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#### ESTIMATING THE RAINFALL-RUNOFF CHARACTERISTICS

OF SELECTED SMALL UTAH WATERSHEDS

by

Clive H. Walker

A thesis submitted in **partial** fulfillment of the requirements for the degree

of

#### MASTER OF SCIENCE

in

Water Resources Engineering

UTAH STATE UNIVERSITY Logan, Utah

#### ACKNOWLEDGMENTS

Nearly all the data used in this thesis has been collected by someone else. This data was made freely available to me by the employees of the Forest Service and the Soil Conservation Service, sometimes at their own inconvenience. This help is deeply appreciated.

376.7ie

The Intermountain Forest and Range Experiment Station made precipitation, runoff, soils and vegetation data available for the watersheds examined in this study. The Soil Conservation Service has provided personal encouragement in this endeavor. Some data was also received from the Watershed Planning Staff of the Utah State Office of the Soil Conservation Service and the Sevier River Basin Planning Party at Richfield, Utah.

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Chine N. Walker

Clive H. Walker

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#### ABSTRACT

Estimating the Rainfall-Runoff Characteristics

of Selected Small Utah Watersheds

by

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Utah State University, 1970

Major Professor: Joel E. Fletcher Department: Water Resources Engineering

Runoff and rainfall data have been taken from three high mountain Utah watersheds and subjected to runoff to rainfall comparisons. The resulting Q/P ratios have been compared to the average volumes of runoff curve numbers (CN) computed from this data for each watershed. Runoff curve numbers were also estimated on the basis of the soils and vegetation data available for the watersheds.

An attempt has been made to estimate the watershed lag characteristics by computing synthetic hydrographs for successively larger values of time to peak estimates until the best fit comparison was achieved between the snythetic and the actual hydrographs. Time lag estimates were also made from the Kirpich method and the Mockus method.

(128 pages)

#### INTRODUCTION

Much of the water supply for the intermountain west comes from small, high mountain watersheds. Most of the runoff volume comes from melting a winter snowpack and the nature of water use below these watersheds depends greatly on the reliability of this supply. However, diversion structures, storage reservoirs, and channels must allow for safe passage of high rates of runoff from high intensity summer rainfall events. The costs of failure of such structures from floods are more than the costs of the structures alone, but include the costs associated with lost water supplies until the structures are rebuilt, as well as damages to properties, crops, transportation, communication, and commerce. Some of these floods cause loss of life and livestock.

Disastrous floods occur from summer rainfall events on these watersheds. In Utah, some of the more notable of these floods have occurred on the Farmington Creek Watershed which included the Halfway Creek and Morris Creek watersheds described in this thesis, on the Pleasant Creek Watershed which is just north of Ephraim Creek, and on the Manti Creek Watershed just south of Ephraim Creek. The Alpine Meadows watershed described herein is part of the Ephraim Watershed. For descriptions of these floods see Woolley (1946), Peterson (1954), and Croft and Bailey (1964).

As potential flood plains become more intensively used, potential damages will greatly increase. New storm drainage systems and flood control structures will be constructed to reduce these damages. The U. S. Forest Service has developed watershed management and land treatment measures that have significantly reduced flood damage potential on several watersheds in the state. On several watersheds this kind of program has been combined with federally supported locally sponsored Small Watershed Protection Projects through the work of the Soil Conservation Service that include flood and debris control structures near the mouth of the mountain canyons. Local governments often find themselves in the flood control business and have constructed flood channels and control structures. Design of downstream flood control structures requires information about peak rates of flow and volumes of runoff from rainstorms from watersheds with and without upstream watershed land treatment and structural flood control measures.

Statistical and regional analysis procedures for evaluating streamflow records have been developed and are used for predicting the frequency of occurrence of specified levels of flood flow for gaged and ungaged watersheds. When this kind of analysis can be made with a reasonable degree of confidence in its accuracy it should provide the basis for hydrologic design of water control structures. This kind of procedure has serious limitations when one must: (a) predict the effect of a watershed change due to land treatment or management measures, (b) predict runoff volumes and peak rates of runoff from an ungaged watershed in an area where there are very few runoff gages, or (c) where the existing runoff gages are affected by diversions and storage to the extent they are not usable for estimating natural runoff conditions. When any of these problems must be solved, some kind of synthetic hydrograph approach must be used to develop design information.

Hydrologists in the Soil Conservation Service (SCS) have developed a synthetic unit-graph procedure to provide for safe design of water supply and flood control structures for high rates of flow. This method requires estimates of the rainfall-runoff and runoff time lag characteristics of the watersheds to which it is applied. The purpose of this study is to develop information about these characteristics for small high mountain watersheds in Utah. The watersheds have been subjected to various watershed management practices.

Morgan and Johnson (1962) tested four synthetic unit-graph methods against actual hydrographs from watersheds in Illinois. Serious deficiencies were found in each method, but the method proposed by the Soil Conservation Service was found to be one of the two better methods. Hydrologists in the Soil Conservation Service use this method throughout the United States. Hydrologists in the Forest Service use this method as part of their standard procedures. Other federal agencies sometimes use this method to obtain information about small watersheds. The Soil Conservation Service method is most recently described in the SCS National Engineering Handbook, Section 4 (Soil Conservation Service, 1964) in an SCS technical paper, SCS-TP-149 (Kent, 1968), and in a handbook edited by Chow (1964). Forest Service procedures for this method are found in a Forest Service handbook (Forest Service, 1959), and in some unpublished information obtained from the regional hydrologist in Ogden, Utah, reproduced in this paper in Table 5. The parts of this method adopted by the Bureau of Reclamation are published in another handbook (Bureau of Reclamation, 1960).

Some of the data originally used to develop the Soil Conservation

Service method are included in publications of the Agricultural Research Service (1959, 1960). Other data that was used remains largely in unpublished form except as included in reports prepared for other purposes by various persons and agencies, as in Croft and Bailey (1964). Research is continuing in projects administered by both the Agricultural Research Service and the Forest Service on several small watersheds throughout the nation. Reports of these studies may be easily found by reference to one of many recently published bibliographies of hydrology and water resources research.

The primary data obtained for this study are contained in unpublished material located in offices of the U. S. Forest Service Intermountain Forest and Range Experiment Station in Logan and Ephraim, Utah. Some soils and vegetation information was obtained from the office of the Sevier River Investigation Staff of the Soil Conservation Service in Richfield, Utah.

#### OBJECTIVES AND SCOPE OF THIS STUDY

#### Objectives

There is some rainfall and runoff data available for summer rainstorm events from a few instrumented small high mountain Utah watersheds. This data will be used to estimate summertime rainfallrunoff characteristics of some of these watersheds. To do this requires reducing the primary stage hydrograph traces into direct runoff volume estimates. Records from recording rain gages must be adjusted to estimates of average rainfall over the watersheds for the same events. Soils and vegetation information must also be obtained to estimate the soil and land use effects of the rainfall-runoff characteristics of the watersheds. The soil and land use effects will be represented by rainfall-runoff curve numbers (CN) to be developed with procedures used by the Soil Conservation Service (1964) as described in Appendix A. If rainfall and runoff events are sufficiently large, CN may also be computed directly from rainfall and runoff data. If the CN obtained by these two methods do not agree or do not describe the rainfall-runoff volume characteristics of the watersheds with accuracy, an alternate method should be presented to represent these characteristics. An attempt will be made to approximate actual runoff hydrographs from rainfall data using the Soil Conservation Service snythetic unit-graph method. This will provide information about the time lag characteristics of the watershed.

#### Scope

This study is limited to a few selected high mountain watersheds in Utah and to measured runoff events that have occurred during the summer rainfall season. The data for this study were obtained in 1966 and 1967 and do not include events more recent than 1965. The watersheds used for this study have been visited and the gaging sites examined. The only field work done was to obtain an estimate of general watershed soils classification where a soil survey was not available and a general familiarization with vegetation types and watershed topography. All quantitative information about precipitation, runoff, and watershed properties has been taken from records developed by the Forest Service, the U. S. Geological Survey, the U. S. Weather Bureau, and the Soil Conservation Service.

There are almost 100 small watersheds in Utah with drainage area smaller than 400 square miles that have some gaged runoff records. Approximately one-fourth of these have drainage areas under 12 square miles, are located in mountains and might have been suitable for this study. Most of these have not been included in this study because there is not a rain gage on or very close to the watershed. Another reason is that runoff records are very short and do not include significant summer runoff events. Many of these watersheds have very limited information about soils and vegetation.

The selected watersheds are experimental watersheds administered by the Forest Service. Each has at least one recording rain gage on or very near the watershed. Some soils and vegetation information has been developed by qualified surveyors for each of them. However, because there is not sufficiently detailed or accurate data, no attempt has been made to probe deeply into the mechanics of the hydrologic cycle on these watersheds.

#### METHODS OF PROCEDURE

#### Obtaining basic data

The physical data of shape, size and slope of the three watersheds have been taken from topographic maps published by the U. S. Geological Survey. The watershed boundaries have been defined as the topographic divide above the runoff gaging sites. Watershed areas have been measured by planimetry of these boundaries. The average watershed slope has been computed by the equation

$$S = 100MN / A$$
 (1)

where S is the slope in per cent, M is the total length of the contour lines in feet, N is the chosen contour interval in feet and A is the total area of the watershed in square feet. Stream lengths and profiles have also been obtained from the topographic maps. This information is given in Appendix B.

Soils information is very sketchy for these watersheds. There is a generalized soils map for the Alpine Meadows Watershed that may be obtained from the Soil Conservation Service. The only soils information found for the Farmington Canyon watersheds was contained in vegetation surveys. These surveys include qualitative estimates of runoff and erosion potential of the soil based primarily on the surface evidence of runoff or erosion occurring before the time of the survey. A field reconnaissance was performed to obtain general estimates of soil textures. These are explained in the descriptions of the individual watersheds in Appendix B. Aerial photographs were used as an aid in defining hydrologic soil group boundaries on the Halfway Creek Watershed. Morris Creek and Alpine Meadows watersheds were assigned one hydrologic soil group for each watershed except for rocky areas. The watershed soils were classified into hydrologic soil groups according to the criteria explained in Appendix A.

Vegetative cover data for the Farmington Canyon watersheds was obtained from surveys made for the Forest Service. The basic data is in unpublished field survey sheets in the possession of the Intermountain Forest and Range Experiment Station. This data has been interpreted and summarized in Appendix B. There is a general range survey available in Forest Service offices that includes the Alpine Meadows Watershed area. This survey identifies important grazing and browse plants, but provides little information about plant cover density. Density has been estimated from a few unrecorded visual observations of different parts of the watershed. Measurements of watershed areas by major vegetative types on the Alpine Meadows Watershed has been taken from aerial photograph interpretation. These data are also summarized in Appendix B.

An estimate of the ability of the soils and vegetal cover of the watersheds to intercept and abstract moisture from rainfall has been obtained using the procedure developed by the Soil Conservation Service and briefly outlined in Appendix A. The estimate is presented in the form of runoff curve number (CN) for each watershed. The CN developed for Halfway Creek Watershed is 50, for Morris Creek is 34, and for Alpine Meadows Watershed is 60. The tabular computations of these CN are in Appendix B.

The runoff data for the summertime rainfall produced runoff events was obtained through a search of existing stage trace hydrographs from

water stage recorders. The more significant runoff events were chosen from the trace. All three watersheds have perennial streams. Personal judgment was used to estimate the beginning of the rise of the hydrograph and the effective end of direct runoff for the event. Recorded stage information was converted to flow estimates using the weir or flume rating tables provided by the agency responsible for the operation of the runoff gaging sites. A computer program was written and used to compute the volume of direct runoff for each hydrograph. Base flow was estimated as a straight line from the beginning of rise to the end of direct runoff. This estimate was subtracted from the recorded total flow at each time-flow coordinate chosen to describe the hydrograph. The differences are adjusted direct flows. These flows were used to compute the mass curve of the volume of runoff in inches for each event.

The rainfall data was obtained from recording rain gage records provided by the Forest Service. These records were searched to find the rainstorm events which most likely caused the runoff events studied. The rainfall and runoff data have been punched on computer cards and are available as computer printout. The rain gage and watershed locations were plotted on map overlays. Storm depths were plotted on these overlays and isohyetal lines were estimated. The average watershed storm rainfall was estimated from these charts. The base maps of watershed and rain gage locations are in Appendix B.

#### Comparison of rain and runoff--Linear base

The first set of rainfall to runoff comparisons is on the basis of volume with straight line plots. These comparisons are given in

Table 1 and in Figures 1, 2 and 3. The computed CN shown in the table and figures were obtained from the equation

$$Q = (P + 2 - 200 / CN)^{2} / (P - 8 + 800 / CN)$$
(2)

with repetitive 'cut and try' estimates of CN. The minimum possible value of CN for each rainfall-runoff event was found from the equation

$$CN = 200 / (P + 2)$$
 (3)

This CN was then repeatedly incremented by a value of 2 until the computed Q was equal to or greater than the actual runoff volume for the event. These CN estimates were summed and averaged for the events studied for each watershed. The computed average CN has been reduced one value to improve the estimate.

Simple linear regression analysis has been performed for the rainfall-runoff data. The resulting equations of best fit straight lines and the computed correlation coefficients for these lines are presented in Table 1. The Q intercepts for Halfway Creek and Morris Creek watersheds are so near zero that the lines obtained from plotting the average Q/P ratios as shown on Figures 1 and 2 are felt to be representative of the rainfall-runoff relationships for these watersheds for the low values of Q encountered. The best fit straight line is retained for the Alpine Meadows Watershed. The line obtained from plotting the average Q/P ratio is also shown in Figure 3 for this watershed. The CN estimated from soils and vegetation data are also shown in the figures combined with runoff values computed for each value of P from the average Q/P ratios. Refer to Appendix B for the development of runoff curve numbers (CN) from soils and vegetation information.

Event Number	Date	Rain Inches	Runoff Inches	Compute CN	d
V 1	7/1/1940	1.50	.0192 -	65	63.25
2	7/12/1942	1.00	.0105 🧹	73	11.66
3	8/19/1945	1.33	.0216	68	6685
4	7/12/1951	. 50	.0036	84	82.92
5	8/1/1952	. 80	.0105	77	1648
6	7/26/1953	.30	.0073 🗸	91	90.54
1	8/4/1954	1.31	.0088 -	66	GANG
1-8	8/19/1959	. 50	.0036	84	82.92
V 19	7/13/1962	1.09	.0147 🖉	71	70,55
/10	9/13/1963	1.00	.0035	71	695
11	9/13/1963	.44	.0030	86	* 84.58
/12	7/18/1965	1.50	.0186 🗸	65	63.25
13	8/12/1965	. 44	.0042 /	86	85-64
14	8/21/1965	.58	.0067 ~	84	81.52

	Table 1.	Rainfall-runoff	comparisons
--	----------	-----------------	-------------

The average runoff to rain Q/P ratio is .011. The equation for the best fit straight line is Q = -.0012 + .0124 P. The correlation coefficient for this line is .82. The average computed CN for this watershed is 75. The CN estimated from soil and vegetation data was 49.

#### Table 1. Continued

Event Number	Date	Rain Inches	Runoff Inches	Computed CN
15	7/1/1940	1.50	.0050	61
16	9/12/1942	1.00	.0022	71
17	8/3/1945	.77	.0032	76
18	8/19/1945	1.36	.0029	64
19	8/13/1946	.30	.0011	89
20	8/10/1947	.76	.0041	76
21	7/27/1951	. 50	.0007	82
22	8/4/1951	.76	.0020	76
23	8/19/1951	.95	.0028	72
24	8/1/1952	.70	.0027	78
25	7/26/1953	.40	.0010	85
26	8/4/1954	1.35	.0036	64
27	7/11/1956	.35	.0005	87
28	7/28/1956	.35	.0008	87
29	8/21/1957	.90	.0072	75
30	8/19/1959	.80	.0011	73
31	7/18/1965	1.20	.0017	64

Morris Creek Watershed

The average runoff to rain Q/P ratio is .0031.

The equation for the best fit straight line is Q = .0003 + .0027 P. The correlation coefficient for this line is .58. The average computed CN for this watershed is 74. The CN estimated from soil and vegetation data was 34.

#### Table 1. Continued

Event Number	Date	Rain Inches	Runoff Inches	Computed CN
32	<ul><li>✓ 7/16/1951</li></ul>	.40	.0027	87
<sup>©</sup> 33	8/3/1951	.30	.0044	91
34	7/28/1952	1.10	.0969	81
\$ 35	7/30/1952	.20	.0119	95
36	7/10/1953	.60	.0345	87
37	7/28/1953	. 47	.0406	91
38	8/31/1953	.79	.0761	86
39	8/16/1955	.61	.0279	85
40	8/3/1961	.77	.0486	84
41	8/7/1961	.38	.0246	92

Alpine Meadows Watershed

The average runoff to rain Q/P ratio is .066.

The equation for the best fit straight line is Q = -.0216 + .1039 P. The correlation coefficient for this line is .92. The average computed CN for this watershed is 87. The CN estimated from soil and vegetation data was 60.



Figure 1. Estimated rainfall-runoff curve for Halfway Creek Watershed --Linear analysis.



Figure 2. Estimated rainfall-runoff curve for Morris Creek Watershed--Linear analysis.



Figure 3. Estimated rainfall-runoff curve for Alpine Meadows Watershed -- Linear analysis.

#### Comparison of rain and runoff--Logarithmic base

The rainfall and runoff volume data may also be analyzed through logarithmic transformation. The resulting rainfall-runoff relationship is of the form

$$Q = a P^{b}$$
(4)

If three events on Halfway Creek Watershed, July 26, 1953, August 4, 1954, and September 13, 1963, are designated mavericks and neglected, the best fit curve has the formula

$$Q = 0.012 P^{1.48}$$
 (5)

with a correlation coefficient of the logarithms of the volumes of 0.98. The best fit curve for the Morris Creek Watershed with no mavericks excluded is

$$Q = 0.0028 P^{1.12}$$
(6)

and for Alpine Meadows Watershed with the event for July 16, 1961, excluded is

$$Q = 0.081 \text{ p}^{1.52} \tag{7}$$

If the exponent of P is held constant at 1.48 and the coefficients (a) chosen so that the curves pass through a mean logarithm of the runoff volumes, the resulting formulas for the best fit curves are

 $Q = 0.012 P^{1.48}$  (8)

$$0 = 0.003 \, \mathrm{P}^{1.48} \tag{9}$$

$$Q = 0.079 P^{1.48}$$
(10)

for the three watersheds in the same order. These curves are plotted in Figures 4, 5 and 6. The CN obtained from soils and vegetation data are plotted on these figures as combined with equations (8), (9) or (10) as appropriate. Figure 7 is a logarithmic plot of the coefficients of P in equations (8), (9) and (10) against the CN obtained from soils and vegetation information.



Figure 4. Estimated rainfall-runoff curve for Halfway Creek Satershed--Logarithmic analysis.



Figure 5. Estimated rainfall-runoff curve for Morris Creek Watershed--Logarithmic analysis.



Figure 6. Estimated rainfall-runoff curve for Alpine Meadows Watershed--Logarithmic analysis.





## Estimating watershed lag factors with synthetic hydrographs

If rainfall and runoff records were always accurately timed and if rainfall came in sudden bursts, but in uniform rates, watershed lag time could be measured directly from plots of hyetographs and hydrographs. Neither condition is met in the basic data used for this study. The construction of synthetic hydrographs with the incremental triangular hydrograph procedure as proposed by the Soil Conservation Service (1964) allows cut and try estimation of watershed lag effects.

When the rain distribution is determined or fixed, the shape of a synthetic hydrograph produced depends on the estimate of watershed lag time as seen in the equation

$$q_{pi} = 484 \text{ A } Q_i / \triangle T_p \tag{11}$$

where  $q_{pi}$  is the peak rate of flow of an incremental triangular hydrograph, A is the watershed area in square miles,  $Q_i$  is the generated incremental estimate of mass runoff and  $\triangle T_p$  is the estimated time of rise of the incremental hydrograph. The relationship of  $T_p$  to watershed lag is defined by the Soil Conservation Service (1964) to be

$$\triangle T_{p} = \triangle D / 2 + L$$
 (12)

where  $\triangle D$  is a chosen increment of time used to divide the mass rainfall curve into short, fairly uniform segments of developing a series of incremental triangular hydrographs, and L is the watershed lag time in hours. The incremental hydrographs may then be combined by addition into a composite synthetic hydrograph. If a natural rainstorm is used for the rainfall mass distribution curve, if the total volume of runoff is predicted correctly, and if the proper  $\bigtriangleup$  D is chosen, the resulting synthetic hydrograph should closely fit the actual runoff hydrograph.

Kent (1968) proposed the use of the approximation

$$\triangle T_{p} = 3 \triangle D \tag{13}$$

to facilitate computer solution of peak flow estimates from incremental triangular hydrographs. This approximation may also be used to develop complete hydrographs. A program written in FORTRAN that will do this is included in Appendix C. This program is written to perform the following operations;

- Compute a synthetic hydrograph from a natural rainstorm adjusted in depth to the estimated average watershed depth.
- Compare this hydrograph to the actual hydrograph and compute the summed squares of the deviations between them.
- 3. If there is only one minimum point in curve of summed squared deviations, the program will search for it by making two adjustments:
  - First, move the actual hydrograph in time to the position of best fit.
  - b. Then, increase  $\triangle D$  and repeat the operation in a.
  - c. Repeat steps a and b until the minimum summed squared deviation is found.

Table 2 is a list of the  $T_p$  estimates computed with this procedure. The minimum squared deviations found and the time that the actual hydrograph was moved from the beginning of rain to the position of best fit are also listed in Table 2. These estimates could be refined somewhat by reducing the size of the  $\triangle D$  increment if desired. Appendix D contains the plotted actual and synthetic hydrographs for

Event Number	Date	Minimum sum of Deviations Squared	T <sub>p</sub> Estimates	Time Move of Actual Hydrograph
1	7/1/1940	8.1723	1.8	1.2
2	9/12/1942	2.45	4.2	7.0
3	8/19/1945	19.7	0.6	0.8
4	7/27/1951	1.9	1.2	1.2
5	8/1/1952	5.5	0.6	0.2
6	7/26/1953	5.1	0.6	0.2
7	8/4/1954	2.1	0.6	0.4
8	8/19/1959	0.3	1.2	0.8
9	7/13/1962	2.8	1.8	1.2
10	9/13/1963	0.1	1.8	0.6
11	9/13/1963	0.1	1.8	1.2
12	7/18/1965	3.6	1.8	1.8
13	8/21/1965	0.1	1.2	0.8
14	8/21/1965	0.3	1.8	1.2
	Total		21.0	18.6
	Average		1.5	1.33
Withou	t Number 2			
	Total		16.8	11.6
	Average		1.3	0.9

Table 2. Summary of goodness of fit data for hydrograph synthesis

Estimate of T from Kirpich procedure = 0.30 hours.

Estimate of  $T_p$  from Mockus (Kent, 1968) procedure = 1.26 hours.

## Table 2. Continued

Event Number	Date	Minimum Sum of Deviations Squared	Tp Estimates	Time Move of Actual Hydrograph
15	7/1/1940	0.012	1.2	2.0
16	9/12/1942	0.036	2.4	0.8
17	8/3/1945	0.008	0.6	0.6
18	8/19/1945	0.021	3.6	4.8
19	8/13/1946	0.006	0.6	0.4
20	8/10/1947	0.125	0.6	0.8
21	7/27/1951	0.001	1.2	0.4
22	8/4/1951	0.006	1.2	1.6
23	8/19/1951	0.050	0.6	0.2
24	8/1/1952	0.106	1.2	1.2
25	7/26/1953	0.004	0.6	0.2
26	8/4/1954	0.019	2.4	2.4
27	7/11/1956	0.002	1.8	1.8
28	7/28/1956	0.002	0.6	0.6
29	8/21/1957	0.437	1.2	0.8
30	8/19/1959	0.002	1.2	2.4
31	7/18/1965	0.014	1.8	2.4
	Total		22.80	23.40
	Average		1.34	1.38

Morris Creek Watershed

Estimate of  ${\rm T}_p$  from Kirpich procedure = 0.19 hours.

Estimate of  ${\rm T}_p$  from Mockus (Kent, 1968) procedure =-3.0 hours.

## Table 2. Continued

Event Number	Date	Minimum Sum of Deviations Squared	T <sub>p</sub> Estimates	Time Move of Actual Hydrograph
32	7/16/1951	0,185	1.2	0.8
33	8/3/1951	0.815	2.4	1.6
34	7/28/1952	334.1	0.6	0.4
35	7/30/1952	13.6	0.6	0.4
36	7/10/1953	30.6	1.2	1.2
37	7/28/1953	447.2	0.6	0.6
38	8/31/1953	120.0	1.2	1.2
39	8/16/1955	140.0	0.6	0.4
40	8/3/1961	143.0	0.6	0.2
41	8/7/1961	11.8	1.8	1.2
	Total		10.80	8.0
	Average		1.08	0.8

Alpine Meadows Watershed

Estimate of  $T_p$  from Kirpich procedure = 0.34 hours.

Estimate of  $T_p$  from Mockus (Kent, 1968) procedure = 0.96 hours.
the events studied. Table 2 also lists  $T_{\mathbf{p}}$  estimates obtained by

$$T_{p} = 0.72 T_{c}$$
 (14)

where  ${\rm T}_{\rm C}$  is estimated from methods presented by Kirpich (1940) and Mockus (Kent, 1968).

## Estimating sources of runoff

The volumes of runoff measured from these watersheds are important in that they may contribute to serious floods when combined with runoff from other watersheds. They also provide a supplemental water supply to depleted summertime streamflows. When converted to inches of runoff from the watershed, however, they are very small. Only two of the events studied approached one tenth of one inch of runoff. These were both on the Alpine Meadows Watershed. There was one event during the study period on Halfway Creek that may have approached this amount of runoff, but it caused a mud flow which submerged the runoff gage. This occurred on August 10, 1947. Other mud flows have occurred from this watershed.

The low volumes of runoff suggest that no significant amount of direct surface overland flow may have occurred from any rain event studied. Suppose all of the runoff occurs from effectively impervious areas. Now estimate the impervious surface area required to produce the amount of runoff to result in the Q/P ratios found in the rainfall-runoff comparisons. This is simply the product of the Q/P ratio and the total area of the watershed. The results are 5.1 acres for Halfway Creek, 0.47 acres for Morris Creek, and 24.8 acres for Alpine Meadows. The Alpine Meadows Watershed has an area of 20 acres or more of snow drift fields and barren areas adjacent to streams. There is also a small wet meadow area just above the runoff gage that would first take rain and runoff to replenish its depletions and then provide almost 100 per cent runoff. This may be the reason for the P axis intercept of best fit regression line in Figure 3. If this area is subtracted from 24.8 acres we still have 4.8 impervious acres.

All three watersheds have perennial streams. Nearly all of the rain that falls on these open streams must leave the watershed as direct flow. Open streams may be detected on aerial photographs. These stream lines may be transferred to the topographic maps and their lengths estimated. The estimates are 16,000 feet for Halfway Creek, 6,000 feet for Morris Creek, and 24,000 feet for Alpine Meadows. If we convert the areas listed above to square feet and divide by the estimated stream lengths, we obtain the estimated average stream widths required to produce the runoff volumes experienced during the period of study. These stream widths are 13.7 feet for Halfway Creek, 3.4 feet for Morris Creek, and 8.7 feet for Alpine Meadows. These figures are too high as estimates of average stream widths, but they are in the correct order of magnitude. The flowing streams tend to keep an area of soil and rocks saturated on their banks. There are rocky and barren areas very close to the streams that will contribute high percentages of rainfall to direct surface runoff or to 'quick return' subsurface flow. This kind of flow is often called interflow and may produce runoff very quickly through an interplay of static and kinetic flows. If these areas are included in the width estimates above they seem to be very realistic.

However, the fairly long lag times experienced and the apparent

relationship shown in Figure 7 between soils and vegetation and the rainfall-runoff relationships may provide additional clues as to the sources of runoff. These factors indicate that the interflow from small source areas may supply more volume than runoff from stream surfaces. The slow lag times estimated for these steep watersheds indicate considerable resistance to flow. This is probably the result of flow through cracks in rocks, over boulders, through porous soils and perhaps, in some areas through humus and litter. The relationship shown in Figure 7 suggests that the CN developed from soils and cover information may indicate the size of small source areas on these high mountain watersheds that have high percentages of runoff. The combination of this relationship with the relationships shown in Figures 4, 5 and 6 might provide a basis for changing the rainfall-runoff curves (CN) shown in Figure 9 in Appendix A to improve estimates of potential runoff from similar watersheds. Further research is needed to confirm these curves or to improve and extend them.

## OBSERVATIONS AND RECOMMENDATIONS

### Observations

The very low volumes of runoff listed in Table 1 strongly indicate that only small parts of the watersheds have contributed runoff for the events studied. The standard runoff curve number (CN) developed as part of the Soil Conservation Service methods for estimating rainfallrunoff relationships appears to be a poor estimator of these low volumes of runoff when it is developed as a weighted average CN for the entire watershed either from soils and vegetation information or from rainfall-runoff comparisons.

The high estimates of lag time obtained from the procedures used in this thesis suggest quite high resistance to water flow. This suggests that more water comes from overland and subsurface flow than from rain falling directly on stream surfaces. This water may be flowing through cracks in rocks, through very porous soils very near the flowing streams, through wet humus or living vegetal matter such as moss and lichens along stream banks, or through litter on the soil surface. Rock outcrops and talus slopes lying close to the stream tributaries are suspected of providing the major source of runoff from the summer rainstorms on these well vegetated small high mountain watersheds.

Figure 7 suggests that the CN obtained from soils and vegetation information for the whole watershed may be an indicator of the size of the small portion of the watershed that actually produces runoff from summer rainstorms. Further research is needed to determine whether this evidence may become the basis for improving the runoff curve number as an estimator of runoff volumes from summer rains on small high mountain watersheds. One possible form these CN might take is shown for each watershed on Figures 4, 5 and 6 as the combined curve.

The consistent underestimates of peak rates of flow by the synthetic hydrograph procedure used in this study suggests a need to change or improve estimates of watershed lag time from the formula given as equation (13) to some shorter time relationship. Holtan and Overton (1963) suggest an "m" value derived from recession curve analysis as an estimator of watershed lag effects. Somewhat more sophisticated programming would be required for computer generation of the synthetic hydrographs to test this value, but its use should be further investigated. Field measurements would also provide a check on lag time estimates.

The excellent agreement of  $T_p$  estimates between those obtained by this study and those obtained from the Mockus procedure for Halfway Creek and Alpine Meadows is most interesting. The fairly poor comparison for the Morris Creek watershed is just as significant. Had the estimate of watershed curve number for Morris Creek also been about 50, the  $T_p$  estimate from the Mockus procedure would have been much closer to the 1.34 hours obtained from the synthetic hydrograph estimate. This suggests that the variations of CN may not be responsible for changing watershed lag characteristics when the volumes of runoff are very low. This finding is consistent with the earlier suggestions made about runoff volumes. The very poor comparison of Kirpich method estimates of  $T_p$  leads to serious doubts as to its usefulness for watersheds and conditions such as those

studied in this thesis.

Though not presented in this thesis, the isohyetal plotting of summer rainfall events led the writer to the following tentative observations. Summer rainstorms in the high mountains are highly variable. Recorded amounts of rain for the same storm could vary from zero at one rain gage to almost an inch at another rain gage less than 3000 feet away. The storm patterns seem to be more significant than orographic effects for these storms. Rain gage catches differ between nearby gages at similar elevations as much as between nearby gages at different elevations. The lower elevation gages catch more rain than the higher gages almost as often as the inverse is true. The data obtained for this study are not sufficient for a true statistical test of these observations.

#### Recommendations

Serious deficiencies have been encountered in records of rainfall and runoff. These are partly a result of human errors, but are more often the result of instrument malfunction. Clocks stop or gain or lose time. Ink traces are lost while the pen is apparently still inked. Ink traces sometimes trace over and over at diminishing levels until some records are obliterated. Clocks between different gages are almost never properly synchronized. Improved instruments and procedures are seriously needed.

The present methods of measurement or sampling of hydrologic data need to be improved. Areal distribution of rain is almost an unknown factor, even when several rain gages exist in a fairly small area. Weirs and flume and sediment traps may significantly disturb time lag factors in runoff measurement. Recent developments in such remote sensing devices as radar offer great promise in reducing areal measurement errors. This kind of measurement should be developed on the experimental watersheds of the Agricultural Research Service and the Forest Service. This would facilitate further study of summer rainstorm patterns in the high mountains.

Range and forest surveys often provide little information about the actual total plant density or the horizontal per cent of area covered with plants. This cover estimate should be developed and related to site condition or other commonly surveyed factors. Soils information obtained with these surveys needs to be improved to allow better estimates of the hydrologic effects of the soil-plant complex.

Field studies of the hydraulic properties of the watersheds included in this study might provide additional valuable information. This study can be expanded with other estimates of rainfall-runoff characteristics and lag times to improve these comparisons.

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APPENDIXES

# Appendix A The SCS synthetic unit-graph method

Theory of direct surface runoff volume. As seen in Figure 8, the general water balance equation for any watershed when ice and snow are not present can be expressed

$$P = (Q_2 - Q_1) + (Z_2 - Z_1) + (W_2 - W_1) + (U_2 - U_1) + (I_2 - I_1)$$
(15)  
where P is precipitation.  $Q_1$  is channel inflow or surface inflow into  
the watershed and  $Q_2$  is surface outflow or channel outflow from the  
watershed.  $Z_1$  is condensation and  $Z_2$  is evapotranspiration.  $W_1$  and  
 $W_2$  are horizontal underground inflow and outflow. The quantity  
 $(U_2 - U_1)$  is the change in both underground and surface storage during  
the specified time interval (Chebotarev, 1966).

For the case of direct surface runoff from a headwater watershed during a short duration storm, some of the variables may be combined or neglected in equation (15).  $Q_1$  is eliminated when the boundary is defined unless artificial importation occurs.  $Z_1$  may be considered a part of P as it occurs during a storm.  $Z_2$  will include evaporation and transpiration during and shortly after the storm and will account for part of the initial abstraction ( $I_a$ ) from P that must occur before any overland runoff takes place.  $W_1$  and  $W_2$  will usually be neglected during a short duration direct runoff event. The storage factor ( $U_2$ - $U_1$ ) as it pertains to underground water can change only as supplied by infiltration or depleted by base flow. Temporary storage on the surface of the watershed during a storm contributes to the effects of lag in time of outflow and supplies water for evaporation and infiltration.  $I_2$  is the source of base flow at the watershed outlet and is estimated separately from the direct runoff





estimates.  $I_1$  is infiltration through the watershed surface.  $I_1$ ,  $W_1$ ,  $W_2$ ,  $U_2$  and  $Z_2$  when combined comprise the total abstractions (S) from P. Part of S is considered to be initial abstraction ( $I_a$ ) and a remainder (F). Equation (15) may now be expressed

$$P = Q_2 + (F + I_a)$$
 (16)

Rearranging and dropping the subscript

$$Q = (P - I_a) - F$$
 (17)

 $(P - I_a)$  is defined as effective rainfall, therefore

$$Q = P_{\rho} - F \tag{18}$$

The SCS method assumes that for any time during the storm after direct runoff begins

$$F / S = Q / P_{\rho}$$
(19)

Equations (18) and (19) combined give

$$Q = P_e^2 / (P_e + S)$$
 (20)

Studies of many small watersheds have led to an average estimate that  $I_a$  = 0.2 S. Resubstituting  $P_e$  = (P - 0.2 S)

$$Q = (P - 0.2 S)^{2} / (P + 0.8 S)$$
(21)

Now define a rainfall-runoff curve number (CN) in terms of S as

$$CN = 1000 / (S + 10)$$
 (22)

Equations (18) and (19) combined produce a second degree equation for the solution of runoff in terms of precipitation and a rainfall-runoff curve number.

$$Q = [P - (200 / CN - 2)]^2 / (P + 800 / CN - 8)$$
(23)

This equation has a distinct minimum point when Q equals zero for each value of CN where

$$P = (200 / CN) - 2$$
 (24)

Equation (23) is meaningless for all values of P less than that

computed in equation (24). When this is true there is assumed to be no overland runoff. Equation (23) can be solved for several values of CN and a chart may be made such as in Figure 9 for ready graphical solution of the equation.

Obtaining factors for the runoff equation. The use of equation (23) or Figure 9 presupposes some reasonable estimate or measurement of CN and P. P may be obtained from actual storm data or from charts of expected rainfall values for various durations and likelihood of occurrence. One source for this information is a set of precipitationfrequency maps for Utah (U. S. Weather Bureau, 1968). Precipitation data used in this study is found in Appendix C.

If precipitation and runoff are both measured quantities, for several events of sufficient magnitude for a small watershed, a value of CN may be estimated for that watershed using equation (23) or Figure 9. CN estimated by this method may be correlated to soil and vegetation information about the watershed. When this is done CN may be estimated for ungaged watersheds on the basis of soil and vegetation information.

The Forest Service has cooperated with the Soil Conservation Service in developing CN values for western forest and range complexes for small watersheds on the basis of information about soils, vegetation, and intensity of land use. CN may be estimated from Figures 10, 11, 12 and 13 which have been adapted from information presented by the Soil Conservation Service (1964). A value for CN for a small watershed is usually estimated as an areally weighted composite of CN values for subdivisions of the watershed. The CN is a practical estimate of the combined interception and infiltration effects in the rainfall-runoff characteristics of a watershed.



Figure 9. Estimation of runoff from rainfall using runoff curve number (CN)



Figure 10. Curve numbers (CN) for forest-range soil cover complexes in the Western United States (SCS, 1964)



Figure 11. Curve numbers (CN) for forest-range soil cover complexes in the Western United States (SCS. 1964)



Figure 12. Curve numbers (CN) for forest-range soil cover complexes in the Western United States (SCS, 1964)



Figure 13. Curve numbers (CN) for forest-range soil cover complexes in the Western United States (SCS, 1964)

<u>Soils in the watershed</u>. From the point of view of the problem discussed in this paper, the most important property of soil is its infiltration or water intake rate. Logic tells us that a soil with a high intake rate will produce less direct runoff than a soil with a low intake rate. The intake rate may be limited by the transmission rate of the saturated soil. The property is called soil permeability.

An exact quantitative analysis of the intake rates and transmission rates of watershed soils would be very difficult, if not impossible. There are many factors that can influence or radically change the intake rate of a soil. A saturated or frozen soil will generally have a lower intake rate than a moist, warm soil. Chemical changes in a soil may affect this property. Biological agents change soil composition, structure, texture and position. However, some approximate analysis of soil properties in the watershed must be made as the basis for runoff prediction from precipitation information.

The SCS method proposes the use of a limit condition of the minimum rate of infiltration obtained for a bare soil when completely wetted for a long time. The determination of this index is empirical and approximate. On this basis, four rather broad hydrologic groupings are given to classify soils. These are simply A, B, C and D. The Forest Service in Utah has also defined groups AB, BC, and CD as hydrologic soil groupings whose characteristics fall between the primary groupings. A description of these groups is given in Table 3, as taken from the SCS handbook, and in Table 4, as taken from an unpublished information sheet obtained from the regional hydrologist's office of the Forest Service in Ogden, Utah.

Hydrologic Soil Group	Description				
А	(Low runoff potential) Soils having high infiltration				
	rates even when thoroughly wetted and consisting chiefly				
	of deep, well to excessively drained sands or gravels.				
	These soils have a high rate of water transmission.				
В	Soils having moderate infiltration rates when thoroughly				
	wetted and consisting chiefly of moderately deep to deep,				
	moderately well to well drained with moderately				
	fine to moderately coarse textures. These soils have a				
	moderate rate of water transmission.				
С	Soils having slow infiltration rates when thoroughly				
	wetted and consisting chiefly of soils with a layer				
	that impedes downward movement of water, or soils with				
	moderately fine to fine texture. These soils have a				
	slow rate of water transmission.				
D	(High runoff potential) Soils have very slow infil-				
	tration rates when thoroughly wetted and consisting				
	chiefly of clay soils with a high swelling potential,				
	soils with a permanent high water table, soils with a				

claypan or clay layer at or near the surface, and shallow soils over nearly impervious material. These soils have a very slow rate of water transmission.

Table 3. Hydrologic soils groups as defined in the SCS handbook

Hydrologic Soil Group	Description				
A	Soils having high infiltration rates even when thoroughly wetted, consisting chiefly of deep ( $\leq 3$ feet) well to excessively drained sands, loamy sands, sandy loams, or gravels. These soils have a high rate of water transmission and a low runoff potential (3.00 inches of infiltration per hour or more).				
	Paralithic soils - granitic soils more than 20 inches deep with a deep decomposed contact zone.				
В	Soils having moderate infiltration rates consisting chiefly of moderately deep (20 inches) moderately well to well drained soils with moderately fine to moderately coarse textures. These soils have a moderate rate of water transmission (1.25 to 3.00 inches per hour).				
	Paralithic soils - granitic soils less than 20 inches deep with a decomposed contact zone.				
С	Soils having slow infiltration rates consisting chiefly of soils with a layer that impedes the downward move- ment of water and soils moderately fine to fine texture and a slow infiltration rate. These soils have a slow rate of water transmission (0.5 to 1.25 inches per hour).				
	<ol> <li>A Change of one or two percolation classes within 10 inches of the surface, depending on roots and structures.</li> </ol>				
	<ol> <li>Moderately fractured limestone at less than 20 inches.</li> </ol>				
	3. The change in permeability is at less than 20 inches.				
	<ol> <li>Soils having moderate compaction in the upper 8 inches of profile.</li> </ol>				
D	Soils having very slow infiltration rates consisting cheifly of clay soils with a high swelling potential, soils with a high permanent water table, soils with a claypan or clay layer near the surface, and shallow soils over nearly impervious materials. These soils have a slow rate of water transmission (less than 0.50 inches per hour).				

Table 4. Hydrologic soils groups as defined by the Forest Service

#### Table 4. Continued

Hydr S G	ologic oil roup		Description
1	D Continu	ed 1.	Less than 12 inches of soil over flat lying sandstones, limestones, etc.
		2.	Lithosols - generally averaging less than l inch per hour infiltration.
		3.	Clay layer or shallow over shale.
		4.	A change of two or more percolation classes in the profile.
		5.	Non-wetting soils, dry silt, sand, etc.
ł	AB H b	ydrolog etween	cic soil group whose characteristics fall primary soil groups.
I	BC S	ame as	above.
(	CD S	ame as	above.

The Soil Conservation Service and other agencies of the U. S. Department of Agriculture have been conducting soil surveys for many years. Much of the agricultural land and some of the public land in this country has been soil surveyed. The minimum infiltration rate of a soil type can be inferred approximately from the soil description written from the soil survey data. A list of over 4000 soil type names with their hydrologic group classification is given in the SCS handbook. This is not a list of all soils, but the local soil scientist can compare the properties of unlisted soils with listed soils with which he is familiar. The soils of the local watershed may be classified from this information. When the watershed soils have not been surveyed, a field reconnaisance may be performed by experienced personnel. Local Forest Service and SCS personnel may be sufficiently familiar with the watershed to make a tentative classification of the soils. The hydrologic soil groups are then outlined on a base map of the watershed and their areas determined by a dot or grid method or they may be planimetered. Impervious areas are kept as separate units. Small areas of differing soils may be included in larger areas if they cover less than three per cent of the combined area.

<u>Watershed vegetation</u>. The following statement can be found in the context of the "Hydrology Handbook".

Of the natural agencies affecting soil structure and its permeability to water none is more important than the impact of rain ... The conditions at the soil surface are often the critical ones that govern infiltration. The arrangement of soil particles in the surface layer may either permit a high rate of intake or cause a very low rate, and is sensitive to great readjustment by either man or nature. An aggregated, porous structure usually is found where a considerable quantity of organic matter is in the soil, where tillage has not been excessive, and where sod forming crops have been included. However, such favorable surface structure may be largely destroyed if the soil, when bare, receives high intensity rains ... The presence of vegetative cover, either as a canopy or as a mulch of dead vegetation, has ... pronounced effects upon the rate of infiltration. The character of surface cover, therefore, is a matter of prime importance. The degree of protection provided is about proportional to its density. (American Society of Civil Engineers, 1949, p 34-37)

Figures 10, 11, 12 and 13 provide estimates of CN according to hydrologic soil group and vegetative cover type and density. Cover density includes litter if present. Adjustments have been made in the basic data to conform to the assumption that  $I_a = 0.2$  S and that average moisture conditons prevail.

Vegetation cover types and density may be estimated from field reconnaissance or from vegetation survey maps and data. On public forest and range land where private use for grazing is permitted, the

Forest Service or the Bureau of Land Management may have range vegetation surveys on file. These are usually maps with the appropriate symbols and keys at a scale of 2 inches per mile. Where commercial timber is available on public lands there may be a timber survey. Much of the state of Utah has been mapped under one of these surveys. Sample data for these surveys is taken from field transect lines or area plots. This sample data is extended with the interpretation of aerial photographs and the judgment of the technician in the field. Some specific watersheds involved in projects authorized under the Small Watersheds Act (Public Law 566) may have guite detailed hydrologic analysis surveys in the files of local, regional, or state offices of the Forest Service or the Soil Conservation Service. There are a number of such watersheds in Utah, but generally these watersheds have very limited precipitation or runoff data. Some general vegetation maps are also on file at the office of the Sevier Basin Study of the Soil Conservation Service, now located in Salt Lake City, Utah.

Where surveys have been made they give information concerning vegetation species, density of cover, canopy and litter, and other information taken from visual measurement or observation, such as slope, aspect, evidence of erosion, surface soil texture, and evidence of grazing and wildlife. Other information may include range site, name and condition or forest site class and index.

From the information gathered in these surveys and the requisite soils information, the watershed can be subdivided into hydrologic complexes so that the soils-cover relationship within each complex is relatively homogeneous. A CN can then be estimated for each complex and these can be combined by weighting by areas into a CN estimate for the watershed.

Antecedent soil moisture. The CN determined from Figures 10, 11, 12 and 13 are given for the average moisture existing for the soils and vegetation or land use conditions described. If soil moisture conditions on the watershed are normally wetter or drier than the average of similar lands in the United States, an adjustment in CN should be made. Tables 5 and 6 (Soil Conservation Service, 1964) give suggested adjustments for three levels of antecedent moisture conditions. AMC II is the average level, AMC III is for the wetter level and AMC I is for the drier level. These adjustments are usually applied after the watershed CN is estimated, but may be applied to watershed soil-vegetation complexes if desired.

matal 5 day antesedant prinfall	
Dormant season	Growing season
Less than 0.5	Less than 1.4
0.5 to 1.1	1.4 to 2.1
Over 1.1	Over 2.1
	Total 5-day antecedent rainfall Dormant season Less than 0.5 0.5 to 1.1 Over 1.1

Table 5. Seasonal rainfall limits for AMC

The triangular hydrograph. The method for using watershed soils, soil moisture and vegetation information to estimate a rainfall-runoff volume relationship for that watershed has been discussed. The runoff volume, however, may tell very little about its rate of flow. The design of spillways or flood control storage requires an estimate of probable peak flows. This may sometimes be obtained from a frequency analysis as previously mentioned, or from some kind of routing

	CN for conditions	
II	I	III
100	100	100
95	87	98
90	78	96
85	70	94
80	63	91
75	57	88
70	51	85
65	46	82
60	40	78
55	35	74
50	31	70
45	26	65
40	22	60
35	18	55
30	15	50

Table 6. Curve numbers (CN) adjustments for wet or dry antecedent moisture conditions

2.17

procedure on larger watersheds or river basins. On smaller watersheds the estimate may be made through some kind of hydrograph analysis.

A basic tool for estimating runoff hydrographs for storm of different amounts and intensities on a watershed is the unit hydrograph. Since many watershed characteristics, such as shape, size and slope are relatively constant from year to year, if two storms are similar in duration and amount, the two runoff hydrographs should be quite similar as well.

The theory of the unit hydrograph as now conceived, is based on the observation that, at a given point on a given stream, not only is the base of the hydrograph of direct runoff resulting from a storm of unit duration constant, regardless of the volumes of rainfall or runoff, but also the ordinates of the hydrograph vary directly as the volume of runoff. ... The unit hydrograph ... may be defined as the discharge hydrograph resulting from one inch of direct runoff generated uniformly over the tributary area at a uniform rate during a specified period of time. (American Society of Civil Engineers, 1949, p. 105)

The unit hydrograph is curvilinear. The SCS method approximates a short duration single peaked runoff hydrograph with a triangular hydrograph according to the principle of the unit hydrograph. Complex runoff hydrographs can be constructed from combining several short duration triangular hydrographs computed for corresponding portions of a rainstorm. This procedure has proven to be reasonably accurate when used on small agricultural watersheds. Useful relationships have been developed for estimating the dimensions of a triangular hydrograph. Figure 14 shows these dimensions.

The volume of runoff is the area under the hydrograph in acre inches per acre or simple inches. Time is in hours, and the rate of flow is in inches per hour. From Figure 14 we obtain

$$q_i = 2Q / T_p + T_r$$
 (25)



Figure 14. The triangular hydrograph.



Figure 15. Development of a composite hydrograph by summing incremental triangular hydrographs,

From evaluating many actual hydrographs the average relationship is devised (Kent, 1968)

$$T_r = 1.67 T_p$$
 (26)

Combining, we have

$$q_i = Q / 1.335 T_p$$
 (27)

Converting  $\textbf{q}_{i}$  in inches per hour to  $\textbf{q}_{p}$  to cubic feet per second (c.f.s.) we have

$$q_{\rm p} = 484 \, {\rm A} \, {\rm Q} \, / \, {\rm T}_{\rm p}$$
 (28)

when A is drainage area in square miles. When equation (26) is not correct, 484 will not be the correct coefficient for equation (28). This coefficient depends on the shape of the hydrograph. From Figure 9 we find that

$$T = D / 2 + L$$
 (29)

where D is the duration of excess rainfall that produces the runoff hydrograph and L is defined as lag time. The relationship

$$L = 0.6 T_{c}$$
 (30)

is an average taken from the study of a large number of watersheds.  $T_c$  is the time of concentration or the average travel time of a particle of direct runoff from the hydraulically most distant point in a watershed to the watershed outlet. If a complex hydrograph is to be constructed from combining incremental hydrographs the portions of an event, the  $\Delta$  D chosen should be less than  $T_c$  and preferably about 0.2 of  $T_p$ . Figure 15 presents a visual concept of the procedure to be followed to obtain a composite synthetic hydrograph from incremental hydrographs.

There are several methods for estimating  $T_c$ . Perhaps the most accurate is to estimate channel hydraulics, for various levels of flow, from channel reaches surveyed in the field. Two approximate methods described in the Soil Conservation Service literature have been developed by Kirpich (1940) and Mockus (Kent, 1968). Figures 16 and 17 are self explanatory nomographs developed for these methods. In Figure 17 watershed slope (S $_w$ ) may be estimated from direct measurement in the field, or the equation

$$S_{w} = 100 \text{ MN} / \text{A}$$
 (31)

may be used where N is the total length of the contours in the watershed in feet, N is the contour interval in feet and A is the area of the watershed in square feet. If the watershed is very small,  $S_w$  may be computed as the ratio of the difference in elevation between the watershed outlet and the most distant ridge to the approximate average length of the watershed.







Figure 17. The Mockus nomograph for estimating time of concentration  $(T_c)$ 

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## Appendix B General description of watersheds

General description. Three small watersheds in two areas have been selected for this study. Table 7 lists their names, planimetered drainage areas, and the stream to which they are a tributary watershed. Each of these larger streams has the same name as the town near its mouth. These are all small high mountain watersheds. Figures 18 and 19 are small maps of these watersheds. All develop an appreciable snowpack with drifts that usually last through May and early June. In some years remnants of larger drifts last through the summer season. Rainfall on these drifts probably contributes nearly 100 per cent of its volume to runoff, but evaluation of the effects of these drifts is very difficult unless information is available about their areal extent, water holding capacity, melting rates and frequency of occurrence. Part of the base flow from a watershed will be melt from these drifts when they exist. All of these streams normally have perennial flow. All of these watersheds are located in the north to south lying chain of the Wasatch Mountains that form the eastern rim of the Great Basin. These mountains supply most of the water for Utah's "fertile cresent."

These watersheds were selected for this study because they are small, located in Utah, high in elevation, have been fairly well instrumented for several years, and are not seriously affected by artificial importation of water. If differences from other studies are noticed in watershed areas, lengths, or other properties, they are the result of independent measurement with tools of limited accuracy.

Halfway Creek Watershed. This watershed is on the lower north slope of Farmington Canyon. It has an average watershed slope of

Name of Watershed	Tributary to	Drainage Area Miles <sup>2</sup>	Drainage Area Acres	Watershed Length Feet	Upper Elevation Feet	Lower Elevation Feet	Average Channel Slope Ft/Ft
Alpine Meadows	Ephraim Creek	0.57	366	5,320	10,444	9,850	0.112
Halfway Creek	Farmington Creek	0.72	464	7,040	9,160	6,200	0.420
Morris Creek	Farmington Creek	0.24	156	4,600	8,307	6,080	0.484

Table 7. List of selected watersheds


Figure 18. Farmington Canyon Watersheds. Estimate of average elevation: Halfway Creek - 7600 ft. Morris Creek - 7200 ft.



Figure 19. Alpine Meadows Watershed, east of Ephraim Utah. Average elevation - 10,000 ft.

about 58 per cent and an average channel slope of 0.42 feet drop per foot of horizontal length. The main channel essentially bisects the watershed from top to south and forms the intersection of two steeply angled planes. The stream flows in a south-southwesterly direction. The extreme upper end of the watershed is more densely vegetated than the west and has less of its area in rocky outcrops. The west side has more area in rock outcrops and talus slopes. This is apparently the result of erosion of the upper ends of granite formations that dip westward.

The hydrologic soil groups on this watershed range from B to C. Most of the watershed is classified in the BC soil group. Much of the lower west side of the watershed has a C soil group and the lower east end of the watershed has a B hydrologic soils group classification. Figure 20 is a map showing estimated hydrologic soils groups and cover types.

The vegetative cover information is from a survey obtained from the Intermountain Forest and Range Experiment Station in Logan, Utah. As far as the writer is aware, this survey is unpublished and untitled. The vegetation varies from a snowbush, sagebrush, grass and weed complex at the upper end to a beautiful tall dense grove of oak and maple at the lower east end of the watershed. Most of the west side is gambel scrub oak with various kinds of brush such as scrub maple, sagebrush, chokecherry and snowbush. The east side is more densely covered with similar vegetation with the tall groves at the lower end and some mapped areas of aspen groves. Table 8 is the computation of the runoff curve number for the watershed from soils and vegetation information for the areas shown in Figure 20.



Figure 20. Estimated hydrologic soils groups and cover types, Halfway Creek Watershed.

Area Number	Vegetation	Cover Density %	Hydrologic Soil Group	CN	Drainage Area Acres	Sum of Products
1	Terraced sage - Grass weed	25	BC	70	28.7	
2	Terraced low brush - Sagebrush	55	BC	56	2.9	
3	Sage - Grass - Weed	25	BC	70	12.9	
4	Low brush - Sagebrush	55	BC	56	84.7	
5	Brush	68	BC	44	28.8	
6	Brush	68	С	48	4.0	
7	Aspen - Maple	47	BC	54	13.8	
8	Aspen - Maple	47	С	57	7.6	
9	Brush - Maple	88	BC	34	19.7	
10	Oak - Maple	76	BC	40	42.3	
11	Oak - Maple	76	C	44	12.5	
12	Oak	76	BC	40	2.9	
13	Aspen	50	BC	52	2.6	
14	Aspen - Brush	70	С	47	0.8	

Table 8. Halfway Creek Watershed CN estimation

Area Number	Vegetation	Cover Density %	Hydrologic Soil Group	CN	Drainage Area Acres	Sum of Products
15	Oak - Sage	51	С	55	56.0	
16	Oak - Sage	51	BC	52	25.4	
17	Oakbrush	70	BC	43	13.1	
18	Brush with Oak and Maple	91	BC	33	31.4	
19	Oakbrush - Scrub Maple - Sage	67	С	48	24.6	
20	Oakbrush - Scrub Maple	67	В	38	4.2	
21	Tall Oak and Maple	66	В	54	45.9	
	Total				464.8	23,475
	Estimated CN					50

Table 8. Continued

Note: Without the terraced area CN would be 49.

The earliest runoff event used for this study occurred July 1, 1940. The latest used occurred August 21, 1965. There are 12 other events used in the study as shown in Appendix C. The highest summer rainfall produced peak flow for these events was 7.9 cubic feet per second above base flow on August 19, 1965. However, there was a mud flow flood on August 10, 1947, that is not included because the runoff gage was covered with debris and sediment.

The location of nearby rain gages are shown in Figure 18, except for the Farmington Warehouse gage located downstream in the town of Farmington. The rain events and data used for this study are listed in Appendix C. The normal precipitation extends from an estimated 25 inches at the mouth to as high as 40 inches at the upper end of the watershed (U. S. Weather Bureau, 1962).

Morris Creek Watershed. This watershed is located on the lower south slope of Farmington Canyon. It is the steepest of the three watersheds with an average watershed slope of about 60 per cent and an average channel slope of 0.48 foot drop per foot of horizontal length. It is narrow at both ends and wide in the middle. The stream has cut a V shaped canyon. The main stream flows in a northnorthwesterly direction. Two tributaries meet the main stream low in the canyon and extend upward into the wide center part of the watershed. The soils on the watershed are fairly uniform and are given a hydrologic soil group classification of BC based on field observation that bare spots will produce a moderately high degree of runoff.

At the extreme upper end of the watershed where drifting occurs the vegetation is sagebrush, snowberry and snowbush, with weeds and wild flowers. This kind of low vegetation extends down the left and

right ridges for some distance. Below this and between the upper ridges scrub oak and maple surround groves of fir and aspen. In the upper center section of the watershed there are extensive fir covered areas mixed with aspen groves. Below this and on the steeper slopes above the fir is more sparse with tall growths of oak, aspen, and maple filling between. A vegetative study in 1953 (Peterson, 1954) found the average plant cover over the watershed to be about 55 per cent. Litter was found to cover about 84 per cent of the watershed. The watershed has shown an increase in vegetative and litter cover since 1939.

Figure 21 is adapted from a vegetative survey map provided by the Intermountain Forest and Range Experiment Station at Logan, Utah. Table 9 is the computation of the runoff curve number (CN) estimate based on the soils and vegetation information.

The earliest runoff event used in this study occurred July 10, 1936. The latest happened on July 18, 1965. The highest summer rainfall produced peak flow during the period was 0.93 cubic feet per second above base flow on August 10, 1947. This was the same event that caused the mud flow which covered the Halfway Creek gaging site. Probable reasons for this low flood record include low grazing use and the north aspect of the watershed which permits a more evenly distributed water supply to the watershed vegetation.

The Halfway Creek and Morris Creek watersheds have been treated as one in plotting isohyetal maps for estimating average watershed rainfall amounts. Figure 18 shows nearby rain gage locations. The rain gage records used for this study are given in Appendix C. Average annual rainfall amounts are similar to those for Halfway Creek, except



Figure 21. Vegetative survey map, Morris Creek Watershed .

Area Number	Vegetation	Cover Density %	Hydrologic Soil Group	CN	Drainage Area Acres	Sum of Products
1	Sage - Snowberry	58	BC	54	2.5	
2	Fir - Aspen - Mixed Browse	90	BC	33	20.2	
3	Large Aspen - Mixed Browse	93	BC	32	19.4	
4	Sage - Weed	69	BC	49	5.0	
5	Brush - Weed	30	BC	67	2.1	
6	Small Aspen	95	BC	31	3.5	
7	Fir - Mixed Browse	90	BC	33	114.0	
	Total				166.7	5,678
	Estimated CN					34

Table 9. Morris Creek Watershed CN estimation

that this is a lower watershed and will not extend to 40 inches of rain at the upper end of the watershed.

Alpine Meadows Watershed. This watershed is located at the upper southeast end of the Ephraim Creek drainage area. It is highest of the three watersheds with an average elevation of about 10,100 feet above sea level. Snow drifts persist later into the season than on the other two watersheds. It is the flattest of these watersheds with an average watershed slope of about 17 per cent and an average channel slope of about 0.11 foot drop per horizontal foot of length. The main stream has many tributaries extending in a radial pattern to the upper ridges. The watershed is shaped like a cut of pie. Figure 19 is a small map of the watershed showing its location by township is section numbers and showing the locations of nearby rain gages. The estimated mean annual precipitation is 35 inches.

The soils on this watershed are fairly uniform and are classified as in hydrologic soil group CD based on field observations that bare spots will produce a high degree of runoff. The predominant soil is tentatively classified by texture as a silty clay loam.

Figure 22 is a vegetation overlay prepared from aerial photograph interpretation. The approximation of per cent of aerial photograph area was used instead of map area to prepare the runoff curve number estimate found in Table 10.

The rainfall and runoff data used for the study of this watershed are listed in Appendix C.



A - Brush - Grass B - Conifer Stands C - Brush Conifer Complex D - Nearly Barren



Area Number	Vegetation	Cover Density %	Hydrologic Soil Group	CN	Aerial Photograph Area - In	Area %	Sum of Products
А	Brush - Grass	60	CD	62	6.67	73.7	
В	Conifer stands	90	CD	42	1.34	14.8	
С	Brush - Conifer complex	70	CD	51	0.57	6.3	
D	Nearly barren - Herbaceous	5	CD	92	0.46	5.1	
	Totals				9.04	100.0	5,980
	Estimated CN						60

Table 10. Ephraim Creek-Alpine Meadows Watershed CN estimation

Appendix C Computer program and basic data

```
1+
             DIMENSION NAME(34) .NDAY(45) .TIME(45) .ADFLO(45) .ADEP(45) .KDAY(30).
            1HOUR(30), PTIME(30), RDEP(30), OTIME(45), RDPTH(30), OTINT(500).
 7*
 3+
            28 HINT ( 50 0) . OFLNT ( 5 00) . RDINT ( 57 0) . OSYN ( 50 0) . OP ( 50 0) . GINC ( 50 0) .
            35 FL OW ( 50 0) . NGAGE (1 0) . SSDEV (50) . SS (50) . HE AD (26) . QSTOR (500) .
 4.4
 5*
            40 85 TP ( 530) + STRFL (500)
 6*
             I COUNTED
 7+
             READ IN RUNOFF DATA, CONVERT TIME TO CONTINUOUS VALUES, AND PRINT.
      C
 .9 *
      C
            PUNOFF DATA WAS PREPARED WITH DECIMAL TIME VALUES.
 9+
         30 READ(5.1) (NAME(1).1 = 1.34)
10+
           1 F OR MAT(3447)
11+
             READ(5.3) N. AREA
12*
             IF(N - 939)117.115.116
        117 APEAK = (AREA / 640.0) + 484.0
13+
14+
             WPITE(6.2) (NAME(I).I = 1.34)
15*
           2 FORMAT( 11.34A2)
16*
           3 FORMAT(5X.13.9X.F9.2)
17*
             WPITE(6.14)AREA
         14 FORMATE SURFACE DPAINAGE AREA = ".F5.0. " ACRES")
13*
19*
             READ(5.23)(HEAD(I).I = 1.26)
20*
          23 F OR MAT(2642)
21+
             D0 66 I = 1.N
2.2 *
             PEAD(5.4) MONTH. NDAY(I). NYEAR. TIME(I), ADFLO(I), ADEP(I)
23*
           4 FORMAT(17.16.16.F5.2.8X.F8.3.F8.5)
74*
             IF(NDAY(I) - NDAY(1))1 30.56.130
25+
         130 TIME(I) = TIME(I) + 24.30
75+
         56 CONTINUE
27*
             ADFLO IS THE NATURAL HYDROGRAPH LESS THE BASE FLOW IN CFS.
      C
28*
             ADEP IS THE CORRESPONDING MASS RUNOFF IN INCHES.
      C
29*
             READLE .25JADJEK . TPEST . TSPAN
30*
         25 FOPMAT(1 7X.F7.2.12X.F8.2.14X.F5.2)
31*
             WRITE (E. 24) MONTH .NDAY (1) .NYEAR
32+
          24 FORMATC DIME FUNDER HYDROGRAPH FOR ". T2. I3. IS. " AS ADJUSTED FOR BA
37+
            ISE FLOW'/' WITH DECIMAL TIME AND FLOW IN CES IS LISTED BELOW.")
34 .
             WPITE ( 5, 15)
```

```
35+
         15 FORMAT(
                        TIME FLOW TIME FLOW TIME FLOW TIME FLOW
36+
           1 TIME FLOW .)
37*
            WEITE(6. B)(TIME(I).ADFLO(I).I = 1.N)
38*
          9 FORMAT(1X+F8-2+F6.2+F8.2+F6.2+F8.2+F6.2+F6.2+F6.2+F6.2+F6.2)
39*
            WRITE (6. 27) ADJPK . TPEST . TSPAN . ADEP(N)
40*
         27 FORMATCY THE APJUSTED PEAK FLOW IS ".F6.2." CFS. "/" THE STORM TIME
           1 OF RISE WAS ".F5.2." HOURS. "/" THE STORM RUNOFF LASTED ".F5.2." H
41*
47*
           20URS. "/" THE ADJUSTED VOLUME OF RUNGEF FOR THIS EVENT IS ".F8.5."
47+
           3INCHES. 1)
44*
           READ RAINFALL DATA, ADJUST TO CONTINUOUS DECIMAL TIME AND PRINT.
      C
45*
            READ(5.5) (NGAGE(I).I = 1.10). M.RBAR
45+
          5 F OR MAT(1042.5X.15.F9.2)
47*
            D0 67 I = 1.M
48*
            READ(5.G)NSTAT. MPM. MNTHR. KDAY(I). KYR. HOUR(I). RTIME(I). RDEP(I)
49*
          F FORMAT(1X.11.12.3X.12.13.15.F3.7.F5.2.F4.2)
50+
            IF(NSTAT - 3) 121, 122, 121
51*
        127 WRITE(6.16)
57*
        16 FORMAT(49H TIME OF DEPTH OF RAIN IS UNKNOWN FOR THIS EVENT )
53+
            GO TO 30
54*
        121 IF(MPM - 2) 123.124.123
55*
        124 HOUP(I) = HOUR(I) + 12.00
56+
            RTIME(I) = RTIME(I) + 12.00
57*
        123 IF(KDAY(I) - KDAY(1)) 125.126.125
58*
         125 HOUR(I) = HOUP(I) + 24.00
55+
             RTIME(I) = RTIME(I) + 24.00
60*
         125 HOUR(I) = ((RTIME(I) - HOUR(I)) / 0.60) + HOUR(I)
51*
         67 CONTINUE
6?*
             WRITE(6,201)(NGAGE(I),I = 1.10)
         201 FORMATI .D. . THE RAIN GAGE RECORD IS FROM . . 10421
53*
64.
             WPITE (5, 22) RBAR
          22 FORMATC THE ESTIMATED AVERAGE WATERSHED RAINFALL IS ".F5.2)
65*
55*
             WPITE (5. 28)
67*
          2º FORMAT( * THE TIME IS DECIMAL. THE DEPTH IS INCHES. *)
68+
             WPITE(6.202)
```

69+ 202 FORMATCY TIME DEPTH TIME DEPTH TIME DEPTH TIME DEPTH 70+ 1 TIME DEPTH\*) 71+ WPITE(6.9)(HOUR(I).RDEP(I).I = 1.M) 72+ C MAKE PAIN AND RUNOFF TIME COMPATIBLE AND MOVE THE RUNOFF HYDRO-73\* GRAPH TO THE PEGINNNING TIME OF RAINFALL . C 74\* IF(NDAY(1) - KDAY(1)) 127.128.129 75\* 127 DO 58 I = 1.M 76\* HOUP(I) = HOUP(I) + 24.0 77\* 52 CONTINUE 78+ GO TO 128 79\* 129 DO 69 I = 1.N 80\* TIME(I) = TIME(I) + 24.00 \$1+ 69 CONTINUE 87\* 128 GTADJ = HOUR(1) - TIME(1): 83\* DO 52 K = 1.N 84\* OTIME(K) = TIME(K) + GTADJ 85\* 52 CONTINUE 26+ ADJUST THE RAINFALL TO THE ESTIMATED AVERAGE AMOUNT BY USING C 87\* C CONSTANT (PX / PT) RATIOS. 88\* DO 53 K = 1.M 89\* ROPTH(K) = RDEP(K) \* (RBAR / RDEP(M)) 90+ 53 CONTINUE RATIO IS THE RATIC OF RUNOFF VOLUME TO ESTIMATED AVERAGE RAINFALL \$18 C 97\* C VOLUME FOR THIS EVENT. 93\* RATIC = ADEP(N) / ROPTH(M) 94\* COMMENCE THE MAJOR LOOP WHICH FINDS THE NEAR OPTIMUM TIME C 05+ INCREMENT FOR DEVELOPING THE COMPOSITE SYNTHETIC HYDROGRAPH FROM C 96\* C TRIANGULAR INCREMENTAL HYDROGRAPHS WHERE TR = 3/8 TB. 97\* DELTA = 0.00 93\* NRUN = 1 99\* 195 DELTA = DELTA + 0.20 100\* IF(NPUN - 30)114 .1 35 .1 35 101\* 175 WPITE(6.21) 102\* 21 FORMATE ODELTA HAS BEEN INCREMENTED 30 TIMES.")

```
103+
             GC TO 115
104*
         114 DO 60 I = 1.500
105+
             0°(I) = 0.00
106*
             OFLNT(I) = 0.00
107*
             QINC(I) = 0.00
108+
             RDINT(I) = D.DO
109*
             SFLOW(I) = 0.00
110*
             QTINT(I) = D_0DD
111+
             QSYN(I) = 0.00
112*
             RHINT(I) = 0.00
113*
          ED CONTINUE
114=
             CALL INTRP (DELTA, GTIME, ADFLO, GTINT, GFLNT, N. NO)
115*
             CALL INTRP (DELTA, HOUR, RDPTH, RHINT, RDINT, M. MR)
             TP = 3.3 + DELTA
116*
117*
       C
             COMPUTE THE SYNTHETIC MASS RUNOFF CURVE. THE RUNOFF FOR EACH DELTA
118*
       С
             TIME INCREMENT. THE ESTIMATED PEAK FLOW FOR EACH INCREMENTAL
             HYDPOGRAPH. AND SUM THE INCREMENTAL HYDROGRAPHS.
119*
       C
120+
             DO 54 I = 1.MR
121*
              QSYN(I) = RDINT(I) * RATIO
122*
          54 CONTINUE
123+
             DO 70 I = 2.MR
124*
              DQ = QSYN(I) - QSYN(I - 1)
125*
              QP(I) = AREAK + DQ / TP
126*
           70 CONTINUE
127*
             LNGTH = MR + 7
128+
              DO 72 J = 1.MR
129*
              QINC(J) = 0.00
130*
              OINC(J + 1) = OP(J + 1) / 3.0
131*
              QINC(J + 2) = 0P(J + 1) + 2.0 / 3.0
132*
              OINC(J + 3) = OP(J + 1)
133*
              9 INC(J + 4) = 0P(J + 1) + 0.80
134 .
              0 INC(J + 5) = 0P(J + 1) + 0.60
135*
              QINC(J + 6) = QP(J + 1) + 0.40
136*
              QINC(J + 7) = 0P(J + 1) + 0.20
```

81

```
137+
             9INC(J + 8) = 0.00
138*
             00 72 I = 1.LNGTH
139+
             SFLOW(I) = SFLOW(I) + DINC(I)
140+
          72 CONTINUE
141*
             BEGIN THE INNER LOOP WHICH FINDS THE NEAR OPTIMUM TIME LOCATION
       C
            FOR THE ACTUAL RUNOFF HYDROGRAPH TO FIT THE SYNTHETIC HYDROGRAPH
142*
       C
             AND COMPUTES THE SUM OF SQUARED DEVIATIONS BETWEEN THE TWO
143*
       C
144*
             HYDROGRAPHS FOR THE TIME INCREMENT SET IN THE MAJOR LOOP.
       C
145*
         145 NREP = 1
145*
             DO 73 I = 1.50
147*
             SSDEV(I) = D.DD
143*
          73 CONTINUE
149*
         109 NJ = NQ + NREP
150*
             DO 59 I = 1.NJ
151+
             DEV = SFLOW(I) - OFLNT(I)
152*
             SSDEV(NREP) = SSDEV(NREP) + (DEV**2)
153*
          59 CONTINUE
154*
             IF(NPEP - 1) 105.105.107
         107 IF(SSDEV(NREP) - SSDEV(NREP - 1)) 106-106-108
155*
             SAVE THIS ACTUAL HYDROGRAPH LOCATION IN OSTOR(I) AND MOVE OFLNT
156+
       C
157*
       C
             ONE DELTA TIME INCREMENT TO THE RIGHT. REPEAT REPITIONS UNTIL THE
158*
       C
             MINIMUM SSDEV IS FOUND.
159+
         106 DO 74 I = 1.N.I
160+
             OSTOP(I) = GFLNT(I)
161*
          74 CONTINUE
162*
             DO 52 I = 1.NG
163*.
             NCNT = NO + NREP - I
164*
             NCNTR = NCNT + 1 ;
165*
             GFLNT(NCNTR) = OFLNT(NCNT)
166+
          62 CONTINUE
167*
             NREP = NREP + 1
168*
             IF(MEEP - 30) 109,109,133
15 ?*
         133 WRITE(5,20)
          20 FORMATC OTHE ACTUAL HYDROGRAPH HAS BEEN MOVED 30 TIMES. GO TO THE
17:1*
```

```
171.
            1 NEXT PUN. ...
17?*
             NRUN = NRUN + 1
173+
             GO TO 146
         108 NPEP = NREP - 1
174+
         146 SS(NEUN) = SSDEV(NREP)
175*
176*
             IF(NRUN - 1) 112.112.111
177*
         111 IF(SS(NRUN) - SS(NRUN - 1)) 11 ?. 112. 113
178*
         112 DO 211 I = 1.NJ
179+
             QRSTR(I) = OSTOR(I)
120+
         211 CONTINUE
191*
             STRDT = DEL TA
             STRP = NREP
132*
183*
             LNSTR = LNGTH
184+
             THOV = STRP . DELTA
185*
             NSAV = NJ - 1
186*
             DO 254 I = 1.LNSTR
187*
             STRFL(I) = SFLOW(I)
188*
         254 CONTINUE
189*
             NPUN = NRUN + 1
190*
             GO TO 105
191*
             RETURN TO THE BEGINNING OF THE RUN LOOP.
       C
192*
         113 CONTINUE
            SET UP TIME COORDINATES.
193*
       C
194*
         115 DO 63 I = 2.500
195+
             AI = I - 1
196*
             OTINT(I) = OTINT(1) + (AI + STRDT)
197*
          E3 CONTINUE
198*
             WPITE ( 5. 8)STRDT . TMOV
199*
           S FORMATI "OTHIS IS THE ACTUAL HYDROGRAPH INCREMENTED TO DELTA = ".
200*
            1F5.2./" AND MOVED ".F6.2." HOURS TO THE TIME OF BEST FIT.")
201*
             WPITE(6.15)
202*
             WRITE(5.9)(OTINT(I).ORSTR(I).I = 1.NSAV)
203*
             WPITE(5.10)STRDI
204+
          10 FORMAT(46HOTHIS IS THE SYNTHETIC HYDROGRAPH WITH DELTA =+ F5-2)
```

205*		WRITE(6,15)
206+		WRITE( $6, 9$ )(OTINT(I), STRFL(I), T = 1.1NSTR)
207*		WRITE (6.11) RATIO
208*	11	FORMATE THE RATIO OF RUNDEF TO RAINFALL FOR THIS EVENT IS
209*		SSO= SS(NRUN - 1).
210* .		WPITE(5, 255) SS0
211*	255	FORMATCE THE MINIMUM SUM OF DEVIATIONS SQUARED IS
212*		GO TO 30
213*	116	STOP
214*		END

END OF UNIVAC 1108 FORTRAN V COMPILATION.

. . .

C \*DIAGNOSTIC\* MESSAGE(S)

```
SUBROUTINE INTRP (DELTA.X.Y.U.W.L.K)
1+
 7.
           DIMENSION U(500) . W(500) . X(45) . Y(45)
 3*
           U(1) = X(1)
 4 *
            W(1) = Y(1)
 5*
           J = L + 1
           X(J) = X(J - 1) + DELTA
 6*
 7*
           Y(J) = Y(J - 1)
*8
           K = 1
9*
           DA 84 I = 2.500
10+
           U(I) = 0.0
11+
           W(I) = 0.0
12+
        84 CONTINUE
13+
           D0 65 I = 2.J
14*
           HOR = X(I) - X(I - 1)
15*
           VERT = Y(I) - Y(I - 1)
15+
           SLOPE = VEPT / HOR
17*
        118 U(K + 1) = U(K) + PELTA
19+
           IF(K - 75)151.150.150
19*
        151 IF(U(K + 1) - X(I)) 119+120+65
20+
        119 W(K + 1) = Y(I - 1) + SLOPE + (U(K + 1) - X(I-1))
21+
           K = K + 1
22*
           GO TO 118
23*
        120 W(K + 1) = Y(I)
24 .
            K = K + 1
25*
         65 CONTINUE
25*
            GO TO 155
27*
        150 WPITE (6. 154)K
28*
        154 FOPMAT( * THERE ARE *. I3. * POINTS ON THE INTERPOLATED CURVE. *)
29*
        155 RETURN
30+
           END
```

END OF UNIVAC 1108 FOPTRAN V COMPILATION. O \*DIAGNOSTIC\* MESSAGE(S)

D FARMINGTON CANYON HALFWAY CREEK WATERSHED SURFACE DRAINAGE AREA = 464. ACRES

FLOW .33 1.79 11 ME 5. ND 6. 43 8. 83 THE ADJUSTED FEAK FLOW IS 2.09 CFS. THE STORM THE DE STES WAS 1.200 HOUSS. THE STORM SUMPER LASTED 11.490 HOUSS. THE ADJUSTED VOLUME OF BUNNOFF FOR THIS EVENT IS 401051 INCHES. THE BURGE HYDORAPIERS 9 12 1942 AS JOUGFED FOR BASE FLOW WITH DECEMPLITHE AND FLOW IN FEAST FLOW THE LITE FOW THE FLOW IN FEAST FLOW THE LITE FOW THE FLOW THE FLOW THE S.S. 1.35 S.S. 1.37 S.S. 1.51 S.103 S.S. 8. 9.400 J.S. 1.400 J.S. 1.51 S.S. 8. 9.400 J.S. 1.400 J.S. 1.51 S.S. 8. 9.400 J.S. 1.400 J.S. 1.51 S.S. 8. 9.400 J.S. 1.400 J.S. 1.400 J.S. 1.400 J.S. 8. 9.400 J.S. 1.400 J.S. 1.400 J.S. 1.400 J.S. 8. 9.400 J.S. 1.400 J.S. 1.400 J.S. 1.400 J.S. 8. 9.400 J.S. 1.400 J.S. 1.400 J.S. 1.400 J.S. 8. 9.400 J.S. 1.400 J.S. 1.400 J.S. 1.400 J.S. 1.400 J.S. 8. 9.400 J.S. 1.400 J.S. 1.400 J.S. 1.400 J.S. 1.400 J.S. 8. 9.400 J.S. 1.400 J.S. 1.400 J.S. 1.400 J.S. 1.400 J.S. 8. 9.400 J.S. 1.400 J.S. 1.400 J.S. 1.400 J.S. 1.400 J.S. 8. 9.400 J.S. 1.400 J.S. 1.400 J.S. 1.400 J.S. 1.400 J.S. 8. 9.400 J.S. 1.400 J.S. 1.400 J.S. 1.400 J.S. 1.400 J.S. 8. 9.400 J.S. 1.400 J.S. 1.4

. 38 TIME DEPTH 5.87 THIS IS THE ACTUAL HYDROGRAPH INCREMENTED TO DELTA = 1.40 THE RAIN GAGE RECORD IS FORM HEAD OF HALFWAY THE RAIN GAGE RECORD IS FORM HEAD OF HALFWAY THE ESTIMATED AVEAGE WATPOSHED RAINAALL IS 1.00 THE THE IS DECTIMAL. THE OFFUH IS RAGES. THE THE OFFUH IS RAGES. 310 THE OFFUH IS RAVES. 32 5.25 30 -50 .12 4.27 .11 4.77 .24 5. 6.48 .40 9.00 .40 16.97 .92 THERE ARE 75 POINTS ON THE INTERPOLATED CURVE.

		Ì				A -TURNEL TURNEL	11011	ALL'SO			
AND	MOVED		1.00	HOUR S	TO THE	TIME OF	BEST F	.11.			
	TIME	1	MO	TIME	FLOW	TIME	FLOW	TIME	FLOW	TIME	FLOW
	.50	•	CO	1.90	• 00	3.33	.00	01.4	.00	6.10	00.
	1.50	•	30	8.90	1.51	10.33	. 90	11.77	.15	13.10	22.
	14.57	•	13	15.90	• 18	17.33	• 03	18.70	cu.	20.10	• 00
THT	S IS I	H	SYNTH	HETIC	HYDROGR	APH WITH	DELTA	1.40			
	TIME	FL	NO	TIME	FLOW	TIME	FLOW	TIME	FLOW	TIME	FLOW
	.50	•	UU	1.90	00.	3.33	.02	4.70	.05	6.10	-13
	7.50	•	20	9.90	.23	10.33	.21	11.70	12.	13.10	56.
	14.50	•	27	15.90	. 30	17.33	. 32	1 8.70	- 30	20.13	. 24
	21.50	•	16	22.90	6U.	24.33	•0.	25.70	10.	27.13	00.
THE	RATIO	10	RUN	OFF TO	RAINFAL	LL FOR TH	HIS FVE	NT IS	12010-		
THE	HIMIH	H	NUS NI	OF DEV	I AT JONS	SQUARED	15	2.45112			

Ω F49MINGTON CANYON HAL™WAY CREEK WATERSHED SUBFACE DRAINAGE ADEA = α.6α. Arore

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HE	RUNDE	F HYDROL	GPAPH F	T 40 -	C+61 1	0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0	STED FO	R BASE	FL OU	
L	H DE CI	MAL TIM	E AND F	ILOW IN	CFS 15	LISTED	BELOW.			
	IMF	FLOW	TIME	FLOW	I IME	FLOW	TIME	FLOW	TIME	FLOW
	3.00	.00	3.60	.12	4.50	. 66	5.00	1.96	5.73	6 . 30
	10.9	3.93	6.25	2.64	6.73	1.38	7.05	1.22	7.50	.70
	A. 7D	. 46	8.50	. 33	6.13	• 20	10.90	.06	12.00	• 00
H	ADJUS	TED PEAN	K FLOW	IS 6.	.30 CFS.					
H	STORM	TIME OF	F RISE	WAS Z	AUOH UT.	۶.				
HH	S TO RM	RUNDEF	LASTEL	00.9 0	HOURS.					
¥	ADJUS	TED VOLI	UME OF	RUNDEF	FOR THI	S EVENT	. SI	12610	INCHES .	
H	NING	GAGE REC	CORD I	FOOM	FADMING	TON WH				
H	ESTIM	ATED AVE	EPAGE 1	ATERSHI	ED RAINF	ALL IS	1.50			
HE	TIME	IS DECIN	HAL. TH	IL DEPT	H IS INC	HES.	1			
	1 IME	DEPTH	TIME	DEPTH	TIME	DEPTH	TIME	DEPTH	TIME	DEPTH
	3.33	.00	3.83	. 08	4.05	.17	4.33	. 30	4.55	. 33
	4.62	.43	4.92	44	5.13	52	5.25	.67	6.08	. 89
	26.9	- 90	7.50	10.1						
H	I SI S	HE ACTU	AL HY DF	ROGRAPH	INCREME	NTED TO	DELTA	- 60		
ND	MOVED	1.20	HOUR S	TO THE	TIME OF	REST F	.11.			
	TIME	FLOW	TIME	FLOW	TIME	FLOW	TIME	FLOW	TIME	FLOW
	3.33	.00	3.93	. 00	4.53	.12	5.13	.48	5.73	1.44
	6.33	- +++	6.93	3.93	7.53	. 1.66	8.13	1.05	8.73	. 56
	9.33	. 35	9.93	.22	10.53	15	11.13	.12	11.73	.01
	12.33	• 03	12.93	.00	1 3.53	.00	14.13	00.		
IH	S IS T	HE SYNTH	HETIC	HYDROGR	APH WITH	DELTA	= .60			
	TIMF	FLOW	TIME	FLOW	I IME	FLOW	TIME	FLOW	TIME	FLOW
	3.33	• 00	3.93	.11	4.53	52	5.13	1.16	5.73	1.97
	6.33	2.42	6.93	2.51	7.53	. 2.18	8.13	1.66	8.73	1.19
	9.33	•63	9.93	. 32	10.53	.17	11.13	• 08	11.73	.00
¥	RATIO	OF RUN	OFF TO	RAINFAL	LL FOR T	HIS EVE	NT IS	.0128	-	
H	HINIW	MNS MN	OF DEV	I AT JONS	SQUARED	SI I	8.17228			

86

D FAPMINGTON CANYON HALFWAY CREFK WATFRSHED SURFACE DRAINAGE AREA = 464. ACRES

FL04 2.19 .08 7 I ME 20.80 TIME 21.73 22.58 \*0.21.66 INCHES THE GUNCE HYPOGGRAPH FOR 3 19 1945 AS ADUNSTED FOR BASE FLOW WITH DECIMALITIE THE FLOW THAT FLOW THAT THAT FLOW THAT FLOW THAT FLOW THAT 1945 - 00 7-000 55 20,113 1279 20,55 4,22 20,8 71,90 - 79 7-20 6,55 20,113 12,79 20,55 4,22 20,8 71,90 - 79 7-20 6,9 20,40 1,12 7,00 -00 70 11,6 2000 THAT 67 12,19 57,00 -00 70 11,6 2000 THAT 67 12,19 55 4005. 06.-DEPTH .05 THIS IS THE ACTUAL HYDROGGAPH INCREMENTED TO DELIA = 7 IME ( 21.58 22.13 FITTINTIC DEFENSES OF ALL CLARACTER ADJUSTED VOLUME OF PUNDEF FOR THIS EVENT IS HEAD OF HALFWAY RAIN SAGE RECORD IS FROM 21.09 THE

	FLOW	2 a 38	3+41	1.000	44.	10.	51.	. fils	cu	40.4
	TIME	21.88	22.88	23.88	24.88	25.88	26.88	27.88	78.88	
	FLOW	• UD	6.12	1.28	* 55	• 30	.18	6J.	× 03	
FIT.	TIME	21.68	22.68	23.69	24.68	25.68	26.68	27.68	28.68	
REST	FLOW	.00	9*93	1.56	.63	. 32	.20	.10	•0.	.01
TIME OF	I INE	21.48	22.48	23.48	84.48	25.48	26.45	27.48	23.48	29.48
TO THE	FLOW	. 00	4.67	1 . 94	• 70	th£ *	• 23	.11	. 05	00 *
HOUR S	TTME	21.28	22+28	23.28	62.42	86.56	26.28	27.28	28.28	82.64
. 80	FLOW	• 00	11.4	c1* c	11.	04*	52.*	-13	.07	.00
AND MOVED	THE PER	21.118	54.7Z	25-0H	80.47	811 ° 57	26.08	27.09	29.08	29. NB

FL04 2.77 5.85 .27 TIME 21.88 22.88 23.88 FL04 .59 7.81 HIS IS THE SWNHETIC HYDONGRAPH WITH MELTA = .20 THE FEOU TIME FEOU TIME FEOU TIME 21.00 -000 -1106 FEOU TIME 2.20 21.00 -000 -21.20 -15 -12.40 -55 -21.40 27.01 -000 -21.20 -171 -21.40 -65 -22.40 27.01 -000 -21.20 -171 -21.40 -65 -22.40 27.01 -000 -21.20 -171 -21.40 -65 -22.40 27.01 -000 -21.20 -21.40 -55 -21.40 -55 -25.40 1.01 -21.01 -21.01 -21.40 -55 -21.40 -55 -25.40 1.01 -21.01 -21.40 -55 -21.40 -55 -21.40 -55 -25.40 HE MINUM SUM FILM SUM FILM SUM FILM 13, 71.40 HE MINUM SUM FILM SUM FILM SUM FILM 13, 71.40 HE MINUM SUM FILM SUM FILM SUM FILM 13, 71.40 HE MINUM SUM FILM SUM

19.72237

D FARMINGTON CANYON HALFWAY CREEK WATERSHED SURFACE DRAINAGE APEA = 4.64. ACRES

. 46 FL04 .58 FLOW 1.93 .11 FL04 .73 .15 1 I ME 2 4.53 74.50 TIME 24.60 26.60 TIME 24.60 26.60 28.60 \*0.03 57 INCHES. The proper write more and a transfer of a scholar of the ast fill and the fill and 0.41 \* .00714 FL0% .69 .72 .02 FL0W .86 TIME DEPTH 74.08 .41 4 INCREMENTED TO DELTA = 2 THE DOR FUST FIT: 2 THE DUN THE FLO 2 5.43 -115 24.20 -1 2 5.43 -115 26.20 -1 2 7.43 -19 1.93915 7 1 I M T 1 I M T 2 4 • 20 2 6 • 20 2 8 • 20 THE RATIO OF RUNDFF TO PAINFALL FOR THIS EVENT IS THE MINIMUM SUM OF DEVIATIONS SQUARED IS 1.933 HEAD OF HALFWAY THE PAIR DEGREGENCIES FON HELFO FALLEN HE EXTENTED AVENCE ANTOHED ANTRELL I HE THE SECTANT IN CONFID ANTRELL I THE GEPTA THE DEFIN IN SIGNES 23.00 05.70 25.0 23.56 231 DELTA . 41 . 38 FLOW HITM H T I ME 23.80 25.80 27.80 THIS IS THE ACTUAL HY 09 05 04 PH IN AND MOVED 1.22 NURSS TO THE TI THE FLOW THE FLOW 25.07 +07 27.40 -07 2 25.07 +07 27.40 -00 2 27.00 -04 27.40 -00 2 1.20 H0495 T0 THE T LOW TIME FLOW .00 27.40 .00 .40 25.40 .20 .04 27.41 .00 FL04 .17 .54 THIS IS THE SYNTHETIC HYD THIE FLOW TIME F 23.00 + 00 23.40 25.00 + 67 23.40 25.00 + 67 25.40 29.00 - 00 29.00 - 01 14E RATIO OF RUNDEF TO RA

I FARMINGTON CANYON HALFWAY CREFK WATERSHED SUPFACE D'AATWAGE APEA = 4.64, ACRES

DEPTH. FLOW 3.22 11.00 12.00 11.0m .0m It.m abuilt the kipuits 4.94 CFS. The Abuilt the of PISE 445 .45 HOURS. The Stope Duncer Lastro #.00 HUUS. The Abuilt fo Volumer of Duncher for THIS EVENT IS .01058 INCHES. T TME 5.18 THE PUTET HYDROCHAF FOR A 1 1925 X2 DUVETED FOR BASE FLOW WITH RECEVELTING AND FLOW IN CFS 15 LISTED RELOW. THE S. TO THE FLOW THE FLOW IN THE FLOW THE FLOW THE S. TO THE FLOW THE FLOW IN THE FLOW THE S. TO THE FLOW THE FLOW IN THE FLOW THE S. TO THE FLOW THE FLOW IN THE FLOW THE S. TO THE FLOW THE FLOW IN THE FLOW THE S. TO THE FLOW THE FLOW IN THE FLOW THE S. TO THE FLOW THE FLOW IN THE FLOW THE S. TO THE FLOW THE FLOW IN THE FLOW THE S. TO THE FLOW THE FLOW IN THE FLOW THE S. TO THE FLOW THE FLOW IN THE FLOW S. TO THE FLOW THE S. TO THE FLOW THE FLOW THE FLOW THE FLOW THE FLOW THE S. TO THE FLOW THE FLOW THE FLOW THE FLOW THE FLOW THE S. TO THE FLOW THE FLOW THE FLOW THE FLOW THE FLOW THE S. TO THE FLOW THE FLO - 2h TIME DEPTH 5.01 .71 THIS IS THE ACTUAL HYDROGRAPH INCREMENTED TO DELTA = .80 PAIN GLGE RECONDIS FROM HEAD OF HALFAAN ESTIMATIO AVENAL CANTENALL IS TIPE IS DECHARL. HIG DEPHILIS INDERS. TIPE IS DECHARL. HIG DEPHILIS INDERS. TIPE OF ON U.73 - 54 U.PC DEPHI S.TO .AD THE

	FLOW	4.53	25.	.29	.19	-12	.08	#C .	.00	
	JHII	5. 33	6. 33	7.32	6. 73	65.9	10.33	11-30	12.30	
	FLOW	4.39	. 73	R.•	. 20	.13	.03	.05	10.	
FIT.	TIME	5.13	6.13	1.17	9.13	ul.6	10.11	11.17	12.10	
REST	FLOW	2.20	1.32	. 38	. 22	.15	6U.	* 0.S	.02	cu
TIME OF	TIME	E6 * 5	5.93	6.93	1.9.1	C6 * 8	66.6	10.93	11.93	12.93
TO THE	FLOW	. 74	2.27	5 4 .	· 73	. 1	.10	. 96	. 0.5	E.0.*
HOUR S	THIT	01. P	5.70	01.1	31.70	0.70	01.0	10.79	11-70	61.71
62.	CLOW	. nn	5.22	147	. 25	.18	11.	10.	. Du	- 111
UD MOVED	TIME	4.50	05*5	6.54	1.2.1	R.5.0	05°F	10.20	11.50	12.50

FLOW 4.55 .45 TIME 5.30 6.30 7.30 

D FARMINGTON CANYON HALFWAY CREEK WATERSHED SURFACE DRAINAGE AREA = 464. ACRES

FL04 .10% .26 .11 .06 110% 11.05 13.05 TIME 6.40 TIME TIME 17.05 18.05 19.05 20.05 21.05 .00732 INCHES. THE BURCH HIPPAGENH FOR 7 26.1953 AS ADUGTED FOR BASE FLOW WITH DECHALTING FOR FLOW IN FYSTED BILDS FLOW THA UPPA FOR 51 11 FLOW ITH FLOW ITH FLOW THA UPPA FOR 510 5-09 5-70 5-21 6-44 7 7-6 - 71 9-20 4-50 FCS THE ADUGTED FEM FLOW 15-00 4-50 FCS THE ADUGTED FEM FLOW 15-00 MORE. G2 · · \* D 24 4D DEPTH 5L0W 3.93 .25 .15 .07 .03 FLOW 3.71 .34 H INCPRENENTED TO DELTA = TIME OF AEST FIT TIME FLOW TIME FLOW TIME FLOW TIME AS TIME A 16.85 HE STIM ALG FORD IS POR HELD OF HELVAN HE STIMATED AFFORE WATFREED ALLYAN HE STIMATED AFFORE WATFREED ALLYAN HE THE IS SOCTAAL, HE DEPH IN STUCKES, 30 HET FORM THE FORM THE FORM THE THE TAS ... THE FORM TO AFFORD AFFORD AT THE TIMF HHIS IS THE SWINFTIC HODORARA HIM PLLA = 23 TIME FLOW TIME FLOW TIME FLOW TIME 16.75 -001 16.45 1.09 16.46 7.33 16.86 17.25 3.401 17.45 1.01 17.45 11.25 3.401 19.45 -703 17.46 1.21 17.45 11.25 4.40 51 10.45 10.414 4.440 15 14.25 4.01 10.45 10.414 4.440 15 14.24 4.40 51 10.414 4.440 15 14.44 5.40 51 10.414 4.440 15 14.44 5.40 51 10.414 10.45 5.4045210 THIS IS THE ACTUAL HYDROGRAPH AND HOVED .20 HOURS TO THE 1145 115.25 115.25 113.25 113.25 20.255 20.255 21.255 21.255

D FARMINGTON CANYON HALFWAY CREEK WATERSHED SURFACE DRAINAGE AREA = 464. ACRES D FARMINGTON CANYON

DEPTH .75 1.19 FL04 .83 11.00 16.00 20.00 TIME 15.12 18.37 TIME 15.22 16.22 17.22 18.22 19.22 20.22 21.22 21.22 INCHES. FLOW THE RUNOFF HYDRIGRAPH FOR 5 4 1954 AS ADJUSTED FOR BASE F WITH OCCIANT THE ADD LIN IS CST LISTS BELOW. THE FLOW THE FLOW THE FLOW THE FLOW THE TOW THE FLOW THE FLOW THE FLOW THE TOW THE FLOW THE FLOW THE FLOW THE TOW THE FLOW THE FLOW THE FLOW THE ADJUST OF THE ADJUST THE STORE AFTER AS AT DHOURS. THE ADJUST OF ATTER AS AT DHOURS. THE STOW RUNOFF LOW THE E ADJUST THE STORE AFTER AS AT DHOURS. • 20 0EPTH .55 1.04 FLOW 1.64 1.74 .15 .15 .15 .15 .15 .15 E TIME OF RECY FIT AL E TIME OF RECY FIT HE 14.00 FLOW TIME 14.00 FLOW TIME 14.00 FLOW TIME 14.00 FLOW TIME 15.00 FLOW DELTA : TIME 1 14.97 17.82 
 RAIN GRE RECOND IS FROM HEAD OF HALFMAY

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 THE CHENT IN THE COFFIL IN THE COFF HOURS TO THE FLOW 14.62 +00 15.62 2.50 15.62 2.50 15.62 -19 17.65 -18 13.65 -18 13.65 -18 13.65 -18 13.65 -18 13.65 -08 23.65 -08 23.65 -08 THIS IS THE ACTUAL HY DROGPAPH THIS IS THE AND MOVED TIME 14.42 15.42 15.42 15.42 17.42 17.42 19.42 20.42 20.42 22.42 TIME ант

7L0W 2.19 .93 .30 .48 TIME 15.22 16.22 17.22 18.22 19.22 20.22 .0.05 75 FL0W 1.39 1.46 .18 .53 .19 TIME TIME 15.07 16.02 17.02 17.02 19.07 20.07 RATIO 14.42 15.42 16.42 17.42 13.42 19.42 20.42 THE

HALFWAY CREEK WATERSHED 4 64. ACRES SURFACE DRAINAGE AREA = O FARMINGTON CANYON

THE         FLO         THE         THE <th>WITH DE CI</th> <th>MAL TIP</th> <th>E AND F</th> <th>LOW IN</th> <th>CFS IS</th> <th>LISTED</th> <th>BELOW.</th> <th></th> <th></th> <th></th>	WITH DE CI	MAL TIP	E AND F	LOW IN	CFS IS	LISTED	BELOW.			
21-00       -00       22-00       -03       23-10       -05       25-00       -03       25-00       -03       31-00       -00         27-00       -08       27-00       -05       20-00       -05       31-00       -00       -05       31-00       -00       -00       31-00       -00       -00       -05       31-00       -00	TIME	FLOW	TIME	FLOW	TIME	FLOW	TIME	FLOW	TIME	FLOW
Z-5-00       -96       Z-00       -95       Z-00       S-00       -95       Z-00       -95       Z-00       -95       Z-00       -95       Z-00       -95       Z-00       -95       Z-00       Z-00 <t< td=""><td>21.00</td><td>• 00</td><td>22.00</td><td>• 03</td><td>23.90</td><td>.05</td><td>23.53</td><td>.73</td><td>23.80</td><td>53.</td></t<>	21.00	• 00	22.00	• 03	23.90	.05	23.53	.73	23.80	53.
Z-00       -08       Z-00       -05       34.00       -05       -00       -00       -00       -00       -00       -00       -00       -00       -00       -00       -00       -00	23.90	.96	24.00	. 93	24.05	. 87	24.50	.37	2 5.00	.19
32.00 - 0.0 3.00 3.00 3.00 - 0.0 3.00 - 0.0 3.00 - 0.0 3.00 - 0.0 3.00 - 0.0 3.00 - 0.0 3.0 - 0	26.00	.08	27.00	• 06	28.03	• 05	30.00	• 05	31.00	• 05
<pre>Int county for that rank for this vert is Int county for that rank for this vert is Int county for volume for this vert is Int county for the volume for this vert is Int county for the vert county for this vert is Int county for the vert county for this vert is Int county for the vert county for this vert is Int county for the vert county for the vert is Int county for the vert county for the vert is Int county for the vert county for the vert is 2000 county for 2000 coun</pre>	32.00	•00	33.00	. 05	34.03	• 03				
THE STORY THE CF FIEX SL 2.00 HUNGS. THE STORY DAVEF LESTED ALLOWER. THE SAUGSTED VOLUPE OF RUNNEF FOR THIS EVENT IS -00356 INCHES. THE RAIN AGE RECORD IS TROM HEAD OF HALVAN THE RAIN AGE RECORD IN THE OFFINITION ACTION AGE 2000 TO 2000 THE TROM THEO REST FILL IS 2.403 THE 2000 THE FLOW THE FLOW THE OFFINITION ACTION AGE 2000 TO 2000 THE TROM THEO REST FILL A - 40 2000 THE FLOW THE FLOW THE OFFINITION ACTION AGE 2000 TO 2000 THE TROM THEO REST FILL A - 40 THE S THE FLOW THE FLOW THE FLOW THE FLOW THE FLOW 2000 THE FLOW THE FLOW THE FLOW THE FLOW THE FLOW 2000 THE STORD TO 2000 THE TROM THE FLOW THE FLOW THE FLOW 2000 THE STORD TO 2000 THE TROM THE FLOW THE FLO	THE ADJUS	TED PE	NK FLOW	SI	.96 CFS.					
THE STORM DWAFF LASTO 11:00 HOLE THE RAUNCE CONTINUE OF RUNNE FOR THIS EVENT IS00356 INCHES. THE RAUNCE VECKOP IS FORM HEAD OF HLLEMAT	THE STOPM	TIME C	JF PISE	WAS 2	AUOH 06.					
THE ADUGIED VOLUPE OF BUNNE FON THIS EVENT IS         -00356         INCHES.           THE RAIN AGE RECORD IF BUNNE FON THIS EVENT IS         -00356         INCHES.           THE RAIN AGE RECORD IF BUNNE FON THAT         -00356         INCHES.           THE FAIL AGE RECORD IF RULME.         -00356         INCHES.           THE FAIL AGE RECORD IF RULME.         -00356         INCHES.           THE FORMATINE REPHA TIME DEFINAT         -111         2.035         -112         2.035         -113         2.035         -113         2.035         -112         2.035         113         2.035         113         2.035         113         2.035         1111	THE STOPM	RUNUFF	LASTED	13.00	HOURS.					
Iffer         State         State <th< td=""><td>THE ADJUS</td><td>TED VOL</td><td>UPE OF</td><td>RUNNEF</td><td>FOR THI</td><td>S EVENT</td><td>SI</td><td>•00356</td><td>INCHES.</td><td></td></th<>	THE ADJUS	TED VOL	UPE OF	RUNNEF	FOR THI	S EVENT	SI	•00356	INCHES.	
THE ESTIMATE OR CONTRACT AND	THE RAIN	GAGE RE	COPD IS	FROM	HEAD OF	HALFWA				
THE TIPE IS DECEMM THE DEFINITION TO THE TOPEN         THE DEFINITION TO THE TOPEN         THE DEFINITION TO THE DUAL	THE ESTIN	ATED AV	LEPAGE W	ATFOSH	ED RATNF	ST LIA	05-			
THE DEFN	THE LIME	IS DECI	MAL . TH	1430 3	H IS INC	HES.				
23.550     .00     70.56     .05     21.57     .07     21.57     .05     21.51     .01     21.51     .01     21.51     .01     21.51     .01     21.51     .01     21.51     .01     21.51     .01     21.51     .01     21.51     .01	TIME	DEPTH	TIME	DEPTH	TIME	DEPTH	TIME	DEPTH	TIME	DFPT
23.2.70       23.2.70       23.2.70       31.2.30       31.2.30       32.2.30       33.2.3.0       33.2.30       33.2.30	20.50	.00	20.65	· 75	21.67	.07	21.82	11.	23.05	- 16
23.472       .3.49       .4.1       30.43       .42         1115       15 Tr Me       FULUL HTNE OF REFT TO DELLA       .40         7114       15 Tr Me       FULUL HTNE OF REFT TO DELLA       .40         7114       17 Tr Me       FULUL HTNE OF REFT TO DELLA       .40         7114       17 Tr Me       FULUL HTNE OF REFT TO DELLA       .40         7114       FULUL HTNE OF REFT TO DELLA       .40       .41         7114       FULUL HTNE OF REFT TO DELLA       .40       .40         7144       FULUL HTNE OF REFT TO DELLA       .40       .41         7140       .40       .41       71.33       .41       71.41       .40         74.40       .41       71.33       .41       71.41       .41       71.41       .41         74.41       .41       71.43       .41       71.41       .41       71.41       .41 </td <td>23.20</td> <td>.28</td> <td>23.42</td> <td>.31</td> <td>27.00</td> <td>• 31</td> <td>28.00</td> <td>. 32</td> <td>28-03</td> <td>E. S</td>	23.20	.28	23.42	.31	27.00	• 31	28.00	. 32	28-03	E. S
<pre>HHIS IS THE ACTUAL HICPOGRAMM INCREMENTED TO DTLIA = .40 A00 NOVED .40 HHUSPOGRAMM INCREMENTED TO DTLIA = .40 A00 NOVED .40 HHUSPOGRAMM INCREMENTED FELSE FILS .40 23.570 .40 23.99 .45 23.13 .41 23.17 .42 23.10 .40 25.570 .43 23.25 11 23.25 11 23.10 .40 25.570 .43 23.25 11 23.25 11 23.10 .40 25.570 .43 23.25 11 23.25 11 23.10 .40 25.570 .43 23.25 11 23.25 11 23.10 .40 25.570 .43 23.25 11 23.25 11 23.10 .40 25.570 .43 23.40 .43 31.33 41 23.10 .40 25.570 .43 23.40 .43 31.33 41 23.10 .40 25.570 .44 24 24 24 24 24 24 24 24 24 24 24 24 2</pre>	29.92	.37	29.98	. 41	30.43	54.				
AND NOVED	THIS IS T	HE ACTL	IAL HY DR	DGRAPH	INCREME	NTED TO	DFLTA			
THE FLOW THE	AND MOVED	. 80	I HOURS	TO THE	TIME OF	BEST F	.11			
27.57         -00         7.90 <th7.90< th="">         7.90         7.90         <th7< td=""><td>TIME</td><td>FLOW</td><td>TIME</td><td>FLOW</td><td>TIME</td><td>FLOW</td><td>TIME</td><td>FLOW</td><td>TIME</td><td>FLON</td></th7<></th7.90<>	TIME	FLOW	TIME	FLOW	TIME	FLOW	TIME	FLOW	TIME	FLON
22-50     -0     23-90     -15     25-10     -16     25-10     -16     25-10     -17     26-10     -17       25-50     -07     75-90     -16     75-10     -15     25-17     15     25-10     -16       25-50     -07     75-90     -16     75-10     -15     25-17     15     25-10     -16       25-50     -07     75-90     -16     75-10     15     25-17     16     26-10     -10       25-50     -07     75-90     -16     31-13     -10     31-10     -10     31-10     -10       25-50     -07     30-90     -16     31-13     -10     31-10     -10     31-10     -10       21-50     -07     30-90     -16     31-13     -10     31-10     -10     -10       21-50     -07     30-90     -16     -14     1	20.50	• 00	20.90	CU.	21.33	10.	21.73	• 02	22.13	. U3
74-57         73-76         74-70 <th< td=""><td>22.50</td><td>+0ª</td><td>27.90</td><td>· 05</td><td>23.33</td><td>.61</td><td>23.70</td><td>ħħ.*</td><td>24.10</td><td>. 71</td></th<>	22.50	+0ª	27.90	· 05	23.33	.61	23.70	ħħ.*	24.10	. 71
2.5.         0.7         7.4.9         0.6         7.5.         1.0         7.6         1.6         7	24.51	• 34	24.90	.19	25.33	. 15	25.70	.10	26.10	.08
7.3         5.450         -05         7.490         -05         7.430         -05         7.430         -05         7.430         -05         7.410         -010         3.410         -010	26.50	.07	26.90	• 76	27.30	. 05	27.70	.05	28.10	.05
30.50         -07         30.40         -03         31.53         -09         31.53         -09         31.53         -09         31.10         -00         34.10         -00         24.10         -01         24.10         -01         24.10         -01         24.10         -01         24.10         -01         24.10         -01         24.10         -01         24.10         -01         24.10         -01         24.10         -01         24.10         -01         24.10         -01         24.10         -01         24.10         -01 <t< td=""><td>28.50</td><td>• 05</td><td>28.90</td><td>• 05</td><td>29.33</td><td>. n5</td><td>2.9.70</td><td>• 05</td><td>30.10</td><td>.06</td></t<>	28.50	• 05	28.90	• 05	29.33	. n5	2.9.70	• 05	30.10	.06
3.3.59 - 05 32.40 - 05 31.33 - 04 31.77 - 04 34.10 - 02 1.5.14E SWHHTIE HYDROGRAPH WITH RELIA = .40 THIS IS THE SWHTIE WITH RELIA = .40 THIS IS THE SWH	30.50	.07	30.90	• 03	31.33	• 03	31.70	· 09	32.10	.08
34.50 .0П НІЗ IS THE SYNINFTIC НУВРАДАРИ WITH RELIA = .40 ТНІЗ IS THE SYNINFTIC НУВРАДАРИ WITH RELIA = .40 21.57 .02 20.30 .05 21.30 .10 21.70 .10 22.10 22.56 .01 20.30 .22 21.30 .10 21.70 .10 26.10 .10 22.56 .01 24.90 .22 21.30 .10 27.70 .00 26.10 .00 22.56 .01 24.90 .11 24.91 .10 27.70 .00 26.10 .00 22.56 .01 24.90 .11 24.91 .10 27.70 .00 26.10 .00 22.56 .13 24.90 .11 24.91 .10 27.70 .00 26.10 .00 23.56 .13 24.90 .11 24.91 .10 24.10 .09 35.10 .01 23.56 .13 24.90 .00 41.41.50 .112 34.70 .09 35.10 .01 23.56 .13 24.90 .11 24.10 .00 35.10 .00 35.10 .01 24.56 .11 24.51 .10 24.10 .00 25.10 .00 25.10 .00 25.56 .13 24.90 .11 24.10 .00 25.10 .00 25.10 .00 25.56 .11 24.51 .10 24.10 .00 25.10 .00 25.10 .00 25.10 .00 25.56 .11 24.51 .10 24.10 .00 25.10 .00 25.10 .00 25.10 .00 25.56 .11 24.51 .10 24.10 .00 25.10 .00 25.10 .00 25.10 .00 25.56 .11 24.51 .10 24.10 .10 24.10 .10 24.10 .00 25.10 .00 25.10 .00 25.56 .13 24.90 .11 24.10 .11 24.10 .10 24.10 .00 25.10 .0	32.50	. Jc	32.90	• 05	33.33	• 04	33.70	ti0.	34.10	.0.
MIS IS THE SWINTIE HYDORGRAPH WITH RELIA	34.50	-01								
THY FLOW THY	THIS IS T	THY S 3H	HETIC H	YDROGR.	APH WITH	DELTA	0.4.0			
Z <sup>11,57</sup>	TIME	FLOW	TIME	FLOW	TIME	FLOW	TIME	FLOW	TIME	FLOW
22:50         -21         27:30         -22         27:40         -22         27:40         -20         27:40         -20         27:40         -20         27:40         -20         27:40         -20         27:40         -20         27:40         -20         27:40         -20         27:40         -20         27:40         -20         27:40         -20         27:40         -20         27:40         -20         27:40         -20         27:40         -20         27:40         -20         26:40         -00         26:40         -00         26:40         -00         26:40         -00         27:40         -00         26:10         -00 <td< td=""><td>20.50</td><td>• 00</td><td>20.30</td><td>• 05</td><td>21.33</td><td>.10</td><td>21.70</td><td>.16</td><td>22.10</td><td>.1.</td></td<>	20.50	• 00	20.30	• 05	21.33	.10	21.70	.16	22.10	.1.
74.50         0.00         84.90         -20         52.57         -19         75.76         -10         75.10         100 <t< td=""><td>22.50</td><td>.21</td><td>22.90</td><td>.22</td><td>23.30</td><td>.30</td><td>23.70</td><td>• 39</td><td>24.10</td><td>. 50</td></t<>	22.50	.21	22.90	.22	23.30	.30	23.70	• 39	24.10	. 50
26-57 - 0.0	24.50	Un.	24.90	• 28	25.33	.19	25.70	• 09	26.10	.00
2.5-57 - 10, 2.4-9 - 11 2.9-1 - 12 2.9-70 - 09 50-10 - 11 2.9-70 - 09 50-10 - 11 2.5-70 - 09 35-10 - 01 2.5-70 - 10 3.7-30 - 10 3.2-10 - 00 32-10 - 00 32-10 - 00 32-10 - 00 12 3.1-10 - 10 12 12 12 12 12 12 12 12 12 12 12 12 12	26.50	.00	26.90	υ.	27.33	.00	27.70	.00	28.10	.04
30-50 13 30-90 49 31,33 42 31,70 409 32,10 401 32-50 61 32,32 90 410 35,33 40 THE RATIO OF RUNOFF TO RATIFALL FOR THIS EVENT IS 40712	28.50	.08	28.90	.11	29.33	.10	29.70	•00	30.10	.11
32.57 .03 32.90 .00 33.50 .00 THE RATIO OF RUNOFF TO RAINFALL FOR THIS EVENT IS .00712	30.50	.13	30.90	. 14	31.33	.12	31.70	•00	32.10	. DE
THE RATIO OF RUNOFF TO RAINFALL FOR THIS EVENT IS "00712	32.50	•03	32.90	.00	33.30	00.				
	THE RATIO	OF RUN	IDFF TO	PAINFA	LL FOR T	HIS EVE	NT IS	.0071	2	

89

2.07504

MINIMUM SUM

D FARMINGTON CANYON HALFWAY CREEK WATERSHED SUPFACE DRAINAGE AREA = 464. ACRES

DEPTH .73 FL0W .99 FL0W 1.79 .32 .14 11.75 12.75 1 7.00 TIME 14.15 17.15 20.15 23.15 THE RUPEL HOREFLY FOR 7 13 JSE2 AS JOUNTED FOR BASE FLOW WITH DECIMULTING RANG FLOW IN CF3 15 JSTED BFLOR WITH CELMULTING RANG FLOW ITHE FLOW ITHE FLOW THE FLOW ITHE FLOW ITHE FLOW ITHE FLOW IN-OTHER FLOW ITHE FLOW ITHE FLOW ITHE FLOW IN-OTHER FLOW ITHE FLOW ITHE FLOW ITHE FLOW ITHE ME STODAY STOPE FLOW ITHE FLOW ITHEFLOW ITHE FLOW ITHEFLOW ITHE FLOW ITHEFLOW ITHEFLO - 60 TIME DEPTH 12.15 .60 1.09 THE RAIN GALE RECOMD IF POWN RAINSTIT THE ESTIMATED AVERAGE MATTOSHED PATIMALL IS THE THE IS DECTMAL IN DEFINING IN INDEC. 11.75 COCTMAL 12.01 12.05 13.1 11.75 COCTMAL 12.77 06 13.10 10.05 12.45 6.46 13.77 06 13.10 10.05 THIS IS THE ACTUAL HYDROGRAPH AND MOVED 11.75 14.75 17.75 20.75 TIME

FLOW 2.37 TIME 14.15 17.15 FL04 3.09 .39 .17 .05 H INCRENENTED TO DELTA = E THE FLOW THAT FLO THE FLOW 1148 FLO 12.95 1.99 1.555 5.0 15.95 4.0 19.55 4.1 19.95 1.9 10.72.55 4.0 24.95 4.07 72.55 4.0 T IMF DELTA FL0W 1.57 HIIM H AP 1 1.20 HOURS TO THE T FLOW TIME FLOW -00 11.35 -00 1.10 15.35 -00 -12 11.35 -73 -12 11.35 -73 -12 21.35 -09 -03 -03 -03 -01 THIS IS THE SYNTHETIC HYDPOGR TIME FLOW 12.35 .70

.013 50 FL04 2.53 2.79016 1 3.55 14.75 1.97 15.35 1.35 15.95 1.78 16.55 17.75 00 19.35 00 THE RATIO OF PUNOFF TO RAINFALL FOP THIS EVENT IS THE MERTHOW SUM OF DEVIATIONS SQUARED IS 2.9010 1 1ME 1 2.95 1 5.95 TIME FLOW 11.75 .00 14.75 1.92 17.75 .00 11.75

• D D 3 54

.12 33 4

D FARMINGTON CANVON HALFWAY CREEK WATERSHED Surface Drainage Area = 464. Acres

-20 -33 06PTH \*25 \*57 -11 -62 -03 FL04 1 IME T I ME 6.05 TIME 7.15 10.15 13.15 TIME 7.15 10.15 13.15 . D D3 54 INCHES. THE DIFFET HYDORADE FOR 9 11 JUSS X ROUNTED OR BASE FLOW WITH DECEMPLITHE ADD FLOW IN FFS IS LISTED RELAVE THE ALTON THE FLOW IN FFS IS LISTED RELAVE THE 4.00 00 6.00 08 710 11 8.00 00 HE ADDUFTED PECK FLOW 1 11 13.00 00 HE ADDUFTED PECK FLOW 1 21 11 13.00 00 HE ADDUFTED PECK FLOW 1 22 00 00055. .12 .47 FL04 .07 .28 .07 FL04 . 15 . 34 TIME DEPTH 11 # WITH PELTA = .60
TIME FLOW TIME
5.95 .07 6.55
5.95 .17 9.55
11.95 .15 12.55 5.62 7.67 10.02 6.55 9.55 12.55 DELTA Really dis Refconds 15 show Hello di Hur Way Estimati Refconds 15 show Hello di Hur Way Time 15 sectimat, he nephi 15 lacets. 1.00 Time 02 hi inte nephi 15 lacets. 1.00 Time 02 hi inte 02 hi 11 states. 1.00 14 states. 100 states. 1.00 states. 1.05 states. 1.00 states. TIME ADJUSTED VOLUME OF RUNDEF FOR THIS EVENT IS I TIME OF REST FIT. TIME FLOW TIM 5.95 .05 6.5 8.95 .10 12.5 11.95 .07 THIS IS THE ACTUAL HYPROGRAPH INCREMENTED TO x D ...60 HOUPS TO THE T FLOW TIME FLOW ...07 5.35 .02 .17 1.35 .27 .42 11.35 .29 .00 14.35 .00 IS THE SYNTHETIC HYDROGRAP 11ME DEPTH 4.93 .03 7.22 .38 9.35 .76 AND MCVED 4.75 7.75 10.75 13.75 IME SIH. ЗНТ НЕ WH1 H1

D FARMINGTON CANYON HALFWAY CREEK WATERSHED Surface Drainage area = 464. Acres

FL04 DEPTH • 24 1 1 . 00 THE 20.0010 DEMONSTOR OF A CFS. THE ADJUSTED PEAK FLOW IS ...AL CFS. THE ADJUSTED PEAK FLOW IS ...AL ADJUSTED ADJ 14.40 THE RUNOFF HYDROGRAPH FOR 9 13 1963 AS ADJUSTED FOR BASE FLOW WITH RECHALTING ADJELVU IN CF 315 10210 0.000 TTHE FLOW 1146 FLOW 1146 FLOW 1147 114.000 -001 14.000 -291 14.53 -880 15.50 -251 17.1 TIME DEPTH 14.28 .18 15.45 .44 \*\*\* THE RAIN GALE RECORD IS FOOM HALO OF HALFWAY THE EXTIMATED AFFORE WAITTSHED PAILFALL IS THE THE TO SECHALL IN DIFTIL IS INCHES, IT AT THE DEFINIT THE DEFINI IT.47 200 14.27 255 14.27 200 1 14.27 255 14.28 200 400

THIS IS THE ACTUAL HYDROGRAPH INCREMENTED TO DELTA = .60

DH D	URS TO THE	TIME OF	HE S T	- 11 -			
JMI	FLOW	3 HI 1	FLOW	TIME	FLOW	TIME	FLOW
.27	. no	14.87	.15	15.47	Lh.	16.07	.64
.27	.19	17.37	.12	18.47	.07	19.07	.07
1.27	50.	20.87	. 05	21.47	•0•	22.07	.03
5.27	· n2	23.87	.02	24.47	CU.	25.07	.00
.27	00.						

TIME 16.07 19.07 FLOW .43 .09 
 THIS IS THE SYMTHTIC HYDROGRAPH WITH DELTA .60

 TIME FLOW
 TIME FLOW
 TIME FLOW
 TIME FLOW
 TIME FLOW

 T1.67
 .00
 14.27
 .01
 14.87
 .24
 15.47

 T1.67
 .00
 14.27
 .01
 14.87
 .24
 15.47

 T6.67
 .05
 17.27
 .33
 17.83
 .21
 18.47

FLOW .51

• 0 05 86 • 09 35 2 19.67 .01 THE RATIO OF RUNOFF TO RAINFALL FOR THIS EVENT IS THE MINIMUM SUM OF DEVIATIONS SOUAPED IS .0355

D FARMINGTON CANYON HALFWAY CREEK WATERSHED SURFACE DRAINAGE AREA = 464. ACRES

TIME DEPTH 20.73 .53 21.27 1.26 FL0W 3.21 .47 .11 FL0W 20.73 TIME 22.32 25.32 28.32 28.32 31.32 THE BURK HYDROSH FOR 7 18 1155 AS ADJUSTED FOR ANSE FLOW WITH DECIMAL THE 200 FLOW IN CT15 ITSED BELOW THE FLOW ITHE FLOW IN CT15 ITSED BELOW 70.55 FLOW 71.07 199 27.07 3.08 124.00 .57 26.00 20.55 FLOW 7.07 199 27.07 3.08 24.00 .57 26.00 20.55 FLOW 7.07 11.50 27.07 20.05 186 20.015 FLOW 17.07 27.07 00.05 186 20.015 FLOW 17.07 20.00 186 20.015 FLOW 17.07 00.005 186 20.010 FLOW 11.50 00005 186 20. . 60 TIME DEPTH 20.50 .31 21.22 1.17 FLON 3.33 .56 .17 .02 
 NUES IS THE AFTUL VENDBORGH INCREFIEND ID INLA

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FL04 3.01 T IME 22.32 25.32 ·01245 FL0W 2.21 3.62 90 5 HIS IS THE SWINFIEL HYDOGGAPH WITH RELTA = .03 THEF ELON THEF ELON THE ELON THE 23-33 -00 23-25 -75 21-12 1-09 21-75 22-32 -66 23-55 -712 21-15 9 21-75 HE 22-92 -66 25-52 -70 HE 21-00 9 400F TO 21-21 1-15 1-59 24-75 HE 21-00 9 400F TO 21-21 1-15 23-42 24-75 HE ATANUM SUM 0F DEVTATIONS SQUARED 15 -3.62920

D FARMINGTON CANYON HALFWAY CREEK WATERSHED SURFACE DRAINAGE APEA = 464, ACRES

FL04 TIME DEPTH 11.05 .28 FLOW 1.08 .14 .05 TIME 11.72 13.72 15.72 15.72 THE BURGEN FURDER BY 11955 AS JOJUSTED FOR BASE FLOW WITH DECEMAL THE RAND FURDER FLOW THE FLOW THAT DECEMAL THE ROW IN FILSTO BELOW THAT SET THE FLOW THE FLOW THE FLOW THAT SET TO A THAT THE FLOW THE FLOW THE THE FLOW THE FLOW THE FLOW THE FLOW THE THE ADURTED FEEK FLOW FOR THE FLOW THE FLOW THE STORE THE FLOW AS TABLES TO AND SATE THE STORE THE FLOW AS TABLES TO AND SATE THE ADURTED FEEK FLOW FLOW THE FLOW TS THE FLOW THE ADURTED FEEK FLOW AS TABLES TO AND SATE 08. TIME DEPTH 10.75 .20 H INCREMENTED TO DELTA = E TINC PENSTETT. INE ELOW TINE FLO 10.92 53 11.22 1 10.92 73 15.12 1 11.92 77 15.12 1 11.92 77 15.22 1 11.92 70 19.22 1 5 5 \* THE PAIN GAE RECORD IS FROM HEAD OF HALFWAY THE ESTIMATED AVERAGE MATFOSHED ATTIFALL IS THE THE SY DECHMEL, HAT DEPHH IS INCHES, THE REPH THE OFFIH IS INCHES, II.0.12, 0.00 10.27, 0.01 10.48, 0.07 11 TIME DFPTH 10.27 .01 11.82 .18 TIME DEPTH 10.12 .00 11.35 .35

FL0W 1.00 TIME 11.72 13.72 FLOW FL0W .72 .72 .72 .72 .72 .72 11.37 TIME T 14.17 .03 14.52 .00 14.92 .01 THE RATIO OF RUNDEF TO RAINALL FOO THE CEVENT TS THE MINIMUM SUM OF DEVIATIONS SQUARED TS .1423 HYDPOGRAPH WITH DELTA FLOW • 55 T IME 10.92 12.92 14.92 TIME FLOW 10.52 .09 6.9 . THIS IS THE SYNTHETIC TIME FLOW TIME 10.12 .00 10.57 12.12 .98 12.53

O FARMINGTON CANYON HALFWAY CPEEK WATERSHED Surface Drainage Area = 464. Acres

FL0W 1.33 .13 .05 .05 FL04 TIME DEPTH 24.13 .35 TIME TIME 25.27 28.27 31.27 34.27 34.27 27.00 .00672 INCHES. THE RUMPE WITHOGGRAPH FOR & 71 1955 AS ADJUSTED FOR BASE FLOW WITH DICTAL THE WIT FLOW IN FLOW THE THE FLOW THE FLOW THE FLOW THE PALAN CONTRACT THE FLOW THE FLOW THE PALAN CONTRACT TO THE FLOW THE PALANCE TO THE THE FLOW THE FLOW THE PALANCE TO THE THE FLOW THE FLOW THE PALANCE FLOW THE THE FLOW THE FLOW THE PALANCE FLOW THE FLOW THE FLOW THE FLOW THE PALANCE PALANCE IS ADDRESS TO THE FLOW THE FLOW THE PALANCE PALANCE IS ADDRESS TO THE FLOW THE FLO TIME DEPTH .00 .33. FL04 .94 .18 .05 .02 
 HIS IS IME ACTUME VERDBORPH INCRECENTED FOLLET =

 AND NUMED
 Image of the properties of the propertes of the properties of the properties of the propertie TIME - 60 .58 THE POINT GAR REFOOD IS FOOM HEAD OF MALFWAY HE ESTIMATED AVEPAGE WATFSMED ANIMALL IS THE THE TS PECTMAL: HE OPFIH IS INCHES. THE OPFIH IME OPFIH IS INCHES. THE OFFIH IME OPFIH IS INCHES. DELTA THIS IS THE SYNTHETIC HYDROGRAPH WITH TIME DEPTH 23.00 .04 24.83 .47 100. 01.45

FL04 . 99 TIME 25.27 28.27 •01159 FLOW 1.08 . 34 86 4 24.67 THE FLOW THE FLOW THE FLOW THE 27-27 0.01 23-47 0.02 25-467 25-467 25-47 0.01 29-47 0.01 40 71-67 25-47 0.01 29-47 0.01 7-40 71-67 THE SATTO OF RUMOFF TO RALIVELE FOR THIS VERNI IS THE MINHOW SUM OF DEVIATIONS SOURCE OF 30466

07 60 0.

D FARMINGTON CANYON MOREIS CREEK WATERSHED SUPFACE DRAINAGE AREA = 156. ACRES

O FARMINGTON CANYON MORRIS CREEK WATERSHED

SURFACE DRAINAGE AREA = 156. ACRES

TIME DEPTH 4.08 .17 12.08 .59 FL04 .05 FL0W .18 FL0W .02 .03 6.10 5.53 TIME 5.53 9.53 13.53 I IME THE RUNCET HEOROGRAPH FIR 9 12 1942 AS 40 USTED FOR BASE FLOW WITH DECTRAL THE AND FLOW TO FST STED BELOW TTRE FLOW THE FLOW THE FLOW THE 3-50 -00 4-00 -03 5-51 -07 5-90 -06 6-11 6-55 201 7-00 -12 7-53 -01 8-00 -06 9-0 10-00 -07 12-00 -00 . 80 TIME DEPTH 3.25 .16 10.92 .51 FLOW FL04 -02 -03 THIS IS THE ACTUAL HYDROGRAPH INCREAMENTED TO DELTA = AND MOVED .8D MOURS TO THE TIME OF BEST FIT. TIME 4.73 
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	HILL	4.58	6.3	8.62	
	DEPTH	.37	1.01	1.38 .	
	TIME	24.43	5.67	8.75	
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SI	T T	• #	5.	7.	
E DEPTH	DEPTH	. n3	. 56	1.20	
IMAL . TH	TIME	3.47	5.00	7.22	
IS DEC	DEPTH	.00	.53	1.17	1.57
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0.4. THIS IS THE ACTUAL HY DROGRAPH INCREMENTED TO DELTA =

	FLOW	00-	.26	10	- 05	.03	00.			FLOW	.02	. 22	.10	- 07	00.	
	TIME	3.90	5. 90	7.90	6. 93	11.90	13.90			TIME	3.90	5.90	7.93	9.90	11.90	
	FLOW	.00	.18	.10	.05	.03	.03			FLOW	.00	.18	.12	.08	.00	.003 36
II.	TIME	3.50	5.50	7.50	9.50	11.50	13.50		04. =	TIME	3.50	5.50	7.50	9.50	11.50	NT IS
PEST F	FLOW	.00	.10	.13	.06'	+0.	CO .	. 00	DELTA	FLOW	.00	.13	.15	.08	.02	IIS EVE
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2.00	FLOW	.00	.03	.27	.10	* D4	.02	.00	HE SYNTH	FLOW	.00	+ 0.4	£2°	60.	.05	OF RUNG
MOVED	TIME	2.30	4.30	6.3n	8.30	10.31	12.30	14.30	S IS TH	TIME	2.30	4.30	6.30	R.30	10.30	RATIO
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THE MINIMUM SUM OF DEVIATIONS SQUARED IS

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THE MINIMUM SUM OF DEVIATIONS SQUARED IS

O FARMINGTON CANYON MORRIS CREEK WATERSHED SURFACE DRAINAGE AREA = 156. ACRES

.05 FLOW .02 TIME DEPTH 19.67 .64 • 64 FL04 FLOW .19 .26 .04 .02 TIME 19.55 20.55 21.55 22.55 TIME 19.55 20.55 21.55 22.55 23.55 71ME THE ADJUSTED PEAK FLOW IS "48 CFS. THE STORM TIME OF RISE WAS 95 HOURS. THE STORM PUNCF LASTED 3.90 HOURS. THE STORM POLUME OF RUMOFF FOR THIS EVENT IS "00329 INCHES. THE RUMPE HYDROGRAPH FOR 8 3 1945 AS ADJUSTED FOR BASE FLOW WITH RECENALITHE AND FLOW IN CFS 15 11510 BELOW THE FLOW THE FLOW IN FLOW THE FLOW THE 18.10 .00 19.05 .48 19.53 .21 19.45 .09 2147 . 20 FL04 -09 -33 -05 -05 TIME DEPTH • 23 FLOW .10 .33 20. . 88 DELTA = 19.38 TIME TIME 19.35 20.35 21.35 22.35 23.35 TIME 19.35 20.35 21.35 22.35 23.35 TT. THIS TS THE ACTUAL HYDROGRAPH INCREMENTED TO DEL AND MOVED ...6D HOURS TO THE TIME OF RESTFIT. TIME FLOW TIME FLOW TIME FLOW TI THE PILY DATE OF EFFORM THE OF THLEE THE EXTENTED ACREACE OF FORM THE OF THE FILL IS THE THE TO FORTHALL IS THE OF THE OF THE OF THE OF THE OF TAPE OF THE OF THE OF THE OF THE TAPE OF THE OF THE OF THE OF THE TAPE OF THE OF THE OF THE OF THE OF THE TAPE OF THE OF THE OF THE OF THE OF THE TAPE OF THE OF THE OF THE OF THE OF THE OF THE TAPE OF THE TAPE OF THE TAPE OF THE O FL04 FL04 .04 .04 .05 .00 .45 .03 .03 HYDROGRAPH WITH 20.15 21.15 22.15 23.15 TIME 20.15 21.15 22.15 23.15 19.15 19.15 , 60 HOURS TO THE FLOW FLOW - 00 - 10 - 10 - 10 . 10 . 13 TIME FLOW 18.10 .00 19.05 .48 22.00 .00 19.95 19.95 21.95 21.95 19.95 20.95 21.95 22.95 56° × I THIS IS THE SYNTHETIC TIME FLOW 18.75 - 00 19.75 - 31 20.75 - 17 21.75 - 04 22.75 - 01 . 30 . 15 . 09 21.75 19.75 20.75 21.75 22.75 19.75 18.75 19.75

D FARMINGTON CANYON MORRIS CREEK WATERSHED SURFACE DRAINAGE AREA = 156. ACRES

FLOW 7 I ME 2 3.00 .002 95 INCHES. 

DEPTH 1.36 22.72 TIME . 69 TIME DEPTH 22.45 THE BAIN GAGE RECOPD IS FROM HEAD OF MILLER THE ESTIMATED AVENDE WATHER PAINELL IS 1.36 THE THE IS DECIMAL. THE OFFINI SINGHES. THE THE IS DECIMAL. THE OFFINI SINGHES. • 59 21.83 +U • 21.05 • 00 20.75

THIS IS THE ACTUAL HYDROGRAPH INCREMENTED TO DELTA = 1.20 AND MOVFD

FL04 .18 TIME 25.55 FLOW TIME 24.35 MOVED 4.80 HOURS TO THE TIME OF REST FIT. TIME FLOW TIME FLOW THE FLOW TI 0.275 4.00 21-95 4.00 23.15 4.00 24.15 4.6.75 4.04 27-95 4.00 23.15 4.00 20.75

FL04 25.55 TIME FL0W • 08 • 01 • 01 74.35 .02056 TIME THIS IS THE SYNTHETIC HYDROGRAPH WITH DELTA = 1.20 20.75 .00 21.95 .01 21.15 .05 74.35 26.75 .07 27.95 .05 29.15 .03 30.35 THE RATIO OF RUNOFF TO RAINFALL FOR THIS EVENT IS FLOW THE MINIMUM SUM OF DEVIATIONS SQUARED IS T IME TIME FLOW TIME FLOW

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EVENT IS

THE RATIO OF RUNOFF TO RAINFALL FOR THIS THE MINIMUM SUM OF DEVIATIONS SQUARED IS

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+ 81 00 +

## O FARMINGTON CANYON MORRIS CREEK WATERSHED SURFACE DRAINAGE AREA = 156. ACRES

FLOW .02 FLOW FLOW .19 .03 TIME 2.50 TIME TIME 1.58 2.58 3.58 1 I ME 1.58 2.59 3.58 ADJUSTED VOLUME DE RUNDEF FOR THIS EVENT IS "DDIII INCHES. FLOW -17 -03 -0370 ..20 HT930 • 39 FL04 .18 .04 LOW TIME LOW TIME 11 1.38 11 2.38 03 3.38 EVENT IS . DELTA = d b. FIT. TIME TIME 2.00 1.38 .30 INCREMENTED TO TIME OF BEST FI S 15 THC SWIMFIC HYDROGRAPH WITH DELTA TTRE FLUM TTPE FLUM TTPE FLUM 1.7" 00 -09 05 1.13 .11 1.7" 11 -01 1.13 .11 2.7" 01 2.0" 06 00 2.1" 11 -01 1.14 2.1" 11 1 TTME FLOW 1.18 .11 7.19 .07 3.18 .03 3.18 .03 CLL FOO THIS EVI FLOW .09 .05 1 IME 1.18 2.18 3.18 HOUPS TO THE TIME FLOW .9P .00 1.99 .08 2.99 .02 THIS IS THE ACTUAL HYDROGPAPH AND MOVED 1.78 1.78 2.78 3.78 THE

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THE RATIO OF PUNDEF THE MINIMUM SUM OF

D FARMINGTON CANYON MORRIS CREEK WATEPSHED SUPFACE DPAINAJE APEA = 156. ACRES

FL0W .43 .19 .05 FLC4 .16 TIME DEPTH 12.58 .68 FL04 - 15 - 15 - 05 - 05 - 03 TIME 12.55 13.55 14.55 15.55 1 3.70 1 3.70 1 8.00 TIME 12.55 13.55 14.55 14.55 16.55 16.55 17.55 .00410 INCHES. THE DULATE HATPOGENA FOR A IT 1947 AS ADJUSTED FOR BASE FLOW WITH DICHALL THE VAP FOR MATTAGE STATED BASE FLOW TTAGENA THE VAP FOR MATTAGE STATED BASE THE DULATE AND THE FLOW TTAGENA THE FLOW TTA 1940 THE ADJUSTED FLOW THE FLOW TTAGENA THE ADJUSTED FLOW THE ADJUSTED FLOW THE FLOW THE FLOW THE FLOW THE ADJUSTED FLOW THE ADJUSTED FLOW THE FLOW THE FLOW THE FLOW THE ADJUSTED FLOW THE ADJUSTED FLOW THE ADJUSTED FLOW THE ADJUSTED FLOW THE FLOW THE FLOW THE FLOW THE ADJUSTED FLOW THE ADJUSTED FLOW THE ADJUSTED FLOW THE ADJUSTED FLOW THE FLOW THE FLOW THE FLOW THE ADJUSTED FLOW THE FLOW THE FLOW THE FLOW THE ADJUSTED FLOW THE ADJUST . 20 TIME DEPTH 12.33 .57 FL04 -03 -26 -07 -07 -07 -01 A MARK CEC. CENTRATED AFEOR CENTRATE MEDIAL. ITYE TS CECHAL. HE OFF. ITYE FORMAL. HE OFF. AD ROTO - AD MOUST OFF. A THE FLOW THE FLOW THE FLOW THE THEFE FLOW THEFE FLOW THEFE FLOW THE THEFE AND THESE AND THE FLOW THEFE THEFE AND THEFE AND THEFE AND THE THEFE AND FLOWER AND THE FLOW THEFE HATED FLOWER AND AND THE FLOWER AND THE WALL THE MANHOW SCHOOL FLOWER AND THE WALLS THE MANHOW SCHOOL FLOW .27 T IME DELTA HYDPOGRAPH WITH THIS IS THE SYNTHFIC

D FARMINGTON CANYON MORRIS CREEK WATERSHED SURFACE DRAINAGE AREA = 156. ACRES THE RUNCE HEROBGAMM FOR 7 27 1951 AS ADUGTED FOR BASE FLOW THE FOLD THE FUNCTION AND FOLD THE FLOW THAT FLOW THE FOLD THE FUNCTION FOR THE FLOW THAT FLOW THAT FOR THE FOLD THE FOLD THAT FOLD THE FOLD THAT FOLD THAT FOLD THAT FOLD THAT FOLD THAT FOLD THE FOLD THAT FOLD THAT FOLD THAT FOLD THAT FOLD THAT FOLD THE FOLD THAT FOL

НИ SIST И К АСТИАН И КОРЛАРАН И КОРСИЕМИЕD ТО DELTA = .40 АНО ИНСЕСТИАН И КОРЛАРАН И КОРСИЕМИЕD ТО DELTA = .40 АНО ИНСЕСТИАН ОНОВУЕТСИИ ТНАЕ ГОЛИ ТНАЕ ГОЛИ ТАТОМ ГООТ ТНАЕ ГОЛИ ТНАЕ ГОЛИ ТНАЕ ГОЛИ ТНАЕ ГОЛИ 27.00 - 00 25.40 - 00 25.46 - 00 26.28 - 00 26.28 - 00

THIS IS THE SYNTHETIC HYDPAGRAPH WITH DELTA = .40 THE FLOW THE FLOW THE FLOW THE FLOW THE FLOW 7.08 -00 7.448 -02 23.48 -02 24.28 -07 24.68 -05 25.08 -00 25.48 -03 25.48 -01 76.28 -00 THE PATTOR FLOW PET TO REINFLY FUELD 10 50 THE RAINAW SUM OF DEVIATIONS 50.4610 15 THE RAINAW

D FARMINGTON CANYON MORPIS CREFK WATERSHED Surface Drainage Area = 156. Acres

TIME DEPTH 10.30 .34 11.68 .87 . 09 . 14 . 06 FL04 -10 -10 T1ME 10.03 12.25 13.50 TIME 10.35 12.35 14.35 TIME 10.35 12.35 14.35 \*00200 INCHES. THE BROKE HYPOGABLEGO A A 1951 AS JOJUSTO FOR BASE FLOW WITH DECTMAL THE AND FEAN INC. A 11510 BEAC FLOW THE FLOW THE FLOW THE FLOW THE FLOW THE ADD SHOT AND THE FLOW THE FLOW THE ADD SHOT AND THE FLOW THE FLOW THE ADD SHOT AND THE FLOW THE FLOW THE FLOW THE PLOT AND THE FLOW THE FLOW THE FLOW THE PLOT AND THE FLOW THE FLOW THE FLOW THE PLOT AND THE FLOW THE FLOW THE FLOW THE PLOT AND THE FLOW TH . 40 .00263 TIME DEPTH 10.13 .27 11.08 .72 FLOW -.01 .01 FL04 .02 .12 • 00 58 9 1 1.13 1 1.08 THIS IS THE ACTUAL HYOROGRAPH INCREMENTED TO DELTA AND MOVED 1.60 HOURS TO THE TIME OF BEST FIL. TIME FLOW TIME FLOW TIME FLOW TIME 9.95 11.95 13.95 11.95 13.95 .76 14.75 .00 THE RATTO OF RUNDEF TO RAINFALL FOR THIS EVENT IS THE MINIMUM SUN OF DEVIATIONS SQUARED IS .0058 
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 -DELTA = FLOW T FLOW T .00 9 .13 11 .13 11 THE PAIN GAGE RECOVD IS FROM HEAD OF MILLER THE ESTIMATED RURCRAGE MATTORNED BAINFALL IS THE TIME IS DECIMAL. HE DEPTH IS INCHES. THE DEPTH THE DEPTH THE DEPTH . 11 -.00 .17 HYDROGRAPH WITH 9.85 T INE 9.55 11.55 13.55 9.55 11.55 13.55 0 1.60 HOURS TO THE T FLOW THE FLOW -. ON 11.15 -.10 -. ON 11.15 -.10 -. ON 11.15 .05 -. ON 15.15 .05 TIME DEPTH 9.53 .08 10.67 .57 FLOW .12 TIME FLOW TIME FLOW 9.00 -000 0.25 000 11.25 11.01 33 17 11.00 01 11.60 11.150 11.1 12.50 23 12.55 02 12.50 EEAK FLOW IS THIS IS THE SYNTHETIC H TIME FLOW TIME 8.75 .00 9.15 11.15 . 69 FL04 .00 .08 8.75 10.38 10.75 8.75 10.75 12.75

D FARMINGTON CANYON MORRIS CREEK WATERSHED SURFACE DRAINAGE AREA = 156. ACRES

DEPTH .75 FLOW .08 .18 .04 FL04 .44 .10 .03 .37 .37 .13 .13 .02 .02 TIME 22.33 23.25 24.50 25.75 TIME 21.67 TIME 21.12 22.12 23.12 24.12 TIME 21.12 22.12 23.12 24.12 24.12 25.12 25.12 26.12 INCHES. 10 BASE FL . 20 FL04 •05 •27 •05 .0 02 88 DEPTH . 85 FL0W .45 .12 .03 FL04 • 15 • 16 • 03 • 03 • 03 • 03 • 03 HE RUNOF HYDROGRAPH FOR 8 19 1951 AS ADUGTED FOR 845 WITH DECLANT THE ADDR FLOW NETS 15 151510 BECUV. 22.07 00 22.08 00 22.18 00 22.55 05 22.67 12 27.67 10 22.23 22.75 00 22.55 22.56 11 27.75 15 24.00 27.75 00 22.55 27.57 11 27.75 10 24.00 22.55 01 25.55 00 24.55 01 27.55 01 25.55 01 25.55 01 26.07 01 27.07 01 22 25.55 01 25.55 01 26.07 01 27.07 01 27.07 01 27.07 01 16 5100H ELATIO FOR FLOY IS 5.55 01 25.55 01 16 5100H ELATIO FOR FLOY IS 5.55 01 25.55 01 16 5100H ELATIO FOR FLOY IS 5.55 01 25.55 01 20.07 10 00005. H INGREMENTED TO DELTA = E THE OF REFIT: THE FLOW TITHE THE FLOW TITHE 20.22 18 20.92 21.22 18 21.92 21.22 18 21.92 21.22 21 22.22 21.22 04 23.92 24.72 01 23.92 TIME [ 21.50 26.00 TIME 11ME 20.92 21.92 22.92 23.92 24.92 25.92 25.92 25.92 .95 
 HE MAIN GAGE RECOPD IS FOOM HEAD OF MILLER

 THE ESTIMATED AVERAGE WATENBOOR DAILARL IS

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D FAPMINGTON CANYON MORPIS CREFK WATERSHED SURFACE DRAINAGE AREA = 156. ACRES

FLOW . 48 . D2	DEPTH	FL04 .43	FL04 .22
FLOW TIME 4.75 6.00 INCHES.	1 I ME	0 TIME 6.18 8.18	T I HE 6.18 8.18 8.18
0R BASE FLOW - 76 - 03 - 00	DEPTH .57	FLOW • 39 • 00	FL04 .25 .0038
85160 F 86L0W- 11ME 94-59 94-59 7-00 7-00	70. 1 IME 5.00	DELTA IT. 11. 5.78 7.78	= .40 TIME 5.78 NT TS NT TS
1952 AS 40JU 1952 IS LISTED 11ME FLOH 1442 43 43 5.53 43 5.53 40 6.75 40 6.75 40 6.75 10 100 100 100 100 100 100 100 100 100	AD OF MILLER PAINFALL IS INCHFS. TIME DEPTH 1.83 .41	CREMENTED TO 4E OF 9EST F TIME FLOW 5-38 - 00 7-38 - 02	WITH DELTA IME FLOW 38 -18 38 -18 38 -05 09 THIS EVE ARED IS
TOR 8 1 FLOW IN CF FLOW 10 -13 -13 -10 -10 -10 -10 -10 -10 -10 -10 -10 -10	5 F D 0 H HE 4 A T F P S H E D 4 E D E P T H I 0 E P T H • 3 I	205PAPH IN T0 THE TI FLOW . 10 . 05 . 10 . 10	ITDPOGRAPH FLOW .09 .11 RAINFALL F RAINFALL F
ITTRE AND F TTRE AND F TTRE AND F TTRE AND F 4.25 20 4.25 20 5.25 20 5.25 20 5.25 20 5.25 20 4.25 20 5.25 20 4.25 20 7.25 20 7.55 20 7	E RECORD IS 0 AVERAGE 5 0 ECIMAL 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1	ACTUAL HYDR 1.20 HOURS 0W TIME 00 4.98 12 6.98 00 8.38	SYNTHETIC H ON TIME CO 4.98 IS 6.98 RUNDEF TO SUM OF DEVI
RUNNFF H H DECIMAL TIME FL 5.00 6.20 6.25 7.00 7.00 7.00 7.00 7.00 7.00 7.00 7.0	PAIN GAG ESTIMATE TIME IS TIME DEP 4.58	> 15 THE MOVED 11ME FL 4.58 8.58	S IS THE TIME FL 4.58 6.58 RATIO OF MINIMUM
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SI EVENT

TO PAINFALL FOR THIS DEVIATIONS SQUARED IS

## D FARMINGTON CANYON MORRIS CREEK WATERSHED SURFACE DRAINAGE AREA = 156, ACRES D FARMINGTON CANYON

.17 .04 FL04 .17 .03 DEPTH THE RUNGE HUDDAGRAPH FAD 7 26 1951 A5 ADJUSTED FOR BASE FLOW WITH DECENT THE ADD FLOW IN FST SILTED RELOW. THHE FLOW THRE FLOW TIME FLOW TIME 13.08 .00 13.55 .00 13.65 .03 13.75 .05 13.00 13.05 .01 13.55 .01 13.65 .01 13.75 .01 13.00 13.55 .01 13.55 .01 13.55 .01 13.75 .01 13.00 13.55 .01 13.55 .01 13.55 .01 13.75 .01 13.00 14 ADJUSTED REK FLOW 17 .01 12.02.55 THE STORM THE OF FLOW 15 .01 10.01 20.25 THE STORM PITEO FLOW 15 .01 10.015 .01 14 STORM RINGE LATED .01 145 .01 145 .01 15 .01 14 STORM RINGE LATED .01 145 .01 145 .01 15 .01 145 .01 14 STORM RINGE LATED .01 145 .01 145 .01 15 .01 145 .01 14 STORM RINGE LATED .01 145 .01 145 .01 15 .01 145 .01 14 STORM RINGE LATED .01 145 .01 145 .01 15 .01 145 .01 14 STORM RINGE LATED .01 145 .01 145 .01 15 .01 145 .01 TIME TIME 41.50 42.50 43.50 + 20 TIME DEPTH FL0W .18 .09 HIS IS THE ACTUAL WY COOGRAPH INCRETENTED TO DELIA = AND MUSE IN THE FISH STSTIT. AND MUSE IN THE FISH STSTIT. AND THE FICA THRE FICATE THRE FIT 0 4 \* THE PRIM GGE RECORD IS FORM HEAD OF WILLER THE EXIMATED AVERAGE WATEVASTOD PAINFALL IS THE FIRE TO EVEN AND AT A THE OPEN THE OPEN THE OF THE OPEN THE OPEN

FL0W .16 HYDROGRAPH WITH DELTA THIS IS THE SYNTHETIC

11ME 41.50 42.50 .002 55 .19 .00 = .20 TIMC 41.30 42.30 FL0W .12 11ME 1.13 1.12 TIME FLOW 40.90 .06 41.90 .09 42.90 .00 .13 TIME FLOW 41.70

2 65 00 . THE RATIO OF RUNOFF TO RAINFALL FOR THIS EVENT IS THE MINIMUM SUM OF DEVIATIONS SOURCED IS \*DU39

MORRIS CREFK WATERSHED D FARMINGTON CANVON MORRIS CREFK SURFACE DRAINAGE APEA = 156. ACRES

		FLOW	00	12	18	.03	11.	104	10								
FLOW		TMF	14.86	15. 33	16.53	17.75	18.83	20.00	21.25					INCHES.			
R RASE		FLOW	. 05	17	- 11	*U*	.10	su.	10.	- 00				0 N3 62			
STED FO	SFLOW -	TIME	14.87	15.25	16.25	17.50	18.75	19.75	21.00	22.25				. SI		22-1	
SULAA SI	ISTED F	FLOW	.05	.15	.14	. 05	.07	. 07	- 02	.00				EVENT	MTILED	IL IS	
4 1954 4	CFS IS L	TIME	14.75	15.15	16.03	17.25	18.53	19.50	20.75	22.00	37 CFS.	OR HOUR	HOURS .	FOR THIS	HEAD DF	D RATNEA	
8 40	NI MOT	FLOW	. 00	. 14	. 20	. 06	• 05	• 73	20.	10.	. 21	WAS 1.	00.8	RUNDEF	FROM	ATFRSHEI	
IGRAPH F	IE AND F	TIME	14.50	15.00	15.75	17.00	10.25	19.25	20.50	21.75	K FLOW	F PISE	LASTED	UME OF	COPD IS	EPAGE W	
F HYTRO	MAL TAM	FLOW	.00	.17	• 30	• 0.7	• 04	.19	* 04	.01	TED PEA	TIME 0	PUNDEF	TED VOL	GAGE RE	ATED AV	
RUNDE	H DE CI	TIME	14.25	14.32	15.50	16.75	13.15	19.00	20.25	21.50	SULUA	STODY	S TORM	ADJUS	NITA	ESTIM	
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TIME DEPTH 15.25 .87 2th \* TIME DEPTH 14.92 TIME IS DECIMAL. THE REPTH IS INCHES. THE DEPTH TIME REPTH IS INCHES. 13.75 -001 14.05 -02 14.58 -04 15.75 -11 16.83 1.09 16.75 1.10 .04 1.14 14.58 16.75 .02 14.05 H

-17 -04 -04 TIME 16.95 20.95 24.95 . 80 .14 .10 THIS IS THE ACTUAL HYDROGRAPH INCREMENTED TO DELTA AND MOVED 2.40 HOURS TO THE TIME OF BEST FIT. THE FLOW TIME FLOW TIME FLOW TIME DELTA 16.15 20.15 24.15 50. HYDROGRAPH WITH 15.35 19.35 23.35 D 2.40 HOURS TO THE T FLOW TIME FLOW - 00 14.55 00 - 03 18.55 01 - 07 22.55 01 THIS IS THE SYNTHETIC TIME 13.75 17.75 21.75

FL0W .16 TIME 16.95 20.95 FLOW .10 16.15 TIME THE FLOW THE FLOW THE FLOW THE 17.75 ON 14.55 ON 15.35 OS 15.15 17.75 JP 15.55 ON 25.45 ON 20.15 17.75 JP 15.55 ON 25.45 THE 84110 OF PUMOFF TO ANIMALL FOR THS VEW TS THE MINHUM SUM OF EVIATIONS SQUARED IS, ODM FLTA : FLOW .05 .09

D FARMINGTON CANYON MORRIS CREEK WATERSHED Surface Drainage Area = 156, acres

FLOW .57 TIME DEPTH FL04 The profession for all 11195 AS ADJUSTED FOR BASE FLOW WITH DECEMALTING FOUN IN EST STERTOR BELOW 19125 FLOW 1914E FLOW TIME FLOW TIME 19125 FLOW 1914E FLOW TIME FLOW TIME 19125 FLOW 19145 FLOW 19150 FLOW 100 2012 20125 FLOW 19150 FLOW 2010 2012 20125 FLOW 19150 FLOW 2010 2012 20135 FLOW 2010 FLOW 100 2013 20135 FLOW 2010 FLOW 100 2013 20145 FLOW 100 FLOW 100 FLOW 115 FLOW 105 NIGHTS-HIE GLOW FLOW 100 HIS STAND 2013 100 FLOW 2010 FLOW 100 FLOW 115 FLOW 105 NIGHTS-19.60 71.53 . 60 • 53 TIME DEPTH HHIS IS THE ACTUAL HY GOOGRAPH INCREMENTED FO DELIA = AND WAYED 1:30 TOURS TO THE THEO FOR STST FIT. THE FOUN THE FLOW THEF FOUNTHEF FOUNT 1:11 AND 14-23 AND 20-33 AND 20-33 22-13 ADD 22-23 AND 23-33 AND 23-33 19.47 • 35 THE RAIN GAGE RECORD IS FROM SUNSET THE ESTIMATED AVERAGE WATFOSHED RAINFALL IS THE TIME IS DECIMAL. THE DEPTH IS INCHES. TIME DEPTH . 48 19.37 + 24 TIME DEPTH .00 19.23 TIME DEPTH 19.17

FLOW TIME 21.53 FL0¥ .05 FLOW ... THIS IS THE SYNTHFIIC HYDROGRAPH WITH DELTA = .60 TIME FIRD TIME FLOW TIME FLOW TIME

D FARMINGTON CANYON MORRIS CREFK WATERSHED SURFACE DRAINAGE AREA = 156. ACRFS

DEPTH .28 100. 200. -00 -01 .12 .04 TIME 73.17 24.00 25.25 1 I HE 2 2 • 97 2 3 • 97 2 4 • 97 2 5 • 97 22.97 23.97 24.97 TIME TIME 22.82 .00089 INCHES. THE RUWRE HYRROGRAMH FOR 7 20 1956 AS ADJUSTED FOR BASE FLOW WITH DECTMAL TIME AND FLOW IN CFS IS LISTED BELOW. TIME FLOW THE FLOW THE FLOW THE FLOW TIME .00254 .16 FL04 .02 .08 .02 FL0W .07 FL04 .04 .09 TIME DEPTH Ĩ. 71MF 73.00 23.75 THIS IS THE ACTUAL HYDROGRAPH INCREMENTED TO DELTA 22.70 22.17 23.17 24.17 25.17 = .20 TIME 22.17 TIME OF REST FIT. TIME FLOW TIME 23.17 (1) 21.97 (1) 21.97 (1) 23.57 (1) 23.77 22.17 (0) 23.17 (0) 24.77 (2) .35 23.25 11 21.21 12 21.21 12 21.29 10 21. 23.26 00 24 25 01 24.29 01 25.40 22 25.60 00 25.40 12 25.40 10 25.40 15 11 25.40 15 15.40 15.40 15 15.40 15 15.40 15 15.40 THE BALL OF A DEFINITION IS FOR A DIVET THE STITUTE OF A DEFINITION OF A DIVET THE THE STOFFAND A DIVET OF A DIVET THE DEFINITION OF A DIVET OF A DIVET 22-37 11 271 06 2245 00 22-32 11 271 06 2245 00 124 06 00 124 0 THIS IS THE SYNTHETIC HYDROGRAPH WITH DELTA 8.18.6 FLOW 10. 22.75 23.53 24.75 22.57 23.57 24.57 25.57 22.57 J INE .60 HOURS TO THE T LOW TIME FLOW .00 22.37 .00 .09 23.37 .01 .05 24.37 .01 25.37 .01 .12 TIME FLOW 22.37 .00 22.50 FL04 .00 .09 .05 .00 TIME FLOW AND MOVED 23.25 24.25 25.50 25.50 22.17 23.17 24.17 25.17 TIME 22.17

99
# MORPIS CREEK WATERSHED SURFACE DRAINAGE AREA = 464. ACRFS O FARMINGTON CANYON

FLOW 1.61 0EPTH .25 .74 1.11 .51 FLOW TIME 1 22.93 23.48 11ME 24.00 27.75 TIME 23.80 25.80 27.83 INCHES. THE RUNCF HYDROGRAPH FOR A 21 1957 AS ADUNSTED FOR BASE FLOW WITH OCCIMENTER AND FOUN IN CFS151 TSTED BLOW THRE FLOW THRE FLOW TIME FLOW TIME 22.40 2017 TSTE 22.51 1111 23.75 99 24.67 24.0 2017 2017 1211 23.75 95 24.75 24.75 24.0 2017 1211 23.75 24.75 04 \* .18 FLOW . 38 . 87 . 06 TIME DEPTH DELTA = 22.67 23.40 TIME 06. TIME OF REST FIT. TIME FLOW TIME 23.00 19 25.40 23.00 1.40 25.40 27.40 1.10 27.44 INCREMENTED TO HAURS TO THE TIME FLOW 22.60 .00 24.50 1.94 26.60 .17 28.60 .00 THIS IS THE ACTUAL HYDROGRAPH AND MOVED TIME 22.20 24.20 26.20 26.20

FL04 1.28 TIME 23.80 25.80 FLOW .81 .34 .00 .0080 24.271 .01 .14. ELON TIME ELON TIME 24.271 1.65 24.61 1.57 25.01 .13 21.04 24.271 1.65 24.61 1.57 25.01 1.31 23.04 1.65.01 .17 25.61 0.01 23.03 23.04 1.65.01 1.77 25.61 0.01 23.03 23.04 1.66.110 0.01 21.01 25 24.04 25 24.04 THIS IS THE SYNTHETIC HYDPOGRAPH WITH DELTA TIME FLOW TIME FLOW

MORRIS CREFK WATERSHED SURFACE DRAINAGE APEA = 156. ACRES D FARMINGTON CANYON

-100 -100 -000 FL04 DEPTH .62 1 6.10 1 9.00 TIME TIME 21.93 21.93 25.93 27.93 29.93 23.18 \*00111 INCHES THE DURFE HYDORDAMH FOR 8 19 1959 AS ADJUSTED FOR BASE FLOW WITH DECEMPLITHE ADJ FLOW IN CFYST 1510 BELOW THE FLOW THRE FLOW ITHE FLOW THR FLOW THR THRE FLOW THRE FLOW THRE FLOW THR FLOW THR 15.00 0.27 17.00 0.09 19.00 19.00 14. FL0W - 00 - 75 - 75 - 75 TIME DEPTH 23.03 .49 
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.80 STORM RUNNFF LASTED 7.60 HOURS. ADJUSTED VOLUME OF RUNOFF FOR THIS EVENT IS ESTIMATED AVEPAGE WATFRSHED PAINFALL IS TIME IS DECIMAL. THE DEFTH IS INCHES. TIME DEFTH TIME DEFTH ITME DEFTH 20-33 &00 20.45 #08 : 22.40 #18 SUNSET THE RAIN GAGE RECORD IS FROM THE ESTIMATED AVERAGE WATFRSHI THE TIME TS DECIMAL . THE DEPTI .67 23.55

FL0W .02 .08 TIME 21.93 23.93 25.93 39 FL04 .02 .05 100. .00193 21.53 23.53 25.53 3WI 1 -22-23 02 22-71 02 23-11 03 23-53 22-33 07 24-73 05 25-13 08 25-53 26-33 00 24-73 09 25-13 08 25-54 HE 8110 0 940FT 0 811FALL F0 H1S EVENT 15 THE AIANUM SUM 0F 0E VILTIONS SOURCED 15 001 H DELTA 10. THIS IS THE SYNTHETIC HYDROGRAPH WITH 21.13 23.13 25.13 I IME TIME FLOW 20.73 .00 22.73 .02 24.73 .05 20.73 22.73 24.73 TIME FLOW 20.33 .00 22.33 .02 20.33

0.8

D FARMINGTON CANYON HORPIS CREEK WATERSHED SURFACE DRAINAGE APEA = 156. ACRES

TIME DEPTH 20.10 .26 20.77 .96 22.03 1.57 FL04 FLOW .12 20.10 20.77 22.03 1 6.00 TIME 21.98 24.98 -00179 INCHES THE BUCKET WITHOOGRAPH FOR 7 18 1955 AS ADUNSTED FOR BASE FLOW WITH DECHALINE HOW FOLD WITHS FLOW THE THE FLOW THE FLOW THE FLOW THE FLOW THE DECAM FLOW THE FLOW THE FLOW THE FLOW THE DECAM FLOW THE FLOW THE FLOW THE FLOW THE DECAM FLOW THE FLOW THE FLOW THE FLOW THE DECAM FLOW THE FLOW THE FLOW THE FLOW THE DECAM FLOW THE FLOW THE FLOW THE FLOW THE DECAM FLOW THE FLOW THE FLOW THE FLOW THE DECAM FLOW THE FLOW THE FLOW THE FLOW THE ADUNTED FLOW THE ADUNTED FLOW THE FLOW THE FLOW THE ADUNTED FLOW THE ADUNTED FLOW . 60 TIME DEPTH 20.00 .12 20.52 .87 21.28 1.41 FL0W .00 
 INIS IS THE ACTUAL HYDROGARPH INCREMENTED TO DELTA =

 AND HYDRO 2.40 MOURS TO THE FILMS FIFT.

 ATTHE FLOW TITHE FLOW TITHE FLOW TITHE FLOW TITHE 219-58

 ATTA 27-58

 ATTA 27-58
THE MAIN GAGE RECOPD IS FOOM BICE THE ESTIMATE NERSAGE MAINENLI IS 1.20 THE TEST DECTMAL. THE DEFTH IS INCHES. THE THE DEFTH THE DEFTH IS INCHES. THE 19-58 -001 19-15 -01 19-45 -00 20-00 20-17 -08 -002 19-15 -01 19-45 -00 20-20 20-17 -08 -002 19-15 -01 19-45 -00 20-20 20-17 -08 -002 19-15 -01 20-12 -00 20-20 20-17 -08 -002 19-16 -01 20 0.12 -00 20-20 20-18 -002 19-10 -002 10-10 -002 20-20 20-18 -002 10-10 -002 10-10 -002 10-10 -002 20-20 20-18 -002 10-100 -002 10-100 -002 10-100 -002 10-1000 -002 10-100 -002 10-1000 -002 10-1000 -002 10-1000 -002

FL0W TIME 21.98 24.98 ·001 49 FLOW .07 HIS 15 THE SWIHFTIC HYDROGRAPH WITH FILT 2 .65 TIME TEQN TIME TEQN TIME FLOW TIME 19-55 .05 70.18 .06 72.79 .03 21.98 22.58 .07 74.18 .06 72.19 .05 24.48 23.58 .07 75.14 .06 25.78 .05 24.48 23.58 .00 75.14 .00 25.78 .05 24.48 HI BATTOO FORDET TO ATMALLE OF HIS EWIN TS

O EPHRAIM CREEK WATERSHED ALPINE CATTLE PASTURE OR MEADOWS SURFACE DRAINAGE AREA = 376. ACRES

FLOU .64 .60 DEPTH TIME 16.50 17.00 19.00 THE .0 02 76 INCHES. THE DURF WORGEAPH FOR 7 16 19:1 AS ADJUSTED FOR BASE FLOW WITH DECEMALTINE AND FOW IN CYSIS 11:51C0 BEAC FLOW THE FLOW THE FLOW THE FLOW THE FLOW THE 15:15 - 00 11:25 - 92 16:20 - 53 16:00 - 50 16:2 15:55 - 64 16:60 - 77 16:22 - 54 15:00 - 113 19:2 17:59 - 64 17:50 - 64 17:55 - 15 16:00 - 113 19:2 20:010 - 05 - 22:00 - 00 TIME DEPTH The Four The Four The Fuer The 0 %\* THE PAIR DAGE RECORD IS FORM ALPINE AREA B THE ESTIMATED ACCAGE WATERSHED CALVERL IS THE THE THE DECHAL, THE OFFIN IS INCHES, THE THE OFFIN IS INCHES, IFON -00 IS.33 -34

FLOW .48 .05 TIME 17.60 19.60 21.63 0.40 FLOW .60 .07 H INCREMENTED TO DELTA = E TIVE OF REST FIT. ITHE FLOW TIME FLU 16.83 63 17.20 4 16.83 63 17.20 4 18.83 13 19.23 4 27.83 03 71.20 4 .80 HUNS TO THE T .80 HUNS TO THE T .00 IS.40 -00 .15 IR.40 -12 .05 20.40 -04 .05 22.40 -00 THIS IS THE ACTUAL HYDROGRAPH TIME FLOW 16.00 .00 18.00 .16 20.00 .05 27.00 .00 AND MOVED

FL04 17.60 FLOW .65 .00 .006 90 . 18 52 4 HIS IS THE SWUTHFIC HYDROGRAPH WITH DELTA = "40 THW FLOW THW FLOW THE FLOW THY IS.M - 300 15.40 - 22 15.40 - 40 17.20 18.40 - 300 15.40 - 25 15.40 - 11.20 19.10 - 300 16.40 - 11.20 - 11.20 - 22 HE RATIO OF BUMOFT ID RAILYALL FOR THIS FWENT IS THE MINIMUM SUM OF DEVIATIONS SQUARED IS

D EPHRAIM CREFK WATERSHED ALPINE CATTLE PASTURE OR MEADUWS SURFACE DPAINAGE APEA = 376. ACPES

FLOW .13 FL0W 1.27 .62 .15 D EPT H FL0W .25 TIME 11.53 15.53 10.00 10.75 12.00 TIME I IME 11.53 ADJUSTED VOLUME OF RUNDEF FOR THIS EVENT IS . DD4 49 INCHES. HLE RUNGE HYDROBEN FOR a 1 1951 AS 42UUSTED FOR BASE FLOW WITH GECTMLITHE AND FLOW IN CS IS 11510 BELOW THE THE FLOW THE FLOW IN CS IS 11510 BELOW THE ALL HILL AND FLOW IN CS IS 11510 BELOW THE ALL HILL AND THE FLOW IN CS IS 11510 BELOW ALL HILL HILL AND FLOW IN CS IS 11510 BELOW ALL HILL HILL AND THE FLOW IN CS IS ALL HILL AND THE FLOW IN CS IS 11510 BELOW IN CONTRACT IN CS IS 1150 BELOW IN CS IS 1500 BELOW IN CS IS 1150 BELOW IN CS IS 1500 BELOW IN CS IS 1150 BELO FLOW •17 •00 •014 97 . 80 FL04 TIME DEPTH HIS IS THE ACTUAL WYROGRAPH INCREMENTED TO DELIA Z AND NOVED 1.40 DIORS TO THE THE OF REST FIT. THE FLOW THE FLOW THE FLOW THE F 9.23 -000 9-43 -000 9-43 -53 10.73 17.33 -000 9-43 -000 9-43 -001 14.73 .30 T IME 10.73 THIS IS THE SWITHETIC HYDPOGRAPH WITH PELTA = .ad THE FLOW THEFLOW THEFLOW THEFLOW THEF 8.33 .000 9.131 .066 9.439 .111 n.73 12.133 .010 13.131 .077 13.93 .031 14.73 HE RATIO OF PUNCFT OF PAILWELLFOW THIS EVENT IS THE MINHUM SUM OF PRIVATIONS SOUNCED IS .administry RAIN GAGE RECMOD IS FOOM ALPINE ADEA B ESTIMATED AREAGE ANTENSIED ARIANGLI IS TITUE IS DECIMAL. THE GEFINI IS INVES. TITUE OFFINI THE OFFINI TIME DEFINI 3.33 AIX 0.50 AU 4DJUSTED PEAK FLOW IS 1.38 CFS. STORM TIME OF RISE WAS 1.74 HOURS. STORM PUNDFF LASTED 5.00 HOURS. .13 THE SHT SHT

· ALPINE CATTLE PASTURE OR MEADOWS D EPHRAIN CREEK WATERSHED - ALPINE SURFACE DRAINAGE AREA = 376. ACRES

FLOW 33.46 4.05 1.39 DEPTH .74 1.13 TIME 11.48 TIME 12.00 13.10 14.50 19.00 \*0969D INCHES. FLOW THE RUNCE HYDORGRAPH F08 7 24 1927 AS ADUNISED F08 BASE WITH DECIMAL THE AND FORMAN AS IS STATED BELOW. THE FLOW THE FORM THE FLOW THE FLOW THE FLOW THE FLOW THE FLOW THE FLOW THE STATE THE FLOW THE FLOW THE FLOW THE STATE AS THE THE FLOW THE FLOW THE STATE AS THE THE FLOW TIME DEPTH 11.38 .67 12.38 1.05 1.10 ADJINTED PEAK FLOW IS 45.46 CFS. STORP TIME OF DIST WAS 50 HOURS. STORP BUNGE 155 FO 9.65 HOURS. ADJUSTED VOLUME OF RUNNEF FOR THIS EVENT IS A MIN GAGE PECOD IS FOON ALPINE APEA B E STIPHED AVEPAGE WATTPENHED BAISWALL IS ITPE FS DECHWAL THE OPPIN IS INCHES. THE FS DECHWAL THE OPPIN IS IN THE OPPIN IS IN THE OPPIN IS IN THE OPPIN IS IN THE OPPIN THE ACTUAL PI SIHT THE 3HL 3HL

FLOW 37.46 4.88 2.50 1.06 .60 .41 .77 .11 .01 FL0W 31.47 6.20 3.16 1.20 1.20 .15 .16 .12 .44 .13 .07 .07 .01 TO DELTA T FIT T IN T HALMYERGEBAPH I HALMYERGEBAPH I TIME TO THE T TIME FLOW AND MOVED 

FL0W 28.28 8.85 1.43 TIME 11.78 12.78 13.78 .0 83 09 A = .20 TIME FLOW 11.59 25.52 12.58 11.50 13.58 2.65 TTHE FLOW TITLE WYDORAPAW WITH RELA = .20 TTHE FLOW TITLE WYDORAPAW WITH FLOW TITL 11.40 - .00 TLL.9 S.20 TL.3 14.27 TL29 2.11.49 11.40 - 2.00 TL.9 S.2.77 TL29 17.25 17.25 17.51 11.40 - 2.40 TL.1 S.2.77 TL29 17.25 - 17.51 11.40 - 4.40 TL41 1.21 - 14.23 - .00 11.40 - 4.40 TL41 1.21 - 14.23 - .00 11.40 - 4.40 TL41 1.21 - 14.23 - .00 11.40 - 4.40 TL41 1.21 - 14.23 - .00 11.40 - 4.40 TL41 1.21 - 14.23 - .00 11.40 - 4.40 TL41 1.21 - 14.23 - .00 11.40 - 6.40 TL41 1.21 - 14.23 - .00 11.40 - 6.40 TL41 1.21 - 14.23 - .00 11.40 - 6.40 TL41 1.21 - 14.23 - .00 11.40 - 6.40 TL41 1.21 - 14.23 - .00 11.40 - 6.40 TL41 1.21 - 14.23 - .00 11.40 - 6.40 TL41 1.21 - 14.23 - .00 11.40 - 6.40 TL41 1.21 - 14.23 - .00 11.40 - 6.40 TL41 1.21 - .00 11.40 - THIS IS THE SYNT TIME FLOW 10.98 7.09 11.98 77.04 12.98 7.45 17.98 .48

THE

OR MEADUWS PASTURF D EPHPAIM CREEK WATERSHED ALPINE CATTLE Subface Drainage Area = 376, Acres

DEPTH .27 FL0W 1.83 5.74 5.74 .45 .18 .18 .11 .08 .05 .05 FL04 4.83 .39 1 1.25 1 2.50 TIME 11.28 1 I 0. 80 1 0. 80 1 1. 80 1 2. 80 \*311 99 INCHES. FLOW FLOW 3.24 1.20 .01 .0395 BASE DEPTH .24 FL0W 4.62 .23 FL04 6.00 .61 .73 .13 .13 .09 .03 HE NUME TERMORFOR TO FOR STATUS STAURTED FOR MASS ATTHE NUME TERMORFORM TO FOR A TO STATUS FOR WEINE FLOW TIME FECTION 10-50 FLOW TO FOR FOR WEINE FLOW TO STATUS STATUS TO THE FORM TO STATUS TILSON 25 11-21 STATUS STATUS TO ALSO TILSON 25 112-21 STATUS STATUS STATUS TILSON 27 142-01 STATUS STATUS STATUS TO STATUS STATUS STATUS STATUS STATUS THE ADJUSTED PERFERSION STATUS STATUS THE STORM TUPE OF STATUS STATUS STATUS STATUS THE STORM TUPE OF STATUS STATUS STATUS STATUS THE STORM TUPE OF STATUS 
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 1 0.42 HULS 15 THE SWITFIE AVDD0GAJDH HITH YELTA = 22 TIME FLOW THEF EVAN THEF FLOW THE FLOW THE TIME YOU THEF EVAN THE 22.11 TOLO TIME WORLD THE 20 STATES THE 22.22 THE TIME WORLD THE STATES THE 22.22 THE TELEW WORLD FLOW STATES THE 22.22 THE TELEW YOU FOR THET FLOW THE STATES THE THE MINIMUM SUM OF DEVLISION TO MARE TO STATES THE THE MINIMUM SUM OF DEVLISION TO MARE TO STATES THE THE MINIMUM SUM OF DEVLISION TO WORLD FOR THE STATES THE THE MINIMUM SUM OF DEVLISION TO WORLD FOR THE STATES THE THE MINIMUM SUM OF DEVLISION TO WORLD FOR THE STATES THE THE MINIMUM SUM OF DEVLISION TO WORLD FOR THE STATES THE THE MINIMUM SUM OF DEVLISION TO WORLD FOR THE STATES THE THE MINIMUM SUM OF DEVLISION TO WORLD FOR THE STATES THE THE MINIMUM SUM OF DEVLISION TO WORLD FOR THE STATES THE THE MINIMUM SUM OF DEVLISION TO WORLD FOR THE STATES THE THE MINIMUM SUM OF DEVLISION TO WORLD FOR THE STATES THE THE MINIMUM SUM OF DEVLISION TO WORLD FOR THE STATES THE THE MINIMUM SUM OF DEVLISION TO WORLD FOR THE STATES THE THE MINIMUM SUM OF DEVLISION TO WORLD FOR THE STATES THE THE MINIMUM SUM OF DEVLISION TO WORLD FOR THE STATES THE THE MINIMUM SUM OF DEVLISION TO WORLD FOR THE STATES THE THE MINIMUM SUM OF DEVLISION TO WORLD FOR THE STATES THE MINIMUM SUM OF DEVLISION TO WORLD FOR THE STATES THE MINIMUM SUM OF DEVLISION TO WORLD FOR THE STATES THE MINIMUM SUM OF DEVLISION TO WORLD FOR THE STATES THE MINIMUM SUM OF DEVLISION TO WORLD FOR THE STATES THE MINIMUM SUM OF THE STATES THE MI •20 DAIN AGE RECOMD IS FORM ALPIVE MEADOWS [STIMATED AVERAGE WATCHARD AD ALVARAL 15 TIME IS DECLAMA, THE OFFM IS INDES. TIME IS DECLAMA, THE OFFM 10.01 AND 10.17 AD 10.28 ALF W W W H

D EPHPAIM CREEK WATERSHED ALPINE CATTLE PASTURE OR MEADOWS SUPFACE DRAIMAGE AREA = 376. ACRES

FL04 12.74 2.55 TIME 16.90 18.90 21.00 HE DUNCT HYDROGRAPH FOR 7 10 1953 AS ADUNSTED FOR BASE FLOW WITH RECHARL THE ADD FLOW AFTS IS ISTED BELOW THE FLOW THE FLOW THE FLOW THAT FLOW THAT IS 5.400 THAT FLOW THAT FLOW THAT IS 4.401 15.10 11.000 4.401 17.45 5.451 17.45 4.401 15.10 11.000 4.401 17.45 5.451 17.45 4.401 15.10 11.000 4.401 17.55 5.451 17.45 4.401 15.10 11.000 4.401 17.5 2.401 1.45 7.401 4.41 12.100 11.000 5.11 14.451 0.152 4.421 17.15 4.401 15.10 11.000 5.11 14.451 0.152 4.451 17.153 4.411 12.100 11.000 5.11 14.451 0.153 5.401 4.11 12.100 11.000 5.11 14.451 0.153 5.401 4.11 12.100 11.000 5.11 14.10 14.15 4.451 0.153 4.115 11.000 5.11 14.10 14.15 4.451 0.153 4.115 11.000 5.411 14.15 4.451 0.153 4.115 11.000 5.411 14.15 4.451 0.153 4.115 11.000 5.411 14.15 4.451 0.153 4.115 11.000 5.411 14.15 4.451 14.15 1

.60 THE

TIME DEPTH 1145 DEPTH 17.83 .60 AT THE ACCOUNT STORM APPLIE CATTLE PAST ESTIMATED AVERAGE MATERVALD RAINFALL IS 460 TIME TY OFCHMAL, HE DEPH IS 14CHF5, 460 TIME REPH AT THE DEPH IS 14CHF5, 147 15.25 400 16.75 35 17.17 45 17.81

Uth \* THIS IS THE ACTUAL HYDROGRAPH INCREMENTED TO DELTA =

	FLO	8.3		~				
	TIME	17.85	19.85	21.85	23.85	25.85	27.85	
	FLOW	8.41	1.20	. 35	18	.07	.01	
FIT.	TIMF	37.45	19.45	21.45	23.45	25.45	27.45	
REST	FLOW	.00	1.81	• 39	. 21	66.	- 02	
TIME OF	TIME	17.05	19.05	21.05	23.05	25.05	27.05	
TO THE	FLOW	00.	7. 44	. 48	hc .	. 10	+0.	
HOUR S	TIME	16.65	12.65	20.65	22.55	24.65	26.65	
1.20	FLOW	• 00	4.94	.65	.28	.13	. 05	- 00
ND MOVED	TIME	16.25	18.25	20.25	22.25	24.25	26.25	28.25

JNM-SOD

FL04 6.22 117.95 19.85 .0 57 62 FLOW 5.53 1.60 THIS IS THE SVNTHFILC HYDROGGARPH WITH RELTA = .40 TIME FLOW TIME FLOW TIME FLOW TIME 16.25 .00 16.65 1.27 17.05 3.19 17.45 16.26 4.87 19.05 3.23 19.45 THE FLOW THM FLOW THM FLOW THM 16.25 -000 16.65 1.27 17.05 3.19 1.06 17.25 .400 16.65 1.27 17.05 3.23 19.45 20.25 .28 20.65 .00 19.05 3.23 19.46 20.25 .28 20.65 .00 145 £EM 15 16.84 10.94 05 FLOM 20.0475 D 15 30.643 15 16.84 10.94 09 05 FLOM 217015 30.0454 20 15 16.84 10.94 09 05 FLOM 217015 30.0475 01 25 10.84 10.94 05 FLOM 217015 30.0475 01 25 10.84 10.94 10.95 10.94 10.95 10.

30.63477

ALPINE CATTLE PASTURE OR MEADOWS SURFACE DPAINAGE APEA = 376. ACRFS EPHRAIM CREEK WATEPSHED c

FL0W 20.76 3.42 .14 DEPTH 41 14.58 TIME 12.10 12.75 14.50 15.00 it? ia.00 00 15.00 it? ia.00 00 the Stowittpe Enk (tou 15: 39.7 CS.) the Stowittpe Enk (tou 15: 39.4004S.) the Stowittpe Enk (15:105, 6.40140S.) the Stowittpe Stowies F ob rids (Stori 15: 0.9063 INCHS.) FLOW . 20 HIR PUNCE HENDORDAND FOR 7 28 1951 A5 ADJUSTED FOR MASE WITH RECENTL THRE AND FLOW HE FOR TASE 1515 EBELOW. THRE FLOW THRE FLO TIME DEPTH 13.75 .40 n L # \* THE PAIN GLGE RECOPD IS FROM ALPINE CATTLE PAST THE ESTIMATED NUEVAGE NATESHED SALMALLIS .44 THE THVE TS DECTMAL. THE OPENH IS INCHES. THE OPENH THME OPENH IS INCHES. THE OPENH THME OPENH INE OPENH THM 11.50 -000 12.00 28 12.75 .39 13.7 15.50 -06 16.00 -07

FL04 23.36 1.64 .12 .08 .08 TIME 12.30 13.30 14.30 14.30 15.30 15.30 15.30 18.30 18.30 FLOW 3.07 3.07 .39 .13 .09 .05 H INCPEMENTED TO DELTA = E TIME OF SET FIL TIME FLOW TIME 11.93 00 12.10 11.93 4.91 13.10 13.93 4.91 15.10 15.93 4.01 15.10 15.93 4.01 15.10 15.93 4.02 10.10 HYDROCRAPH WITH 
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FLOW 10.59 3.19 .36 1.25 .89 .89 TIME 112.30 113.30 114.30 114.30 115.30 115.30 117.30 FL0W 8.61 5.23 .40 .94 1.19 1.19 DELTA FLOW 4.59 7.58 7.58 .67 .64 1.43 1.43 TIME 11.93 12.93 13.93 14.93 15.93 FL04 1.53 9.31 1.17 1.17 1.57 1.57 TIME 11.70 12.70 13.70 13.70 13.70 15.70 IIS IS THE SYNTHFIE THE FLOW TIM II.FR FLOW TIM II.FR FLOW TIM II.FR 101.52 II.FR 101.52 II.FR 11.57 I .58 .00

.0 86 45 THE MINIMUM SUM OF DEVIATIONS SQUARED IS 947.19478 THE RATIO OF RUNDFF TO RAINFALL FOR THIS EVENT IS

n EPHRATM CREEK WATERSHED ALPINE CATTLE PASTURE OR MEADOWS Surface Drainage area = 376, acres

THE RUNDEF HYDROGRAPH FOR 8 31 1953 AS ADJUSTED FOR BASE FLOW WITH DECEMAL TIME AND STATE 2000 2000

FLOW 3.655 0.355 0.355 1.499 1.499 1.09 1.09 1.09 1.09 1.09
TIME 19.55 1 20.00 1 21.00 2 22.50 25.75 26.75 33.00
FLOW 110.33 10.33 8.50 .41 .41 .31 .17
8ELOW. 19.40 19.40 19.40 20.78 22.09 24.25 25.50 26.50 30.00 30.00
LISTED FLOW 12.955 10.57 .57 .57 .57 .55 .55 .55 .55 .55
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LOW IN FLOW 13.33 13.33 13.33 13.33 1.41 1.41 1.41
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Appendix D Graphical comparisons of hydrographs

























## VITA

### Clive H. Walker

### Candidate for the Degree of

# Master of Science

Thesis: Estimating the Rainfall-Runoff Characteristics of Selected Small Utah Watersheds

Major Field: Water Resources Engineering

Biographical Information:

- Personal Data: Born at Nampa, Idaho, November 28, 1935, son of Clive Stevenson and Ardeth Hansen Walker; lived on a small farm until 1953; married Margaret Vay Broadbent, June 8, 1956; three children--Annette, Vida Maria and Clive Steven.
- Education: Attended elementary school near Meridian, Idaho; attended Central Junior High School, Nampa, Idaho; graduated from Nampa High School in 1953; honors include attendance at Idaho Boys State; received a Bachelor of Science degree from Utah State University in Civil Engineering, with academic background in history, education, language and biological science, in 1960; completed the requirements for a Master of Science degree in Water Resources Engineering, 1970.
- Professional Experience: 1962-present, employed by the Soil Conservation Service as a civil or hydraulic engineer; presently employed as hydraulic engineer on the watershed and river basin planning staff at Casper, Wyoming; previous experiences include a training position in hydraulic engineering at Portland, Oregon, work unit engineer at Logan, Utah, and Richfield, Utah; temporary assignments have included flood damage reporting in Oregon and watershed project construction inspection in Arizona and Utah; 1962-1966, served as supply officer, executive officer, and company commander of various Utah National Guard and U. S. Army Reserve artillery and engineering units in Utah; 1960-1962, served as engineer combat platoon leader, supply officer, personnel officer and Assistant adjutant in the U. S. Army Corps of Engineer units at Fort Lewis, Washington; 1960, research engineering technician, Agricultural Research Service, Brawley, California; 1957-1960, civil engineering trainee, Agricultural Research Service, Logan, Utah.