA+1)-PWKSUM(TWJA))*THAA WATR 566
      TWHFA(1)=TWHB(1)                             WATR 567
      PWK(1)=(TWKSA(2)-TWKSA(1))/(TWHHA(2)-TWHHA(1)) WATR 568
      GO TO 360                                       WATR 569
350      PWK(1) = (PWKSUM(TWJA+1)-PWKSUM(TWJA)) / (PMH(TWJA+1)
1 -PMH(TWJA))                                       WATR 570
360 CONTINUE                                         WATR 571

```

Determine PWK(1) just as PWK(I) for I = 2 to PMN—2 were determined in lines 494-517. The only difference is that water content is not averaged in line 556 as it was in line 495 (averaging surface node led to nasty oscillations).

```

      TWSFC=PWK(1)*TWDI*(TWHB(1)-TWHB(2)+9.833E-4*CHDXX(1))/CHDXX(1) WATR 574

```

Calculated surface flux equals conductivity times time-step times SWP gradient [the term $9.833 \times 10^{-4} \times \text{CHDXX}(1)$ is contribution due to gravity].

```

      IF(TWSF.LT.0.0)GOTO370                         WATR 575
      IF(TWSFC.GE.TWSF)GOTO390                       WATR 576
      GOTO380                                         WATR 577
370 IF(TWSFC.LE.TWSF)GOTO390                         WATR 578

```

Check if potential surface flux condition can be met.

```

380 CONTINUE                                         WATR 580
      TWIN=.FALSE.                                   WATR 581
      TWSFD=TWSF-TWSFC                              WATR 582
      TWSF=TWSFC                                     WATR 583
      GOTO400                                         WATR 584

```

If we reach here, they can't be met. Set flag announcing this, determine difference between potential and calculated fluxes, set surface flux equal to calculated value and go on.

```

390 CONTINUE                                         WATR 587
      TWIN=.TRUE.                                    WATR 591
      TWSFD=0.0                                      WATR 592
400 CONTINUE                                         WATR 594

```

If we reach here, potential surface flux condition can be met. Set flag announcing this, set difference between potential and actual equal to zero and go on.

```

      IF(.NOT. TWFRZN) GO TO 420                     WATR 600
      EO410I=1,PMN-2                                WATR 601
410 PWK(I)=PWK(I)*1.E-6                             WATR 602
420 CONTINUE                                         WATR 603

```

If soil is frozen, reduce hydraulic conductivity by a factor of a million and stop soil water movement.

```

EO 430 I=1, PMN-1
TWA(I) = PWK(I) / CHDXX(I)
IF(I .EQ. PMN-1) GO TO 430
TWHB(I) = (2.*PWNC(I)*CHDX(I)) / TWDT
430 CONTINUE
DO 440 I=2, PMN-3
TWA(I) = TWA(I+1)
TWC(I) = TWA(I)
TWH(I) = TWHB(I)+TWA(I)+TWC(I)
440 TWD(I) = TWA(I+1)*(TWHB(I)-TWA(I)-TWC(I)) + TWA(I)*TWC(I)
1 + TWA(I+2)*TWA(I) + 2.*(PWK(I) - PWK(I+1)) *9.833E-4
2 -2.*TWTSPR(I)/TWDT
TWA(1) = TWA(2)
TWC(1) = 0.
TWH(1) = TWHB(1) + TWA(1)
TWD(1) = TWA(2)*(TWHB(1)-TWA(1)) + TWA(3)*TWA(1)
1 + TWSF*2./TWDT -2.*PWK(2)*9.833E-4
2 -2.*TWTSPR(1)/TWDT
TWA(PMN-2) = 0.
TWC(PMN-2) = TWA(PMN-2)
TWH(PMN-2) = TWHB(PMN-2) + TWC(PMN-2) + TWA(PMN-1)
TWD(PMN-2) = TWA(PMN-1)*(TWHB(PMN-2)-TWC(PMN-2)-TWA(PMN-1))
1 + TWA(PMN-2)*TWC(PMN-2) + 2.*TWHB(PMN)*TWA(PMN-1)
2 +2.*(PWK(PMN-2)-PWK(PMN-1)) *9.833E-4
3 -2.*TWTSPR(PMN-2)/TWDT

```

```

WATR 641
WATR 642
WATR 643
WATR 644
WATR 645
WATR 647
WATR 648
WATR 649
WATR 650
WATR 651
WATR 652
WATR 653
WATR 655
WATR 656
WATR 657
WATR 658
WATR 659
WATR 660
WATR 661
WATR 662
WATR 663
WATR 664
WATR 665
WATR 666
WATR 667

```

Generate matrix elements of tridiagonal matrix. These elements are coefficients of SWP in the system of difference equations (one for each layer) we need to solve to determine SWP of each layer. For more information see "Derivation of Difference Equations." These matrix elements can be determined only after we've determined: 1) hydraulic conductivity; 2) specific water capacity; 3) transpiration; 4) top surface flux.

```
CALL FDM(TWA, TWH, TWC, TWD, TWU, PMN-2)
```

```
WATR 677
```

TDM is the subroutine which solves the tridiagonal matrix, generated above, for PMN—2 values of SWP. These values are returned to WATER in array TWU. For more details, see the Detailed Program Description of TDM.

```

DO450 I=2, PMN-1
TWTHB(I)=TWA(I)+PWNC(I-1)*(TWU(I-1)-TWA(I))
IF(TWTHB(I) .GT. PWT(PMN)) TWTHB(I)=PWT(PMN)
IF(TWTHB(I) .LT. PMLLIM) TWTHB(I)=PMLLIM
450 CONTINUE

```

```

WATR 690
WATR 691
WATR 692
WATR 693
WATR 694

```

Calculate TWTHB, theta of nodes at end of TWDT, from values of TWU just returned from subroutine TDM. Equation used is

$$(\text{theta}) = (\text{specific water capacity}) \times [\Delta(\text{SWP})].$$

Ensure that values obtained are within range.

```

EO460 I=2, PMN-1
TWJA=TWTHB(I)/PWDELW+1.
IF(TWJA .GE. PNM) TWJA=PNM-1
TWAA=(TWTHB(I)-PWT(TWJA))/PWDELW
460 TWHB(I)=PWH(TWJA)+(PWH(TWJA+1) - PWH(TWJA))*TWAA

```

```

WATR 697
WATR 698
WATR 699
WATR 700
WATR 701

```

Calculate TWHB, SWP of nodes at end of TWDT, from values of theta just calculated (same method as that starting in lines 495 and 556). This is roundabout in that we could have set TWHB(I) = TWU(I—1) for I = 2, PMN—1. By using the specific water capacity, we find theta much faster than by hunting backwards through array PWH, but we

run the risk of the two values not exactly coinciding because of the various averages used. By determining SWP from theta rather than directly from TWU, then, small errors might occur, but theta and SWP will be consistent with each other.

```

DO470I=2,PMN-1
S=ABS(TWTHB(I)-TWTHA(I))
IF(S.GT.PWTLIM)GOTO480
470 CONTINUE
GO TO 510

```

```

WATR 705
WATR 706
WATR 707
WATR 708
WATR 709

```

Check if any of the delta thetas just calculated are larger than the limit PWTLIM. If yes, see if TWDT can be reduced. If no, go on to check of water balance.

```

4*0 IF(TWJM.GE.2)GOTO510
TWDTQ=TWDT
TWDT=TWDT*0.9*PWTLIM/S
TWJM=TWJM+1
IF(TWDT.LT.0.02083)TWDT=0.02083
R=TWDT/TWDTQ
490 DO500I=2,PMN-1
TWHB(I)=TWHA(I)+(TWHB(I)-TWHA(I))*R
500 TWTHB(I)=TWTHA(I)+(TWTHB(I)-TWTHA(I))*R
GO TO 130
510 CONTINUE

```

```

WATR 710
WATR 711
WATR 712
WATR 713
WATR 715
WATR 716
WATR 717
WATR 718
WATR 719
WATR 720
WATR 721

```

Reduce TWDT if: 1) it hasn't been reduced two or more times already (limit the number of iterations to cut down on computer time); and 2) if it isn't already at its lower limit of 30 min (0.02083 days). In line 712, reduce TWDT; the factor 0.9 is to be on the safe side. Then increment the iteration counter and make sure TWDT is not less than 30 min. Finally, reduce the estimates of SWP and theta just calculated above by making the change proportional to the ratio of new TWDT to old TWDT, and go back to statement 130 (line 375) which is the start of the loop to calculate SWP once TWDT is set.

```

TWSFB=-PWKCAL*TWDT*((TWHB(PMN-1)+TWHA(PMN-1))*0.5
1 +9.833E-4*CHDXX(PMN-1)-TWHB(PMN))/CHDXX(PMN-1)

```

```

WATR 726
WATR 727

```

Calculate water leaving bottom of soil profile (remember that convention is: in is +). This is basically the same equation as line 574.

```

CALL WBAL(TWSPR,TWSF,TWSFB,TWTHA,TWTHB,CHOX,PMN,
1 TWMIN,TWSTRD,R)

```

```

WATR 730
WATR 731

```

Call subroutine WBAL to check if the sum of the amounts of water into or out of the profile (rain, evaporation, transpiration, water through caliche layer) is equal to the change in the water stored in the profile. For more detail see "Subroutine WBAL," below.

```

IF(TWDT.LT.0.05)GOTO520
IF((R.GT.99.)OR.(ABS(TWMIN-TWSTRD).LT.0.006)
1 .OR.(ABS(R-1.).LT.0.01))GOTO520
TWDT=0.5*TWDT

```

```

WATR 733
WATR 736
WATR 737
WATR 738

```



```

R=0.5
GO TO 490
520 CONTINUE

```

```

WATR 739
WATR 740
WATR 741

```

If the amount of water into or out of the profile is equal to the change in water stored, TWDT remains unchanged. If they are not equal, TWDT may be halved. TWDT will not be halved if its value is currently small (less than .05). Also, it will not be halved if one of the following conditions is met: 1) if R, the ratio of water in to water stored, is greater than 99 (to avoid a divide by zero, WBAL sets R = 99.9999 if TWSTRD is less than 0.001); 2) if TWWIN is sufficiently close to TWSTRD (within 0.006); 3) if R is sufficiently close to 1.0 (within 0.01).

```

CWINF=CWINF+TWSF
TWSFOT = TWSFBT + TWSFB

```

```

WATR 778
WATR 782

```

Accumulate top and bottom surface fluxes. Neither flux includes transpiration.

```

IF(.NOT. TWRFRZ) GO TO 570
TWSTND=TWSTND+TWR
TWR=0.0
GO TO 620
570 CONTINUE

```

```

WATR 785
WATR 786
WATR 787
WATR 788
WATR 789

```

If soil is frozen, add precipitation to standing water and skip the normal standing water calculation.

```

IF(TMIN) GO TO 580
TWSTND=TWSFD*10.*PWRNRF**TWDT
IF (TWSTND .LT. 0.0) TWSTND=0.0
GO TO 590
580 TWSTND=0.0
590 CONTINUE

```

```

WATR 792
WATR 794
WATR 795
WATR 796
WATR 797
WATR 798

```

If potential surface flux conditions are met there will be no standing water. If unmet, then in line 794 calculate new level of standing water using parameter PWRNRF, which provides a crude estimate of runoff or runoff and will have a value near 1. (A value of 1.25 for Curlew Valley, e.g., made water continually build up on a surface once the spring melting occurred.) TWSFD is the potential surface flux minus the actual surface flux; the factor 10 changes units to millimeters; and PWRNRF to the TWDT power is a compound interest type factor.

```

600 TWR=TWR-ZRINT+TWDT*24.
610 IF(TWR .LT. 0.0) TWR=0.0

```

```

WATR 801
WATR 802

```

Determine the amount of precipitation remaining to infiltrate this PMDT. Make sure it doesn't get negative.

```

XWSTND=TWSTND

```

```

WATR 805

```

Set state variable XWSTND equal to TWSTND.

```

DO 640 I=1,PMN-2
XWTHTA(I)=TWTHB(I+1)
640 CWPSI(I)=TWHB(I+1)

```

```

WATR 826
WATR 827
WATR 828

```

Set state variable XWTHTA, theta of layers, equal to theta of appropriate node at end of PMDT (remember layer 1 is centered on node 2). Set communication variable CWPSI, SWP of layers, equal to SWP of appropriate node at the end of PMDT.

```

ENTRY WINIT

```

```

WATR 858

```

This entry point is called once, from MAIN. It is used only for initialization purposes: 1) to read in WATER variables; 2) to initialize miscellaneous flags and variables; 3) to perform calculations which need to be done only once.

```

READ(5,670)TWCCC
WRITE(6,680)
WRITE(6,690)TWCCC

```

```

WATR 870
WATR 871
WATR 872

```

Read and write a comment which may say, e.g., INITIALIZATION DATA FOR WATER.

```

READ(5,670)TWCCC
WRITE(6,690)TWCCC
READ(5,7)PWDELW
WRITE(6,710)PWDELW

```

```

WATR 875
WATR 876
WATR 877
WATR 878

```

Read and write a comment which describes variable PWDELW, which is then read and written. In similar manner, variables PWH, PWKCAL, PWKIN, PWHDRY, PWHWET, PWM, PWTLIM, PWLLIM, PWRNRF, TWSTND and TWTHB are read and written.

```

DO 750 I=1, PWM
TWI = I-1
750 PWT(I)=TWI+PWDELW

```

```

WATR 999
WATR1000
WATR1001

```

Set up array PWT, containing PWM values of theta, each PWDELW apart. The conductivity in PWKIN(I) and the SWP in PWH(I) correspond to a theta value of PWT(I).

```

DO 760 I=1,PMN
TWJA=TWTHB(I)/PWDELW + 1.
IF(TWJA .GE. PMN) TWJA=PMN-1
TWAA=(TWTHB(I)-PWT(TWJA)) / PWDELW
760 TWHB(I)=PWH(TWJA)+(PWH(TWJA+1)-PWH(TWJA))*TWAA

```

```

WATR1004
WATR1005
WATR1006
WATR1007
WATR1008

```

Calculate initial SWP values from theta values just read in TWTHB. Approach used is same as in line 483.

```

PWKSUM(1)=(PWKIN(1)+PWKIN(2))*(PWH(2)-PWH(1))*0.5
DO 770 I=2,PMN-1
PWKSUM(I)=(PWKIN(I)+PWKIN(I+1))*(PWH(I+1)-PWH(I))*0.5 +PWKSUM(I-1)
PWKSUM(PMN)=2.*PWKSUM(PMN-1)-PWKSUM(PMN-2)

```

```

WATR1011
WATR1012
WATR1013
WATR1017

```

Calculate PWKSUM, sum of conductivity times delta SWP. This is used in line 504 in average conductivity calculations. Each element in PWKSUM equals the previous element plus an increment. Take PWKSUM(I), e.g. It equals PWKSUM(I-1) plus the increment, which is the

average conductivity for theta between values of PWT(I) and PWT(I+1), times the change in SWP for theta between values of PWT(I+1) and PWT(I). In line 1017, the last increment is taken equal to the second to last increment.

```

WRITE(6,760)
700 FORMAT(' PWKSUM, SUM OF CONDUCTIVITY TIMES DELTA PRESSURE')
WRITE(6,720) PWKSUM

```

```

WATR1020
WATR1021
WATR1022

```

The array PWKSUM, which was just calculated, is written out.

```

READ(5,670)TWCCC
WRITE(6,690)TWCCC

```

```

WATR1025
WATR1026

```

Read and write a comment which says, e.g., END INITIALIZATION FOR WATER.

SUBROUTINE WBAL

WBAL is a small, simple subroutine which acts as an overall check on the accuracy of the calculations performed in WATER. For the time period in question, either TWDT or PMDT, it calculates WIN, amount of water into or out of the soil profile by transpiration, and top and bottom surface flux. It also calculates WSTRD, change in water stored during time period, from the changes in theta of each layer and from layer thickness. WBAL then determines the ratio of these two quantities. The calling program, WATER, in this case, then determines if the ratio is close enough to 1.

```

DO100I=1,PMN=2
A=A+TSPR(I)
100 WSTRD=WSTRD+(THTAF(I+1)-THTAI(I+1))*DX(I)

```

```

WBAL 16
WBAL 17
WBAL 18

```

This is a DO loop through layers. Variable A accumulates transpiration, WSTRD accumulates changes in water held in layers.

```

WIN = SF - A + SFB

```

```

WBAL 19

```

Calculate WIN, change in water in profile from surface fluxes (top and bottom) and transpiration (convention is that water in is positive).

```

IF(ABS(WSTRD),LT,1,E=3)RATIO=99,9999

```

```

WBAL 20

```

Set RATIO equal to a strange large number if WSTRD is small (avoid divide by zero or near zero).

```

IF(ABS(WSTRD),GE,1,E=3)RATIO=WIN/WSTRD

```

```

WBAL 21

```

Actually calculate RATIO if WSTRD is large enough.

FUNCTION WTIME

The purpose of this function is to determine total elapsed CPU time since the beginning of the run. Because we use the variable name TIME for other purposes in the model, the Burroughs Intrinsic Function TIME must be put in a subprogram of some kind in order for it to be used. The Z in the argument of WTIME is a completely dummy variable -- a FORTRAN function needs an argument whether it is used or not.

```
WTIME=0.016667*TIME(2)
```

```
WTM 08
```

TIME(2) is CPU time in 60ths of a second. The factor $0.016667 = 1/60$ changes the right-hand side to seconds. CPU time is useful for debugging purposes.

COMPLETE PROGRAM LISTING

SUBROUTINE WATER

```

SUBROUTINE WATER                                WATR 001
C                                               WATR 002
C WRITTEN BY PAUL W. LOMMEN                     WATR 003
C   DESERT BIOME - ECOLOGY CENTER UMC 52       WATR 004
C   UTAH STATE UNIVERSITY                     WATR 005
C   LOGAN, UTAH 84322                        WATR 006
C                                               WATR 007
C   AUGUST 1976                              WATR 008
C                                               WATR 009
C THIS SUBMODEL OF THE WATER RESPONSE MODEL   WATR 010
C POTENTIAL IN UP TO 8 SOIL LAYERS.  INPUTS   WATR 011
C NEEDED ARE PRECIPITATION,                   WATR 012
C POTENTIAL EVAPORATION, TRANSPIRATION,      WATR 013
C SOIL HYDRAULIC CONDUCTIVITY -              WATR 014
C SOIL WATER CONTENT RELATIONSHIP, AND SOIL  WATR 015
C WATER CONTENT RELATIONSHIP.                WATR 016
C                                               WATR 017
C VVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVV  WATR 018
C VVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVVV  WATR 019
C LIST OF VARIABLES FOR WATER
C                                               WATR 020
C CHD(10)  DEPTHS, CM, OF NODES BELOW SURFACE (A NODE IS A POINT WATR 021
C          WHERE A WATER POTENTIAL IS CALCULATED) STARTING AT SURFACE WATR 022
C                                               WATR 023
C CHDX(10) THICKNESS OF REGION CENTERED ON NODE WHERE COUNTING WATR 024
C          STARTS AT FIRST NODE BELOW SURFACE AND GOES DOWN WATR 025
C                                               WATR 026
C CHDXX(10) CHDXX(I)=CHD(I+1) - CHD(I) WATR 027
C                                               WATR 028
C CVTSPR(10) MM OF WATER TAKEN FROM EACH LAYER DURING PMDT WATR 029
C                                               WATR 030
C CWINF    NET TOP SURFACE FLUX FOR PMDT, CM, DOES WATR 031
C          NOT INCLUDE TRANSPIRATION. WATR 032
C                                               WATR 033
C CWPSI(10) SOIL WATER POTENTIAL, BARS, VALUE AT END OF PMDT WATR 034
C                                               WATR 035
C DBGFLG   IF=.TRUE. WRITE DEBUGGING INFORMATION THIS TIME STEP WATR 036
C                                               WATR 037
C DEBUG1   NUMBER OF TIME STEPS BETWEEN DEBUGGING OUTPUTS WATR 038
C                                               WATR 039
C FLGFLG   IF=.TRUE. WRITE DEBUGGING INFORMATION THIS TWDI WATR 040
C                                               WATR 041
C PMDT     TIME STEP DETERMINED IN MAIN, DAYS WATR 042
C                                               WATR 043
C PMN      NUMBER OF NODES, SAME AS NUMBER OF NODES IN HEAT, WATR 044
C          COUNTING STARTS AT SURFACE WATR 045
C                                               WATR 046
C PWDELH   INCREMENT IN VMC, TYPICALLY 0.01-0.02 WATR 047
C                                               WATR 048
C PWH(60)  TABLE OF HYDRAULIC PRESSURE HEAD VERSUS THETA, BARS, WATR 049
C          USING SAME THETA SCALE AND SPACING AS PWKIN WATR 050
C                                               WATR 051
C PWHLRY   ALLOWABLE LOW PRESSURE LIMIT FOR ANY LAYER WATR 052
C                                               WATR 053
C PWHHET   ALLOWABLE HIGH PRESSURE LIMIT FOR ANY LAYER WATR 054
C                                               WATR 055
C PWK(10)  HYDRAULIC CONDUCTIVITY, CM2 BAR-1 DAY-1, PWK(I) IS WATR 056
C          CONDUCTIVITY AVERAGED OVER TWDI, IN REGION BETWEEN NODES WATR 057
C          I AND I+1, WHERE COUNTING STARTS AT SURFACE. WATR 058
C                                               WATR 059
C PWKCAL   CONDUCTIVITY OF CALICHE LAYER WATR 060
C                                               WATR 061
C PWKIN(60) TABLE OF CONDUCTIVITY VERSUS THETA, CM2 BAR-1 DAY-1, WATR 062
C          STARTING WITH VALUE AT THETA = 0.0, READ IN IN WINIT. WATR 063
C                                               WATR 064
C PWKSUM(60) ARRAY USED TO CALCULATE AVERAGE K, (CM-BARS)/DAY WATR 065
C                                               WATR 066
C PWLLIM   LOWER LIMIT OF THETA, DIMENSIONLESS WATR 067
C          CALCULATED IN WINIT AND = SUM OF CONDUCTIVITY X DELTA H WATR 068
C                                               WATR 069
C PMN      NUMBER OF ENTRIES IN CONDUCTIVITY AND PRESURE TABLES WATR 070
C                                               WATR 071
C PWRNRF   RUNON TO RUNOFF RATIO PER DAY WATR 072
C                                               WATR 073
C PWT(60)  ARRAY CONTAINING UNIFORMLY DISTRIBUTED VALUES OF THETA WATR 074
C                                               WATR 075
C PWTLM    LARGEST CHANGE IN THETA ALLOWED FOR NODES 2 TO PMN-1 WATR 076
C          DURING ANY TIME STEP TWDI. TYPICAL VALUE .01 TO .02 WATR 077
C                                               WATR 078
C PWHC(10) WATER CAPACITY, BAR-1, OF REGIONS SURROUNDING NODES, WATR 079
C          COUNTING STARTS AT FIRST NODE BELOW SURFACE WATR 080
C                                               WATR 081
C R        A VERY TEMPORARY VARIABLE - USED IN REFINED TWDI ESTIMATE, WATR 082
C          TRANSPIRATION SECTION, REDUCTION OF TWDI IF DELTA THETA WATR 083
C          TOO LARGE, AND AS RATIO OF WATER IN TO WATER STORED WATR 084

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C		AS CALCULATED IN MBAL.	WATR 085
C			WATR 086
C	S	A VERY TEMPORARY VARIABLE - USED IN REFINED TWDT ESTIMATE,	WATR 087
C		TRANSPIRATION SECTION, CONDUCTIVITY SECTION, AND REDUCTION	WATR 088
C		OF TWDT IF DELTA THETA TOO LARGE.	WATR 089
C			WATR 090
C	T	A VERY TEMPORARY VARIABLE - USED IN REFINED TWDT ESTIMATE.	WATR 091
C			WATR 092
C	TDM	A SUBROUTINE USED BY WATER AND HEAT TO SOLVE A SET OF	WATR 093
C		SIMULTANEOUS DIFFERENCE EQUATIONS, THE LEFT HAND SIDES OF	WATR 094
C		WHICH FORM A TRI-DIAGONAL MATRIX.	WATR 095
C			WATR 096
C	TIME	CURRENT JULIAN DATE - BOOKKEEPING DONE IN MAIN PROGRAM.	WATR 097
C			WATR 098
C	TSTART	JULIAN DATE OF START OF RUN - INITIALIZED IN MAIN PROGRAM	WATR 099
C			WATR 100
C	TWAA	FRACTION OF PWDELW INTERVAL, USED IN INTERPOLATING	WATR 101
C			WATR 102
C	TWAA(10)	PARAMETER A FOR TRI-DIAGONAL MATRIX ROUTINE (TDM ROUTINE)	WATR 103
C			WATR 104
C	TWAAA(10)	ARRAY CONTAINING VALUES OF TWTHA(I) FOR I=2,PMN-1	WATR 105
C			WATR 106
C	TWB(10)	PARAMETER B FOR TDM ROUTINE	WATR 107
C			WATR 108
C	TWBE(10)	CONTAINS ORDERED VALUES OF TWAAA.	WATR 109
C			WATR 110
C	TW	A VERY TEMPORARY VARIABLE - USED IN SURFACE FLUX	WATR 111
C		CALCULATION, TRANSPIRATION SECTION, AND CONDUCTIVITY	WATR 112
C		SECTION.	WATR 113
C			WATR 114
C	TWC(10)	PARAMETER C FOR TDM ROUTINE	WATR 115
C			WATR 116
C	TWCCC	USED TO READ AND WRITE COMMENTS IN INPUT DATA	WATR 117
C			WATR 118
C	TWO(10)	PARAMETER D FOR TDM ROUTINE	WATR 119
C			WATR 120
C	TWDT	ACTUAL TIME STEP WATER IS USING	WATR 121
C			WATR 122
C	TWO7HP(1)	CHANGE IN THETA DURING LAST TWDT	WATR 123
C			WATR 124
C	TWDTP	LENGTH OF PREVIOUS WATER TIME STEP	WATR 125
C			WATR 126
C	TWDTQ	HOLDS CURRENT VALUE OF TWDT AS IT IS BEING REDUCED	WATR 127
C		BECAUSE DELTA THETA IS TOO LARGE.	WATR 128
C			WATR 129
C	TWEVAP	EVAPORATIVE DEMAND, MM, WATER TO EVAPORATE FROM SURFACE	WATR 130
C		DURING REMAINDER OF PMDT	WATR 131
C			WATR 132
C	TWFRZN	EQUALS .TRUE. IF SOIL PROFILE FROZEN, EQUALS	WATR 133
C		.FALSE. IF NOT FROZEN.	WATR 134
C			WATR 135
C	TWFI	TEMPORARY VARIABLE USED TO SET UP PWT ARRAY.	WATR 136
C			WATR 137
C	TWHA(10)	PRESSURE HEAD AT BEGINNING OF TWDT, BARS	WATR 138
C			WATR 139
C	TWHB(10)	PRESSURE HEAD AT END OF TWDT, BARS	WATR 140
C			WATR 141
C	TWHHA	INTERPOLATED VALUES OF PWH	WATR 142
C			WATR 143
C	TWIN	EQUALS .TRUE. IF SURFACE PRESSURE IS WITHIN LIMITS	WATR 144
C			WATR 145
C	TWJA	INTEGER USED IN PICKING VALUES OF K, H OFF TABLES	WATR 146
C			WATR 147
C	TWJB	INTEGER USED IN PICKING VALUES OF K, H OFF TABLES	WATR 148
C			WATR 149
C	TWJM	INTEGER COUNTER USED TO LIMIT NUMBER OF TWDT HALVINGS	WATR 150
C			WATR 151
C	TWKSA	INTERPOLATED VALUE OF PWKSUM	WATR 152
C			WATR 153
C	TWONE	EQUALS .TRUE. IF FIRST DELTA THETA APPROXIMATION HAS	WATR 154
C		ALREADY BEEN DONE DURING CURRENT TWDT.	WATR 155
C			WATR 156
C	TWR	RAIN, MM, WATER TO INFILTRATE, LEAVE STANDING OR	WATR 157
C		RUNOFF DURING REMAINDER OF PMDT.	WATR 158
C			WATR 159
C	TWRAIN	EQUALS .TRUE. IF RAIN OCCURS DURING TWDT	WATR 160
C			WATR 161
C	TWRP	EQUALS .TRUE. IF TWRAIN WAS TRUE LAST TWDT.	WATR 162
C			WATR 163
C	TWSF	SURFACE FLUX OF WATER OR VAPOR, CM	WATR 164
C			WATR 165
C	TWSFB	FLUX OF WATER MOVING INTO PROFILE DURING TWDT THROUGH	WATR 166
C		CALICHE, CM.	WATR 167
C			WATR 168
C	TWSFBT	FLUX OF WATER MOVING INTO PROFILE DURING PMDT THROUGH	WATR 169
C		CALICHE, CM.	WATR 170
C			WATR 171
C	TWSFC	CALCULATED SURFACE FLUX USING TWHB(1) AND PWK(1).	WATR 172
C			WATR 173

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C TMSFD      EQUALS TMSF-TMSFC                      WATR 174
C TMSIND     STANDING WATER, MM, AFTER RUNON-RUNOFF ASSUMPTIONS HAVE WATR 175
C            BEEN CONSIDERED.                       WATR 176
C TMSTRD     CHANGE DURING TIME STEP IN WATER STORED IN LAYERS OF WATR 177
C            PROFILE, CM. (SHOULD EQUAL TWIN)        WATR 178
C TMTA(10)   WATER AVAILABLE IN LAYERS, DIMENSIONLESS WATR 179
C TMTAA(10)  VOLUMETRIC WATER CONTENT AT BEGINNING OF PMDT WATR 180
C TMTA(10)   VOLUMETRIC WATER CONTENT (VMC) AT BEGINNING OF TWOT WATR 181
C TMTAB(10)  VMC AT END OF TWOT                     WATR 182
C TMTAC(10)  VMC AT END OF TWOT                     WATR 183
C TMTBC(10)  VMC AT END OF TWOT                     WATR 184
C TMTDC(10)  VMC AT END OF TWOT                     WATR 185
C TMTDC(10)  VMC AT END OF TWOT                     WATR 186
C TMTDC(10)  VMC AT END OF TWOT                     WATR 187
C TMTDC(10)  VMC AT END OF TWOT                     WATR 188
C TMTDC(10)  VMC AT END OF TWOT                     WATR 189
C TMTDC(10)  VMC AT END OF TWOT                     WATR 190
C TMTDC(10)  VMC AT END OF TWOT                     WATR 191
C TMTDC(10)  VMC AT END OF TWOT                     WATR 192
C TMTDC(10)  VMC AT END OF TWOT                     WATR 193
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C TMTDC(10)  VMC AT END OF TWOT                     WATR 261
C TMTDC(10)  VMC AT END OF TWOT                     WATR 262
C TMTDC(10)  VMC AT END OF TWOT                     WATR 263

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      T=0.0
      DO110 I=2, PMN-1
      U=ABS(TWDTHP(I))
      IF(U.LT.R) GOTO 110
      IF(U.GT.S) GOTO 120
      IF(U.GT.T) T=U
110  CONTINUE
      IF(T.LT.R) T=R
      IF(T.LE.0.0) GOTO 120
      TWDT=.9*TWDT+(PMTLIH/T)
C
C 120 IF(TWDT .GT. (PHDT-TWTTOT)) TWDT=PHDT-TWTTOT
C   FND INITIAL STAB AT TWDT
C
C
      TWJM = 0
C   TWJM IS A COUNTER WHICH ALLOWS A LIMITED NUMBER OF TWDT
C   HALVINGS
C
C
C 170 CONTINUE
C   BEGINNING OF LOOP TO CALCULATE PRESSURES ONCE TWDT IS SET
C
C
C   CHECK FOR FROZEN SOIL
      IF(TWFRZN) GO TO 150
C   SURFACE FLUX CONSIDERS RAIN, STANDING WATER AND EVAPORATION (BUT NOT
C   TRANSPIRATION).
C   WATER IN IS PLUS, OUT IS MINUS. THIS IS POTENTIAL FLUX.
C
C   BEGIN MAIN ITERATION LOOP WITHIN TWDT
C   TWSF NOT TWSFC IS USED FOR SURFACE FLUX IN EQUATIONS SECTION.
C   THE ALTERNATIVE IS TO USE TWSFC AND PASS ON A CORRECTION = TWSF -
C   TWSFC TO NEXT TIME STEP.
C
C   TWR, TWSTND, TWEVAP IN MM.
      TW=AMIN1(TWR,(ZREINT*TWDT*24.))
      TWSF=TW+TWSTND-TWEVAP+TWDT/PHDT
C   BOTH TWSTND AND TWSF ARE IN MM AT THE MOMENT
C   REALLY WANT SURFACE FLUX IN CM FOR TWDT NOT MM
140  TWSF=TWSF*0.1
      GO TO 160
150  TWSF=0.0
160  CONTINUE
C   FND INITIALIZATION SECTION
C ITTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTT
C
C
C
C ITTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTT
C   TRANSPIRATION SECTION
C
C   WITHDRAW WATER FROM A LAYER PROPORTIONAL TO WATER AVAILABLE AND
C   FRACTION OF ROOTS IN THAT LAYER. TWTK IS PROPORTIONALITY CONSTANT
C   FOR FUNCTIONAL GROUP BEING CALCULATED.
C   TWTSPR(I) IS TRANSPIRATION FROM LAYER I DURING TWDT, CM.
C
      R=TWDT/PHDT
      DO170 I=1, PMN-2
170  TWTSPR(I)=0.0
C   BUTER DO LIMIT GOES TO 6 RATHER THAN PMFGPS IN ORDER TO
C   INCLUDE ANNUALS IN PLACES 5, 6 IN ARRAYS CVTSPR AND CVRDST
      DO 210 J=1,6
      TWTK=0.0
      DO180 I=1, PMN-2
C   TWTA IS WATER AVAILABLE IN LAYER, DIMENSIONLESS
      TWTA(I)=TWTHA(I+1)-PMLLIH
180  TWTK=TWTK+TWTA(I)+CVRDST(J,I)
C   FOLLOWING LINE ADDED SO LACK OF FUNCTIONAL GROUP OR LACK OF ROOTS
C   SOMEWHERE DOESN'T SCREW TRANSPIRATION CALCULATION.
      IF(TWTK .LE. 1.E-8) GO TO 190
      TWTK=R+TWTK(J)/TWTK
190  CONTINUE
      DO200 I=1, PMN-2
200  TWTSPR(I)=TWTSPR(I)+TWTK+TWTA(I)+CVRDST(J,I)
210  CONTINUE
C
C   FIND LARGEST DELTA THETA THIS WOULD RESULT IN FOR THIS TWDT.
      TW=0.0
      DO230 I=1, PMN-2
      S=TWTSPR(I)/CHDX(I)
C   NEVER TAKE ALL THE AVAILABLE WATER.
      IF(S.LE.(TWTA(I)+0.5)) GOTO 220
      TWTSPR(I)=TWTA(I)+0.5*CHDX(I)
      S=TWTSPR(I)
220  CONTINUE
      IF(TW.LT.S) TW=S
230  CONTINUE

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WATR 354
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C IF TW IS TOO LARGE, REDUCE TWDOT ACCORDINGLY           WATR 444
  IF((TW.LE.PMTLIM).OR.(TWMH.GE.2).OR.(TWDOT.LE.0.02083))GOTO240  WATR 445
  TWMH=TWMH+1                                           WATR 446
  TWDOT=TWDOT+0.9*PMTLIM/TW                          WATR 447
  IF(TWDOT.LE.0.02083)TWDOT=0.02083                 WATR 448
  GO TO 130                                           WATR 449
240 CONTINUE                                           WATR 450
C                                                         WATR 450
C FND TRANSPIRATION SECTION                             WATR 451
C CCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCC WATR 452
C C C C C C C C C C C C C C C C C C C C C C C C C C C C WATR 453
C C C C C C C C C C C C C C C C C C C C C C C C C C C C WATR 454
C C C C C C C C C C C C C C C C C C C C C C C C C C C C WATR 455
C C C C C C C C C C C C C C C C C C C C C C C C C C C C WATR 456
C CCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCC WATR 457
C GENERATE VALUES OF CONDUCTIVITY AND SPECIFIC WATER CAPACITY WATR 458
C FROM VALJES OF VWC                                 WATR 459
C C C C C C C C C C C C C C C C C C C C C C C C C C C C WATR 460
C IF CONDITIONS ARE ROUGHLY THE SAME THIS TWDOT AS LAST, WATR 461
C APPROXIMATE CHANGES IN THETA BY ASSUMING FOR NOW THAT CHANGES WATR 462
C THIS TWDOT EQUAL CHANGES LAST TWDOT. IF CONDITIONS ARE NOT THE SAME WATR 463
C ASSUME FOR NOW THAT BEGINNING AND END VALUES ARE THE SAME. WATR 464
C GO THROUGH THIS SECTION ONLY ONCE FOR EACH TWDOT   WATR 465
  IF(TWONE) GO TO 280                                 WATR 466
  TWONE=.TRUE.                                       WATR 467
  IF((TWRP .AND. TWRAIN) .OR. ((.NOT. TWRP) .AND. (.NOT. TWRAIN))) WATR 468
    GO TO 250                                        WATR 469
  GO TO 280                                          WATR 470
250 CONTINUE                                          WATR 471
  DO270I=2,PMN                                     WATR 472
C C C C C C C C C C C C C C C C C C C C C C C C C C C C WATR 473
C S=TWDOT/TWDOTP                                     WATR 474
  IF(S.GT.1.)S=SQRT(S)                             WATR 475
  TW=TWTHB(I)+S*TWDTHP(I)                          WATR 476
  IF(TW .GE. PMT(PWM)) GO TO 260                   WATR 477
  IF(TW .LT. PMLLIM) TW=PMLLIM                   WATR 478
  TWTHB(I)=TW                                       WATR 479
  IF(I.GT.2)GOTO270                                 WATR 480
C RECALCULATE PRESSURE VALUE OF TOP LAYER           WATR 481
C PWT(TWJA) AND PWT(TWJA+1) BRACKET TWTHB(I)     WATR 482
  TWJA=TWTHB(I) / PWDELW + 1.                     WATR 483
C TWAA IS FOR INTERPOLATING                         WATR 484
  TWAA=(TWTHB(I)-PWT(TWJA))/PWDELW                WATR 485
  TWTHB(I)=PWH(TWJA)+(PWH(TWJA+1)-PWH(TWJA))*TWAA WATR 486
  GO TO 270                                         WATR 487
260 TWTHB(I)=PWT(PWM)                              WATR 488
  TWTHB(I)=PWH(PWM)                              WATR 489
270 CONTINUE                                       WATR 490
280 CONTINUE                                       WATR 491
C C C C C C C C C C C C C C C C C C C C C C C C C C C C WATR 492
C C C C C C C C C C C C C C C C C C C C C C C C C C C C WATR 493
C DO 320I=2,PMN-1                                   WATR 494
  TWTHC=(TWTHA(I)+TWTHB(I)) / 2.                  WATR 495
  IF(TWTHC .GT. PMT(PWM)) TWTHC=PWT(PWM)         WATR 496
  IF(TWTHC .LT. PMLLIM) TWTHC=PMLLIM            WATR 497
C PWT(TWJA) AND PWT(TWJA+1) BRACKET TWTHC        WATR 498
  TWJA=TWTHC/PWDELW + 1.                          WATR 499
  IF(TWJA .GE. PMN) TWJA=PMN-1                   WATR 500
C TWAA IS FOR INTERPOLATING                         WATR 501
  TWAA=(TWTHC-PWT(TWJA))/PWDELW                  WATR 502
C TWKSA AND TWHHA ARE PART OF AVERAGE CONDUCTIVITY CALCULATION WATR 503
  TWKSA(I)=PWKSUM(TWJA)+(PWKSUM(TWJA+1)-PWKSUM(TWJA))*TWAA WATR 504
  TWHHA(I)=PWH(TWJA)+(PWH(TWJA+1)-PWH(TWJA))*TWAA WATR 505
C CALCULATE WATER CAPACITY                          WATR 506
C IT IS SIMPLY THE RECIPROCAL OF THE SLOPE OF PRESSURE VS. THETA WATR 507
  PWK(I-1) = PWDELW/(PWH(TWJA+1)-PWH(TWJA))      WATR 508
  IF(I.LE.2)GOTO310                                WATR 509
  IF(TWJB .EQ. TWJA) GO TO 300                    WATR 510
  PWK(I-1)=(TWKSA(I)-TWKSA(I-1))/(TWHHA(I)-TWHHA(I-1)) WATR 511
  GO TO 310                                         WATR 512
300 PWK(I-1)=(PWKSUM(TWJA+1)-PWKSUM(TWJA)) / (PWH(TWJA+1)-PWH(TWJA)) WATR 513
310 TWJB=TWJA                                       WATR 514
C TWJB IS USED TO HOLD ON TO THE VALUE OF TWJA FOR ONE MORE PASS WATR 515
C THROUGH DO LOOP                                   WATR 516
320 CONTINUE                                       WATR 517
C C C C C C C C C C C C C C C C C C C C C C C C C C C C WATR 518
C FND CONDUCTIVITY (EXCEPT FOR TOP LAYER) AND WATER CAP. CALCULATIONS WATR 519
C CCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCC WATR 520
C C C C C C C C C C C C C C C C C C C C C C C C C C C C WATR 521
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C C C C C C C C C C C C C C C C C C C C C C C C C C C C WATR 526
C BOUNDARY CONDITION SECTION                       WATR 527
C C C C C C C C C C C C C C C C C C C C C C C C C C C C WATR 528
C FOR THE BOTTOM BOUNDARY CONDITION MUST BE ABLE TO SET CONDUCTIVITY IN WATR 529
C THE REGION IMMEDIATELY ABOVE THE BOTTOM NODE. WE MUST ALSO SET WATR 530
C PRESSURE HEAD AT BOTTOM NODE. PRESSURE HEAD WILL BE CALCULATED WATR 531
C FOR NODES 2 THROUGH PMN-1                       WATR 532
C PWK(PMN-1) AND TWHB(PMN) WERE READ IN IN INIT. THEY REPRESENT WATR 533

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C PMN-1 AND PMN AND IS READ IN AND REPRESENTS THE CONDUCTIVITY OF THE WATR 624
C CALICHE LAYER. THE WATER CAPACITY PWWC(I) IS AN AVERAGE OVER TWDT WATR 625
C FOR A REGION CENTERED ON NODE I+1. VALUES ARE NEEDED FOR NODES WATR 626
C 2 THROUGH PMN-1. THE THICKNESS OF A REGION CENTERED ON A NODE I+1 WATR 627
C IS CHDX(I). VALUES ARE NEEDED FOR NODES 2 THROUGH PMN-1. WATR 628
C THERE IS FURTHER POSSIBILITY FOR CONFUSION BECAUSE EQUATION I IN WATR 629
C TDM IS FOR NODE I+1. WATR 630
C WATR 631
C IT HAS BEEN DECIDED (NOT BY ME OBVIOUSLY) THAT THE UNITS FOR WATR 632
C PRESSURE IN THIS SUBROUTINE ARE BARS. THUS, CONDUCTIVITY HAS UNITS WATR 633
C OF CM+2 BAR-1 DAY-1. WATER CAPACITY IS IN BAR-1. PRESSURE HEADS WATR 634
C DUE TO HEIGHT DIFFERENCES MUST BE MULTIPLIED BY 9.833E-4 BAR/CM WATR 635
C TO GET BARS. TO CONVERT TO PRESSURE UNITS IN CM MUST MULTIPLY WATR 636
C CONDUCTIVITY BY 9.833E-4 BAR/CM AND GET UNITS OF CM/DAY. CAN USE WATR 637
C PRESSURE DUE TO HEIGHT DIFFERENCES DIRECTLY IN CM. THE CONVERSION WATR 638
C FACTOR IS 9.833E-4 BAR/CM OR 1017 CM/BAR. WATR 639
C WATR 640
C DO 430 I=1,PMN-1 WATR 641
C TWHA(I) = PWK(I) / CHDX(I) WATR 642
C IF(I .EQ. PMN-1) GO TO 430 WATR 643
C TWHB(I) = (2.*PWWC(I)+CHDX(I)) / TWDT WATR 644
470 CONTINUE WATR 645
C WITH TWHA AND TWHB GENERATE ABCD'S FOR TDM WATR 646
DO 440 I=2, PMN-3 WATR 647
C TWA(I) = TWHA(I+1) WATR 648
C TWC(I) = TWHA(I) WATR 649
C TWB(I) = TWHB(I)+TWA(I)+TWC(I) WATR 650
440 TWD(I) = TWHA(I+1)*(TWHB(I)-TWA(I)-TWC(I)) + TWHA(I)+TWC(I) WATR 651
1 + TWHA(I+2)*TWA(I) + 2.*(PWK(I) - PWK(I+1)) *9.833E-4 WATR 652
2 -2.*TWTSPR(I)/TWDT WATR 653
C CALCULATE FIRST AND LAST VALUES WATR 654
TWA(1) = TWHA(2) WATR 655
TWC(1) = 0. WATR 656
TWB(1) = TWHB(1) + TWA(1) WATR 657
TWD(1) = TWHA(2)*(TWHB(1)-TWA(1)) + TWHA(3)+TWA(1) WATR 658
1 + TWSF*2./TWDT -2.*PWK(2)*9.833E-4 WATR 659
2 -2.*TWTSPR(1)/TWDT WATR 660
TWA(PMN-2) = 0. WATR 661
TWC(PMN-2) = TWHA(PMN-2) WATR 662
TWB(PMN-2) = TWHB(PMN-2) + TWC(PMN-2) +TWHA(PMN-1) WATR 663
TWD(PMN-2) = TWHA(PMN-1)*(TWHB(PMN-2)-TWC(PMN-2)-TWHA(PMN-1)) WATR 664
1 + TWHA(PMN-2)*TWC(PMN-2) + 2.*TWHB(PMN)*TWHA(PMN-1) WATR 665
2 +2.*(PWK(PMN-2)-PWK(PMN-1)) *9.833E-4 WATR 666
3 -2.*TWTSPR(PMN-2)/TWDT WATR 667
C WATR 668
C FND EQUATION SECTION WATR 669
C WATR 670
C WATR 671
C EEEEEEEEEEEEEEEEEEEEEEEEEEEEEEEEEEEEEEEEEEEEEEEEEEEEEEEEEEEEEEE WATR 672
C WATR 673
C WATR 674
C TTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTT WATR 675
C CALL TDM TO CALCULATE PMN-2 PRESSURES FOR NODES 2 THROUGH PMN-1 WATR 676
CALL TDM(TWA, TWB, TWC, TWD, TWU, PMN-2) WATR 677
C TTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTTT WATR 678
C WATR 679
C WATR 680
C WATR 681
C WATR 682
C MMMMMMMMMMMMMMMMMMMMMMMMMMMMMMMMMMMMMMMMMMMMMMMMMMMMMMMMMMMMMM WATR 683
C HERE ARE SOME MISCELLANEDUS CHECKS, DECISIONS AND CALCULATIONS WATR 684
C WATR 685
C CALCULATE THETA FOR NODES 2 TO PMN-1 USING PRESSURE VALUES JUST WATR 686
C OBTAINED FROM TDM. USE WATER CAPACITY A LA HANKS. DELTA THETA= WATR 687
C DELTA PRESSURE TIMES WATER CAPACITY. WATR 688
C WATR 689
DO450I=2,PMN-1 WATR 690
C TWTHB(I)=TWTHA(I)+PWWC(I-1)*(TWU(I-1)-TWHA(I)) WATR 691
IF(TWTHB(I) .GT. PWT(PMN)) TWTHB(I)=PWT(PMN) WATR 692
IF(TWTHB(I) .LT. PWLLIM) TWTHB(I)=PWLLIM WATR 693
450 CONTINUE WATR 694
C WATR 695
C CALCULATE TWHB FROM TWTHB WATR 696
DO460I=2,PMN-1 WATR 697
C TWJA=TWTHB(I)/PWDELW+1. WATR 698
IF(TWJA .GE. PMN) TWJA=PMN-1 WATR 699
TWA=(TWTHB(I)-PWT(TWJA))/PWDELW WATR 700
460 TWHB(I)=PWH(TWJA)+(PWH(TWJA+1) - PWH(TWJA))*TWA WATR 701
C WATR 702
C WATR 703
C CHECK IF ANY DELTA THETA'S TOO LARGE. IF YES, REDUCE TWDT WATR 704
DO470I=2,PMN-1 WATR 705
S=ABS(TWTHB(I)-TWTHA(I)) WATR 706
IF(S.GT.PWTLIM)GOTO480 WATR 707
470 CONTINUE WATR 708
GO TO 510 WATR 709
480 IF(TWJM.GE.2)GOTO510 WATR 710
TWDT=TWDT WATR 711
TWDT=TWDT*0.9*PWTLIM/S WATR 712
TWJM=TWJM+1 WATR 713
C DON'T LET TWDT GET SMALLER THAN 30 MINUTES WATR 714

```

```

          IF(TWDT.LT.0.02083)TWDT=0.02083
          R=TWDT/TWDTQ
490   LOSOQ=2,PMN-1
          TWHB(I)=TWHA(I)+((TWHB(I)-TWHA(I))*R
500   TWTHB(I)=TWTHA(I)+((TWTHB(I)-TWTHA(I))*R
          GO TO 130
510   CONTINUE
C
C
C CHECK WATER BALANCE
C
          TWSFB=-PWKCAL*TWDT*((TWHB(PMN-1) + TWHA(PMN-1)) + .5
          1 +9.833E-4*CHDXX(PMN-1) - TWHB(PMN)) / CHOXX(PMN-1)
C
C
          CALL WBAL( TWSPR, TWSF, TWSFB, TWTHA, TWTHB, CHDX, PMN,
          1 TWIN, TWSTRO, R)
C
          IF(TWDT.LT.0.05)GOTO520
C WATER IS SOMETIMES NOT CONSERVED
C WHEN THIS IS THE CASE, HALVE TWDT.
          IF((R.GT. 99.) .OR. (ABS(TWIN-TWSTRO) .LT. 0.006)
          1 .OR.(ABS(R-1.) .LT.0.01))GOTO520
          TWDT=0.5*TWDT
          R=0.5
          GOTO490
520   CONTINUE
C
C
C
C
C
C
C
C
C
C THIS IS THE WRAP UP SECTION. ONCE WE'RE HERE WE EITHER RETURN TO
C CALLING PROGRAM OR DO ANOTHER TWDT.
C
C
C
C
C
C
C
C
C
C
C
          DEBLGGIN
          IF((.NOT. FLGFLG) .OR. (.NOT. DBGFLG)) GO TO 550
          Y=TIME(Z)
          WRITE(1,530) TWDT,TWR,TWRP,TWRAIN,TWONE,TWIN,TWSTND,TWSF,
          1 TWJM,TWIN,TWSTRO,R,Y
          WRITE(1,540) TWTHB,TWHA
          530  FORMAT(' M ',2E9.3,4L1,F5.2,E9.3,I3,' TWIN, TWSTRO, RATIO',
          1 3F9.4,' CPU TIME=',F7.3)
          540  FORMAT(' M TWTHB=', 10F5.3, ' TWHA=', 10F5.2)
          550  CONTINUE
C
C
C
C
C
C
C
C
C
C
C
          SET SOME VALUES UP FOR NEXT TWDT.
          TWDTP=TWDT
          TWRP=TWRAIN
          TWONE=.FALSE.
          DO 560 I=1,PMN-1
          TWHA(I)=TWHB(I)
          TWDTHP(I)=TWTHB(I)-TWTHA(I)
          TWTHA(I)=TWTHB(I)
          560  CONTINUE
C
C
C
C
C
C
C
C
C
C
C
          CWINF IS THE NET TOP SURFACE FLUX - DOES NOT INCLUDE TRANSPIRATION
          CWINF=CWINF+TWSF
C
C
C
C
          TWSFBT IS NET BOTTOM SURFACE FLUX FOR PMO, OUT IS MINUS
          TWSFBT = TWSFBT + TWSFB
C
C
C
C
          CHECK FOR FROZEN SOIL
          IF(.NOT. TWFRZN) GO TO 570
          TWSTND=TWSTND+TWR
          TWR=0.0
          GO TO 620
          570  CONTINUE
C
C
C
C
          ANY STANDING WATER?
          IF(TWIN) GO TO 580
C TWSFD = RAIN + STANDING WATER - EVAPORATION - ACTUAL SURFACE FLUX
          TWSTND=TWSFD*10.*PWRRRF**TWDT
          IF (TWSTND .LT. 0.0) TWSTND=0.0
          GO TO 590
          580  TWSTND=0.0
          590  CONTINUE
C
C
C
C
          DETERMINE AMOUNT OF RAIN REMAINING TO INFILTRATE.
          600  TWR=TWR-ZRINT*TWDT*24.
          610  IF(TWR .LT. 0.0) TWR=0.0
C
          620  CONTINUE

```

WATR 715
 WATR 716
 WATR 717
 WATR 718
 WATR 719
 WATR 720
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 WATR 723
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 WATR 796
 WATR 797
 WATR 798
 WATR 799
 WATR 800
 WATR 801
 WATR 802
 WATR 803
 WATR 804

```

      XWSTND=TWSTND
C
C   CALCULATE SUM OF TRANSPIRATION, CH
      DO 630 I=1,PMN=2
        6*0 TWHTS(I)=TWHTS(I)+TWTSPR(I)
C
C   RESET TIMING
      TWTTOT=TWTTOT+TWDI
C
      IF(ABS((TWTTOT-PMDT)/PMDT) .GT. 0.002) GO TO 60
C
C   FND WRAP UP SECTION FOR TROT
CXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXX
C
C
C
C
C   FFFFFFFF
C THIS IS THE FINAL SECTION
C DO THE LAST COUPLE OF THINGS NECESSARY BEFORE RETURNING .
C
      DO640I=1,PMN=2
        XWHTA(I)=TWHTA(I+1)
        640 CWPSI(I)=TWHB(I+1)
C
C
C   DEBUGGING
      IF(.NOT. D0GFLG) GO TO 660
C   CHECK WATER BALANCE FOR THIS PMDT
      CALL WBALC(TWHTS, CWINF, TMSFBI, TWTHAA, TWHB, CHDX, PMN,
        1 TWWIN, TWSTRD, R)
      WRITE(1,650) XWHTA,CWPSI,TWHTS,CWINF,TWWIN,TWSTRD,R
        650 FORMAT(' WM XWHTA=',10F5.3, ' CWPSI=',10F6.2,/, ' WM TWHTS='
          1,10F6.4, ' CWINF=',F8.4, ' TWWIN,TWSTRD,RATIO',3F8.4)
        660 CONTINUE
C
C   FND DEBUGGING
C
C
C   FND FINAL SECTION
C   FFFFFFFF
C
C
C   RETURN
C
C   CEEFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFF
C   CEE*FFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFF
C   CEEFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFFF
C
      ENTRY WINIT
C
C   CRRRRRRRRRRRRRRRRRRRRRRRRRRRRRRRRRRRRRRRRRRRRRRRRRRRRRRRRRRRRRRRRRRRR
C THIS ENTRY POINT IS CALLED ONCE FROM MAIN. VARIABLES PECULIAR TO
C WATER ARE READ IN HERE AND PRINTED OUT FROM HERE BOTH TO GET THESE
C DATA LISTED AS PART OF THE PROGRAM AND TO SERVE AS A DATA CHECK.
C Y NEED TO READ: PWDELW, K-THETA TABLE, H-THETA TABLE, PWKCAL,
C PMH, TWHB(10), TWHB(10), PWHORY, PWHMET,
C CONSTANT PRESSURE OF CALICHE READ IN AS TWHB(PMN)
C MUST ALSO READ IN TWSTND, PMTLIM, PWRNRF
C
C   PEAL AND WRITE DATA AND ACCOMPANYING COMMENTS
      READ(5,670)TWCCC
      WRITE(6,680)
      WRITE(6,690)TWCCC
      WRITE(6,700)
C
      READ(5,670)TWCCC
      WRITE(6,690)TWCCC
      READ(5,/)PWDELW
      WRITE(6,710)PWDELW
      WRITE(6,700)
C
      READ(5,670)TWCCC
      WRITE(6,690)TWCCC
      READ(5,/)PMH
      WRITE(6,720)PMH
      WRITE(6,730)
C
      READ(5,670)TWCCC
      WRITE(6,690)TWCCC
      READ(5,/)PWKCAL
      WRITE(6,720)PWKCAL
      WRITE(6,700)
C
      READ(5,670)TWCCC
      WRITE(6,690)TWCCC

```


FUNCTION WTIME

```
      FUNCTION WTIME(Z)                                WTM 0100
C      MARCH 1976      PAUL LOMMEN                    WTM 0200
C THIS FUNCTION SIMPLY DETERMINES CPUTIME IN SECONDS.  TIME(2) IS AN WTM 0300
C INTRINSIC FUNCTION,                                  WTM 0400
C I HAD TO DO IT THIS WAY BECAUSE TIME IS A VARIABLE NAME WE'VE WTM 0500
C ALREADY USED AND IS IN THE COMMON BLOCK,           WTM 0600
C TIME(2) IS IN 60THS OF A SECOND.  0.016667#1/60   WTM 0700
      WTIME=0.016667*TIME(2)                          WTM 0800
      RETURN                                           WTM 0900
      END                                              WTM 1000
```

DERIVATION OF DIFFERENCE EQUATIONS

Soil nodes are introduced into the soil profile as reference points for the set of difference equations WATER will generate and then solve. This derivation will proceed very much as the derivation of difference equations for HEAT, since both Equations B-1 and C-1 are diffusion equations.

For a thin layer of soil we rather arbitrarily divide water movement into four components, for convenience:

$$\begin{aligned} & \text{(water in through top)} - \text{(water out through bottom)} \\ & - \text{(transpiration)} = \text{(water stored)}. \end{aligned} \quad (\text{C-2})$$

For layer i :

$$\text{(water in through top)}_i = -K \frac{\partial H}{\partial z}, \quad (\text{C-3})$$

where

the right-hand side is evaluated at the top of layer i ;

water moving down is a positive flow;

counting of layers starts at soil surface and goes down;

each side of Equation C-3 has units of $\text{cm} \cdot \text{day}^{-1}$;

K = soil hydraulic conductivity, $\text{cm}^2 \cdot \text{bar}^{-1} \cdot \text{day}^{-1}$;

H = soil hydraulic potential head, bars (includes matric potential plus gravitational potential);

z = depth in soil, cm.

Now fill out Equation C-3 as a difference equation:

$\text{(water in through top)}_i$

$$\begin{aligned} & = (-K_{i-1/2}) \left\{ \left[\frac{h_i - h_{i-1}}{\Delta z_{i-1/2}} + \frac{h_i^* - h_{i-1}^*}{\Delta z_{i-1/2}} \right] \left(\frac{1}{2} \right) - \frac{(G) (\Delta z_{i-1})}{2 \Delta z_{i-1/2}} \right\}; \\ & = (-K_{i-1/2}) \left[\frac{h_i - h_{i-1} + h_i^* - h_{i-1}^* - 2G(\Delta z_{i-1})}{2 \Delta z_{i-1/2}} \right], \end{aligned} \quad (\text{C-4})$$

where

h_i = soil matric water potential (hereafter SWP) of layer i at beginning of time-step of length Δt days, located at the node in the center of layer i ;
 h_{i-1} = SWP of layer $i-1$ at beginning of Δt ;
 h_i^* = SWP of layer i at end of Δt ;
 h_{i-1}^* = SWP of layer $i-1$ at end of Δt ;

$K_{i-1/2}$ = hydraulic conductivity of soil in region from center of layer $i-1$ to center of layer i (i.e., from node i to node $i+1$, see Fig. C-2);

G = constant to convert gravitational contribution to total soil hydraulic potential head from cm water potential to bars, = $9.833 \times 10^{-4} \text{ bar} \cdot \text{cm}^{-1}$;

$\Delta z_{i-1/2}$ = distance from node i to node $i+1$.

Similarly:

$\text{(water out through bottom)}_i$

$$= (-K_{i+1/2}) \left[\frac{h_{i+1}^* - h_i + h_{i+1} - h_i^* - 2G(\Delta z_{i+1/2})}{2 \Delta z_{i+1/2}} \right]. \quad (\text{C-5})$$

Transpiration in $\text{cm} \cdot \text{day}^{-1}$ is

$$\text{(transpiration)}_i = \frac{T_i}{\Delta t}, \quad (\text{C-6})$$

where

T_i = transpiration removed from layer i during Δt , centimeters.

Water stored in $\text{cm} \cdot \text{day}^{-1}$ is

$$\begin{aligned} \text{(water stored)}_i &= \frac{(\Delta \theta)_i (\Delta x_i)}{\Delta t}; \\ &= \frac{(h_i - h_i) (S_i) (\Delta x_i)}{\Delta t}, \end{aligned} \quad (\text{C-7})$$

where

$\Delta \theta_i$ = change in theta of layer i during Δt , dimensionless (θ_i = cm water in layer i ÷ cm soil in layer i);

Δx_i = thickness of layer i , cm;

S_i = specific water capacity of layer i during Δt , bar^{-1} , $S_i = \Delta \theta_i / \Delta h_i$.

In the same manner as in the HEAT submodel, combine Equations C-4, C-5, C-6 and C-7 with Equation C-2. After several lines of algebra we get

$$-C_i h_{i-1}^* + B_i h_i^* - A_i h_{i+1}^* = D_i, \quad (\text{C-8})$$

where

$$C_i = \frac{K_{i-1/2}}{\Delta z_{i-1/2}}; \quad (\text{C-9})$$

$$A_i = \frac{K_{i+1/2}}{\Delta z_{i+1/2}}; \quad (\text{C-10})$$

$$B_i = A_i + C_i + \frac{2S_i \Delta x_i}{\Delta t}; \quad (C-11)$$

$$D_i = h_i \left(\frac{2S_i \Delta x_i}{\Delta t} - A_i - C_i \right) + h_{i-1} C_i + h_{i+1} A_i + 2G \left(K_{i-1} - K_{i+1} \right) - 2 \frac{T_i}{\Delta t}. \quad (C-12)$$

For top and bottom layers, adjustments must be made in Equations C-9 through C-12. The water in through the top of the top layer is the surface flux, as calculated in lines 537 through 594 of WATER, divided by Δt . Then, A_i , B_i , C_i and D_i are determined by going through the same sequence of steps which resulted in Equations C-9 through C-12. For the bottom layer the difference which must be accommodated is that the SWP below the layer is fixed.

Now summarize the equations for A, B, C, D for all layers. It is first convenient to define:

M = number of layers;

$$Y_i = K_{i-1} / \Delta z_{i-1}; \quad (C-13)$$

$$R_i = 2S_i \Delta z_i / \Delta t. \quad (C-14)$$

The result is a set of M equations of the form shown in Equation C-8, which form a tridiagonal matrix, where

$$A_i = Y_{i+1} \text{ for } i = 1, \dots, M-1; \\ = 0 \text{ for } i = M; \quad (C-15)$$

$$C_i = 0 \text{ for } i = 1; \\ = Y_i \text{ for } i = 2, \dots, M; \quad (C-16)$$

$$B_i = Y_2 + R_1 \text{ for } i = 1; \\ = Y_i + Y_{i+1} + R_i \text{ for } i = 2, \dots, M; \quad (C-17)$$

$$D_i = h_1(R_1 - Y_2) + h_2 Y_2 + \frac{2F}{\Delta t} - \frac{2GK_3}{2} - \frac{2T_1}{\Delta t} \text{ for } i = 1; \\ = h_1(R_1 - Y_i - Y_{i+1}) + h_{i-1} Y_i + h_{i+1} Y_{i+1} + 2G \left(K_{i-1} - K_{i+1} \right) - \frac{2T_i}{\Delta t} \text{ for } i = 2, \dots, M-1; \\ = h_M(R_M - Y_M - Y_{M+1}) + h_{M-1} Y_M + h_{M+1}^{**} (2 Y_{M+1}) + 2G \left(K_{M-1} - K_{M+1} \right) - \frac{2T_M}{\Delta t} \text{ for } i = M, \quad (C-18)$$

where

F = surface flux into top layer during Δt , cm;
 h_{M+1}^{**} = fixed potential of region below bottom layer, bars.

Finally, translate Equations C-15 through C-18 into computer code with the aid of Table C-2.

Table C-2. Transformation of algebraic to FORTRAN variables

Algebraic Variable	FORTRAN Variable (or Constant)
i	I
Δt	TWDT
F	TWSF
G	9.833E-4
M	PMN-2
K_{i-1}	PWK(I)
h_{i+1}	TWHA(I+2)
h_{i+1}^*	TWNB(I+2)
h_{M+1}^{**}	TWNB(PMN)
Δz_{i-1}	CHDXX(I)
Δx_i	CHDX(I)
S_i	PWNC(I)
T_i	TWTSR(I)
Y_i	TWNA(I) = PWK(I)/CHDXX(I)
R_i	TWNB(I) = 2.*PWNC(I)*CHDX(I)/TWDT
A_i	TWA(I)
B_i	TWB(I)
C_i	TWC(I)
D_i	TWD(I)

Equation C-15 becomes

$$TWA(I) = TWNA(I+1) \text{ for } I = 1, \dots, PMN-3; \\ = 0 \text{ for } I = PMN-2. \quad (C-19)$$

Equation C-16 becomes

$$TWC(I) = 0 \text{ for } I = 1; \\ = TWNA(I) \text{ for } I = 2, \dots, PMN-2. \quad (C-20)$$

Equation C-17 becomes

$$TWB(I) = TWNA(2) + TWNB(1) \text{ for } I = 1; \\ = TWNA(1) + TWNA(I+1) + TWNB(I) \\ \text{for } I = 2, \dots, PMN-2. \quad (C-21)$$

And finally, Equation C-18 becomes

$$TWD(I) = TWHA(2)*(TWNB(1) - TWNA(2)) \\ + TWHA(3)*TWNA(2) + 2.*TWSF/TWDT \\ - 2.*9.833E-4*PWK(2) - 2.*TWTSR(1)/TWDT \\ \text{for } I = 1;$$

```

= TWHA (I+1)*(TWWB(I) - TWWA(I) - TWWA(I+1))
+TWHA(I)*TWWA(I) + TWHA(I+2)*TWWA(I+1)
+2.*9.833E-4*(PWK(I) - PWK(I+1))
-2.*TWTSPR(I)/TWDI
for I = 2, . . . , PMN-3;

= TWHA(PMN-1)*(TWWB(PMN-2) - TWWA(PMN-2) - TWWA(PMN-1))
+TWHA(PMN-2)* TWWA(PMN-2) + 2.*TWHB(PMN)* TWWA(PMN-1)
+2.*9.833E-4*(PWK(PMN-2) - PWK(PMN-1)) -2.*TWTSPR(PMN-2)/TWDI
for I = PMN-2.
(C-22)

```

Equations C-19 through C-22 are the FORTRAN versions of the difference equations used to solve for SWP of the soil layers. They are found in the program in lines 647 through 667 and immediately precede the call to the tridiagonal matrix subroutine which solves them.

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D. SUPPORT PROGRAM TDM

P. W. Lommen

GENERAL DESCRIPTION OF SUBROUTINE TDM

This subroutine solves a set of N linear equations in N unknowns if the coefficients of the unknowns form a tridiagonal matrix. Difference equation approximations to the diffusion equation are typically of the tridiagonal type. The method used here is from Richtmyer (1957, page 103).

Briefly, this routine goes forward through the equations once, eliminating the $U(J)$ for the smallest J at each step. At the $N-1$ equation, one is left with two equations in two unknowns. Then $U(J)$ is solved for and the routine goes backwards through the equations solving for a value of U at each step.

$-A, B, -C$ are the coefficients of the U 's to the right of the main diagonal, on the diagonal and to the left of the diagonal, respectively; the D 's are the constant, right-hand sides of the equations. The U 's are the solutions. A, B, C, D must be evaluated in the routine which calls TDM.

PROGRAM DESCRIPTION

To illustrate the method of solution, first write out the equations:

$$\begin{aligned} B_1 U_1 - A_1 U_2 &= D_1, & \text{(D-1.1)} \\ -C_2 U_1 + B_2 U_2 - A_2 U_3 &= D_2, & \text{(D-1.2)} \\ -C_3 U_2 + B_3 U_3 - A_3 U_4 &= D_3, & \text{(D-1.3)} \\ &\vdots \\ &\vdots \\ -C_N U_{N-1} + B_N U_N &= D_N. & \text{(D-1.N)} \end{aligned}$$

$$\begin{aligned} E(1) &= A(1) / B(1) \\ F(1) &= D(1) / B(1) \end{aligned}$$

TDM 31
TDM 32

If Equation D-1.1 is divided by B_1 , we get

$$U_1 - (A_1/B_1)U_2 = D_1/B_1, \quad \text{(D-2)}$$

transposing we get

$$U_1 = (A_1/B_1)U_2 + (D_1/B_1) = E_1 U_2 + F_1. \quad \text{(D-3)}$$

```

DO 10 I=2, N-1
  DD = B(I) - C(I)*E(I-1)
  E(I) = A(I) / DD
10 F(I) = ( D(I) + C(I)*F(I-1) ) / DD

```

TDM 33
TDM 34
TDM 35
TDM 36

To see how these equations came about, substitute Equation D-3 into Equation D-1.2 and solve for U_2 :

$$\begin{aligned} -C_2[E_1 U_2 + F_1] + B_2 U_2 - A_2 U_3 &= D_2, \\ (B_2 - C_2 E_1)U_2 &= D_2 + A_2 U_3 + C_2 F_1, \\ \text{or } U_2 &= E_2 U_3 + F_2, & \text{(D-4)} \end{aligned}$$

where

$$E_2 = A_2 / (B_2 - C_2 E_1), \quad (D-5)$$

and

$$F_2 = (D_2 + C_2 F_1) / (B_2 - C_2 E_1). \quad (D-6)$$

Continuing in this manner, we get

$$E_i = A_i / DD, \quad (D-7)$$

$$F_i = (D_i + C_i F_{i-1}) / DD, \quad (D-8)$$

where

$$DD = (B_i - C_i E_{i-1}), \quad (D-9)$$

and

$$U_i = E_i U_{i+1} + F_i. \quad (D-10)$$

The quantity DD is not subscripted because it is not needed again.

$$U(N) = (C(N) * F(N-1) + D(N)) / (B(N) - C(N) * E(N-1))$$

TDM 37

The last equation generated in the DO loop just discussed is

$$U_{N-1} = E_{N-1} U_N + F_{N-1}. \quad (D-11)$$

Substitute Equation D-11 into Equation D-1.N:

$$\begin{aligned} -C_N [E_{N-1} U_N + F_{N-1}] + B_N U_N &= D_N, \\ U_N &= (C_N F_{N-1} + D_N) / (B_N - C_N E_{N-1}), \end{aligned} \quad (D-12)$$

which is what we have in line 37.

Since all the quantities on the right in Equation D-12 have been evaluated in the calling program or in TDM, U_N is thus determined.

```

I=N
20 I=I-1
  U(I) = U(I+1) * E(I) + F(I)
  IF ( I .GT. 1 ) GO TO 20

```

TDM 38
TDM 39
TDM 40
TDM 41

Once we know U_N from Equation D-12, then from Equation D-11 we get U_{N-1} . Proceeding backwards using Equation D-10 for $I = N-2, N-3, \dots, 2, 1$, we determine the remainder of the U 's.

COMPLETE PROGRAM LISTING

```

SUBROUTINE TDM(A, B, C, D, U, N)                                TDM 01
C                                                                TDM 02
C                                                                TDM 03
C   MARCH   1976    PAUL LOHMEN                                TDM 04
C                                                                TDM 05
C   THIS SUBROUTINE SOLVES A SET OF N LINEAR EQUATIONS IN N UNKNOWNS TDM 06
C   IF THE COEFFICIENTS OF THE UNKNOWNS FORM A TRI-DIAGONAL MATRIX, TDM 07
C   DIFFERENCE EQUATION APPROXIMATIONS TO THE DIFFUSION EQUATION ARE TDM 08
C   TYPICALLY OF THE TRI-DIAGONAL TYPE, THERE ARE SEVERAL MEANS OF TDM 09
C   SOLVING SUCH A SET OF EQUATIONS, THE MOST STRAIGHTFORWARD, AND THE TDM 10
C   ONE THAT I USE HERE AND HANKS USES IN SEVERAL OF HIS MODELS COMES TDM 11
C   FROM, ROBERT D. RICHTMYER, 1957, DIFFERENCE METHODS FOR TDM 12
C   INITIAL-VALUE PROBLEMS, INTERSCIENCE PUBLISHERS, INC., NEW YORK, TDM 13
C   SEE PAGE 103. NOTATION USED HERE IS RICHTMYERS'. TDM 14
C                                                                TDM 15
C   BRIEFLY, THIS ROUTINE GOES FORWARD THROUGH THE EQUATIONS ONCE, TDM 16
C   ELIMINATING THE U(J) FOR THE SMALLEST J AT EACH STEP, AT THE TDM 17
C   N=1 EQUATION ONE IS LEFT WITH 2 EQUATIONS IN 2 UNKNOWNS, THEN TDM 18
C   U(N) IS SOLVED FOR AND THEN THE ROUTINE GOES BACKWARDS THROUGH THE TDM 19
C   EQUATIONS SOLVING FOR A VALUE OF U AT EACH STEP. TDM 20
C                                                                TDM 21
C   *A, C =C, ARE THE COEFFICIENTS OF THE U'S TO THE RIGHT OF THE MAIN TDM 22
C   DIAGONAL, ON THE DIAGONAL AND TO THE LEFT OF THE DIAGONAL TDM 23
C   RESPECTIVELY. THE D'S ARE THE CONSTANT, RIGHT HAND SIDES OF THE TDM 24
C   EQUATIONS. THE U'S ARE WHAT IS SOLVED FOR. A, B, C, D MUST TDM 25
C   BE EVALUATED IN THE ROUTINE WHICH CALLS TDM. TDM 26
C                                                                TDM 27
C                                                                TDM 28
C                                                                TDM 29
C                                                                TDM 30
C   DIMENSION A(20), B(20), C(20), D(20), E(20), F(20), U(20) TDM 31
C   E(1) = A(1) / B(1) TDM 32
C   F(1) = D(1) / B(1) TDM 33
C   DO 10 I=2, N-1 TDM 34
C     DD = B(I) - C(I)*E(I-1) TDM 35
C     E(I) = A(I) / DD TDM 36
C 10 F(I) = ( D(I) + C(I)*F(I-1) ) / DD TDM 37
C   U(N) = ( C(N)*F(N-1) + D(N) ) / ( B(N) - C(N)*E(N-1) ) TDM 38
C   I=N TDM 39
C 20 I=I-1 TDM 40
C   U(I) = U(I+1) * E(I) + F(I) TDM 41
C   IF ( I .GT. 1 ) GO TO 20 TDM 42
C   RETURN TDM 43
C   END

```

LITERATURE CITED

RICHTMYER, R. D. 1957. Difference methods for initial value problems. Interscience Publ., New York.