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BSAM: Basin Simulation Assessment Model Documentation and User Manual

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BASIN SIMULATION ASSESSMENT MODEL, BSAM,
DOCUMENTATION AND USER MANUAL

Introduction

The river basin hydrologic simulation model is an evolutionary version of the model originally developed by the UWRL for simulating the streamflow of the Bear River basin in Utah. Variations of the same basic model have been used to simulate many watersheds and river basins in the Western United States, Mexico, and South American. The underlying concept of the model is that of conservation of mass for a monthly time interval. This is represented by the equation of continuity,

$$\text{Inflow} - \text{Outflow} = \text{Change in storage.}$$

A mathematical description of the various hydrologic components of the continuity equation and the routing processes utilized in the model is included in Appendix A. The computer implementation of the model also incorporates a modified pattern search algorithm to

aid in the calibration phase of the modeling process. A description of this algorithm as applied to a river basin model is given in the paper, "A Self-Verifying Hybrid Computer Model of River-Basin Hydrology," by R. W. Hill, E. K. Israelsen, A. L. Huber, and J. P. Riley at the Seventh American Water Resources Conference, Washington, D. C., October 24-28, 1971.

Figure 1 shows the structure and interrelationship between the main program, BSAM, and the major subprograms. Detailed instructions for using the program and preparing the input data are given in the following sections. A card listing of the input data for a typical run and the resulting output is given in Appendix B. The FORTRAN computer program is listed in Appendix C.

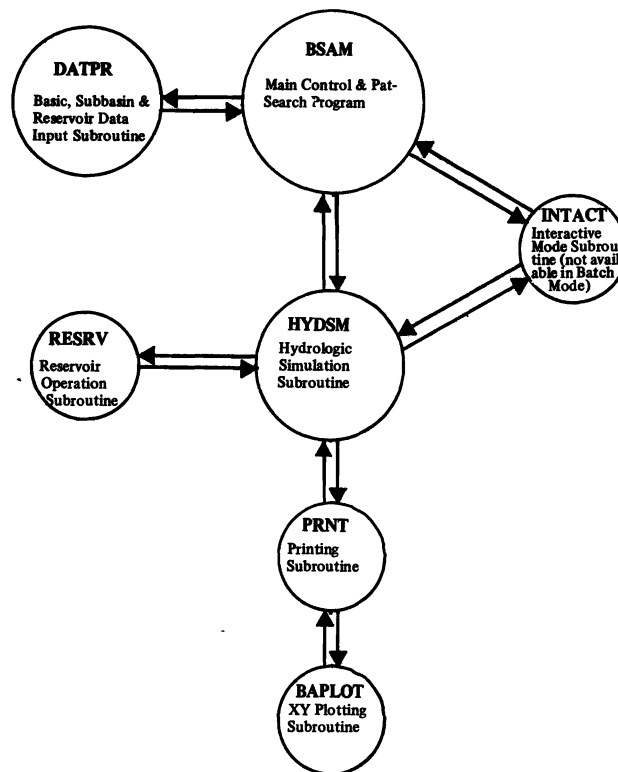


Figure 1. Structure and Interrelationship between BSAM and major subprograms.

General control and I/O structure

After the computer model has been compiled and stored in an executable form, operational control and I/O options are determined by the four variables, ITY, IPL, IPRT, and ICOL punched in the first card of the input data deck. Execution control options are shown schematically in Figure 2 and a deck setup for a typical

simulation run is shown in Figure 3. Table 1 gives instructions for preparing the input data. Table 2 describes all of the model parameters and Table 3 identifies each line of the output tables. Figure 4 shows keypunch coding sheets prepared for a typical simulation run.

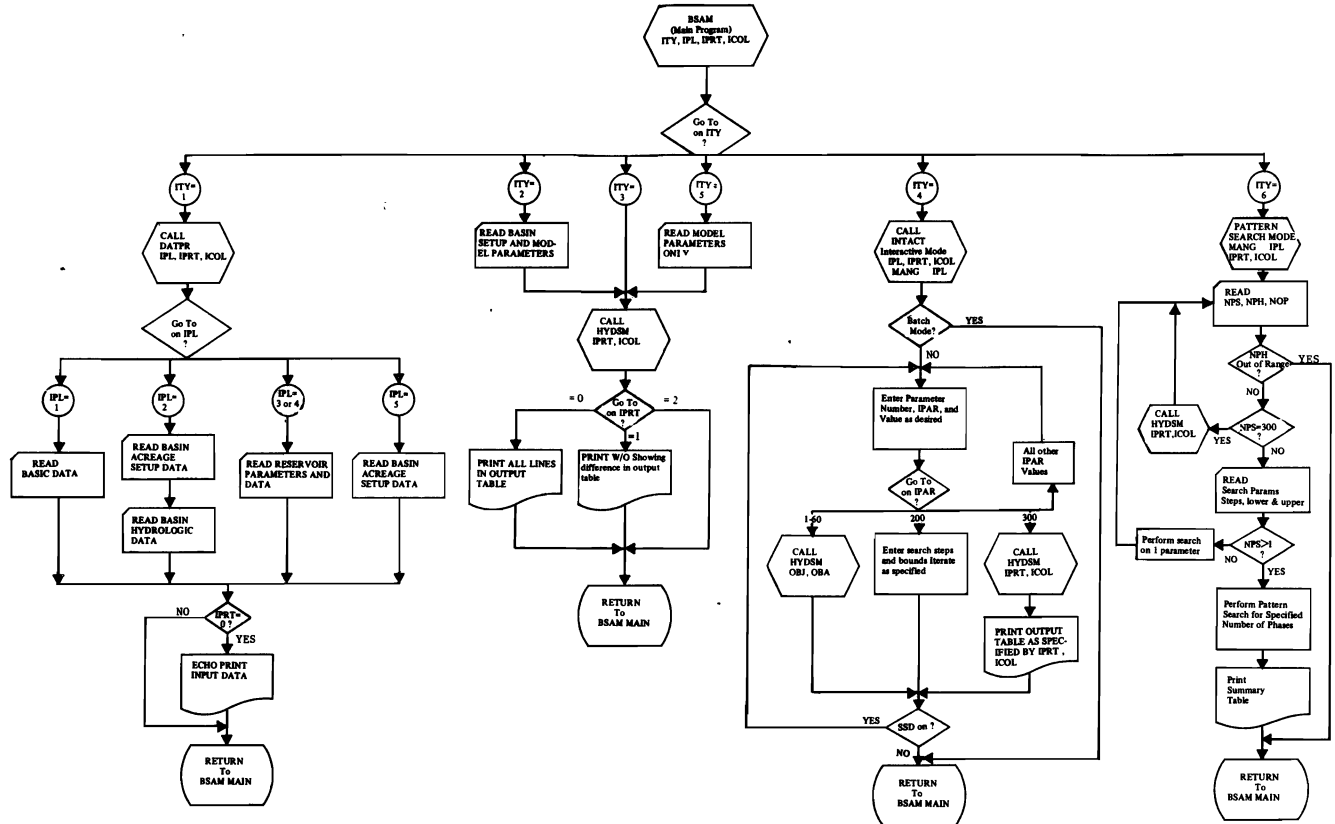


Figure 2. Schematic Diagram of the Execution Control Options for BSAM

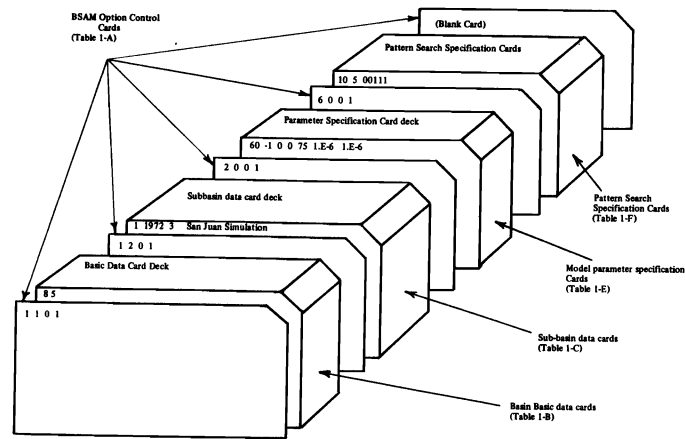


Figure 3. Card Deck Setup for a Typical Simulation Run

Table 1. Input data preparation.

A. Option control card in FORMAT (4I3)

Col.	Name	FORMAT	Description
1-3	ITY	I3	Program control option (1-6)
4-6	IPL	I3	Specification option subordinate to ITY (0-5)
7-9	IPRT	I3	Printing option subordinate to ITY, (0-2)
10-12	ICOL	I3	Line printer column length specification 1 = 80 column printer 2 = 132 column printer

B. Basic data (Called with ITY = 1, IPL = 1)

Card	Column	Name	FORMAT	Description
1	1-3	NCROP	I3	Number of crops
	4-6	NPHRT	I3	Number of phreatophytes
2	1-60	HDC _i	15A4	Headings for the column of the output tables
3-8	1-8	LABL _i	10A8	60 element vector of labels, each 8 characters long, for up to 60 lines of output from the model
9	3-10	DLAB	A8	Label for M & I consumptive use coefficients card
10	11-15...	PKCMI _i	12F5.3	12 element vector of monthly M & I consumptive use coefficients
	3-10	DLAB	A8	Label for proportion of daylight hours card
11	11-15...	PDL _i	12F5.3	12 element vector of monthly proportion of daylight hours for the modified Blaney-Criddle Crop CU method
	3-10	DLAB	A8	Label for SCS modified B-C crop consumptive use coefficients
	11-15...	CPKC _{g,i}	12F5.3	12 element vector of BC monthly CU coefficients for crop or phreatophyte _g

C. Subbasin data cards (Called with ITY = 1, IPL = 2 or 5)

Card	Column	Name	FORMAT	Description
1	1-2	ISB	I2	Subbasin number (-9, 99)
	3-7	IYRB	I5	Beginning year of simulation
	8-12	NYR	I5	Number of years of simulation
	21-80	BASID	15A4	Descriptive name of basin (<= 60 columns)
2...	11-13	I _i DCA _i	(I3,F7.0)	Crop number and area in acre for crop i. Can put up to 7 sets/ card and requires a set for each crop and phreatophyte
3	11-20	URLND	F10.0	Urban land area
	21-30	UNDLND	F10.0	Undeveloped land area
				Return if IPL = 5
4	1-3	N _i	15I3	15 element vector of number of stations for each of 15 types of hydrologic data
5	1-40	FMT ₁	10A4	Format for temperature data which follows
6	...	T ₁	FMT ₁	Monthly temperature data in °F format specified by FMT ₁ .
7	...	FMT ₂	10A4	Format for precipitation data in inches which follows
8	...	PPT ₂	FMT ₂	Monthly precipitation data in format specified by FMT ₂
9	...	FMT _k	10A4	Format for kth data (k = 3 to 15)
10	...	IX	FMT _k	Power of 10 by which all data on this card are multiplied
	...	DAT _k	FMT _k	Monthly values of kth hydrologic data in acre-feet. Do not include a format and data card if N _k = 0
11				Continue with format and data pairs for all data necessary. The data is specified by k as follows: if N _k = 0 that format and data pair should not be included. k = 3 is River inflow, RIV = 4 is Tributary inflow, TRIB = 5 is Ungaged correlation inflow, QCOR = 6 is Canal division, CNL = 7 is Groundwater inflow, QGI = 8 is Gaged outflow, GAG = 9 is Pumped water from the shallow aquifer, QPUM = 10 is M & I diversion, ΣMIDIV = 11 is Reservoir inflow, RIN = 12 is Required minimum reservoir releases to stream, REL = 13 is Inbasin release, ARD = 14 is Reservoir exports from the basin, REXP = 15 is Basin export from the stream channel, BEXP

Table 1. continued.

D. Reservoir parameter cards (Called with ITY = 1, IPL = 3 or 4). These cards are not included if IRES=0, which indicates that there is no reservoir in the subbasin.

Card	Column	Name	FORMAT	Description	
1-2	11-20	RES ₁	(10X7F10.0)	14 element vector of reservoir parameters	
Reservoir parameters are indexed as follows:					
i = 1	STI	Initial storage in acre-feet		i = 9	A ₂ Reservoir area at zero storage for equation above BPS
= 2	CSTI	not used		=10	C ₃ Coefficient in area vs. storage equation
= 3	STMN	Minimum useable storage		=11	C ₄ Exponent in area vs. storage equation
= 4	STMX	Maximum allowable storage		=12	DSPD Change in gaged storage for the calibration period
= 5	A ₁	Reservoir area at zero storage		=13-14	Not used
= 6	C ₁	Coefficient in area vs. storage equation			
= 7	C ₂	Exponent in area vs. storage equation			
= 8	BPS	Break point storage value between equation defined by A ₁ , C ₁ and C ₂ and the one defined by A ₂ , C ₃ and C ₄			

E. Parameter specification cards (Called when ITY=2*)

Card	Column	Name	FORMAT	Description
1	1-3	NPR	I3	Number of parameters (60)
	4-6	MANG	I3	Management option specification for canal diversions MANG = -1 Calibration mode - use QCNL recorded but limit to water available for diversion, WAD = 0 Calibration mode - Use QCNL as recorded = 1 Management mode - Calculate QCNL and use without limit to satisfy PET = 2 Management mode - Calculate QCNL but limit to WAD = 3 Management mode - Calculate QCNL but put leaching water to zero and limit QCNL to WAD
	7-0	IRES	I3	Reservoir option specification IRES = 0 No reservoir = 1 Tributary or upstream reservoir and requires reservoir inflow, RIN, to be read as input data = 2 Downstream reservoir - sets the reservoir inflow to the computed simulated surface runoff
	10-12	IQGO	I3	Groundwater outflow limiting option: IQGO = 0 Limit routed groundwater outflow, QGO, to be non negative = 1 Allow QGO to take on any value
	13-15	ITMX	I3	Iteration limit for calculating canal seepage
	16-25	TOLF	E10.3	Error criteria for indicating convergence in calculating canal seepage
2	26-35	WH	E10.3	Multiplier for weighting the objective function calculations
	3-10	DLAB ₁	A8	Label for canal diversion option vector
	11-34	IDTM ₁	12I2	Vector of diversion option controls - one for each month. If IDTM = 0 Do not allow diversion = 1 Allow diversion
3	3-10	DLAB ₂	A8	Label for allowable soil moisture storage card
	11-70	CMS ₁	12F5.3	Vector of soil moisture level, CMS, that must be maintained in calculating QCNL when in the management mode
* 4-9	1-80	PR ₁	(10F8.3)	The 60 element vector of model parameters (see Table 2)
*	If ITY = 5, enter here and read only the 6 parameter cards.			

F. Pattern Search Specification Cards (Called when ITY = 6)

Card	Column	Name	FORMAT	Description
1	1-3	NPS	I3	Number of parameters to be searched or if NPS = 300 print the entire output table
	4-6	NPH	I3	Number of phases for pattern search (1-5)
	7-21	NOP ₁	5I3	Vector of options for resetting the initial parameter vector at the completion of each phase. If NOP ₁ = 0 Reset to the original initial parameter vector = 1 Reset to the best local minimized objective function parameter vector
2	1-4	L	I4	Parameter number to be searched
(NPS+1)	4-8	NLL	I4	Number of steps in the search
	9-18	PLL	E10.5	Lower boundary for parameter L
	19-28	PHL	E10.5	Upper boundary for parameter L

Note: If NPS = 1 then will only read one type 2 card and return to read another type 1 card. If NPH is outside the range of (1-5) then will return to read a program option control card.

Table 2. Glossary of parameters used in BSAM.

<u>Par</u>	<u>Name</u>	<u>Description</u>	<u>Units</u>
1	SNI	Initial snow storage	Inches
2	SNK	Snow melt rate coefficient	
3	TSM	Threshold temperature above which snow melts	Deg F
4	TPR	Threshold temperature above which precipitation is rain	Deg F
5	COR	Ungaged flow correlation with correlating stream	Ac-ft/ac-ft
6	CSM	Ungaged flow correlation with snow melt	Ac-ft/inch
7	CRN	Ungaged flow correlation with precipitation as rain	Ac-ft/inch
8	RTH	Runoff threshold for ungaged flow from rain	Inches
9	CTP	Ungaged flow correlation with total precipitation	Ac-ft/inch
10	PTH	Runoff threshold for ungaged flow from total precipitation	Inches
11	SMI	Initial soil moisture level	Inches
12	CSM	Critical soil moisture to limit evapotranspiration	Inches
13	SMC	Soil moisture holding capacity	Inches
14	DTA	Delay time for routing DP	Months
15	DPI	Initial DP rate for routing DP	Inches
16	DPK	DP routing rate coefficient	
17	CDP	Proportion of soil moisture above CSM that goes to DP	
18	PSP	Proportion of spills from canal diversions	
19	ECV	Canal conveyance efficiency	
20	EAP	Irrigation application efficiency	
21	CNI	Initial rate for routing canal seepage to GW	Ac-ft/month
22	CNK	Canal seepage rate routing coefficient	
23	CRT	Proportion of canal seepage that returns to stream	
24	C1	Influent flow coefficient for QRIV = 1.	Ac-ft/month
25	C2	Influent flow coefficient	
26	C3	Influent flow limiting value of $C_1 - C_2 \text{ Log}_{10} (\text{QRIV})$	
27	GW1	Initial rate of groundwater outflow	Ac-ft/month
28	GWK	Groundwater routing rate coefficient	
29	CGO	Proportion of GW outflow from basin	
30	TAJ	Adjusting coefficient on temperature	
31	PAJ	Adjusting coefficient on precipitation	
32	CNA	Adjusting coefficient on canal diversions	
33	GWA	Adjusting coefficient on groundwater inflow	
34	DMIA	Adjusting coefficient for M & I diversion	
35	CUMIA	Adjusting coefficient for M & I CU	
36	CUA	Adjusting coefficient for irrigated land CU	
37	CGW	Change in GW storage during the calibration period	
38	CUPGW	Proportion of CU for phreatophytes from GW	
39	CUUR	Coefficient for CU on urban land	
40	CUUN	Coefficient for CU on undeveloped land	
41	PURGW	Proportion of urban undeveloped land runoff to GW	
42	SCNR	Ratio of actual inbasin reservoir releases to canals	
43	STH	Snowmelt threshold for ungaged flow from snowmelt	Inches
44	PBST	Proportion of bank storage for reservoir operation	
45-48		Not used currently	
49-60	AGW	Values of delayed DP - one for each month of delay up to 12 (driving function for routed DP (ARF))	Inches/month

Table 3. Information printed in the output array

Line	Variable name used in program	Description	Units
1	T	Temperature	Deg F
2	PPT	Precipitation	Inches
3	SNMT	Snowmelt	Inches
4	SNWI	End of month snow storage	Inches
5	PHET	Phreatophyte evapotranspiration	Inches
6	ETP	Potential crop evapotranspiration	Inches
7	AET	Actual crop evapotranspiration	Inches
8	SML	End of month soil moisture storage	Inches
9	RIV	River inflow	Acre-feet
10	TRB	Tributary inflow	Acre-feet
11	QUNG	Ungaged inflow	Acre-feet
12	RMPH	Phreatophyte rain plus snow melt	Acre-feet
13	URSF	Urban and undeveloped land surface runoff	Acre-feet
14	QPUM	Pumped water from the shallow aquifer	Acre-feet
15	STGW	Influent water from the stream channels	Acre-feet
16	PHSF	Phreatophyte use from the surface supply	Acre-feet
17	WAD	Surface water available for diversion	Acre-feet
18	EMIDIV	Municipal and industrial diversion	Acre-feet
19	EMIRF	Municipal and industrial return flows	Acre-feet
20	CNL	Canal diversion	Acre-feet
21	GWNSV	Canal seepage	Acre-feet
22	GWCN	Proportion of the routed canal seepage that goes to groundwater	Acre-feet
23	SEEP	Canal seepage return flow that can be rediverted	Acre-feet
24	SPILL	Canal conveyance and operating spills	Acre-feet
25	QCV	Water delivered to the farm laterals by the canal system	Acre-feet
26	TWTR	Surface runoff from the farms (tailwater return flow)	Acre-feet
27	WAPP	Total water applied to the irrigated area including rain plus snow melt	Acre-feet
28	AETT	Actual crop evapotranspiration	Acre-feet
29	XSM	End of month soil moisture storage	Acre-feet
30	DP	Deep percolation from soil moisture storage	Acre-feet
31	ARF	Routed deep percolation	Acre-feet
32	GWIN	Groundwater inflow	Acre-feet
33	URGW	Groundwater inflow from the urban and undeveloped land area	Acre-feet
34	PHGW	Phreatophyte consumptive use from the groundwater	Acre-feet
35	QGO	Routed groundwater	Acre-feet
36	GEF	Effluent groundwater	Acre-feet
37	CHGW	Change in groundwater storage	Acre-feet
38	GWO	Groundwater outflow	Acre-feet
39	BEXP	Basin surface export	Acre-feet
40	SRF	Surface runoff which is also reservoir inflow if modeling a downstream reservoir (IRES = 2)	Acre-feet
41	QSO or ST	Computed surface outflow or the end of month reservoir storage if modeling a downstream reservoir (IRES = 2)	Acre-feet
42	GAG	Gaged or measured surface outflow or measured end of month reservoir storage if modeling a down- stream reservoir (IRES = 2)	Acre-feet
43	DIFF	Difference between computed and gaged values (Lines 44-53 are printed only if IRES > 0)	Acre-feet
44	EVP	Reservoir evaporation	Inches
45	RIN	Reservoir surface inflow	Acre-feet
46	AV	Average area of the reservoir water surface during the month	Acres
47	RQRL	Required releases to be met if possible	Acre-feet
48	REL	Actual release to channel	Acre-feet
49	ARD	Canal diversion from reservoir for in basin irrigation	Acre-feet
50	REXP	Reservoir export for trans basin use	Acre-feet
51	RSR	Total reservoir releases	Acre-feet
52	ST	End of month reservoir storage	Acre-feet
53	DS	Monthly change in reservoir storage	Acre-feet

Model Requirements

The fundamental requirements of a computer model of a hydrologic system are:

1. It simulates on a continuous basis all important processes and relationships within the system it represents.
2. It is non-unique with respect to space. This implies that it can be easily applied to different geographic areas with existing hydrologic data.
3. It is capable of answering questions concerning perturbations in the system or of accurately predicting outputs resulting from varying input and process parameters.

The Conceptual Model

The basis of the hydrologic model is a fundamental and logical mathematical representation of the various hydrologic processes and routing functions. These physical processes are not specific to any particular geography, but rather are applicable to any hydrologic unit. Experimental and analytical results can be used to assist in testing and establishing some of the mathematical relationships included within the model. Under a model verification procedure, equation constants are established which calibrate or fit the model for a particular drainage area. Average values of hydrologic quantities needed for model verification were estimated in one of three ways: (1) From available data, (2) by statistical correlation techniques, and (3) through verification of the model.

A flow diagram of the hydrologic system is shown by Figure A-1. As this flow chart indicates, the total input to a subbasin is the combination of surface and subsurface inflows of water obtained by summing river and tributary inflows, precipitation, groundwater inflows, and imports from other basins. Depletions from the subbasins occur through evapotranspiration, municipal and industrial consumption, and exports. The residual quantity is a combination of surface and subsurface outflow of water from the area. Subsurface flows may undergo various time delays as they move through the system.

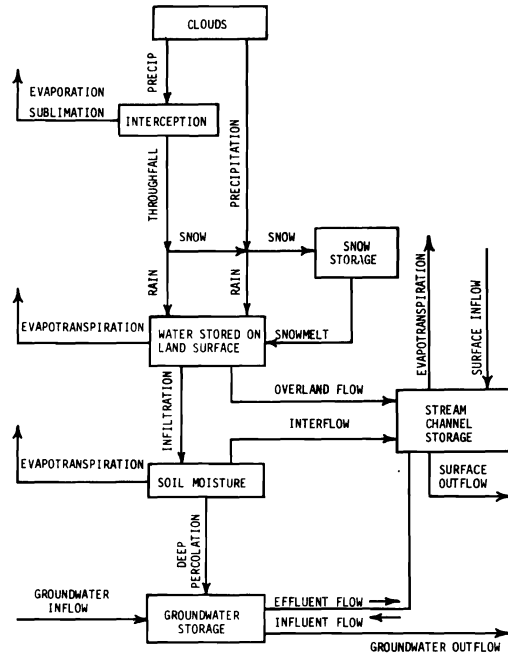


Figure A-1. Flow Diagram of the Hydrologic System

The Hydrologic Balance

A dynamic system consists of three basic components, namely the medium or media acted upon, a set of constraints, and an energy supply or driving force. In a hydrologic system, water in any one of its physical states is the medium of interest. The constraints are applied by the physical nature of the hydrologic basin, and the driving forces are supplied by direct solar energy, gravity, and capillary potential fields. The various functions and operations of the different parts of the system are interrelated by the concepts of continuity of mass and momentum. Unless relatively high velocities are encountered, such as in channel flow, the effects of momentum are negligible, and the continuity of mass becomes the only link between the various processes within the system.

Continuity of mass is expressed by the general equation:

$$\text{Input} - \text{Output} = \text{Change in Storage}$$

A hydrologic balance is the application of this equation to achieve an accounting of physical, hydrologic measurements within a particular unit. Through this means and the application of appropriate translation or routing functions, it is possible to predict the movement of water within a system in terms of its occurrence in space and time.

The concept of the hydrologic balance is pictured by the block diagram in Figure A-1. The inputs to the system are precipitation and surface and ground-water inflow, while the output quantity is divided among surface outflow, groundwater outflow and evapotranspiration. As water passes through this system, storage changes occur on the land surface, in the soil moisture zone, in the groundwater zone, and in the stream channels. These changes occur rapidly in surface locations and more slowly in the subsurface zones.

In the course of model development, each of the system processes must be described mathematically as completely as possible. The flow chart of Figure A-1 is a schematic representation of the system processes and storage locations and their relationship to each other. In the model each box and connecting line is represented by a mathematical expression.

Time and Space Increments

Practical data limitations and problem constraints require that increments of time and space be considered during model design. Data, such as temperature and precipitation readings, are usually available as point measurements in terms of time and space; and integration in both dimensions is usually accomplished by the method of finite increments.

The complexity of a model designed to represent a hydrologic system largely depends upon the magnitude of the time and spatial increments utilized in the model. In particular, when large increments are applied, the scale magnitude is such that the effect of phenomena which change over relatively small increments of space and time is insignificant. For instance, on a monthly time increment, interception rates and changing snowpack temperatures are neglected. In addition, the time increment chosen might coincide with the period of cyclic changes in certain hydrologic phenomena. In this event net changes in these phenomena during the time interval are usually negligible. For example, on an annual basis, storage changes within a hydrologic system are often insignificant, whereas on a monthly basis, the magnitude of these changes are frequently appreciable and need to be considered. As time and spatial increments decrease, improved definition of the hydrologic processes is required. No longer can short-term transient effects or appreciable variations in space be neglected, and the mathematical model, therefore, becomes increasingly more complex with an accompanying increase in the requirements of data, computer capacity, and computer capability.

For this report, a monthly time increment and large space units (subbasins) were adopted. Selection of the subbasins was based on hydrologic boundaries and points of data collection. It was felt that the selection of the subbasins and the monthly time increment would satisfy the requirements of a general planning-management model for the basin.

System Processes

Surface Inflows

The basic inflow or input of water into any hydrologic system originates as a form of precipitation. However, for simulation models of valley floor areas, direct precipitation input to the system is greatly overshadowed by river and tributary inflows.

Streamflow is defined as that portion of the precipitation which appears in streams and rivers as the net or residual flow collected from all or a portion of a watershed. When unaffected by the activities of man, such runoff is referred to as "natural or virgin" flow. Artificial diversions and regulatory action in lakes and reservoirs affect the regimes of nearly every stream within the basin.

The surface water inflow component consists of flow traveling over the ground surface and through channels to enter a stream. At the stream surface runoff usually combines with other flow components to form the total surface runoff hydrograph. Within the runoff cycle surface runoff begins to occur when the capacities of vegetative interception, infiltration, and surface retention are reached. Continued precipitation beyond this point serves as a source for surface runoff. Small basins have different runoff characteristics than do large watersheds, and the characteristics peculiar to each basin must be evaluated on an individual basis.

For each subbasin a limiting rate of surface runoff exists for any particular time period. Surface runoff is assumed to occur when the threshold or limiting rate of surface water supply, consisting of snowmelt, rainfall, canal diversions, or any combination of these, is exceeded.

This concept of surface runoff is particularly important when precipitation is considered as the initial water input to the watershed. Riley and Chadwick (1966) indicate that for particular conditions there exists a limiting or threshold rate of surface supply, R_{tr} , at which surface runoff, S_r , begins to occur. This relationship can be written:

$$S_{wr} = W_{gr} - R_{tr}, (S_{wr} \geq 0) \dots (1)$$

in which

- S_{wr} = rate of surface runoff during a particular time
- W_{gr} = rate at which water is available at the soil surface
- R_{tr} = limiting or threshold rate of surface water supply at which surface runoff begins to occur

When considered for a model time increment of one month, an average value of the threshold surface runoff rate, R_{tr} , is probabilistic in nature, depending essentially upon soil surface conditions, soil moisture, storm characteristics, and rate of available water, W_{gr} .

In this study only the valley bottom lands are considered in the model, and it is assumed that no surface runoff from precipitation occurs from these relatively flat areas. Under this assumption, the rate at which precipitation is available at the soil surface at no time exceeds the threshold rate for surface runoff to occur. Thus,

$$S_{wr} = 0, (W_{gr} \leq R_{tr}) \dots (2)$$

The model does provide for surface runoff from agricultural lands due to irrigation application rates which exceed soil infiltration rates. The runoff quantity constitutes a portion of the irrigation return flow.

Surface runoff from the surrounding watershed areas is concentrated in stream channels, and therefore, enters the model (valley bottom) as tributary flow. That part of the inflow rate which is measured or gaged is designated as $Q_{is}(m)$.

Unmeasured surface inflows to the model are estimated by a correlation technique which considers three hydrologic parameters, namely a gaged tributary inflow rate, precipitation rate, and snowmelt rate. Thus, in functional form:

$$Q_{is}(u) = \int \{q_{is}(m), P_r, W_{sr}\} \dots (3)$$

in which

- $Q_{is}(u)$ = estimated rate of unmeasured surface inflow
- $q_{is}(m)$ = measured rate of surface inflow from a particular tributary area
- P_r = gaged precipitation rate in the form of rain on the valley floor
- W_{sr} = estimated snowmelt rate in terms of water equivalent

If empirical correlation factors are included in the preceding equation, the expression becomes:

$$Q_{is}(u) = k_u q_{is}(m) + k_a P_r + k_b W_{sr} \dots (4)$$

in which k_u , k_a , and k_b are correlation factors re-

lating ungaged surface inflow rate to, respectively, a gaged surface inflow rate, precipitation rate, and snowmelt rate. Each of these factors is established through the model verification process for a particular subbasin.

With reference to the measured tributary inflow rate, $q_{is}(m)$, used in Equation 4, this quantity might refer to either the total measured tributary inflow or a specific stream within the area. The main criterion for selecting the gaged area is that the watershed exhibit the same general runoff characteristics as that of the ungaged area.

The second independent term in Equation 4 refers to the rates of precipitation occurring on the valley floor in the form of rain. Because it is assumed that the influence of snow upon the surface runoff is restricted to the melt period, only rainfall is considered by the equation. Generally, a direct plot of rainfall against runoff for individual storms yields a low correlation because of the variable nature of the factors affecting runoff (Chow, 1964). However, when mean monthly values of precipitation and runoff are considered, many of the transient processes are smoothed and reasonably good correlation of runoff with precipitation is achieved.

The third independent term of Equation 4 considers the influence of snowmelt upon surface runoff. Snowmelt rates on the valley floor are predicted in the model by Equation 8. This process is discussed in further detail later in this chapter.

The total surface inflow rate to the model (valley floor) is estimated by summing the measured rate and estimated ungaged rate from Equation 4. Thus,

$$Q_{is} = Q_{is}(m) + Q_{is}(u) \dots (5)$$

in which Q_{is} refers to the total surface inflow rate, and the two independent quantities are as previously defined.

Interflow

Interflow is defined as the lateral movement of moisture through the plant root zone. The process is discussed in further detail at a later point in this chapter. Interflow rate, N_r , is not treated as a separate identity in the hydrologic model of the valley bottoms, but is considered as being a part of the surface runoff from irrigation. In most cases, small interflow rates are encountered in flat lands. Furthermore, for a model time increment of one month, interflow usually produces an insufficient delay time to enable this quantity to be distinguished from surface runoff.

Groundwater Inflow

Groundwater or subsurface inflow refers to those waters which enter the model area or valley floor beneath the ground surface. Much of this water is subsequently discharged as effluent flow into the main channel of the valley, and thus provides a "base flow" for the stream. Discharge from the groundwater basin of the valley floor also occurs by way of spring flows, pumped waters, and consumptive use by phreatophytes.

Essentially, all groundwater is constantly in motion though velocities may range from several feet per day to only a few inches per year. This groundwater movement is basically confined to permeable geologic formations called aquifers which serve as transmission conduits. Movement and volume of groundwater runoff may be calculated through application of Darcy's Law, providing adequate data are available. However, if subsurface flow data are sparse, time and spatial distribution of groundwater flows can be estimated by an empirical approach through the model verification procedure. For the steep watersheds near the headwaters of major drainage divisions, groundwater inflows to the valley floors were usually sufficiently small to be neglected. At downstream subbasins it may become apparent through the time distribution of the water inputs to the model in relation to the recorded outflow that groundwater inflow rates are appreciable. Correlation procedures and transport delays can be used to estimate and simulate groundwater movement into the subbasin. This water is then distributed with time through use of long transport delay networks on the computer. The required delay setting of these networks is established during the model verification process. Hence, the rate of groundwater inflow is described as follows:

$$Q_{ig}(u) = k_c q_{is} \dots \dots \dots (6)$$

in which

- $Q_{ig}(u)$ = rate of total unmeasured inflow to the groundwater system
- k_c = coefficient relating the rate of unmeasured groundwater inflow to a measured surface runoff rate
- q_{is} = rate of surface runoff (either total measured tributary inflow or measured inflow from a representative tributary)

For some subbasins a subsurface groundwater movement under the outflow gage in the streambed alluvium occurs. The time and spatial distribution of this outflow forms a component of the groundwater inflow or input to the adjacent downstream subbasin.

Total Inflow

Total inflow rate to the valley bottoms consists of the summation of the surface and groundwater inflow rates. The surface inflows for the most part have already been summed and are concentrated in stream channels as they enter the valley floor or agricultural areas. Gaged surface inflow rates are available from surface water records published by the U. S. Geological Survey.

The remaining two components of total inflow, namely unged surface inflow and groundwater inflow, are estimated from Equation 4 and 6, respectively. Therefore, the total inflow, Q_i , to a given subbasin is given by the following expression:

$$Q_i = Q_{is} + Q_{ig} \dots \dots \dots (7)$$

in which the terms Q_{is} and Q_{ig} are given by Equations 5 and 6, respectively.

Precipitation

The ultimate source of water input to any hydrologic unit is precipitation in one form or another. Precipitation is considered to be any moisture which emanates from the atmosphere and falls to the earth.

Precipitation input to the hydrologic system varies with respect to both time and space and it is therefore, necessary to convert point measurements from climatological stations into an integrated or averaged monthly value over a finite area. Some precipitation data are available for all subbasins used in this study.

Two forms of precipitation, rain and snow, are considered in this study. Air temperature is used as the criterion for establishing the occurrence of these two forms. This criterion is not an ideal index for determining the form of precipitation since no one temperature exists below which it always snows and above which it always rains. However, the surface air temperature appears to be the most suitable single index of precipitation form now available.

Based on investigations by the U. S. Army (1956) at a surface air temperature of 35°F, there is a 50 percent chance that precipitation will be as snow, whereas at 32°F the probability is 95 percent of precipitation falling as snow. In this study a double standard was used because average monthly temperatures provided the criteria. Snow was assumed to fall below a threshold temperature, and snowmelt was assumed to occur above another threshold temperature. This assumption provided for the occurrence of snowfall and snowmelt during the same month. These threshold temperatures varied in the different subbasins.

Temperature

Air temperature is an important parameter in a hydrologic system because it can be utilized as a criterion for establishing the form of precipitation, and as an index of the energy available for the snowmelt and evapotranspiration processes. Temperature varies according to both time and space. To obtain average temperature values for the valley floor, or a portion thereof, within a particular subbasin, requires that point measurements be utilized for estimating an effective or average temperature value for an area. In this study average temperatures for a particular area and a given time period (one month) are estimated by an arithmetic average of temperature measurements taken in the subbasin.

Snowmelt

Although rational formulas which include the various factors involved in the snowmelt process have been developed, data limitations often prohibit a strictly analytical approach to the process. Rational models include fundamental processes, such as those which relate to energy transfer, and requirements for input data are high. An additional restriction to the analytical approach for snowmelt computation is a large modeling time increment such as a month. Many of the short-term, transient phenomena which occur within a snowpack cannot be represented in a macroscopic model of this scale. An empirical relationship was, therefore, adopted for this study model. Riley et al. (1966) proposed a relationship which states that the rate of melt is proportional to the available energy and the quantity of precipitation stored as snow. As a differential equation the relationship is written:

$$\frac{d W_s(t)}{dt} = - k_s (T_a - T_b) \frac{RI_s}{RI_h} W_s(t) \quad . \quad . \quad (8)$$

in which the undefined terms are:

- k_s = a constant
- T_a = surface air temperature in degrees F
- RI_s = the radiation index on a surface possessing a known degree and aspect of slope
- RI_h = the radiation index for a horizontal surface at the same latitude as the particular watershed under study
- W_s = snow storage in terms of water equivalent
- T_b = assumed base temperature in degrees F at which melt begins to occur.

Riley et al. (1966) report reasonable agreement between predicted snowmelt rates from Equation 8 and observed values. They used a value of k_s equal to 0.10 based on studies using data from several snow courses in the Rocky Mountain area where average snow depths are high. It has been found, however, that the value of k_s is somewhat inversely dependent upon snowpack depth. In other words, as the snow depth decreases pack melt rates increase for a given energy input. Thus, k_s is relatively larger for areas of shallow snowpack depth and relatively smaller for areas where depths tend to be large. The radiation index parameter allows adjustment to be made for variation of the total insolation with land surface slope and aspect. However, since only the valley floors are included in the modeling area, it is assumed that the topographic surface of the area is essentially horizontal. This assumption simplifies Equation 8 in that RI_s becomes equal to RI_h and their ratio goes to unity.

The independent variables on the right side of Equation 8 can be expressed as either continuous functions of time or as step functions consisting of mean constant values for a given time increment. For this study a time increment was utilized and integration was performed in steps over each successive time period. Hence, the final values of $W_s(t)$ at the end of a particular time period became the initial value for the integration process over the following period. On this basis, and setting the ratio RI_s/RI_h equal to unity (assuming the agricultural area is flat), the differential form of Equation 8 becomes:

$$\int_{W_s(0)}^{W_s(1)} \frac{dW_s}{W_s} = - k_s (T_a - T_b) \int_0^1 dt \quad . \quad . \quad (9)$$

$$W_s(1) = W_s(0) \exp \{-k_s (T_a - T_b)\} \quad . \quad . \quad (10)$$

Canal Diversions

Canal diversions profoundly affect the time and spatial distribution of water within an irrigated area. A portion of this diverted water is returned directly to the stream as spill and portion is evaporated directly to the atmosphere. Another portion enters the soil profile through canal seepage and the infiltration on the irrigated lands, and the remainder returns to the source as overland flow or evaporates. Some of the water which enters the soil profile is lost through plant consumptive use. The remainder either percol-

ates downward to the groundwater basin or is intercepted by drainage systems. Irrigation practices, therefore, alter the distribution characteristics of a hydrologic system. The irrigation efficiency factors used in this study include both the conveyance and application efficiencies. A composite irrigation efficiency factor is given by the following expression:

$$\text{Eff} = (W_{\text{tr}} - W_{\text{sp}}) E_c E_a \dots \dots \dots (11)$$

in which

- Eff = water conveyance and application efficiency in percent
- W_{sp} = the spills
- E_c = the conveyance efficiency
- E_a = application efficiency
- W_{tr} = total rate at which water is diverted from the stream or reservoir

Records of water diversion to the agriculture lands within each subbasin were found to be lacking. Adjustment of these records was necessary in many cases to get a realistic application rate for the irrigated acreage.

As already indicated, a portion of the water diverted for irrigation returns to the streams as overland flow and interflow. Although the large time increment allows this water to be treated in the model as a single identity, it is important to distinguish between the two components. Overland flow (often termed tailwater) is surface return flow or runoff from the end of the field resulting from the application of water to the irrigated land at rates exceeding the infiltration capacity of the soil. Interflow is defined as that part of the soil water which does not enter the groundwater basin, but rather which moves largely in a lateral direction through the upper and more porous portion of the soil profile until it enters a surface or subsurface drainage channel. Both the overland flow and the interflow return to the stream channels in short distances and times consisting of usually only a few days.

Available Soil Moisture

The usual definition of available soil moisture capacity is the difference between the field capacity and the wilting point of the soil. Water within this range is available for plant use, and is termed available soil moisture. The field capacity is defined as the soil moisture content after gravity drainage has occurred. Most of the gravity water drains rapidly from the soil thus affording plants little opportunity for its use. The wilting point represents the soil

moisture content when plants are no longer able to abstract water in sufficient quantities to meet their needs, and permanent wilting occurs. Available soil moisture can be expressed in several units but in this report it carries the unit of depth in inches.

Sources of available water. Basically, moisture in the soil is derived from infiltration, which is the passage of water through the soil surface into the soil profile. The water available for infiltration at the soil surface is derived from three sources, namely, precipitation in the form of rain, P_r , snowmelt, W_{sr} , and irrigation water, W_{dr} . As springtime temperatures rise to the point at which melting occurs, all snow cover on the land is assumed to melt and enter the soil mantle through the infiltration process. In the case of irrigated crops, the most important source of available soil moisture is water which is diverted to the agriculture lands. Thus, the total water available for infiltration into the soil, W_{gr} , can be written as:

$$W_{\text{gr}} = W_{\text{dr}} + P_r + W_{\text{sr}} \dots \dots \dots (12)$$

in which all quantities are as previously defined.

Available soil moisture quantities. The maximum quantity of water in a soil available for use by plants is a function of the moisture holding capacity of the soil and the average rooting depth or extraction pattern of the plant.

The basic forces involved in the absorption of water by plants are osmotic, imbibitional, metabolic, and transpiration pull (Thorne and Peterson, 1954). These forces basically define the soil moisture tension or "pull" that must be exerted by the plant to remove water from the soil. Of these forces the principal one is the osmotic pressure created within plant root cells. Opposing these forces are those exerted on the moisture by the soil particles. The forces exerted by the plants vary with different plants, soils, and climates, but the average maximum force which plants can exert in obtaining sufficient water for growth is approximately 15 atmospheres of pressure. At field capacity where water is readily available for plant use, the average soil moisture tension is only about 0.1 atmosphere. However, the soil moisture tension or "pull" plants exert in their quest for water is in itself no indication of the amount of available water contained by the soil. The actual amount of water held by the soil at any given tension value is a function of the soil type.

Determination of the soil depth effectively utilized by a plant is based on the average rooting depth or the average moisture extraction pattern. The soil moisture available for extraction depends on the mois-

ture holding capacity of the soil and the extraction pattern. The typical agricultural crop extracts 70 percent of its moisture from the upper 50 percent of the soil penetrated by the plant roots. Average or typical rooting depths for various plants are reported by McCulloch et al. (1967). Illustrative depths include 4 to 6 feet for alfalfa, 4 feet for grains and corn, and 2 to 3 feet for pasture. The average available soil moisture capacity of the irrigated lands was estimated for each subbasin.

Under normal circumstances, additions to available soil moisture storage occur through the infiltration process, F_r . Abstractions or depletions from available soil moisture storage occur through evapotranspirational losses, ET_r , and deep percolation, G_r . Thus, the soil moisture storage existing at any time, t , can be stated:

$$M_s(t) = (F_r - ET_r - G_r) dt \quad . \quad . \quad . \quad (13)$$

Each of the three terms on the right side of this equation is discussed in the following sections.

Infiltration

As already indicated, additions to available soil moisture occur through the process of infiltration, F_r . Factors which influence the infiltration rate include various soil properties and surface characteristics. A moisture gradient induced by the adhesive properties of the soil particles also influence infiltration rate.

In this study, the rate of infiltration into the soil is given by the following equations

$$F_r = W_{gr}, (W_{gr} \leq R_{tr}) \quad . \quad . \quad . \quad (14)$$

and

$$F_r = R_{tr}, (W_{gr} > R_{tr}) \quad . \quad . \quad . \quad (15)$$

for which all terms were previously defined. The quantity W_{gr} in Equation 14 is given by Equation 12.

Evapotranspiration

The second term on the right side of Equation 13 represents depletion from the soil moisture storage through the evapotranspiration process, ET_r . Consumptive use, or evapotranspiration is the sum of all water used and lost by growing vegetation due to transpiration through plant foliage and evaporation from the plant and surrounding environment such as adjacent soil surfaces. Potential evapotranspiration is defined as that rate of consumptive use by actively growing plants which occurs under conditions of complete crop cover and non-limiting soil moisture supply.

The evapotranspiration process depends upon many interrelated factors whose individual effects are difficult to determine. Included among these factors are type and density of crop, soil moisture supply, fertilizer, soil salinity, and climate. Climatological parameters usually considered to influence evapotranspiration rates are precipitation, temperature, daylight hours, solar radiation, humidity, wind velocity, cloud cover, and length of growing season. Numerous relationships have been developed for estimating the potential evapotranspiration rate.

Perhaps one of the most universally applied evapotranspiration equations is that proposed by Blaney and Criddle (1950). This equation is written as follows:

$$U = kf \quad . \quad . \quad . \quad . \quad . \quad . \quad . \quad . \quad (16)$$

in which

U = monthly crop potential consumptive use in inches

k = monthly coefficient which varies with type of crop and

F = monthly consumptive use factor and is given by the following equation:

$$f = \frac{tp}{100} \quad . \quad . \quad . \quad . \quad . \quad . \quad . \quad . \quad (17)$$

in which

t = mean monthly temperature in degrees F

p = monthly percentage of daylight hours of the year

A modification of the Blaney-Criddle formula was proposed by Phelan et al. (1962), wherein the monthly coefficient, k , is subdivided into two parts, a crop coefficient, k_c , and a temperature coefficient, k_t . The relationship describing k_t is an empirical one, depending upon only temperature, and is expressed as:

$$k_t = (0.0173T_a - 0.314) \quad . \quad . \quad . \quad . \quad (18)$$

where T_a is the mean monthly temperature in degrees F. The crop coefficient, k_c , is basically a function of the physiology and stage of growth of the crop. Curves of k_c are available for many agriculture crops (Soil Conservation Service, 1964).

Thus, the modified Blaney-Criddle equation for estimating potential evapotranspiration rates is written as follows:

$$ET_{cr} = k_c k_t \frac{T_a p}{100} \quad . \quad . \quad . \quad . \quad . \quad . \quad . \quad . \quad (19)$$

Because of its simplicity, low data requirements (only surface air temperature is needed), and applicability to the irrigated areas of the Western United States, Equation 19 was adopted for this study model. Since the time increment selected for use was one month, the variables on the right of Equation 19 represent mean monthly values although these parameters could

be expressed as continuous functions instead of the indicated step functions. Thus, Equation 19 estimates the mean potential evapotranspiration rate during each month.

Evapotranspiration losses from the agriculture area during the non-cropping season were estimated from Equation 19. For many crops it was necessary to extend the k_c curves to include the non-growing season (West, 1959). Because the k_c curve for grass pasture seems to represent a reasonable set of values for native vegetation (Riley et al., 1967), this curve was used as a guide in the development of a similar k_c curve for phreatophytes.

Effects of soil moisture on evapotranspiration.

As was discussed earlier, as the moisture content of a soil is reduced by evapotranspiration, the moisture tension which plants must overcome to obtain sufficient water for growth is increased. It is generally conceded that some reduction in the evapotranspiration rate occurs as the available quantity of water decreases in the plant root zone. Studies by the U. S. Salinity Laboratory in California (Gardner and Ehlig, 1963) indicate that transpiration occurs at the full potential rate through approximately the first one-third of the available soil moisture range, and that thereafter the actual evapotranspiration rate lags the potential rate. When this critical point in the available moisture range is reached, the plants begin to wilt because soil moisture becomes a limiting factor.

The actual evapotranspiration rate is approximated by the two following equations which take the above phenomenon into account:

$$ET_r = ET_{cr}, \{M_{es} < M_s(t)\} \dots (20)$$

and

$$ET_r = ET_{cr} \left(\frac{M_{cs} - M_{es}}{M_{cs}} \right), \{0 \leq M_s(t) \leq M_{es}\} \dots (21)$$

in which

- ET_r = actual evapotranspiration rate
- ET_{cr} = potential evapotranspiration rate
- M_{es} = limiting or threshold content of available water within the root zone below which the actual becomes less than the potential evapotranspiration rate
- $M_s(t)$ = quantity of water available for plant consumption which is stored in the root zone at any instant of time
- M_{cs} = root zone storage capacity of water available to plants

Deep Percolation

The final independent term, G_r , of Equation 13 represents the rate of deep percolation. Percolation is simply the movement of water through the soil. Deep percolation is defined as water movement through the soil from the plant root zone to the underlying groundwater basin. The dominant potential forces causing water to percolate downward from the plant root zone are gravity and capillary forces. Water is removed quickly by gravity from a saturated soil under normal drainage conditions. Thus, the rate of deep percolation, G_r , is most rapid immediately after irrigation when the gravity force dominates, and decreases constantly, continuing at slower rates through the unsaturated conditions. Because the capillary potential applies through all moisture regimes, deep percolation continues, though at low rates, even when the moisture content of the soil is less than field capacity (Willardson and Pope, 1963).

Because of a lack of data in the study area regarding deep percolation rates in the unsaturated state, and in order to simplify the model, the assumption was made that deep percolation occurs at a continual small percent, k_p , of soil moisture storage between saturation and the limiting threshold storage. Thus, for this model, the deep percolation rate is expressed as:

$$G_r = F_r - ET_{cr}, \{M_s(t) = M_{cs}\} \dots (22)$$

$$G_r = k_p M_{sg} \{M_{es} < M_s(t) < M_{cs}\} \dots (23)$$

$$G_r = 0, \{M_s(t) < M_{cs}\} \dots (24)$$

in which all terms are as previously defined.

River Outflow

Using the continuity of mass principle the hydrologic balance is maintained by properly accounting for the quantities of flow at various points within the system. The appropriate translation or routing of inflow water through the system in relationship to the chronological abstraction and additions occurring in space and time concentrates the water at the outlet point as both surface and subsurface outflow. As mentioned earlier, active network delays on the computer simulate the long transport time necessary for groundwater inflows and deep percolating waters to be routed to the outflow gaging station.

Thus, the total rate of water outflow from a sub-basin is obtained through the summation of various quantities as follows:

$$Q_o = Q_{is} - W_{tr} + OF_r + Q_{ob} - Q_e \dots (25)$$

in which

- Q_o = total rate of outflow from the system
- Q_{is} = rate of total surface inflow to the subbasin including both measured and unmeasured flows
- W_{tr} = total rate at which water is diverted from the stream or reservoir
- OF_r = total of overland flow and interflow rates
- Q_{ob} = rate of outflow from the groundwater basin of routed deep percolating waters and subsurface inflows to the subbasin
- Q_e = rate of water diversions from surface sources for use outside the boundaries of the subbasin. Exports to other drainage basins fall within this category

If subbasins are selected such that there exists no flow of subsurface water past the gaged outflow point, the hydrograph of surface outflow, Q_{so} , is given by Equation 25. This situation is assumed to exist at reservoir sites within the basin because of construction measures taken to eliminate subsurface flows under the dams which create the reservoir. For this reason, whenever possible, subbasins were terminated at the outfall of a reservoir. These sites thus enabled a check to be made on groundwater inflow rates to the subbasin as predicted from verification studies involving models for one or more upstream subbasins.

For many subbasins the termination or outlet point was taken at a Geological Survey gaging station, and in several of these cases groundwater flow occurs in the streambed alluvium beneath the surface channel. For these basins, the total system outflow can be written as:

$$Q_{to} = Q_{so} + Q_{go} \dots \dots \dots (26)$$

in which

- Q_{so} = rate of surface outflow from the subbasin
- Q_{go} = rate of subsurface or groundwater outflow from the subbasin

Surface outflow rates, Q_{so} , can be compared to the recorded values, but subsurface outflow rates, Q_{go} , are unmeasured and must be predicted or estimated.

In the model described in this report, the groundwater outflow is estimated as a proportion of the total groundwater flux. The groundwater flux is computed by keeping track of all inflows and effluent flows that interface with the shallow groundwater aquifer. These include influent river channel flow, canal seepage, deep percolation, pumped water and groundwater inflow. The net effect of

these variables provides the driving function for a linear reservoir routing subroutine. Flow routing through a linear reservoir is derived as follows from the continuity and linear storage equations:

$$\frac{dS}{dt} = I - O \dots \dots \dots (27)$$

and

$$S = k_s O \dots \dots \dots (28)$$

where

- $\frac{dS}{dt}$ = change in storage
- I = inflow
- O = outflow
- S = storage
- k_s = the linear routing rate coefficient

By substituting equation 28 into 27 and assuming I is constant $\frac{dO}{I-O} = \frac{1}{k_s} dt$ is obtained which may then be integrated to give

$$\int_0^1 \frac{dO}{I-O} = \frac{1}{k_s} \int_0^1 dt$$

or

$$\ln (I-O) \Big|_0^1 = - \frac{1}{k_s} t \Big|_0^1 \dots \dots \dots (29)$$

By evaluating at the limits of integration, exponentiating both sides and solving for O_1 the final routing equation is obtained:

$$O_1 = I + (O_0 - I) e^{-\frac{1}{k_s} t} \dots \dots \dots (30)$$

The routed value of groundwater flux Q_{ob} , is then separated into an effluent portion, Q_{ef} , which makes up part of the surface outflow and the groundwater outflow, Q_{go} , by the equation:

$$Q_{go} = k_{gw} Q_{ob} \dots \dots \dots (31)$$

and

$$Q_{ef} = Q_{ob} - Q_{go} \dots \dots \dots (32)$$

where k_{gw} is the proportion of groundwater flux that remains as groundwater flow.

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Appendix B. Input to and Output from a Typical Simulation Run

Listing of Input Data Cards

```

1 1 0 1
8 5
VAR OCT NOV DEC JAN FEB MAR APR MAY JUN JUL AUG SEP ANN
FHR TEMP PRECIPSNOW MLTSNOW STR PHR ETCROP PETCROP AET SM STRRIVER IN TRIB IN
UNG INPHR RPSM URB SRF PUMP INRIVER GW PHR SUR WTR AVL M&I DIV M&I RET CNL DIV
CNL SEEP CNL GWSEEP RTN SPILLFARM DELTAIL WTR TOT APLCROP AET SM STR DP
ROUT DP GW INURBGW IN PHR GW ROUT GW EFPL GWCH GWSTR GW OUTCHNL EXPSUR RNOF
COMP OUTGAGE OUT DIFFRES EVAP RES INRES AREA REQ REL ACT REL CNL RELRES EXPT
TOT RELEOM STRGCHG STRG
PKCMI 0.1 0.1 0.1 0.1 0.1 0.1 0.1 0.1 0.1 0.1 0.1 0.1
PDL .0783,0688,0674,0692,0682,0834,0887,0985,0989,1005,0944,0837
1ALFALFA 91 78 64 63 73 86 99 108 113 111 106 99
2PAS+OHAY 85 73 60 56 65 80 92 99 103 102 99 93
3SPR GRN 25 25 25 25 25 25 25 52 105 128 57 25
4CORN+SOR 58 25 25 25 25 25 25 36 58 92 110 100
5DRY BEAN 25 25 25 25 25 25 25 25 64 106 100 48
6POTATOES 25 25 25 25 25 25 25 31 55 108 136 131
7TR CROPS 25 25 25 25 25 25 25 31 64 81 80 59
8COV ORCH 90 78 66 63 73 86 98 109 113 111 106 99
C1 HWC 1 185 185 185 185 185 185 185 185 185 185 185
C2 HWG 2 125 102 75 65 80 113 136 141 142 142 141 136
C3 HWG 3 81 76 96 96 53 53 55 61 68 77 82 83
C4 LWGT 4 50 47 42 35 33 33 34 38 42 48 51 51
B1 OPWT 5 100 100 100 100 100 100 100 100 100 100 100 100
1 2 0 1
7 1972 1 SAN JUAN ABOVE BLUFF
SUB7 CROPS 1 29400. 2 25200. 3 7000. 4 5600. 5 700. 6 0. 7 700.
CROP7 CONT 8 1400. 9 0. 10 0. 11 0. 12 0. 13 2818.
UR-UN LAND 0. 0.
2 2 1 4 1 1 1 0 0 0 0 0 0
(20X,12F5,0)
72 CORTEZ, COLO. 48.5 35.9 25.9 29.8 35.0 44.7 48.3 56.0 66.8 72.8 69.5 62.8
72 BLUFF, UTAH 51.4 38.3 25.9 30.8 38.3 49.7 54.6 61.1 72.0 78.0 75.7 67.2
(20X12F5,0)
72 CORTEZ 2.12 0.84 1.74 0.00 0.00 0.15 0.13 0.14 0.67 0.33 0.45 1.91
72 BLUFF, UTAH 1.39 1.22 1.66 0.00 0.00 0.00 0.00 0.00 0.35 0.50 0.09 0.60
(6X,12,12F6,0)
723680 3 61.29 44.29 95.51116.30104.70126.10 62.39 83.20 99.63 13.64 26.21 46.08
(6X,12,12F6,0)
723685 0 878. 863. 447. 353. 345. 2100. 1800. 3270. 2820. 670. 390. 570.
723690 0 170. 185. 150. 120. 140. 1200. 490. 720. 390. 98. 53. 105.
723695 0 52. 82. 62. 46. 40. 620. 280. 250. 110. 52. 11. 28.
72 IMP 8800. 8800. 5500. 7400.17000.29700.24400.10600. 6400. 5000.
(6X,12,12F6,0)
723710 2120 1500 2060 1010 1480 1090 94 79 287 0 76 328
(6X,12,12F6,0)
72 CD7 0 17000 850 810 760 760 8730 35935 77375 65170 44140 24253 20695
(6X,12,12F6,0)
720H6 4311. 296. 166. 136. 113. 60. 1417. 5296. 3739. 3222. 3944. 4571.
(6X,12,12F6,0)
723795 3 99.53 59.2 110.2 119.2 109.5 119.5 64.61 81.10118.1 17.06 30.60 56.33
2 0 0 1
60 -1 0 0 75 1. 1.E-06
IDTM 1 1 1 1 1 1 1 1 1 1 1 1 1
CMS 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0. 0.
0. .48 28. 30. 3. 5000. 20000. 0. 1.2 0. 0.
4.0 3.0 6. 1. 0. 0. .05 0. .6 0.
.5 1.0 1. 0. 0. 0. 0. 0. 0. 1.
1. 1. 1. 1. 1. 1. 0. .3 0.
0. 0. 0. 0. 0. 0. 0. 0. 0.
0. 0. 0. 0. 0. 0. 0. 0. 0.
6 -1 0 1
3 2 0 0
5 10 0.0 5.0
6 20 0.0 10000.
7 20 15000. 25000.

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BLANK CARD

Output Produced from the Foregoing Input Data

VAR	OCT	NOV	DEC	JAN	FEB	MAR	APR	MAY	JUN	JUL	AUG	SEP	ANN
1 FHR TEMP	2	PRECIP	3	SNOW MLT	4	SNOW STR	5	PHR ET	6	CROP PET	7	CROP AET	8
13 URB SRF	14	PUMP IN	15	RIVER IN	16	PHR SUR	17	WTR AVL	18	M&I DIV	19	M&I RET	20
25 FARM DEL	26	TAIL WTR	27	TOT APL	28	CROP AET	29	SM STR	30	DP	31	ROUT DP	32
37 CH GWSTR	38	GW OUT	39	CHNL EXP	40	SUR RNOF	41	COMP OUT	42	GAGE OUT	43	DIFF	44
49 CNL REL	50	RES EXPT	51	TOT REL	52	EOM STRG	53	CHG STRG	54		55		56
KCMI	.100	.100	.100	.100	.100	.100	.100	.100	.100	.100	.100	.100	.100
PDL	.078	.068	.067	.069	.068	.083	.088	.098	.098	.100	.094	.083	
1ALFALF	.910	.780	.640	.630	.730	.860	.990	1.080	1.130	1.110	1.060	.990	
2PAS+OH	.850	.730	.600	.560	.650	.800	.920	.990	1.030	1.020	.990	.930	
3SPR GR	.250	.250	.250	.250	.250	.250	.250	.520	1.050	1.280	.570	.250	
4CORN+S	.580	.250	.250	.250	.250	.250	.250	.360	.580	.920	1.100	1.000	
5DRY BE	.250	.250	.250	.250	.250	.250	.250	.640	1.060	1.000	.480		
6POTATO	.250	.250	.250	.250	.250	.250	.250	.310	.550	1.000	1.360	1.310	
7TR CRO	.250	.250	.250	.250	.250	.250	.250	.310	.640	.800	.590		
8COV DR	.900	.780	.660	.630	.730	.860	.980	1.090	1.130	1.110	1.060	.990	
9 HWMC	1.850	1.850	1.850	1.850	1.850	1.850	1.850	1.850	1.850	1.850	1.850	1.850	
10 HWGW	1.250	1.020	.750	.650	.800	1.130	1.360	1.410	1.420	1.420	1.410	1.360	
11 MWGT	.810	.760	.960	.560	.530	.530	.550	.610	.680	.770	.820	.830	
12 LWGT	.500	.470	.420	.350	.330	.330	.340	.380	.420	.480	.510	.510	
13 OPWT	1.000	1.000	1.000	1.000	1.000	1.000	1.000	1.000	1.000	1.000	1.000	1.000	

SUB 7 1972 TO 1972 SAN JUAN ABOVE BLUFF

LAND AREAS	CROP LAND	PHR LAND	URB LAND	UND LAND	TOTAL LAND							
	70000.	2818.	0.	0.	72818.							
INCH TO AC-FT CONVERSION FACTORS	5833.333007	234.833312	.000000	.000000	6868.168815							
CROP AND PHREATOPHYTE ACRES	1 29400.	2 25200.	3 7000.	4 5600.	5 700.	6 0.	7 700.	8 1400.	9 0.	10 0.	11 0.	12 0.
	13 2818.											
PROP CROP AND PHR	1 .419999	2 .359999	3 .099999	4 .079999	5 .010000	6 .000000	7 .010000	8 .020000	9 .000000	10 .000000	11 1.000000	12 .000000
WEIGHTED CROP AND PHR COEFFICIENTS	.782	.656	.548	.528	.605	.716	.816	.918	1.032	1.075	.985	.886
	1.000	1.000	1.000	1.000	1.000	1.000	1.000	1.000	1.000	1.000	1.000	1.000

HYDROLOGIC DATA	1972	1972	1972	1972	1972	1972
1 1972	49.95	37.10	25.90	30.30	36.65	47.20
	51.45	58.55	69.40	75.40	72.60	65.00
	619.49					
2 1972	1.75	1.03	1.70	.00	.00	.07
	.06	.07	.51	.41	.27	1.25
	7.14					
3 1972	61289.99	44289.99	95509.98	116299.98	104699.98	126099.98
	82389.99	83199.98	99829.98	13639.99	26289.99	46079.99
	879339.00					
4 1972	9900.00	9930.00	6159.00	519.00	525.00	11380.00
	20370.00	33940.00	27720.00	11420.00	6854.00	5703.00
	144420.00					
5 1972	2120.00	1590.00	2060.00	1010.00	1480.00	1090.00
	94.00	79.00	287.00	.00	76.00	328.00
	10214.00					
6 1972	17000.00	850.00	810.00	760.00	760.00	8730.00
	35935.00	77375.00	65170.00	44140.00	24253.00	20695.00
	296478.00					
7 1972	4311.00	296.00	166.00	136.00	113.00	60.00
	1417.00	5296.00	3739.00	3222.00	3944.00	4571.00
	27271.00					
8 1972	99529.98	59199.99	110199.98	119199.98	109500.00	119500.00
	64809.98	81099.98	118099.98	17079.99	30599.99	56329.99
	984949.12					

Output (Continued)

SAN JUAN ABOVE BLUFF

PAR	1	2	3	4	5	6	7	8	9	10
	.000	.400	28.000	30.000	3.000	5000.000	20000.0	1.200	.000	.000
	4.000	3.000	6.000	1.000	.000	.000	.050	.000	.600	.800
	.500	1.900	1.000	.000	.000	.000	.000	.000	.000	1.000
	1.000	1.000	1.000	1.000	1.000	1.000	.000	.300	.000	.000
	.000	.000	.000	.000	.000	.000	.000	.000	1.000	.000
	.000	.000	.000	.000	.000	.000	.000	.000	.000	.000
IDTM	1	1	1	1	1	1	1	1	1	1
CMS	.000	.000	.000	.000	.000	.000	.000	.000	.000	.000

SAN JUAN ABOVE BLUFF WATER

YEAR 1978

VAR	OCT	NOV	DEC	JAN	FEB	MAR	APR	MAY	JUN	JUL	AUG	SEP	ANN
1 PHR TEMP	49.840	37.899	88.899	38.388	38.649	47.189	81.449	88.549	88.399	73.399	72.899	64.899	51.624
2 PRECIP	1.788	1.838	1.788	.000	.000	.078	.000	.518	.518	.418	.278	1.255	7.144
3 SNOW MLT	.000	.000	.000	1.882	.000	.000	.000	.000	.000	.000	.000	.000	1.700
4 SNOW STR	.000	.000	1.700	.077	.000	.000	.000	.000	.000	.000	.000	.000	.000
5 PHR ET	2.151	.836	.823	.029	.799	1.078	2.829	4.838	8.888	7.888	6.455	4.489	38.935
6 CROP PET	1.883	.848	.888	.332	.484	1.417	2.148	3.781	6.281	6.039	3.718	2.182	29.643
7 CROP AET	1.883	.848	.888	.332	.484	1.417	2.148	3.781	6.281	6.039	3.718	2.182	29.643
8 SM STR	8.888	8.888	8.888	8.888	8.888	8.888	8.888	8.888	8.888	8.888	8.888	8.888	8.888
9 RIVER IN	61289	44288	88888	116289	104689	126889	82389	83189	98629	13839	88289	48879	879339
10 TRIS IN	8888	8888	8188	818	825	11388	88378	33848	27788	11428	8854	5783	144428
11 UNG IN	17488	4778	8188	8142	7788	3378	882	237	881	881	228	284	51341
12 PHR RPHM	412	241	8	248	154	22	18	18	119	97	83	8	1877
13 URB BRP	8	8	8	8	8	8	8	8	8	8	8	8	8
14 PUMP IN	8	8	8	8	8	8	8	8	8	8	8	8	8
15 RIVER GW	8	8	8	8	8	8	8	8	8	8	8	8	8
16 PHR SUR	88	88	88	88	23	389	421	881	918	1188	1818	518	5174
17 MTR AVL	81387	88818	188848	128879	113834	142337	88888	133491	147888	43184	47883	65785	1178287
18 MBI DIV	8	8	8	8	8	8	8	8	8	8	8	8	8
19 MBI RET	8	8	8	8	8	8	8	8	8	8	8	8	8
20 CNL DIV	17888	888	818	788	788	8738	38838	77378	88178	43184	24233	28895	295442
21 CNL DEEP	8888	388	384	384	384	3492	14374	38888	28888	17241	9781	8278	18178
22 CNL GW	8	8	8	8	8	8	8	8	8	8	8	8	8
23 DEEP RTN	2783	1783	1188	825	811	1788	8838	18788	28872	18288	15318	12437	18822
24 OPILL	8	8	8	8	8	8	8	8	8	8	8	8	8
25 PARM DEL	18189	818	488	488	488	5238	21588	48424	38181	28882	14518	12418	177888
26 TAIL MTR	2848	182	87	81	81	1847	4312	8888	7828	5172	2918	2483	35453
27 TOT APL	18387	8418	388	6328	4182	4758	17888	37848	34258	23118	13218	17254	183491
28 CROP AET	8882	3882	1874	1948	2824	8267	12823	21888	38841	48479	21893	12282	172828
29 SM STR	38888	32838	38854	34258	34579	28872	33572	34888	38827	18878	2888	8878	8878
30 OP	1211	888	788	884	1847	1881	1588	14831	1781	2887	328	81	28833
31 ROUT DP	8833	1211	888	788	884	1847	1881	1588	14831	1781	2887	328	81
32 SM IN	4311	288	188	138	113	88	1417	8888	3738	3222	3844	4871	27271
33 URB IN	8	8	8	8	8	8	8	8	8	8	8	8	8
34 PHR GW	27	8	38	8	18	132	188	279	382	499	435	222	2217
35 ROUT GW	18118	1887	1189	828	1887	874	2387	8822	17877	5318	3888	4878	37728
36 EPFL GW	18118	1887	1189	828	1887	874	2387	8822	17877	5318	3888	4878	37728
37 CH GWTR	8	8	8	8	8	8	8	8	8	8	8	8	8
38 GW OUT	8	8	8	8	8	8	8	8	8	8	8	8	8
39 CHNL EXP	8	8	8	8	8	8	8	8	8	8	8	8	8
40 SUR RNDP	8884	81878	188345	128137	118882	138838	88888	71823	188384	18482	31541	52282	888827
41 COMP OUT	8884	81878	188345	128137	118882	138838	88888	71823	188384	18482	31541	52282	888827
42 GAGE OUT	88888	89189	118189	118189	188888	118888	64888	81888	118888	118888	38888	58329	884848
43 DIFP	-13888	2378	-884	8837	4482	18138	-4344	-8178	-8888	-8887	841	-4877	-18821

OBJ = 761.97 OAH = -18821.8

PAR	NL	PL	PH	DF
5	18	.00000	3.00000	5.00000
6	28	.00000	18888.00000	588.00000
7	28	18888.00000	28888.00000	588.00000

PAR	1	2	3	4	5	6	7	8	9	10
	.000	.400	28.000	30.000	3.000	5000.000	20000.0	1.200	.000	.000
	4.000	3.000	6.000	1.000	.000	.000	.050	.000	.600	.800
	.500	1.900	1.000	.000	.000	.000	.000	.000	.000	1.000
	1.000	1.000	1.000	1.000	1.000	1.000	.000	.300	.000	.000
	.000	.000	.000	.000	.000	.000	.000	.000	1.000	.000
	.000	.000	.000	.000	.000	.000	.000	.000	.000	.000
IP	LV	PAR	OBJ	OAH	GRAD					
5	1	.000	902.194	-47563.882	.000					
5	2	.500	858.979	-42456.882	-86.429					
5	3	1.000	823.782	-37349.882	-70.554					
5	4	1.500	796.363	-32242.878	-54.677					
5	5	2.000	776.982	-27135.878	-38.882					
5	6	2.500	765.498	-22028.886	-22.927					
5	7	3.000	761.972	-16921.886	-7.852					
5	8	3.500	766.384	-11814.886	8.823					
5	9	4.000	778.733	-6707.886	24.698					
5	10	4.500	799.828	-1688.886	48.574					
6	1	.000	695.384	-25421.855	.000					
6	2	500.000	698.649	-24571.855	.006					
6	3	1800.000	702.731	-23721.855	.008					
6	4	1500.000	707.552	-22871.855	.009					
7	1	1500.000	844.173	-19971.878	.000					
7	2	1550.000	835.253	-19866.886	-.017					
7	3	1600.000	826.489	-19361.886	-.017					
7	4	1650.000	817.888	-19056.886	-.017					
7	5	1700.000	809.426	-18751.886	-.016					

Output (Continued)

7	6	17500.000	801.128	-18446.886	-.016
7	7	18000.000	792.986	-18141.886	-.016
7	8	18500.000	784.999	-17836.886	-.015
7	9	19000.000	777.168	-17531.886	-.015
7	10	19500.000	769.492	-17226.886	-.015
7	11	20000.000	761.972	-16921.886	-.015
7	12	20500.000	754.807	-16616.886	-.014
7	13	21000.000	747.398	-16311.886	-.014
7	14	21500.000	740.344	-16006.886	-.014
7	15	22000.000	733.446	-15701.886	-.013
7	16	22500.000	726.704	-15396.886	-.013
7	17	23000.000	720.116	-15091.886	-.013
7	18	23500.000	713.685	-14786.886	-.012
7	19	24000.000	707.409	-14481.886	-.012
7	20	24500.000	701.288	-14176.886	-.012
7	21	25000.000	695.323	-13871.886	-.011

SAN JUAN ABOVE BLUFF WATER

YEAR 1978

VAR	OCT	NOV	DEC	JAN	FEB	MAR	APR	MAY	JUN	JUL	AUG	SEP	ANN
1 FHR TEMP	49.949	37.899	25.899	38.300	36.649	47.199	51.449	68.549	69.399	75.399	72.899	64.999	51.624
2 PRECIP	1.755	1.038	1.708	.000	.000	.075	.000	.078	.518	.415	.278	1.255	7.144
3 SNOW HLT	.000	.000	.000	1.022	.055	.021	.000	.000	.000	.000	.000	.000	1.708
4 SNOW STR	.000	.000	1.708	.877	.000	.000	.000	.000	.000	.000	.000	.000	.000
5 PHR ET	2.151	1.536	.523	.629	.799	1.078	2.029	4.038	6.085	7.505	6.455	4.409	38.835
6 CROP PET	1.083	.548	.286	.332	.484	1.417	2.148	3.781	6.281	8.074	6.364	3.987	35.229
7 CROP AET	1.083	.548	.286	.332	.484	1.417	2.148	3.781	6.281	8.074	6.364	3.987	35.229
8 SH STR	5.282	8.845	5.289	5.872	5.927	5.138	5.785	6.888	8.147	1.884	3.718	2.194	4.494
9 RIVER IN	61289	44289	95889	118299	184899	128899	82399	83199	98829	13839	28889	48879	878339
10 TRFD IN	9989	9938	6159	519	525	11389	28378	33848	27728	11428	6854	5783	144428
11 UNG IN	28235	4778	8188	3838	4448	3278	282	237	861	8	228	2359	48892
12 PHR RPSM	412	241	0	248	154	22	15	10	119	97	63	294	1877
13 URB BRP	0	0	0	0	0	0	0	0	0	0	0	0	0
14 PUMP IN	0	0	0	0	0	0	0	0	0	0	0	0	0
15 RIVER GW	0	0	0	0	0	0	0	0	0	0	0	0	0
16 PHR SUR	65	0	86	0	23	38	421	691	916	1185	1816	518	5174
17 WTR AVL	94148	68818	188948	128768	118253	142231	89588	133491	147888	43184	47593	68888	1164837
18 HAI DIV	0	0	0	0	0	0	0	0	0	0	0	0	0
19 HAI RET	0	0	0	0	0	0	0	0	0	0	0	0	0
20 CNL DIV	17889	888	818	788	788	8738	38935	77375	65178	43184	24253	28885	285442
21 CNL SEEP	6888	348	324	384	384	3492	14374	38888	28888	17241	9781	8278	118178
22 CNL SH	0	0	0	0	0	0	0	0	0	0	0	0	0
23 SEEP RYN	2783	1783	1188	825	611	1798	8938	18788	28872	18288	15318	12437	188222
24 SPILL	0	0	0	0	0	0	0	0	0	0	0	0	0
25 FARM DEL	18199	518	488	458	458	5238	21588	48424	38181	28882	14551	12416	177285
26 TAIL WTR	2848	182	97	91	91	1847	4312	9888	7828	5172	2818	2485	35453
27 TOT APL	18397	6418	388	6328	4192	4752	17828	37548	34258	23118	13818	17254	183491
28 CROP AET	9822	3282	1874	1948	2884	8287	1528	21889	38841	48478	81893	12282	172828
29 SH STR	38898	38938	38884	34259	34878	29878	33572	34889	38827	18878	2888	6878	6878
30 DP	1211	988	788	884	1847	1881	1585	14531	2587	1781	329	91	28833
31 ROUT DP	5833	1211	988	788	884	1847	1881	1585	14531	2587	1781	329	28833
32 GW IN	4311	288	188	138	113	68	1417	5888	3739	3222	3844	4871	27271
33 URB GW IN	0	0	0	0	0	0	0	0	0	0	0	0	0
34 PHR GW	27	0	38	0	18	132	188	279	392	489	438	222	2217
35 ROUT GW	18118	1887	1188	828	1887	874	2327	6822	17877	5318	5288	4878	57728
36 EPFL GW	18118	1887	1188	828	1887	874	2327	6822	17877	5318	5288	4878	57728
37 CH GW STR	0	0	0	0	0	0	0	0	0	0	0	0	0
38 GW OUT	0	0	0	0	0	0	0	0	0	0	0	0	0
39 CNL EXP	0	0	0	0	0	0	0	0	0	0	0	0	0
40 SUR RNOF	89299	61578	189345	121824	118871	135823	68888	71923	188394	18482	31541	52527	982577
41 COMP OUT	89299	61578	189345	121824	118871	135823	68888	71923	188394	18482	31541	52527	982577
42 GAGE OUT	98299	58199	118199	118199	189588	119588	64888	81889	118899	17879	38899	56329	984848
43 DIFF	-18238	2378	-854	1824	1171	1823	-4344	-8178	-8785	-6597	941	-3882	-22371

		PHASE 2 PMIN= 695.384									
PAR	1	2	3	4	5	6	7	8	9	10	
	.000	.400	28.000	38.000	3.000	.000	2500.0	1.200	.000	.000	
	4.000	3.000	6.000	1.000	.000	.000	.050	.000	.600	.800	
	.500	1.900	1.000	.000	.000	.000	.000	.000	.000	1.000	
	1.000	1.000	1.000	1.000	1.000	1.000	.000	.300	.000	.000	
	.000	.000	.000	.000	.000	.000	.000	.000	1.000	.000	
	.000	.000	.000	.000	.000	.000	.000	.000	.000	.000	
IP	LV	PAR	OBJ	GAH	GRAD						
5	1		793.851	-53013.851	.000						
5	2	.500	746.474	-47906.851	-94.753						
5	3	1.000	707.835	-42799.851	-78.878						
5	4	1.500	675.533	-37692.859	-63.002						
5	5	2.000	651.970	-32585.855	-47.127						
5	6	2.500	636.344	-27478.855	-31.251						
5	7	3.000	628.655	-22371.855	-15.376						
5	8	3.500	628.985	-17264.855	.498						
5	9	4.000	637.892	-12157.855	16.374						
5	10	4.500	653.217	-7050.855	32.250						
6	1	.000	628.655	-22371.855	.000						
6	2	500.000	632.000	-21521.855	.006						
6	3	1000.000	636.883	-20671.855	.008						
6	4	1500.000	640.984	-19821.855	.009						
7	1	15000.000	777.585	-28471.847	.000						
7	2	15500.000	768.585	-28166.855	-.017						
7	3	16000.000	759.821	-27861.855	-.017						
7	4	16500.000	751.212	-27556.855	-.017						
7	5	17000.000	742.758	-27251.855	-.016						

Output (Continued)

7	6	17500.000	734.460	-26946.855	-.016
7	7	18000.000	726.318	-26641.855	-.016
7	8	18500.000	718.331	-26336.855	-.015
7	9	19000.000	710.500	-26031.855	-.015
7	10	19500.000	702.824	-25726.855	-.015
7	11	20000.000	695.304	-25421.855	-.015
7	12	20500.000	687.939	-25116.855	-.014
7	13	21000.000	680.730	-24811.855	-.014
7	14	21500.000	673.676	-24506.855	-.014
7	15	22000.000	666.778	-24201.855	-.013
7	16	22500.000	660.036	-23896.855	-.013
7	17	23000.000	653.448	-23591.855	-.013
7	18	23500.000	647.017	-23286.855	-.012
7	19	24000.000	640.741	-22981.855	-.012
7	20	24500.000	634.620	-22676.855	-.012
7	* 21	25000.000	628.655	-22371.855	-.011

SAN JUAN ABOVE BLUFF WATER

YEAR 1972

VAR	OCT	NOV	DEC	JAN	FEB	MAR	APR	MAY	JUN	JUL	AUG	SEP	ANN
1 PHN TEMP	49.948	37.889	25.000	36.388	36.640	47.199	51.440	58.540	69.399	75.399	72.599	64.999	51.624
2 PRECIP	1.755	1.038	1.700	.000	.000	.075	.000	.078	.318	.415	.270	1.255	7.144
3 SNOW MLY	.000	.000	.000	1.922	.056	.051	.000	.000	.000	.000	.000	.000	1.700
4 SNOW STR	.000	.000	1.700	.077	.021	.000	.000	.000	.000	.000	.000	.000	.000
5 PHN ET	2.151	.030	.923	.029	.799	1.078	2.029	4.038	6.085	7.585	6.455	4.489	36.035
6 CROP PET	1.603	.548	.206	.332	.484	1.417	2.146	3.781	6.281	8.074	6.364	3.987	35.229
7 CROP AET	1.603	.548	.206	.332	.484	1.417	2.146	3.781	6.281	8.074	6.364	3.987	35.229
8 SM STR	5.202	5.845	5.209	5.072	5.027	5.138	5.755	6.000	5.147	1.854	.354	1.194	4.454
9 RIVER IN	61289.	44289.	95589.	110289.	104889.	120889.	92389.	83189.	99289.	13839.	26289.	46879.	87339.
10 TRID IN	9989.	9938.	6159.	519.	585.	11389.	28379.	33949.	27729.	11429.	6854.	5783.	144429.
11 UNG IN	88235.	4778.	6189.	3838.	4448.	3278.	282.	237.	881.	0.	228.	2359.	45892.
12 PHN RPSM	412.	241.	0.	248.	154.	22.	15.	19.	119.	97.	63.	294.	1677.
13 URB BRP	0.	0.	0.	0.	0.	0.	0.	0.	0.	0.	0.	0.	0.
14 PUMP IN	0.	0.	0.	0.	0.	0.	0.	0.	0.	0.	0.	0.	0.
15 RIVER GM	0.	0.	0.	0.	0.	0.	0.	0.	0.	0.	0.	0.	0.
16 PHN SUR	89.	0.	89.	0.	83.	389.	421.	691.	916.	1165.	1816.	518.	5174.
17 MTR AVL	94142.	68818.	108948.	128768.	118253.	142231.	89588.	133491.	147888.	43184.	47593.	86888.	1164837.
18 MBI DIV	0.	0.	0.	0.	0.	0.	0.	0.	0.	0.	0.	0.	0.
19 MBI REY	0.	0.	0.	0.	0.	0.	0.	0.	0.	0.	0.	0.	0.
20 CNL DIV	17889.	856.	819.	768.	788.	8738.	38938.	77378.	65178.	43184.	24233.	2869.	28442.
21 CNL SEEP	6889.	348.	324.	384.	384.	3492.	14374.	38958.	28668.	17241.	9781.	8278.	118176.
22 CNL GM	0.	0.	0.	0.	0.	0.	0.	0.	0.	0.	0.	0.	0.
23 SEEP RTN	2793.	1793.	1189.	829.	611.	1798.	6939.	10769.	28572.	19289.	15318.	12437.	108222.
24 SPILL	0.	0.	0.	0.	0.	0.	0.	0.	0.	0.	0.	0.	0.
25 FARM DEL	18189.	518.	488.	488.	488.	5838.	81588.	46424.	39181.	25882.	14591.	12418.	177265.
26 TAIL MTR	2848.	188.	97.	91.	91.	1847.	4312.	9888.	7828.	5172.	2918.	2483.	35433.
27 TOT APL	18397.	8418.	388.	8328.	4182.	4758.	17888.	37548.	34258.	23118.	13218.	17254.	183491.
28 CROP AET	9822.	3282.	1674.	1948.	2884.	8887.	18523.	21589.	36641.	48479.	21693.	12262.	172928.
29 SM STR	38696.	32938.	38854.	34259.	34579.	29972.	33872.	34999.	38827.	18876.	2889.	6978.	6978.
30 DP	1211.	988.	798.	984.	1847.	1891.	1895.	14931.	2587.	1781.	329.	91.	26933.
31 ROUT DP	8833.	1211.	988.	798.	984.	1847.	1891.	1888.	14531.	2587.	1781.	329.	32873.
32 GM IN	4311.	898.	188.	138.	113.	68.	1417.	5888.	3739.	3222.	3944.	4371.	27271.
33 URBGM IN	0.	0.	0.	0.	0.	0.	0.	0.	0.	0.	0.	0.	0.
34 PHN GM	27.	0.	36.	0.	18.	132.	188.	279.	392.	499.	435.	222.	2217.
35 ROUT GM	18116.	1587.	1189.	988.	1887.	974.	2327.	6882.	17877.	5318.	5288.	4678.	57729.
36 EPFL GM	18116.	1587.	1189.	988.	1887.	974.	2327.	6882.	17877.	5318.	5288.	4678.	57729.
37 CM GMSTR	0.	0.	0.	0.	0.	0.	0.	0.	0.	0.	0.	0.	0.
38 GM OUT	0.	0.	0.	0.	0.	0.	0.	0.	0.	0.	0.	0.	0.
39 CHNL EXP	0.	0.	0.	0.	0.	0.	0.	0.	0.	0.	0.	0.	0.
40 SUR ROP	89299.	61579.	189348.	121824.	118871.	138923.	68888.	71823.	188394.	18488.	31541.	52527.	982577.
41 COMP OUT	89299.	61579.	189348.	121824.	118871.	138923.	68888.	71823.	188394.	18488.	31541.	52527.	982577.
42 GAGE OUT	99229.	59199.	118199.	118199.	189588.	118988.	64888.	81889.	118889.	17872.	38589.	56329.	984849.
43 DIFF	-18838.	2378.	-854.	1824.	1171.	18823.	-4344.	-8178.	-9788.	-8897.	941.	-3882.	-22371.

INITIAL VECTORS

PHASE OBJ	1	2	3	4	5	6
761.972						
PAR						
1	.000	.000	.000			
2	.400	.400	.400			
3	28.000	28.000	28.000			
4	30.000	30.000	30.000			
5	3.000	3.000	3.000			
6	5000.000	.000	.000			
7	20000.000	25000.000	25000.000			
8	1.200	1.200	1.200			
9	.000	.000	.000			
10	.000	.000	.000			
11	4.000	4.000	4.000			
12	3.000	3.000	3.000			
13	6.000	6.000	6.000			
14	1.000	1.000	1.000			
15	.000	.000	.000			
16	.000	.000	.000			
17	.050	.050	.050			
18	.000	.000	.000			
19	.600	.600	.600			
20	.800	.800	.800			

Output (Continued)

21	.500	.500	.500
22	1.900	1.900	1.900
23	1.000	1.000	1.000
24	.000	.000	.000
25	.000	.000	.000
26	.000	.000	.000
27	.000	.000	.000
28	.000	.000	.000
29	.000	.000	.000
30	1.000	1.000	1.000
31	1.000	1.000	1.000
32	1.000	1.000	1.000
33	1.000	1.000	1.000
34	1.000	1.000	1.000
35	1.000	1.000	1.000
36	1.000	1.000	1.000
37	.000	.000	.000
38	.300	.300	.300
39	.000	.000	.000
40	.000	.000	.000
41	.000	.000	.000
42	.000	.000	.000
43	.000	.000	.000
44	.000	.000	.000
45	.000	.000	.000
46	.000	.000	.000
47	.000	.000	.000
48	.000	.000	.000
49	1.000	1.000	1.000
50	.000	.000	.000
51	.000	.000	.000
52	.000	.000	.000
53	.000	.000	.000
54	.000	.000	.000
55	.000	.000	.000
56	.000	.000	.000
57	.000	.000	.000
58	.000	.000	.000
59	.000	.000	.000
60	.000	.000	.000

Appendix C. FORTRAN Listing of the Computer Program BSAM

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C GENERAL HYDROLOGY MODEL *BSAM* 7/15/78
COMMON/BLK1/LABL(80),AGW(24),CMS(12),PKCHI(12),IDTM(12),BASID(15),
1RES(14),PR(120),PDL(12),CKC(12),CPKC(12),CPKC(25,12),WDG(20),
2OUT(8,13)/BLK2/MANG,NCRPH,ISB,NSB,IYRB,NVR,IYRN,KR,IRES,IQGO,W,H,
3XMN,KXV,IMH,IMX,SHAX,SMIN,CONV,CONPV,CONUR,CONUN,TOLF,DUM(13),ITMX
4/BLK3/HD(15,12,1)
DOUBLE PRECISION LABL,DLAB1,DLAB2
DIMENSION XIN(6,100),XMN(100),XPM(100),DF(100),OBJ(6),NOP(6),
5IPL(100),PH(100),NL(100),NP(100)
KR=6
1 READ(KR,200)ITY,IPL,IPRT,ICOL
200 FORMAT(5I3,2E10,3)
IF(ITY.LE.0.OR.ITY.GT.6)GO TO 99
3 GOTO(4,5,8,15,8,25)ITY
C EXPLANATION OF OPTIONS
C ITY=1 READ DATA WITH DATPR
C ITY=2 SUBBASIN ONLY IPRT=0 ECHO PRINT INPUT
C IPRT=1 SUPPRESS PRINT
C ITY=3 REPEAT WITH SAME DATA
C IPRT AND ICOL SAME AS UNDER ITY=2
C ITY=4 INTERACTIVE SEARCH MODE
C IPL = MANG OPTION
C IPRT AND ICOL SAME AS ITY =2
C ITY=5 READ PARAMETERS ONLY=OPERATE AND PRINT
C IPL NOT USED
C IPRT AND ICOL SAME AS ITY=2
C ITY=6 PATTERN SEARCH MODE
C IPL=MANG OPTION
C IPRT AND ICOL SAME AS ITY=2
C READ DATA WITH DATPR
4 CALL DATPR(IPL,IPRT,ICOL)
GO TO 1
C READ SUBBASIN DATA=PARAMETERS=OPERATE AND PRINT
5 READ(KR,200)NPR,MANG,IRES,IQGO,ITMX,TOLF,W,H
WRITE(6,227) (BASID(I),I=1,15)
227 FORMAT(14I20A4)
WRITE(6,200) NPR,MANG,IRES,IQGO,ITMX,TOLF,W,H
READ(KR,201)DLAB1,(IDTM(I),I=1,12)
201 FORMAT(2XAB,12I2)
READ(KR,202)DLAB2,(CMS(I),I=1,12)
202 FORMAT(2XAB,12F5,3)
C READ PARAMETERS ONLY OPTION ITY=3
6 READ(KR,203) (PR(L),L=1,NPR)
203 FORMAT(10F8,3)
WRITE(6,226) (I,I=1,10)
226 FORMAT(1X3HPAR14,9I8)
WRITE(6,204) (PR(I),I=1,NPR)
204 FORMAT(1X10F8,3)
WRITE(6,201)DLAB1,(IDTM(I),I=1,12)
WRITE(6,202)DLAB2,(CMS(I),I=1,12)
8 DO 10 I=1,NPR
XIN(1,I)=PR(I)
10 CONTINUE
CALL HYDSM(IPRT,ICOL,OBJ,DAH)
IF (MANG.LE.0)WRITE(6,206)OBJ,DAH
GO TO 1
C INTERACTIVE SEARCH MODE
15 MANG=IPL
16 CALL INTACT(NPR,NL,IPRT,ICOL,PR)
GO TO 1
C PATTERN SEARCH MODE
25 MANG=IPL
CALL HYDSM(4,1,OBJ,DAH)
OBJ=OBJ
NLM=0
C INITIALIZE PH, PL, AND NL
DO 26 I=1,NPR
PH(I)=0.
PL(I)=0.
NL(I)=1.
26 CONTINUE
27 READ(KR,100)NPS,NPH,(NOP(L),L=1,NPH)
100 FORMAT(25I3)
IF(NPH.LE.0.OR.NPH.GT.5) GO TO 1
NPS IS NO OF PARAMETERS TO BE SEARCHED
NPH IS NO OF PHASES = NOP IS RESETTING OPTION FOR EACH PHASE
IF NOP = 0 RESET TO INITIAL, IF = 1 RESET TO BEST LOCAL
IF(NPS.NE.500) GO TO 277
CALL HYDSM(IPRT,ICOL,OBJ,DAH)
WRITE(6,206) OBJ,DAH
206 FORMAT(5H0BJ *F15.2,5X5HDAH *F15.1)
GO TO 27
277 DO 28 I=1,NPS
READ(KR,101) L,NLL,PLL,PHL
101 FORMAT(2I4,2E10,3)
IF(L.LE.0.OR.L.GT.60) GO TO 27
NP(I)=L
NL(L)=NLL
PL(L)=PLL
PH(L)=PHL
28 CONTINUE
OBJ(1)=OBJ
AH = DAH
C INITIALIZE MINIMUM CONDITIONS
PHMN=OBJ
PRMN=OBJ
DO 30 L=1,NPR
RL=NL(L)
XMN(L)=XIN(1,L)
XPM(L)=XIN(1,L)
DF(L)=(PH(L)-PL(L))/RL
30 CONTINUE
C WRITE OUT HEADINGS FOR PAR NL PL PH AND DF
WRITE(6,102)
102 FORMAT(14H 4X,3HPAR,3X2HNL10X2HPL13X2HPH13X2HDF)
DO 31 L=1,NPS
L=NP(L)
WRITE(6,103) L,NL(L),PL(L),PH(L),DF(L)
103 FORMAT( 1X17,I5,3F15.5)
35 CONTINUE
IF(NPS.GT.1) GO TO 39
J=NP(1)
NLM=NL(J)+1
IF(NLO.LE.0.) GO TO 27
WRITE(6,106)
106 FORMAT(3X2HIP3X2HVL4X3HPAR12X3HOBJ12X3H0AH11X4HGRAD)
DO 38 I=1,NLO
XI=I-1
IF(DF(J),EQ.0.)GO TO 27
PR(J)=PL(J)+DF(J)*XI
OBJ=OBJ
CALL HYDSM(4,1,OBJ,DAH)
GRAD=0.
IF(I.EQ.1)GO TO 37
IF(DF(J),EQ.0.) GO TO 37
GRAD=(OBJ-OBJ1)/DF(J)
C CHECK FOR CONSECUTIVE IMPROVEMENT AND IF NOT IMPROVED FOR 3 TIMES
C THEN TERMINATE SEARCH
IF(GRAD.LT.0.)GO TO 36
NLM=NLM+1
GO TO 37
36 IF(NLM.GT.0)NLM=0
37 WRITE(6,107) J,I,PR(J),OBJ,DAH,GRAD
107 FORMAT(2X13,2X13,F10.3,3F15.3)
IF(NLO.EQ.2) GO TO 27
IF(NLM.GT.3)GO TO 27
38 CONTINUE
GO TO 27
C BEGIN PHASE LOOP
39 DO90 K=1,NPH
C TAKE NEW PAGE WRITE PHASE ONE INITIAL VECTOR
WRITE(6,104K)PHMN,(I,I=1,10)
104 FORMAT(14I20X5HPHASE13,2X5HPMIN=F15.3/1X3HPAR12,9I8)
WRITE(6,105) (XIN(K,L),L=1,NPR)
105 FORMAT(1X10F8,3)
WRITE(6,106)
C BEGIN PAR LOOP
DO 80 JJ=1,NPS
J=NP(JJ)
NLM=NL(J)+1
IF(NLO.LE.2) GO TO 80
D1=XIN(K,J)+.0005
D2=XIN(K,J)-.0005
C BEGIN INCR LOOP
NLM=0
DO 70 I=1,NLO
IF(I.GT.1) GO TO 40
OBJ=OBJ
GRAD=0.
DS=DF(J)
IF(D3.EQ.0.)GO TO 60
XIN(I-1)
PR(J)=PL(J)+DS*XI
D3=PR(J)
IF(D3-D2)40,41,41
41 IF(D3-D1)42,42,40
42 IF(NOP(K))40,60,40
C OPERATE MODELS AND DETERMINE OBJECTIVE FUNCTION
49 CALL HYDSM(4,1,OBJ,DAH)
GRAD=0.
IF(I.EQ.1)GO TO 47
GRAD=(OBJ-OBJ1)/DS
IF(GRAD.LT.0.)GO TO 46
NLM=NLM+1
GO TO 47
46 IF(NLM.GT.0)NLM=0
47 WRITE(6,107)J,I,PR(J),OBJ,DAH,GRAD
GO TO 65
60 OBJ=OBJ(K)
GRAD=0.
IF(I.EQ.1)GO TO 64
GRAD=(OBJ-OBJ1)/DS
IF(GRAD.LT.0.)GO TO 63
NLM=NLM+1
GO TO 64
63 IF(NLM.GT.0)NLM=0
64 WRITE(6,108)J,I,XIN(K,J),OBJ,DAH,GRAD
108 FORMAT(2X13,2H *I3,F10.3,3F15.3)
65 OBJ=OBJ
C IF NEW PAR, INITIALIZE LOCAL MIN
IF(I.GT.1) GO TO 67
PRMN=OBJ
XMN(J)=PR(J)
C CK LOCAL AND PHASE MINS
GO TO 51
67 IF(OBJ-PRMN)50,51,51
50 PRMN=OBJ
XMN(J)=PR(J)
51 IF(OBJ-PRMN)52,55,55
52 PRMN=OBJ
DOB3L=1,NPR
53 XPM(L)=PR(L)
C IF NO IMPROVEMENT IN OBJ FOR 3 CONSECUTIVE TIMES, GO TO NEXT

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Program Listing (Continued)

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C   PARAMETER
55 IF(NLM,GE,3)GO TO 68
70 CONTINUE
C   RESFY PR(J) TO FIXED LEVEL FOR NEXT PAR.
68 IF(NOP(K)-1)71,72,72
71 PR(J)=XIN(K,J)
GO TO 88
72 PR(J)=XMN(J)
80 CONTINUE
NPOP=POP(K)
IF(NPOP,EQ,1) GO TO 83
81 DOB2L=1,NPR
82 PR(L)=XMN(L)
CALL HYDSM(4,1,OBJ,DAH)
PRMN=CRJ
C   SELECT BEST VECTOR FOR NEXT PHASE
83 IF(PRMN-PHMN)84,86,86
84 DOB5L=1,NPR
XIN(K+1,L)=XMN(L)
85 PR(L)=XMN(L)
GO TO 88
86 DOB7L=1,NPR
XIN(K+1,L)=XPM(L)
87 PR(L)=XPM(L)
88 CALL HYDSM(IPRT,ICOL,OBJ,DAH)
OBJ(K+1)=OBJ
AM=DAH
90 CONTINUE
WRITE(OUT)INITIAL VECTOR TABLE
NHP=NHP+1
WRITE(6,109)(L,L=1,6),(OBJ(L),L=1,NHP)
109 FORMAT(1M127X15MINITIAL VECTORS/1X3PHASE18,5I13/1X3HOBJ5F13.3,
1F12.3)
WRITE(6,110)
110 FORMAT(12X3MPAR/)
DOO1 L=1,NPR
91 WRITE(6,111)L,(XIN(M,L),M=1,NHP)
111 FORMAT(1X13,8F12.3)
GO TO 1
99 STOP
END

C   INTERACTIVE CALIBRATION SUBROUTINE
SUBROUTINE INTACT(NPR,NL,IPRT,ICOL,PR)
DIMENSION NL(1),PR(1)
CALL HYDSM(4,1,OBJ,DAH)
3 NPR1=NPR
IL=8
C   SSM OPTIONS
D ON STAY IN INTERACTIVE MODE
C   IF SSM E ON SUPPRESS PRINT
OCT 23410
J .2
J .1
2 WRITE(6,111)(I,I=1,10),(PR(I),I=1,NPR)
111 FORMAT(/1X3MPAR14,9I8/(1X10F8.3))
WRITE(6,112)OBJ,DAH
112 FORMAT(5H OBJ=F15.5,10X4HDAH=F15.1)
1 TYPE 100
100 FORMAT(28HENTER PARAMETER NO AND VALUE/)
ACCEPT 101,L,D
101 FORMAT(15,F15.5)
IF(L,LE,0,OR,L,GT,NPR1)GO TO 15
C   IF SSM E ON SUPPRESS PRINT
OCT 23440
J .5
J .10
5 WRITE(6,107)L,PR(L),D
107 FORMAT(11M CHANGE PARI3,5M FROMF15.5,3M TOF15.5)
10 TYPE 102,L,PR(L),D
102 FORMAT(10HCHANGE PARI3,5M FROMF15.5,3M TOF15.5)
PR(L)=D
IL=IL+1
NL(IL)=L
GO TO 1
15 IF(L,EQ,200) GO TO 68
IF(L,NE,300)GO TO 20
CALL HYDSM(IPRT,ICOL,OBJ,DAH)
C   IF L = 300 PRINT ALL TABLES WHEN RUNNING HYDSM
GO TO 90
20 IF(IL,LE,0) GO TO 1
IF SSM E ON SUPPRESS PRINT
OCT 23410
J .25
J .35
25 WRITE(6,110)(NL(I),I=1,IL)
110 FORMAT(1X8MPAR CHNG24I3)
35 TYPE 112,OBJ,DAH
45 CALL HYDSM(4,1,OBJ,DAH)
IL=8
46 TYPE 112,OBJ,DAH
C   IF SSM E ON SUPPRESS PRINTING
OCT 23410
J .50
J .1
50 WRITE(6,112) OBJ,DAH
GO TO 1
C** ITERATION MODE
60 TYPE 103
103 FORMAT(47HENTER PAR, NO STEPS, INIT VALUE AND FINAL VALUE/)
ACCEPT 104,L,N,P1,P2
104 FORMAT(2I5,2F15.5)

IF(L,LE,0,OR,L,GT,NPR1) GO TO 68
IF(N,LE,0) GO TO 1
EN=N
NF=N+1
DX=(P2-P1)/EN
C   IF SSM E ON SUPPRESS PRINT
OCT 23410
J .65
J .70
65 WRITE(6,105)
105 FORMAT(8H PAR STEPSX3MVAL14X3HOBJ17X3HDAH)
70 TYPE 105
DO 80 I=1,NF
E11=I-1
PR(L)=P1+E11*DX
CALL HYDSM(4,1,OBJ,DAH)
C   IF SSM E ON SUPPRESS PRINT
OCT 23410
J .74
J .75
74 WRITE(6,106)L,I,PR(L),OBJ,DAH
75 TYPE 106,L,I,PR(L),OBJ,DAH
106 FORMAT(2I4,F9.3,2F20.3)
IF(N,LE,1) GO TO 1
80 CONTINUE
GO TO 1
C   IF SSM D ON CONTINUE IN INTERACTIVE MODE
90 OCT 23420
J .99
J .5
99 RETURN
END

C   INPUT OF BASIC DATA **DATPR**
SUBROUTINE DATPR(IPL,IPRT,ICOL)
COMMON/BLK1/LABL(60),AGM(24),CMS(12),PKCHI(12),IDTM(12),BASID(15),
1RES(14),PR(120),POL(12),CCK(12),CPKC(25,12),HDG(15),
2OUT(60,13)/BLK2/MANG,NCRPH,ISB,NSB,IYRB,NYR,IYRN,KR,IRES,IOGO,WH,
3KMN,KMY,JMN,JMX,SMAX,SMIN,CONV,CONPV,CONUR,CONUN,TOLF,DUM(13),ITM
5/BLK3/HD(15,12,1)
DOUBLE PRECISION LABL,DLAB
DIMENSION N(15),CAC(25),I1(25),DCA(25),PCAP(25),FMT(10),DD(12)
C ** IPL=1 BASIC DATA ONLY
C ** IPL=2 SUBBASIN DATA ONLY
C ** IPL=3 OR 4 RESERVOIR DATA ONLY
C ** IPL=5 ACREAGE AND PARAMETERS ONLY
IF(ICOL,LE,0,OR,ICOL,GT,2) ICOL=1
GOTO(10,30,200,200,30) IPL
10 READ(KR,500)NCRPH,NPHRT
500 FORMAT(25I3)
NCRPH=NCRPH+NPHRT
15 READ(KR,501)(HDG(I),I=1,15)
501 FORMAT(20A4)
IF(IPRT,EQ,0) WRITE(6,502)(HDG(I),I=1,15)
502 FORMAT(1M120A4)
READ(KR,503)(LABL(I),I=1,60)
503 FORMAT(10A8)
IF(IPRT,EQ,0)WRITE(6,504)(I,LABL(I),I=1,60)
504 FORMAT(1X6(I4,1XAB))
READ(KR,505)DLAB,(PKCHI(I),I=1,12)
505 FORMAT(2XAB,12F5.3)
IF(IPRT,EQ,0)WRITE(6,506)DLAB,(PKCHI(I),I=1,12)
506 FORMAT(1XAB,12F6.3)
C***READ IN PROPORTION DAYLIGHT HOURS AND USE COEFFICIENTS
READ(KR,505)DLAB,(POL(I),I=1,12)
IF(IPRT,EQ,0)WRITE(6,506)DLAB,(POL(I),I=1,12)
DO 20 L=1,NCRPH
READ(KR,505)DLAB,(CPKC(L,I),I=1,12)
IF(IPRT,EQ,0)WRITE(6,507)L,DLAB,(CPKC(L,I),I=1,12)
507 FORMAT(1X12,A6,12F6.3)
20 CONTINUE
C *** WRITE INITIAL DATA
GO TO 300
C *** READ SUB-BASIN DATA
30 READ(KR,508)ISB,IYRB,NYR,(BASID(I),I=1,15)
508 FORMAT(12,2I5,8X15A4)
IYRN=IYRB+NYR=1
IF(IPRT,EQ,1) GO TO 35
WRITE(6,507)
WRITE(6,509)ISB,IYRB,IYRN,(BASID(I),I=1,15)
509 FORMAT(1X3HSUBI2,15,3M TOI5,2X15A4)
C** READ CROP AND PHREATOPHYTE ACREAGES
35 DO 40 I=1,25
CAC(I)=R.#
40 CONTINUE
SPAC=0.0
SCAC=0.0
READ(KR,106)(I1(I),DCA(I),I=1,NCRPH)
106 FORMAT((10X7(I3,F7.0)))
DO 50 I=1,NCRPH
L=I1(I)
IF(L,LE,0) GO TO 50
CAC(L)=DCA(I)
IF(L,GT,NCRPH) GO TO 45
SCAC=SCAC+CAC(L)
GO TO 50
45 SPAC=SPAC+CAC(L)
50 CONTINUE

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Program Listing (Continued)

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C   REAT URBAN LAND AREA AND UNDEVELOPED LAND AREA
   READ(KR,50) URLND,UNDLND
510  FORMAT(1X2F12.0)
   TOTL=URLND+UNDLND+SCAC+SPAC
C *** COMPUTE PROPORTIONS
   DO 60 I=1,NCRPH
   IF(I.GT,NCROP) GO TO 55
   PCAP(I)=CAC(I)/SCAC
   GO TO 60
55  PCAP(I)=CAC(I)/SPAC
60  CONTINUE
C *** COMPUTE SCALE FACTORS
   CONVP=CAC/12.0
   CONPV=SPAC/12.0
   CONUR=URLND/12.0
   CONUN=UNDLND/12.0
   CONTOT=TOTL/12.0
C ** COMPUTE WEIGHTED USE COEF.
   DO 80 I=1,12
   SPKC=0.0
   SKCC=0.0
   DO 75 L=1,NCRPH
   SCP=CPKC(L,I)*PCAP(L)
   IF(L.GT,NCROP) GO TO 70
   SKCC=SKCC+SCP
   GO TO 75
70  SPKC=SPKC+SCP
75  CONTINUE
   CKC(I)=SKCC
   PKC(I)=SPKC
80  CONTINUE
C *** WRITE OUT DATA UP TO THIS POINT
   IYRN =IYRB +NYR-1
   IF(IPRT,EQ,1) GO TO 100
   WRITE(6,511)
511  FORMAT(1X10LAND AREAS)
   WRITE(6,512)
512  FORMAT(8X9HCROP LAND8XBHPRH LAND8XBHURB LAND8XBHMUND LAND
16X10TOTAL LAND)
   WRITE(6,513)SCAC,SPAC,URLND,UNDLND,TOTL
513  FORMAT(1X5F16.0)
   WRITE(6,514)
514  FORMAT(1X32INCH TO AC-FT CONVERSION FACTORS)
   WRITE(6,515)CONV,CONPV,CONUR,CONUN,CONTOT
515  FORMAT(1X5F16.0)
   WRITE(6,517)
517  FORMAT(1X27HCROP AND PHREATOPHYTE ACRES)
   WRITE(6,518) (I,CAC(I),I=1,NCRPH)
518  FORMAT((1X6(I5,F8.0)))
   WRITE(6,519)
519  FORMAT(1X17HPROP CROP AND PHR)
   WRITE(6,520) (I,PCAP(I),I=1,NCROP)
520  FORMAT((1X6(I3,F10.0)))
   K=NCROP+1
   WRITE(6,520) (I,PCAP(I),I=K,NCRPH)
   WRITE(6,521)
521  FORMAT(1X34HWEIGHTED CROP AND PHR COEFFICIENTS)
   WRITE(6,522) (CKC(I),I=1,12)
522  FORMAT(1X12F6.3)
   WRITE(6,522) (PKC(I),I=1,12)
100  IF(LEQ,5)RETURN
C *** INPUT HYDROLOGIC DATA
C   DATA IS READ IN THE FOLLOWING ORDER AND STORED IN HD(I,K,J)
C   I = 1 IS TEMP, T
C   2 IS PRECIP, PPT
C   3 IS RIVER INFLOW, RIV
C   4 IS TRIBUTARY INFLOW, TRIB
C   5 IS UNGAGED CORRELATION INFLOW, QCOR
C   6 IS CANAL DIVERSIONS, CNL
C   7 IS GROUNDWATER INFLOW, GGI
C   8 IS GAGED OUTFLOW, GAG
C   9 IS PUMPED WATER FROM THE SHALLOW AQUIFER, GPUM
C   10 IS M AND I DIVERSION, EMIDIV
C   11 IS RESERVOIR INFLOW, RIN
C   12 IS REQUIRED MINIMUM-RESERVOIR RELEASES TO STREAM, REL
C   13 IS RESERVOIR IN BASIN RELEASE FOR CANAL, ARD
C   14 IS RESERVOIR EXPORTS FROM THE BASIN, REXP
C   15 IS BASIN EXPORTS FROM THE STREAM CHANNEL, BEXP
   READ(KR,500) (N(I),I=1,15)
   DO 120 I=1,15
   DO 120 K=1,12
   DO 120 J=1,NYR
   HD(I,K,J)=0.
120  CONTINUE
   DO 170 I=1,15
   NN=N(I)
   IF(NN.LE,0)GOTO170
   IF(I.GT,2)GOTO140
C ** INPUT TEMP AND PPT
   XC=NN
   CDL=1.0/XC
   READ(KR,501) (FMT(L),L=1,10)
   DO 130 L=1,NN
   DO 125 J=1,NYR
   READ(KR,FMT) ((D(K),K=1,12)
   DO125 K=1,12
   HD(I,K,J)=HD(I,K,J)+DD(K)*CDL
125  CONTINUE
130  CONTINUE
   GO TO 170
C ** INPUT STREAMFLOW DATA
140  READ(KR,501) (FMT(L),L=1,10)
   DO 150 L=1,NN
   DO145 J=1,NYR
   READ(KR,FMT) IX, ((DD(K),K=1,12)
   DO 145 K=1,12
   HD(I,K,J)=HD(I,K,J)+DD(K)*IX
145  CONTINUE
150  CONTINUE
C *** WRITE DATA BY TYPE AND YEAR
C ** HYDROLOGIC **
   IF(IPRT,EQ,1) GO TO 300
   WRITE(6,523)
523  FORMAT(1X15HYDROLOGIC DATA)
   DO 190 I=1,15
   NN=N(I)
   IF(NN.LE,0)GO TO 190
   DO 185 J=1,NYR
   SS=0.
   DO 180 K=1,12
   SS=SS+HD(I,K,J)
180  CONTINUE
   LYRP=IYRB+J-1
   IF(ICOL,GE,2) GO TO 183
   WRITE(6,524)I,LYRP, (HD(I,K,J),K=1,12),SS
524  FORMAT(1X13,I5,6F12.2,(9X6F12.2))
   GO TO 185
183  IF(I.GT,2) GO TO 184
   WRITE(6,529)I,LYRP, (HD(I,K,J),K=1,12),SS
529  FORMAT(1X13,I5,12F9.3,F12.3)
   GO TO 185
184  WRITE(6,528)I,LYRP, (HD(I,K,J),K=1,12),SS
528  FORMAT(1X13,I5,12F9.0,F12.0)
185  CONTINUE
190  CONTINUE
   GO TO 300
C   READ RESERVOIR PARAMETERS
200  READ(KR,525) (RES(I),I=1,14)
525  FORMAT(10X,7F10.0)
   IF(IPRT,EQ,1) GO TO 300
   WRITE(6,526)
526  FORMAT(1X20HRESERVOIR PARAMETERS)
   WRITE(6,527) (I,RES(I),I=1,14)
527  FORMAT((1X5(I3,F10.2)))
300  RETURN
   END

C   PRINT SUBROUTINE
   SUBROUTINE PRNT(J,IOF,ICOL)
   COMMON/BLK1/LABL(60),AGW(24),CMS(12),PKCMI(12),IDTM(12),BASID(15),
1RES(14),PR(120),PDL(12),CKC(12),CPKC(12),CPKC(25,12),HDG(15),
2OUT(60,13)/BLK2/HANG,NCRPH,18B,NSB,IYRB,NYR,IYRN,KR,RES,IOG,WH,
3KMN,KMX,JHN,JNX,SHAX,SHIN,CONV,CONPV,CONUR,CONUN,TOLF,DUM(13),ITMX
4/BLK3/HD(15,12)
   DOUBLE PRECISION LABL
   LYRP=IYRB+J-1
C   IF IOF=0 PRINT GAGED + DIFF
C   =1 SUPPRESS PRINTING GAGED + DIFF
C   =2 OR GREATER SUPPRESS PRINTING = GO TO BAPLOT
C   IF ICOL = 1 OR COLUMN PRINTER
C   =2 132 COLUMN PRINTER
C   SET UP FOR WATER OR WATER + RESERVOIR
C   IF (IDF,GE,2) GOTO 199
C   INITIALIZE ON FIRST CALL
   IF(J.GT,1) GOTO 5
   NT=1
   NT1=8
   NT2=9
   NT3=43
   NT4=44
   NT5=45
   NT6=53
   NDF=NT3-1
   IF(ICOL.LT,1,OR,ICOL.GT,2) ICOL=1
   WRITE(6,100) (BASID(I),I=1,15),LYRP
   WRITE(6,101)
100  FORMAT(1X15A4,5X4HEARIS)
101  FORMAT(1X5HWATER/)
   GO TO (10,150), ICOL
C   80 COL PRINT LINE
C   IPG=0 FOR FIRST PAGE AND 1 FOR 2ND PAGE
10  IPG=0
   WRITE(6,115) (HDG(I),I=1,7)
115  FORMAT(6XA4,8XA4,5(7XA4)/)
   GO TO 21
20  WRITE(6,116) (HDG(I),I=8,15)
116  FORMAT(/,1XA4,4XA4,6(7XA4)/)
21  DO 25 I=NT,NT1
   IF(IPG,EQ,1) GO TO 23
   WRITE(6,118) I,LABL(I), (OUT(I,L),L=1,6)
118  FORMAT(1X13,1XA8,6F11.3)
   GO TO 25
23  WRITE(6,119) I, (OUT(I,L),L=7,13)
119  FORMAT(1X13,7F11.3)
25  CONTINUE
   DO 30 I=NT2,NT3
   IF(I,GE,NDF,AND,IOF,GT,0) GOTO 35
   IF(IPG,EQ,1) GOTO 33
   WRITE(6,120) I,LABL(I), (OUT(I,L),L=1,6)
120  FORMAT(1X13,1XA8,6F11.0)
   GOTO 30
33  WRITE(6,121) I, (OUT(I,L),L=7,13)
121  FORMAT(1X13,7F11.0)

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Program Listing (Continued)

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38 CONTINUE
35 IF (IRES,GE,1) GOTO 38
   IF (IPG,GE,1) GOTO 199
   GOTO 65
C   OUTPUT RESERVOIR DATA
38 WRITE(6,108)
108 FORMAT(/1X9HRESERVOIR)
   I=NT4
   IF (IPG,EQ,1) GOTO 48
   WRITE(6,118)I,LABL(I),(OUT(I,L),L=1,6)
   GOTO 45
48 WRITE(6,119)I,(OUT(I,L),L=7,13)
45 DO 58 I=NT5,NT6
   IF (IPG,EQ,1) GOTO 48
   WRITE(6,120)I,LABL(I),(OUT(I,L),L=1,6)
   GOTO 50
48 WRITE(6,121)I,(OUT(I,L),L=7,13)
50 CONTINUE
68 IF (IPG,GE,1) GOTO 105
65 IPG=1
   WRITE(6,122) LYRP
122 FORMAT(1M1I4)
   GOTO 20
C   132 COLUMN PRINTER
150 WRITE(6,105) (HDG(I),I=1,7), (HDG(I),I=9,15)
105 FORMAT(6XA4,6XA4,11(5XA4),7XA4/)
   DO 155 I=NT1,NT1
   WRITE(6,106)I,LABL(I),(OUT(I,L),L=1,13)
106 FORMAT(1X13,1XA8,12F9.3,F11.3)
155 CONTINUE
   DO 156 I=NT2,NT3
   IF (I,GE,NDF,AND,IPG,GT,0) GOTO 157
   WRITE(6,107)I,LABL(I),(OUT(I,L),L=1,13)
107 FORMAT(1X13,1XA8,12F9.0,F11.0)
156 CONTINUE
157 IF (IRES,EQ,0) GOTO 199
C   OUTPUT RESERVOIR DATA
   WRITE(6,108)
   I=NT4
   WRITE(6,106)I,LABL(I),(OUT(I,L),L=1,13)
   DO 168 I=NT5,NT6
   WRITE(6,107)I,LABL(I),(OUT(I,L),L=1,13)
168 CONTINUE
165 IF (J,NE,NYR) GOTO 199
C   WRITE RESERVOIR EXTREMES
   IKMN=1
   IKMX=1
   IF (KMN,GT,6) IKMN=2
   IF (KMX,GT,6) IKMX=2
   WRITE(6,109)
109 FORMAT(/1X10HRESERVOIR EXTREMES)
   KI=KMX+IKMX
   WRITE(6,110) SMAX,HDG(KI),JMX
110 FORMAT(1X15.0,7H AC=FT A4,I6)
   KI=KMN+IKMN
   WRITE(6,110)SMIN,HDG(KI),JMN
199 CALL BAPLOT(J,NYR,OUT)
   RETURN
   END
C *** SUBROUTINE DOUT ***
   SUBROUTINE DOUT(K,L,XD,KTYPE)
   COMMON/BLK1/LABL(60),AGW(24),CMS(12),PKCHI(12),IDTM(12),BASID(15),
   IRES(14),PR(120),PDL(12),CKC(12),CPKC(25,12),HDG(15),
   2OUT(60,13)/BLK2/MANG,NCRPH,158,NSB,IYRB,NYR,IYRN,KR,IRES,IOGO,WH,
   3KMN,KMX,JMN,JMX,SMAX,SMIN,CONV,CONPV,CONUR,CONUN,TOLF,DUM(13),ITMX
   5/BLK3/MD(15,12,1)
   DOUBLE PRECISION LABL
   KTYPE =0 RETURN
C   1 SUM
C   2 AVERAGE
C   3 LAST MONTH
   IF (KTYPE,LE,0,OR,KTYPE,GT,3) GOTO 99
   OUT(L,K)=XD
   GOTO (80,85,90),KTYPE
80 OUT(L,13)=OUT(L,13)+XD
   GOTO 99
85 OUT(L,13)=OUT(L,13)+XD/12.
   GOTO 99
90 OUT(L,13)=OUT(L,K)
99 RETURN
   END
C**** HYDROLOGIC SIMULATION ***HYDOSH***
   SUBROUTINE HYDOSH(IPRT,ICOL,OBJ,DAH)
   COMMON/BLK1/LABL(60),AGW(24),CMS(12),PKCHI(12),IDTM(12),BASID(15),
   IRES(14),PR(120),PDL(12),CKC(12),CPKC(25,12),HDG(15),
   2OUT(60,13)/BLK2/MANG,NCRPH,158,NSB,IYRB,NYR,IYRN,KR,IRES,IOGO,WH,
   3KMN,KMX,JMN,JMX,SMAX,SMIN,CONV,CONPV,CONUR,CONUN,TOLF,DUM(13),ITMX
   5/BLK3/MD(15,12,1)
   DOUBLE PRECISION LABL
   EQUIVALENCE (PR(1),SNI),(PR(2),SNK),(PR(3),TSM),(PR(4),TPR),
   1(PR(5),COR),(PR(6),CSN),(PR(7),CRN),(PR(8),RTH),(PR(9),CTP),
   2(PR(10),PTH),(PR(11),SMI),(PR(12),CSM),(PR(13),SNC),(PR(14),DTA),
   3(PR(15),DPI),(PR(16),DPK),(PR(17),CDP),(PR(18),PSP),(PR(19),ECV),
   4(PR(20),EAP),(PR(21),CNI),(PR(22),CNK),(PR(23),CRT),(PR(24),C1),
   5(PR(25),C2),(PR(26),C3),(PR(27),GWI),(PR(28),GWR),(PR(29),CGO),
   6(PR(30),TAJ),(PR(31),PAJ),(PR(32),CNA),(PR(33),GWA),(PR(34),DHIA),
   7(PR(35),CUMIA),(PR(36),CUA),(PR(37),CGW),(PR(38),CUPGW),(PR(39),
   8CUUR),(PR(40),CUUN),(PR(41),PURGW),(PR(43),STH),(PR(44),PBST)

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Program Listing (Continued)

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UNTP=CTP*(PPT-PTH)
IF (UNTP,LT,P) UNTP=0
QUNGCOR=QCR-UNSM+UNRN+UNTP
C*** CALCULATE INFLUENT GW
QRIV=RTV+TRB+QUNG
IF (QRIV,GT,1) GO TO 87
STGW=0
GO TO 85
R# CINFC1=C2+ALOG10(QRIV)
IF (CINFC1,GT,3) CINFC1=3
STGW=CINFC1*QRIV
C*** CALCULATE RUNOFF AND CONSUMPTIVE USE FROM LAND AREA OF URBAN AND
UNDEVELOPED LAND AND PROPORTION TO GW
85 PCPURU=RPSM*(CONUR+CONUN)
URUNCU=RPSM*(CONUR+CONUR+CONUN+CONUN)
URSF=PCPURU-URUNCU
URGW=URSF+PURGW
URSF=URSF-URGW
QIN=QRIV+QPUM+URSF+STGW
CALCULATE PHREATOPHYTE USE
PHSF=PHET*CONPV+RMPH
IF (PHSF,GT,0) GO TO 90
QIN=QIN-PHSF
PHSF=0
PHGW=0
GO TO 100
90 PHGW=PHSF+QUPGW
PHSF=PHSF-PHGW
IF (QIN,GE,PHSF) GO TO 100
PHSF=QIN
IF (PHSF,LT,0) PHSF=0
C 100 QIN=QIN-PHSF
CALCULATE CANAL DIVERSIONS FOR MANAGEMENT MODE
IF (MANG,LE,0) GO TO 110
CALL CANAL (CNL,MANG,ITM(K),ETP,EAP,ECV,PSP,SM1,CHS(K),RPSM,CONV)
110 KC2=0
SEEP1=GWCNIC*CRT
SEEP=SEEP1
CNL=CNL
C RESERVOIR OPERATION
IF (IRES,NE,1) GO TO 120
IF IRES = 0 NO RESERVOIR
C = 1 THEN TRIBUTARY OR UPSTREAM RESERVOIR
C = 2 THEN DOWNSTREAM RESERVOIR
ARD=CNL*3CNR
CNL=CNL-ARD
CALL RESRV(J,K,IRES,OUT,HD,RES,PR,EVP,PPT,RIN,ARD,REL,ST,KMN,
JMX,JMN,JMX,SMIN,SMAX,IYRB)
CNL=CNL-ARD
C ITERATE HERE TO ABOUT 130 FOR SEEPAGE IN MANAGEMENT MODE
120 WAD=QIN+SEEP+REL+ARD-RIN
IF (WAD,LT,0) WAD=0
CALL URBEMI(WAD,EMIDIV,CFEMI,EMICU,EMIRF)
C IF MANG=-1 LIMIT CNL TO WAD
C IF MANG= 0 OR 1 USE AS RECORDED
C IF MANG = 2 OR 3 LIMIT TO WAD
IF (MANG,EQ,0,OR,MANG,EQ,1) GO TO 125
IF (CNL,GT,WAD) CNL=WAD
125 GWCN=CNL*(1-ECV)
GWCNSV=GWCN
C ROUTE SEEPAGE WATER FROM CANALS
CALL GWROUT(GWCN,GWCNIC,CNK,GWCN)
SEEP=GWCN*CRT
IF (ABS(SEEP-SEEP1),LE,TOLF) GO TO 130
CNL=CNL1
KC2=KC2+1
SEEP1=SEEP
IF (KC2,LE,ITMX) GO TO 120
WRITE(0,900) J,K,SEEP1,SEEP
900 FORMAT(1X10#SEEP ITMAX FOR YEARIS,2X5#MONTHI3,2F10.1)
130 SPILL=CNL+PSP
QCV=CNL-GWCNSV-SPILL
GWCN=GWCN-SEEP
TWTR=QCV+PTH
WAGS=QCV-TWTR
QIGS=RPBM+WAGS/CONV
STR=WAD-CNL+TWTR+SPILL
WAPP=QIGS+CONV
C WAGS IS DELIVERED WATER BY IRRIGATION CANALS IN ACRE-FT
C QIGS IS TOTAL WATER APPLIED TO SOIL IN INCHES
C WAPP IS TOTAL WATER THROUGH THE SOIL IN ACRE-FT
C COMPUTE SOIL MOISTURE LEVEL
DPO=0
SM1=SM1+QIGS
IF (SM1,GT,CSM) DPO=(SM1-CSM)*CDP
SM1WDP=SM1-DPO
SM2=SM1WDP-ETP
IF (SM2,LT,0) SM2=0
DP=SM2-SMC
IF (DP,LT,0) GO TO 140
SM2=SMC
GO TO 145
140 DP=0
IF (SM2,GE,CSM) GO TO 145
ET IS LESS THAN POTENTIAL
ETT=SM1WDP-CSM
IF (ETT,LT,0) ETT=0
ETB=ETP-ETT
C ETB IS AMT OF ETP THAT MUST COME FROM BELOW CRITICAL
AET=ETT+ETB+SMSTRC
SM2=SM1WDP-AET
IF (SM2,GT,0) GO TO 150
SM2=0
AET=SM1WDP
GO TO 150
145 AET=ETP
150 DP=ADP+DPO
CROPLAND DP ROUTED TO GW
KIDTA=K-IDTA
AGW(KIDTA)=DP+PO+AGW(KIDTA)
AGW(KIDTA+1)=AGW(KIDTA+1)+DP+P1
CALL GWROUT(ARF2,ARF1,DPK,AGW(K))
ARF=ARF2+CONV
C*** CALCULATION OF SPRING FLOW AND RETURN FLOW DIVERTED TO CANALS
DUMSP=GWCN+ARF+STGW+URGW
ARF1=ARF2
SM1V=(SM1+SM2)*0.5
SM1=SM2
GWCNIC=GWCN+SEEP
C*** ROUTINE OF GW THROUGH BASIN
GGW=GWIN+DUMSP-QPUM-PHGW
CALL GWROUT(Q02,GW1,GWK,GGW)
C*** LIMIT Q02 TO BE GE 0
IF (Q02,NE,0) GO TO 160
IF (Q02,LT,0) Q02=0
160 Q02=Q02
GW1=Q02
CHGW=GGW-Q02
GWO=Q02+CGO
GEF=Q02-GWO
C PROPORTION OF GAGED SURFACE RUNOFF SRF
SRF=STR+GEF-BEXP
C IF DOWNSTREAM RESERVOIR THEN SRF BECOMES INFLOW TO RESERVOIR
C AND OBM AND OAM ARE BASED ON DS, THE CHANGE IN RESERVOIR STORAGE
CALL RESRV(J,K,IRES,OUT,HD,RES,PR,EVP,PPT,SRF,ARD,REL,ST,KMN,
JMX,JMN,JMX,SMIN,SMAX,IYRB)
QSO=ST
GO TO 170
165 QSO=SRF
170 DIFF=QSO-GAG
OBM=OBM+DIFF+DIFF
AETT=AET+CONV
YSN=SM1+CONV
DP=DP+CONV
C 21 CALL DOUT ( K, 1, T,2)
CALL DOUT ( K, 2, PPT,1)
CALL DOUT ( K, 3, SNMT,1)
CALL DOUT ( K, 4, SNM1,3)
CALL DOUT ( K, 5, PHET,1)
CALL DOUT ( K, 6, ETP,1)
536 CALL DOUT ( K, 7, AET,1)
CALL DOUT ( K, 8, SM1,2)
CALL DOUT ( K, 9,RTV,1)
CALL DOUT ( K,10,TRB,1)
CALL DOUT ( K,11,QUNG,1)
CALL DOUT ( K,12, RMPH,1)
CALL DOUT ( K,13, URSF,1)
CALL DOUT ( K,14, QPUM,1)
CALL DOUT ( K,15, STGW,1)
CALL DOUT ( K,16, PHSF,1)
CALL DOUT ( K,17, WAD,1)
CALL DOUT ( K,18,EMIDIV,1)
CALL DOUT ( K,19, EMIRF,1)
CALL DOUT ( K,20, CNL,1)
CALL DOUT ( K,21,GWCNSV,1)
CALL DOUT ( K,22, GWCN,1)
CALL DOUT ( K,23, SEEP,1)
CALL DOUT ( K,24, SPILL,1)
CALL DOUT ( K,25, QCV,1)
CALL DOUT ( K,26, TWTR,1)
CALL DOUT ( K,27, WAPP,1)
CALL DOUT ( K,28, AETT,1)
CALL DOUT ( K,29, XSM,3)
CALL DOUT ( K,30, DP,1)
CALL DOUT ( K,31, ARF,1)
CALL DOUT ( K,32, GWIN,1)
CALL DOUT ( K,33, URGW,1)
CALL DOUT ( K,34, PHGW,1)
CALL DOUT ( K,35, QGO,1)
CALL DOUT ( K,36, GEF,1)
CALL DOUT ( K,37, CHGW,1)
CALL DOUT ( K,38, GWO,1)
CALL DOUT ( K,39, BEXP,1)
CALL DOUT ( K,40, SRF,1)
IF (IRES,LT,2) GO TO 180
CALL DOUT (K,41, ST,3)
CALL DOUT (K,42, GAG,3)
GO TO 185
180 CALL DOUT ( K,41, QSO,1)
CALL DOUT ( K,42, GAG,1)
185 CALL DOUT ( K,43, DIFF,1)
190 CONTINUE
C COMPUTE OBJECTIVE FUNCTION
OAH=OAH+OUT(43,13)
CHGWS=CHGWS+OUT(37,13)
195 CALL PRNT (J,IPRT,ICDL)
200 CONTINUE
DX=CHGWS-CGW
OBI=(OBM+DX+DX)*WH
OBI=OBH
RETURN
END

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Program Listing (Continued)

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C*****SUBROUTINE HSP ****
SUBROUTINE HSP(D,J,K,I)
COMMON/BLK1/LABL(80),AGW(24),CMS(12),PKCMI(12),IDTM(12),BASID(15),
1RES(14),PR(120),PDL(12),CKC(12),CPKC(25,12),HOG(15),
2OUT(80,13)/BLK2/MANG,NCRPH,ISB,NSB,IYRB,NYR,IYRN,KR,IRES,IGGO,MM,
3KMM,KMX,JMN,JMX,SMAX,SMIN,CONV,CONPV,CONUR,CONUN,TOLP,DUM(13),ITMX
5/BLK3/HD(15,12,1)
DOUBLE PRECISION LABL
D=HD(I,K,J)
RETURN
END
C
C BLANEY CRIDDLE CU CALCULATION
SUBROUTINE BCCU(T,PDL,CKC,EKT,ETF,PET)
EKT=0.0173*T-0.314
IF(EKT.LT.0.3)EKT=0.3
ETF=EKT*T+PDL
PET=CKC*ETF
RETURN
END
C
C CANAL DIVERSIONS FOR MANAGEMENT
SUBROUTINE CANAL(CNL,MANG,IDTM,ETP,EAP,ECV,PSP,EMS,CMS,RPSM,CONV)
C*** IF MANG=-1 LIMIT QCNL TO WAD BUT USE GAGED RECORDS
C      0 USE QCNL AS RECORDED
C      1 CALC QCNL AND USE WITHOUT LIMIT TO SATISFY 7ET
C      2 CALC QCNL BUT LIMIT TO WAD
C      3 PUT LEACHING WATER TO ZERO AND LIMIT TO WAD
IF(MANG.GE.3) CNL=0.
IF(IDTM.GT.0) GO TO 10
CNL=0.
GO TO 50
10 ETN=ETP-(RPSM*(EMS-CMS))
IF(ETN.LT.0.) ETN=0.
CNL=CNL+(ETN*CONV)/(EAP*(ECV-PSP))
50 RETURN
END
C
C SUBROUTINE FOR M&I CU
SUBROUTINE URBEMI(WAD,DIV,CF,CU,RF)
IF(DIV.GT.WAD) DIV=WAD
CU=DIV*CF
RF=DIV*CU
WAD=WAD*CU
20 RETURN
END
C
C GROUND WATER ROUTING SUBROUTINE
SUBROUTINE GWROUT(Q02,Q01,XKG,QI)
IF(XKG.LE..012) GO TO 1
Q02=QI*(Q01-QI)*EXP(-1./XKG)
GO TO 2
1 Q02=QI
2 RETURN
END

C** RESERVOIR SIMULATION ** RESRV***
SUBROUTINE RESRV(J,K,IOP,OUT,HD,RES,PR,EVP,PPT,RIN,ARD,REL,ST,KMN,
1KMX,JMN,JMX,SMIN,SMAX,IYRB)
DIMENSION OUT(80,1),HD(15,12,1),RES(1),PR(1)
C
C IOP=IRES= 1 IS FOR TRIBUTARY OR UPSTREAM RESERVOIR AND
C RIN AND ARD ARE READ INTO HD BY DATPR
C
C 2 IS FOR DOWNSTREAM RESERVOIR AND RIN IS
C PASSED FROM HYDSM
C
C J=YR.,K=MNTH,ETP=MOD BC TEMP FACTOR
C ARD=ACTUAL CANAL RELEASE (WATER OR SALT)
C RSR=TOTAL RELEASES (WATER OR SALT)
C
C REXP=RES EXPORT
C RESERVOIR PARAMETERS, RES, ARE AS FOLLOWS
C RES(1) = INITIAL STORAGE (AC-FT)
C 2 = NOT USED (WAS INITIAL SALT CONCENTRATION (PPM))
C 3 = MINIMUM USEABLE STORAGE (AC-FT)
C 4 = MAXIMUM ALLOWABLE STORAGE (AC-FT)
C 5 = RESERVOIR AREA AT ZERO STORAGE INTERCEPT (ACRES)
C 6 = C1 COEFFICIENT IN AREA VS STORAGE EQUATION
C 7 = C2 EXPONENT IN AREA VS STORAGE EQUATION
C 8 = STORAGE LEVEL SEPARATING TWO AREA VS STORAGE EQUATIONS
C 9 = RESERVOIR AREA AT ZERO STORAGE - INTERCEPT FOR UPPER EQ
C 10 = C3 COEFFICIENT FOR UPPER EQUATION
C 11 = C4 EXPONENT FOR UPPER EQUATION
C 12-14 = NOT USED
10 JK=J*K
IF(JK.NE.0)GOTO 13
INITIALIZE FIRST MNTH, FIRST YEAR
CBST=1.-PR(44)
STI=RES(1)
CSTI=RES(2)
SMAX=STI
SMIN=STI
JMN=IYRB-1
JMX=JMN
KMN=12
C
C SET UP MONTHLY DATA
13 CALL HSP(REL,J,K,12)
CALL HSP(REXP,J,K,14)
IF(IOP.GE.2) GOTO 14
CALL HSP(RIN,J,K,11)
CALL HSP(ARD,J,K,13)
C
C ** BEFORE OPERATIONS
C RIN IS RESERVOIR INFLOW
C REL IS MINIMUM REQUIRED RESERVOIR RELEASE TO THE STREAM SHANNEL
C ARD IS IN BASIN CANAL RELEASE
C REXP IS RESERVOIR EXPORT DESIRED
C ** AFTER OPERATIONS
C RQRL IS REQUIRED MINIMUM RELEASE
C REL IS ACTUAL RELEASE
C REXP IS ACTUAL EXPORT
C AND IS ACTUAL CANAL RELEASE
C DS IS CHANGE IN STORAGE
C ST IS EOM STORAGE
14 RQRL=REL
DSEP=PPT-EVP
CALL AREA(STI,AT,RES)
OPERATE RESERVOIR
DSPOS=(RES(4)-STI)/CBST
DSNEG=(RES(3)-STI)/CBST
EXRS=0.
RSR=REL+ARD+REXP
DST=RIN-RSR+DSEP+AI/12.
16 IF(DST.GT.0.) GOTO 20
C CASE WHEN DST IS NEGATIVE
C IF(DST.GT.DSNEG) GOTO 17
C IF(DSNEG.GT.0.) GOTO 30
DST=DSNEG*CBST
GOTO 30
17 DST=DST+CBST
GOTO 30
C CASE WHEN DST IS POSITIVE
20 IF(DST.LT.DSPOS) GOTO 17
DST=DSPOS*CBST
30 ST=STI+DST
IF(ST.LT.1.)ST=1.
CALL AREA(ST,AE,RES)
AV=(AI+AE)*.5
DST=RIN-RSR+DSEP+AV/12.
C CHECK EOM STORAGE AGAINST MAX AND MIN
32 IF(DST.GT.DSPOS) GOTO 33
ST=STI+DST*CBST
IF(ST.LT.RES(3)) GOTO 34
GOTO 40
C CASE OF EXCESS WATER - SPILLS FROM RES
33 EXRS=DST-DSPOS
RSR=RSR+EXRS
REL=REL+EXRS
ST=RES(4)
GOTO 40
34 STCK=ST+RSR-RES(3)
IF(STCK)35,35,37
35 ST=ST+RSR
RSR=0.0
REL=0.0
ARD=0.0
REXP=0.0
IF(ST.LT.0.0)ST=0.0
GOTO 40
37 RSR=STCK
ST=RES(3)
IF(STCK.GT.REL)GO TO 38
REL=STCK
ARD=0.0
REXP=0.0
GO TO 40
38 RSRX=STCK-REL
IF(RSRX.GT.ARD)GO TO 39
ARD=RSRX
REXP=0.0
GO TO 40
39 REXP=RSRX-ARD
C INITIALIZE NEXT MONTHS STORAGE
40 DS=ST-STI
STI=ST
C
41 CALL DOUT(K,44, EVP,1)
CALL DOUT(K,45, RIN,1)
CALL DOUT(K,46, AV,2)
CALL DOUT(K,47, RQRL,1)
CALL DOUT(K,48, REL,1)
CALL DOUT(K,49, ARD,1)
CALL DOUT(K,50, REXP,1)
CALL DOUT(K,51, RSR,1)
CALL DOUT(K,52, ST,3)
CALL DOUT(K,53, DS,1)
C CHECK EXTREMES
50 EMX=ST-SMAX
IF(EMX.LE.0.0)GOTO 51
SMAX=ST
JMX=IYRB+J-1
KMX=K
GOTO 52
51 EMN=SMIN-ST
IF(EMN.LE.0.0)GO TO 52
SMIN=ST
JMN=IYRB+J-1
KMN=K
52 RETURN
END
C RESERVOIR AREA SUB **AREA**
SUBROUTINE AREA(S,A,RES)
DIMENSION RES(14)
IF(S.LT.0.0) GO TO 10
IF(S.LT.RES(8)) GO TO 1
C4=RES(11)
A=RES(9)+RES(10)+S*.C4
GOTO 12
1 C2=RES(7)
A=RES(5)+RES(6)+S*.C2
GOTO 12
10 A=RES(5)
12 RETURN
END

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Program Listing (Continued)

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C ** BASIN PLOT SUBROUTINE ** BAPLOT
SUBROUTINE BAPLOT(J,NYR,OUT)
DIMENSION OUT(60,1)
DATA IPX/0/
C*** PLOT INSERT
C****SSW A ON PLOT DATA *****
999 OCT 023500
J .81
IF(IPX.EQ.1) GO TO 21
998 CALL QSHYIN(IERR,500)
CALL QSC(1,IERR)
C** SET ALL DAC'S TO ZERO AND RAISE PEN
CALL QNLBB('30000,IE)
DO 20 I=1,8
KK=I-1
CALL QWJDAR(0.,KK,IE)
20 CONTINUE
IPX=1
21 TYPE 201
201 FORMAT(32HSET B ON TO READ PLOT PARAMETERS/20HTURN B OFF TO GO TO
NEXT YEAR/)
C ** IF SSW B IS ON READ PLOT PARMETERS **
C ** IF SSW B IS OFF GO TO NEXT YEAR **
OCT 025000
OCT 023500
J .80
C RAISE PEN
CALL QNLBB('030000,IE)
READ(6,203) IPLT, IDASH, SC, NXPYR, ID, LD, ILF, ILW, KX, KY
C** KX AND KY ARE THE DAC CHANNELS FOR PLOTTING
203 FORMAT(I2,I3,F10.0,I0I5)
C** IPLT IS VAR NO TO BE PLOTTED, IDASH IS NO OF DASHES TO BE DRAWN
C** PER TIME PERIOD, SC IS SCALE FACTOR IN UNITS/VOLT, NXPYR IS NO.
C** OF TIME PERIODS PER YEAR, ID AND LD ARE DELAY TIMES TO ALLOW FOR
C** USING DIFFERENT ANALOG PLOTTERS, ILF IS DELAY TIME FOR RAISING
C** THE PEN AND ILW IS DELAY TIME FOR LOWERING THE PEN.
TYPE 202, IPLT
202 FORMAT(12HPLT VAR NO,I3/)
SC=SC*10.
XXC=NYR
XDX=J-1
XDI=- (XDX/XXC)
XNVR=NXPYR
XD=1./ (XXC*XNVR)
XD=XD
PDX1=0.0
XP=XDI
CALL QWJDAR(XP,KX,IE)
CALL QWJDAR(PDX1,KY,IE)
LAD=ID*LD
CALL QSDLY(LAD)
C LOWER PEN
CALL QNLBB('102000,IE)
CALL QSDLY(ILW)
IF(IDASH.LE.1) GOTO23
ISPAC=IDASH-1
IDS=ISPAC*IDASH
LDD=LD/IDS
XDASH=3*IDASH-1
XDASH=XDASH/2.0
XL=XD/XDASH
XS=XL/2.0
23 DO 30 I=1,NXPYR
PDX=OUT(IPLT,I)/SC
PDX=-PDX
CALL QWJDAR(PDX,KY,IE)
CALL QSDLY(ID)
IF(IDASH.LE.1) GO TO 26
24 XP=XP+XL
CALL QWJDAR(XP,KX,IE)
CALL QSDLY(LDD)
DASH LINE LOOP
DO 25 L=1,ISPAC
CALL QNLBB('103000,IE)
CALL QSDLY(ILF)
XP=XP+XS
CALL QWJDAR(XP,KX,IE)
CALL QSDLY(LDD)
CALL QNLBB('102000,IE)
CALL QSDLY(ILW)
XP=XP+XL
IF(XP.LE.-1.) XP=-.999999
CALL QWJDAR(XP,KX,IE)
CALL QSDLY(LDD)
25 CONTINUE
J .28
26 XP=XP+XD
IF(XP.LE.-1.) XP=-.999999
CALL QWJDAR(XP,KX,IE)
CALL QSDLY(LD)
28 CONTINUE
30 CONTINUE
C RETURN PEN TO INITIAL LOCATION
CALL QWJDAR(PDX1,KY,IE)
CALL QSDLY(ID)
CALL QNLBB('103000,IE)
CALL QSDLY(ILF)
CALL QWJDAR(XDI,KX,IE)
CALL QSDLY(ID)
J .21
80 IF(J.NE.NYR) GOTO01
CALL QWJDAR(PDX1,KX,IE)
C** END OF PLOT INSERT
81 RETURN
END

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