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Quaternary Evolution of the Colorado River at Lees Ferry, Arizona

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QUATERNARY EVOLUTION OF THE COLORADO RIVER AT LEES FERRY,

ARIZONA

by

William Scott Cragun

A thesis submitted in partial fulfillment of the requirements for the degree

of

MASTER OF SCIENCE

in

Geology

Approved:

UTAH STATE UNIVERSITY Logan, Utah

2007

ABSTRACT

Quaternary Evolution of the Colorado River at Lees Ferry, Arizona

by

William Scott Cragun, Master of Science

Utah State University, 2007

Major Professor: Dr. Joel L. Pederson Department: Geology

A well-exposed suite of Colorado River fill terraces preserved at Lees Ferry records the oscillating history of this major river superimposed on its overall downcutting of the Colorado Plateau. Detailed mapping, sedimentology, cross-sectional surveys, and the use of two geochronometers have been used in order to establish a detailed chronostratigraphy for the area. Eight distinct deposits have been identified along the Colorado River (Ml -M7, and S3), and four deposits have been identified along the Paria River $(P1-P4)$.

Geochronology of six of these deposits using optically stimulated luminescence and cosmogenic ¹⁰Be exposure techniques indicates a long-term average bedrock incision rate of 290 to 470 m/my. These incision rates are approximately two to three times higher than others reported in Grand Canyon and the upper Colorado River basin, but are similar to the recently reported high incision rates near Glen Canyon and along the Fremont River. These results suggest that there is a region of faster incision along the Colorado River in the central Colorado Plateau in the vicinity of Lees Ferry and Glen

Canyon. This apparent increase in central plateau Pleistocene incision rates may be caused by either epeirogenic uplift due to tectonics and erosional isostatic rebound, or transient waves of incision in response to original drainage integration.

In addition to recording the incision history of the Colorado River, the wellpreserved Pleistocene fluvial terraces provide evidence regarding the timing and processes of terrace formation at Lees Ferry. Chronostratigraphic analysis indicates that aggradation was occurring at \sim 20 ka (M2), \sim 70 to 40 ka (M3), \sim 115 to 90 ka (M4), and \sim 130 ka (M5). Aggradation and incision along the Paria River appears to be occurring at the same time as that on the Colorado River. Deposits at Lees Ferry are generally younger than correlative deposits in headwater catchments and in eastern Grand Canyon. In addition, the most prominent deposit in the Lees Ferry area (M4) correlates to MIS stage 5b-c , a time in which no glaciations have been reported in headwater drainages. Data from this study indicate that fluvial responses at Lees Ferry are a complicated integration of signals from climate change in headwater catchments and sediment production from local hillslopes and tributaries .

(l 92 pages)

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I am indebted to the Glen Canyon National Recreational Area for permission to access the study area and to Richard Hereford for his detailed explanations of continuing research efforts along the Colorado and Paria rivers. I would also like to thank my field assistants, Ben DeJong, Rob Mackley, Kevin Hadder, Alan Hidy, Kelly Mitchell, Gary O'Brien, and my little brother Ben, for sharing the quiet magic of the Lees Ferry landscape. This project could not have been completed without financial support from the NSF grant EAR-0346054 awarded to Joel Pederson, the Four Corners Geological Society, the Rocky Mountain Section of the Society for Sedimentary Geology, and the Utah State University Geology department. Finally, I would like to thank my wife, Yvette, and my five little boys, Jonah, Enoch, Asher, Samuel, and Hiram, for their love and support through my endless educational endeavors. Thank you for standing patiently by my side and giving me the motivation that I need to always be better.

W. Scott Cragun

 \mathbf{V}

PREFACE

This research explores the geochronology and sedimentology of the wellpreserved Colorado and Paria River terraces at Lees Ferry, Arizona. This thesis is organized into four chapters . Chapter 1 introduces the research problems and outlines the methods used to address these questions. Chapter 2 is a short manuscript written for journal submission, which discusses incision rates calculated at Lees Ferry and the differential incision of the central Colorado Plateau. Chapter 3 is a longer manuscript written for journal submission. This chapter addresses the timing of aggradation and degradation episodes at Lees Ferry in relation to local and distant climatic forcing. Chapter 4 reviews the results and conclusions discussed in Chapters 2 and 3. The appendices contain the complete data collected throughout the course of this research endeavor.

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CHAPTER l

INTRODUCTION

The timing and cause for incision of the Colorado Plateau by the Colorado River and its tributaries have been debated since the late l 800's when geologists first began to explore the deep canyons and excellent exposures of the region. Research through the past century has continued to shed light on the complex history of the Colorado River through the Colorado Plateau. Incision is generally thought to have begun \sim 6 Ma (Lucchitta, 1966) and been nearly complete by ~ 0.4 to 1.2 Ma (Hamblin, 1994; Fenton et al., 2004). However, new stratigraphic and chronologic research on Pleistocene river deposits throughout the Colorado Plateau indicates that significant incision has continued through the Quaternary (e.g. Lucchitta et al., 2000; Pederson et al., 2002).

The few river gravels and terraces that record the incision history of the Colorado River are generally preserved in locations where erosion of softer strata or structural controls have created wide valleys along the river corridor. Recent studies from sites along the Colorado River in western Grand Canyon, eastern Grand Canyon, Glen Canyon, Westwater Canyon, Glenwood Canyon, and along the tributary San Juan and Fremont Rivers have reported variable bedrock incision rates that indicate differential incision along the Colorado River drainage through the Quaternary. Lees Ferry also contains a well-exposed suite of Colorado River terraces that lie in the center of this region, and while the Holocene history has been explored in depth by Hereford et al. (2000), the Pleistocene terraces have remained relatively unstudied.

This study presents an analysis of the fluvial terraces at Lees Ferry in order to provide insight into two important questions: I) are incision rates at Lees Ferry consistent

with others in the region; and 2) is the timing of aggradation and incision at this central location consistent with climate changes in glaciated headwaters, or rather with changes in the non-glaciated landscapes of the immediate area? These questions have been addressed through various field observations, including detailed sedimentologic descriptions, clast counts, sand petrology , geologic mapping of Quaternary deposits, and topographic cross-sectional surveys of terrace heights and valley profiles. In addition, two dating methods have been employed in order to quantify the timing of aggradation and subsequent terrace abandonment: optically stimulated luminescence (OSL) and 10 Be terrestrial cosmogenic nuclide (TCN) dating. Results of this study have refined the regional picture of varying incision rates along the length of the Colorado and contributed to the debate on the timing of aggradation-degradation cycles in relation to local and distant climate forcing.

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CHAPTER 2

PLEISTOCENE GEOCHRONOLOGY AND INCISION RATES OF THE COLORADO RIVER AT LEES FERRY: TOWARDS SOLVING THE MYSTERY OF DIFFERENTJAL INCISION ON THE COLORADO PLATEAU'

ABSTRACT

Chronostratigraphic analysis of the well-preserved Pleistocene fluvial terraces along the Colorado River at Lees Ferry indicate a long-term average bedrock incision rate of 290 to 470 m/my. These incision rates are approximately two to three times higher than others reported downstream in Grand Canyon and upstream in the upper Colorado River basin. In contrast, our incision rates at Lees Ferry are similar to the recently reported high incision rates near Glen Canyon and along the Fremont River. Our results suggest that there is a region of faster incision along the Colorado River in the central Colorado Plateau in the vicinity of Lees Ferry and Glen Canyon. We propose two possible mechanisms for this increase in central plateau incision rates: I) epeirogenic uplift due to tectonics and erosional isostatic rebound; or 2) transient waves of incision in response to original drainage integration.

INTRODUCTION

The Colorado River flows through the heart of the Colorado Plateau and has carved a landscape of steep canyons that are commonly several hundred meters deep. Debate about the timing and causes of this large-scale incision has existed since the late

¹ Coauthored by W. Scott Cragun, Joel L. Pederson, and Tammy M. Rittenour

l 800's when early geologists such as John Wesley Powell (1875), Clarence Dutton (1882), and William Morris Davis (1901) first explored the deep canyons and excellent rock exposures of the region. Additional research during the past century has continued to shed light on the complex history of the Colorado River through the Colorado Plateau (e.g. Hunt, 1969; McKee and McKee, 1972) ; however , the timing and driving forces of late-Cenozoic incision are still debated.

Incision of the Colorado Plateau region by the Colorado River and its drainages is thought to have occurred as the result of a large-scale reversal in drainage direction that occurred over the late-Cenozoic (Lucchitta, 1972; Pederson et al. , 2002a). Paleodrainage networks flowed approximately northwest off the Laramide highlands of central Arizona and were disrupted by Basin and Range extension and the opening of the Gulf of Mexico in Miocene time (e.g. Lucchitta, 1972; Young and McKee, 1978). The ancestral Colorado River was captured as developing drainages became integrated throughout the lower Colorado River region (Lucchitta, 1972), ultimately disrupting flow direction and driving incision throughout the plateau by lowering baselevel. Analysis of sediments deposited in the Grand Wash Trough, where the Colorado River exits the Grand Canyon, indicates that the Colorado River had established its present-day course and begun to incise the Grand Canyon \sim 5.6 Ma (Lucchitta, 1966; Faulds et al., 2001). Newer stratigraphic and chronologic research on Quaternary river deposits throughout the Colorado Plateau indicates that significant incision has continued through the Quaternary (e.g. Lucchitta et al., 2000; Pederson et al., 2002b).

In general, incision is driven by baselevel change through tectonics, eustacy, or drainage integration and may result in the formation of fluvial terraces within a river

valley (e.g. Merritts et al., 1994). Glacial-interglacial climate cycles also influence the formation of terraces; however, climatic controls generally occur over shorter time scales and are generally superimposed on the long-term downcutting of a region (e.g. Hancock and Anderson, 2002). Analysis of fluvial terrace geometries and sedimentology of associated deposits, coupled with a geochronologic framework, provides valuable information for reconstructing the timing and causes of incision through the past. Here we present new geochronologic data and rigorous incision rates from the well-preserved fluvial terrace record of the Colorado River at Lees Ferry in order to explore the patterns and causes of late-Cenozoic incision of the Colorado Plateau.

BACKGROUND

With the advent of new geochronologic tools, recent research has focused on quantifying Quaternary incision rates at several localities along the length of the Colorado River and its tributaries (Figures 2.1 and 2.2). Although these studies have provided valuable results in specific locations along the profile of the Colorado River , a complete picture of regional incision throughout the Colorado Plateau has yet to be established. We review several of these studies in order to place our new results from Lees Ferry in the context of incision reported throughout the Colorado Plateau and provide a framework for interpreting the potential causes of differential incision throughout the region.

Figure 2.1 (A) Map of the central Colorado River and its drainages, showing locations of studies that have reported incision rates throughout the Colorado Plateau. (8) Map showing the distribution of Quaternary deposits at Lees Ferry.

Figure 2.2. Longitudinal profile of the Colorado River from Kremling, CO to the Mexican border, including profiles of the San Juan and Fremont Rivers. Vertical exaggeration \sim 300. Incision rates from studies throughout the Colorado Plateau are plotted along the profiles and along the bottom of the figure in order to highlight the relatively high incision rates in the Lees Ferry and Glen Canyon reach. The San Juan incision rate shown with grey dot in order to highlight its relatively far distance from the trunk Colorado River and to indicate that it is a minimum rate.

Incision rates reported in the Granite Park area in western Grand Canyon over the past 300 to 600 ky are \sim 70 to 90 m/my (Lucchitta et al., 2000; Pederson et al., 2002b) . In contrast, incision rates in the Furnace Flats area in eastern Grand Canyon , upstream of the Hurricane-Toroweap fault zone, are consistently \sim 140 m/my during approximately the same time period (Pederson et al., 2002b; Pederson et al., 2006).

Incision rates reported upstream in the tributary Fremont River over the past \sim 200 ky are 430 m/my (Marchetti and Cerling, 2005) . Similar , yet slightly higher, incision rates of 400 to 700 m/my have been reported in the Glen Canyon region over the past \sim 500 ky (Garvin et al., 2005). In contrast to these relatively high rates within the greater Glen Canyon reach, Wolkowinsky and Granger (2004) report an incision rate along the tributary San Juan River of 110 m/my over the past 1.36 My. The different timescale used in this last calculation may be important in the interpretation of these data (Gardner et al., 1987). Adjustment using the scaling methods of Gardner et al. (1987) to the significantly longer-term San Juan River record (by an order of magnitude) yields an incision rate of \sim 170 m/my, which is more comparable to those reported downstream in Grand Canyon and in the upper Colorado River basin than to those in the Glen Canyon area.

Incision rates calculated in the upper Colorado River basin in Westwater Canyon using the Lava Creek B tephra (-600 ky) are -180 m/my (Willis and Biek, 2001), significantly lower than rates reported from the Glen Canyon region. Similar incision rates $(\sim 150 \text{ m/my})$ have been reported from the upper Colorado River basin along the Green River and its major tributaries also using the Lava Creek B tephra (e.g. Reheis et al., 1991). Additional incision rates reported near the headwaters of the Colorado River

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in Glenwood Canyon over the past \sim 1.5 My are significantly higher (\sim 240 m/my) than others in the upper Colorado River basin (Bryant et al., 2002).

The incision rates reported throughout the Colorado Plateau indicate differential incision along the Colorado River and its drainages through the Quaternary , especially in the Glen Canyon region (see bottom curve in Figure 2.2). Lees Ferry lies in the heart of this region at an inflection in long-profile gradient and contains a well-exposed suite of Pleistocene deposits and terraces (Hereford et al., 2000). Analysis of this fluvial record at Lees Ferry provides an independent datum between Glen and Grand canyons wherein the regionally variable incision rates from these sites can be investigated .

Lees Ferry is situated in a wide valley that marks the end of Glen Canyon and the start of Marble Canyon (Figure 2.1). Paleozoic and Mesozoic sedimentary rocks dominate the bedrock geology of the area and are well-exposed from river level to the surrounding escarpment of the Vermillion Cliffs (Phoenix, 1963). Differential erosion of the easily eroded Moenkopi Formation at river level has created a particularly wide valley wherein the preservation potential of fluvial deposits has been increased. The Paria River, a major tributary to the Colorado River, enters at this relatively open spot in the landscape. The Lees Ferry area is generally semiarid, although climate varies from the arid low-lying areas including terrace surfaces , to the subhumid higher elevations of surrounding high plateaus.

METHODS

We mapped Quaternary deposits in the Lees Ferry area in order to identify the sedimentary characteristics of fluvial deposits and correlate terrace levels. In addition,

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. cross-sectional valley profiles have been surveyed using total station equipment in order to measure terrace heights. These field observations and measurements have been combined with two geochronologic tools, optically stimulated luminescence (OSL) and terrestrial cosmogenic nuclide 10 Be exposure (TCN) dating, in order to reconstruct the aggradation /incision history of the Colorado River and calculate incision rates at Lees Ferry. Although both dating methods have been used on tluvial deposits, ages calculated from OSL and TCN require different interpretations; OSL ages indicate the timing of aggradation, whereas TCN ages represent the timing of renewed incision.

Ten OSL sand samples were collected in aluminum tubes (5 cm x 20 cm) from recently exposed road or stream cuts. OSL sample preparation and analysis was performed at the University of Nebraska Luminescence Laboratory using the singlealiquot regenerative method on a RISO TL/OSL-DA-15B/C reader with blue-green light stimulation (470 nm, Hoya U340 filter). Additionally, four TCN surface samples were collected from terrace treads that exhibited long-term stability in the form of wellestablished desert pavements. Surface samples composed of \sim 100-200 quartzite pebbles (1-2 cm in diameter) were crushed and analyzed in a single amalgamated sample after the methods of Repka et al. (1997). A 220 cm depth-profile consisting of six quartz-rich sand samples (collected every 30 to 40 cm) was used to account for pre-depositional inheritance. This single good profile from the most prominent deposit in the area was used for all surface ages, with the assumption that the Colorado River has deposited sediment with a consistent amount of inheritance during each episode of aggradation in its cyclic history. TCN samples were prepared at Dalhousie University and accelerated mass spectrometer analysis was performed at Lawrence Livermore National Laboratory.

RESULTS

Seven distinct deposits (M1-M7) and ten terrace levels (M4y, M5m, and M5y are erosional fill-cut terraces) have been identified along the Colorado River in the Lees Ferry area (Figure 2.3). Deposits generally range from 10 to 30 m in thickness, and the most prominent of these deposits (M4) exhibits an irregular basal contact wherein a "false" strath is significantly higher than the true basal strath. Terrace levels range in height from 5.4 m (M1) to \sim 180 m (M7) above the reference river stage (380 m³/s), and the treads of the M2-M5 terraces exhibit moderately- to well-developed desert pavements. Pavement surfaces have not formed on the M1 terrace due to recent or active deposition, whereas the M6 and M7 terraces are exhumed remnants and do not have preserved planar surfaces or well developed pavements . The deposits at Lees Ferry display three distinct sedimentary facies interpreted as: I) a mixture of far-traveled (quartzite, volcanic porphyry) and local clasts (sandstone, chert, limestone) deposited in mainstem fluvial environments ; 2) more massive debris flow beds dominated by local clast lithologies ; and 3) overbank Colorado River sand.

Integrating the positions and ages of geochronologic samples (Table 2.1; Appendix D), we are able to reconstruct the fluvial history of the Colorado River at Lees Ferry and calculate a long-term bedrock incision rate. Maximum and minimum incision rates are reported due to the complex nature of aggradation and incision episodes represented by the fill terraces at Lees Ferry (Pederson et al., 2006). The minimum rate was calculated using elevation data and geochronology results from samples collected near the middle of the deposit (4-10 m above respective straths). The maximum rate was calculated using elevation data and geochronology results from samples near the terrace treads. lncision rates were calculated through linear regression of height vs. age plots of samples collected at similar positions within fluvial cycles (Figure 2.4). These are long-term average incision rates $(>100 \text{ ky})$ that integrate pulses of incision, periods of stability, and aggradation episodes and should be comparable to other carefully calculated long-term rates reported throughout the Colorado Plateau.

DISCUSSION

Middle-late Pleistocene incision rates of 290 to 470 m/my at Lees Ferry are approximately two to three times higher than rates reported downstream along the Colorado River in eastern and western Grand Canyon (Pederson et al. , 2002b; Lucchitta ct al., 2000) (Figure 2.2) . Similarly , incision rates at Lees Ferry are significantly higher than rates reported upstream in Westwater Canyon $\left(\frac{180 \text{ m/my}}{\text{y}}\right)$; Willis and Biek, 2001), in the greater Green River basin $\left(\frac{150 \text{ m/my}}{\text{m}}\right)$; Reheis et al., 1991), and along the San Juan River $(\sim]10 \text{ m/m}$ y; Wolkowinsky and Granger, 2005). In contrast, our range is similar to, yet slightly lower, than those reported just upstream in the greater Glen Canyon reach (Garvin et al., 2005), including the tributary Fremont River (Marchetti and Cerling, 2005). Incision rates of \sim 240 m/my near the headwaters in Glenwood Canyon reported by Bryant et al. (2002) are also slightly lower than those that we report from Lees Ferry.

Figure 2.3. Schematic cross-sectional valley profile of the fill terraces at Lees Ferry, showing heights, geometries, and ages of deposits. Sec Table 2.1 for error calculations on individual ages.

Figure 2.4. Curve representing the height of the Colorado River channel bed through time . Stippled pattern represents fill deposits, whereas the gray represents bedrock. Alluvial aggradation cycles are superimposed on overall downcutting at Lees Ferry. Solid bold lines are regressions approximating minimum and maximum incision rates at Lees Ferry. See Table 2.1 for error calculations on individual ages.

Data Point Description	Height (m) ¹	Age $(ka)^2$	Dating Method
Colorado River			
Holocene flood deposits (M1)	$0 - 6$		Various
Surveyed M2 terrace tread	14.4		
M ₂ sand lens in fill	10.7	21 ± 1	OSL^5
M3y sand lens in fill	12.8	39 ± 3	OSL^5
Surveyed M3 strath	15.3		
Surveyed M3 terrace tread	25.6		
M3 desert pavement	25.6	38 ± 3	TCN ⁴
M3 sand lens in fill	19.2	72 ± 5	OSL^5
M3 sand lens in fill	22.2	71 ± 5	OSL^5
S3 sandy unit in fill	51.9	70 ± 6	OSL^5
S3 sandy unit in fill	60.2	40 ± 3	OSL^5
Surveyed M4 strath (lower)	15.8		
Surveyed M4 strath (upper)	30.8		
Surveyed M4y terrace tread	41.8		
M4y desert pavement	41.8	87 ± 7	TCN ⁴
Surveyed M4o terrace tread	45.8		$\overline{}$
M4o desert pavement	45.8	86 ± 7	TCN ⁴
M4 sand lens in fill	34.9	97 ± 8	OSL^5
M4 sand lens in fill	18.3	114 ± 8	OSL^5
Surveyed M5 strath (lower)	42.5		
Surveyed M5 strath (upper)	60.5		
M5y sand lens in fill (lower)	45.0	121 ± 12	OSL^5
M5 sand lens in fill (upper)	62.5	138 ± 10	OSL^5
Surveyed M5y terrace tread	64.3		
Surveyed M5m terrace tread	66.7		
M5m desert pavement	66.7	129 ± 10	TCN ⁴
Surveyed M5o terrace tread	78.7	$- -$	
Referenced to a local river stage of 380 m^3/s (13,000 cfs) 22δ errors reported for all ages ⁵ Fine-grained Holocene overbank sediment of the Colorado River (Hereford et al., 2000) ⁴ Terrestrial cosmogenic-nuclide date of surface sample corrected for inheritance with depth-profile, minimum surface age (see			

Table 2.1. Survey and geochronologic data for Colorado River deposits.

Appendix E for complete data)

⁵ OSL samples represent ages of deposition (see Appendix F for complete data)

Our data confirm that Quaternary incision of the central Colorado Plateau is relatively fast and indicate that a reach of faster incision may exist in the greater Lees Ferry and Glen Canyon region (Figure 2.2). Variable incision rates downstream of Lees Ferry in eastern and western Grand Canyon have been well explained by localized movement along active Quaternary faults (Pederson et al., 2002b); however, the cause of differential incision in the upper Colorado River and its drainages remains unresolved. We propose that differential incision of the Colorado Plateau, especially in the Lees Ferry and G len Canyon region, may be caused by: 1) epeirogenic uplift enhanced by erosion exhumation of the central plateau; or 2) waves of incision (knickzones) passing through the region.

Data from the southern edges of the Colorado Plateau suggest that epeirogenic uplift rates of the region since \sim 5 Ma are \sim 220 m/my (Sahagian et al., 2002). Although these rates are broadly consistent with incision rates reported throughout the Colorado Plateau, epeirogenic uplift could only result in the regional pattern we report if a localized portion of the central plateau were uplifting faster than surrounding areas. This study will be the first to test this hypothesis in the greater Glen Canyon region .

One established positive feedback of epeirogenic rock uplift reported on the Colorado Plateau is isostatic rebound driven by erosional exhumation (Pederson et al., 2002a) . Although epeirogenic uplift of a central portion of the Colorado Plateau remains untested, higher erosional exhumation in the Henry Mountain and Glen Canyon region could produce a localized isostatic response at a wavelength of \sim 400 km (Roy et al., 2005). Significantly higher incision rates within the potentially affected region would be expected, particularly in the Lees Ferry and Glen Canyon areas. In contrast, the

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apparently low incision rates reported near this region from Westwater Canyon \sim 180 m/my) and the San Juan River (\sim 110 m/my) argue against this hypothesis. However, these latter incision rates have been calculated from areas that are located on the flanks of this region of high total exhumation (Pederson et al., 2002a) and would be expected to have slightly lower incision rates.

Alternatively, the higher incision rates along the Lees Ferry and Glen Canyon reach may indicate that a transient knickzone has passed through the region. Jn tenns of faulting that might drive such a knick point, the Hurricane and Toroweap faults in western Grand Canyon are the only active faults downstream of Lees Ferry. However , the downto-the-west motion on these normal faults primarily results in hangingwall subsidence, not footwall uplift, and can only be geometrically responsible for dampening some downstream incision rather than driving upstream downcutting (Pederson et al., 2002b).

Another possible source of baselevel fall that would result in a transient knickzone is the drainage integration of the Colorado River off the Colorado Plateau . This resulted in a geologically instantaneous >1000 m baselevel drop at \sim 6 Ma that must have been transferred upstream through the Grand Canyon and into the upper Colorado River basin (Pederson et al., 2002b). The question remains: would this ancient drainage capture event still have a transient signal in the central plateau region today? Bryant et al. (2002) report incision rates that increase from the late Miocene and early Pliocene (24 m/my) to the Pleistocene (241 m/my) in Glenwood Canyon (see also Larson et al., 1975). This may indicate the upper portion of the Colorado River basin had already felt at least a portion of this baselevel fall by \sim 1.5 Ma. However, the Colorado River and its tributaries could reasonably still be responding to the original drainage integration today and a

partial signal of that initial event may still have a transient signal in the region. In particular, hard bedrock may act to delay or hold up a knickzone as it passes through a system and diffuses (Gardner, 1983; Crosby and Whipple, in press), slowing incision and creating locally steep reaches. In contrast, weak bedrock may locally enable incision and result in low gradient reaches (e.g. Stock et al., 2005; Mackley , 2005). The steeper gradient but relatively low incision rates in Grand Canyon may indicate that part of the signal from the original drainage capture is held up by the harder bedrock of the area . In contrast, the lower gradient, soft bedrock, and higher incision rates reported from Lees Ferry up through Glen Canyon is consistent with a signal that may have passed through the area quickly.

In conclusion, as our knowledge of the evolution of the Colorado River system improves, a complex picture of river incision is beginning to emerge. On the scale of the entire drainage basin, incision of the Colorado Plateau is driven by the baselevel fall that occurred at the onset of drainage integration 5-6 Ma. However, a zone of higher incision rates in the greater Lees Ferry and Glen Canyon reach has become evident with the addition of this research. These faster mid-late Pleistocene incision rates of the central Colorado Plateau may be caused by: 1) localized epeirogenic uplift due to tectonics and isostatic rebound; or 2) transient knickzones resulting from drainage integration \sim 6 Ma, which are moderated by variations in bedrock resistance.

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CHAPTER 3

PATTERNS OF FLUYIAL AGGRADATION AND DEGRADATION RELATED TO CLIMATE CHANGE ALONG THE COLORADO RIVER AT LEES FERRY, ARIZONA²

ABSTRACT

A well-exposed suite of Colorado River fill terraces preserved at Lees Ferry records the oscillating history of this major river superimposed on its overall downcutting of the Colorado Plateau. Detailed mapping , sedimentology, and the use of two geochronometers has been undertaken in order to establish a detailed chronostratigraphy for the area. Seven distinct deposits have been identified along the Colorado River, and four deposits have been identified along the Paria River. Geochronology of six of these deposits using optically stimulated luminescence and cosmogenic ¹⁰Be exposure dating techniques indicates that river aggradation occurred at \sim 20 ka (M2), \sim 70 to 40 ka (M3), \sim 115 to 90 ka (M4), and \sim 130 ka (M5). Aggradation and incision along the Paria River appears to have occurred in concert with that on the Colorado River, at least at its mouth where baselevel affects may be a strong control. Sedimentologic and stratigraphic examination of the deposits suggests that far-traveled sediment has been mixed with locally derived fluvial and debris-flow facies, with the proportion of these not detectably changing through time. Deposits at Lees Ferry are generally younger than correlative deposits in headwater catchments and in eastern Grand Canyon. In addition, deposition of the M4 occurred \sim 115 to 90 ka during a time in which no glaciations have been

² Coauthored by W. Scott Cragun, Joel L. Pederson, and Tammy M. Rittenour

reported in headwater drainages. Our data indicate that fluvial responses at Lees Ferry are a complicated integration of signals from climate change in headwater catchments and sediment production from local hillslopes and tributaries.

INTRODUCTION

The influence of climate change on fluvial processes and the evolution of river systems has been debated since the early 1900's when Penck and Brückner (1909) recognized the connection between glacial moraines in the Alps and fluvial deposits in the Alpine foreland. Research on this subject has been revitalized in the past several decades with the advent of innovative dating techniques such as cosmogenic exposure and optically stimulated luminescence (e.g. Repka et al., 1997; Aitken, 1998; Forman et al., 2000; Zreda and Phillips , 2000). Detailed fluvial chronologies have been reported from headwater rivers within or near glaciated terrain that suggest the timing of aggradation and degradation cycles closely matches glacial-interglacial oscillations observed throughout the Quaternary (e.g. Chadwick et al., 1997 ; Pan et al., 2003). Deposition and incision along large continental rivers also occurs in response to glacialinterglacial scale climate change (e.g. Blum and Tornqvist, 2000 ; Straffin et al., 2000; Tornqvist et al., 2000; Anders et al., 2005); however , our understanding of how regional rivers integrate climate signals from the diverse bio-climatic and geomorphic terrains within their drainage basins remains incomplete.

In general, local patterns of aggradation and degradation along the length of large fluvial systems are determined by climate change, sea-level fluctuation, and baselevel change in response to drainage integration or tectonic movement (e.g. Blum and

Tomqvist, 2000; Wallinga et al., 2004). Climate change resulting in the advance and retreat of glaciers creates signals that migrate downstream from headwater sources, whereas baselevel change in response to fluctuating sea-level generally affects only the lower portion of fluvial systems (Blum and Törnqvist, 2000). Drainage integration and tectonic movement may also affect the responses of large fluvial systems within specific reaches where these processes are active. In addition, changes in hillslope sediment production and hydrology as a result of local climate change may create unique responses within specific reaches and will become integrated with both upstream and downstreamsourced signals. Sorting out the interplay between these spatially and temporally variable factors is not simple. However, analysis of terrace geometries and sedimentology, coupled with a geochronologic framework, provides a valuable first step in deciphering the various controls on fluvial responses within these complex systems.

The Colorado River has a drainage area of $\sim 640,000 \text{ km}^2$ and passes through seven U.S. states before reaching the Pacific Ocean in the Gulf of Mexico. The \sim 2300 km course of this continental-scale river begins in the alpine headwaters of the Rocky Mountains and traverses the semiarid and arid regions of the Colorado Plateau and Great Basin. Lees Ferry lies in the center of the Colorado River drainage in an open valley between the generally steep landscapes of Glen and Grand canyons (Figure 3.1). The Lees Ferry area features a well-exposed suite of fill terraces that record an integrated signal of changes occurring within its large and varied catchment.

In this paper, we present a robust chronostratigraphy of the fill terraces along the Colorado and Paria rivers at Lees Ferry using optically stimulated luminescence (OSL) and terrestrial cosmogenic nuclide 10 Be exposure (TCN) dating techniques in conjunction

with stratigraphic and sedimentologic analyses. The chronostratigraphy of these deposits is presented elsewhere in order to derive long-term incision rates (Chapter 2). Here we focus on the relation between sedimentology, stratigraphy, and geochronology of deposits in the greater Lees Ferry area in order to document the timing and patterns of aggradation and incision of this large river and begin to interpret its relation to distant and local climatic forcing.

BACKGROUND

Factors controlling the aggradation and degradation of fluvial systems have been debated for many years and continue to be explored by geomorphologists. Long-term regional erosion typically occurs as a result of baselevel fall through large-scale tectonic events. In contrast, shorter-term aggradation and degradation cycles, which are the subject of this paper, may be superimposed on overall downcutting of a region and occur in response to three major forcing factors: I) climate change; 2) sea level fluctuation ; and 3) baselevel controls such as tectonic activity or drainage integration.

Figure 3.1. Shaded relief map of the central Colorado Plateau, showing the Colorado River and its major tributaries.

In general, climate change influences the balance between hydrology and sediment supply within a fluvial system. The extent of alpine and continental glaciers, weathering rates, topography, and vegetation cover within headwater catchments ultimately govern the amount of water and sediment delivered to a system through time. The ratio between the size and quantity of sediment and the ability of a stream to transport that sediment determines whether a stream will aggrade, incise , or remain in equilibrium (Lane, 1955; Bull, 1991). Aggradation occurs as flow competence within a fluvial system decreases , either as a result of decreasing discharge or increasing quantity of sediment delivered to the system. Greater sediment production in drainage basins during times of particularly intense physical or chemical weathering (i.e. glacial environments or changes in vegetation cover) can overwhelm a fluvial system with more sediment than can be carried. Decreases in effective discharge will also reduce the ability of a river to move sediment and may occur in response to lower precipitation or greater infiltration within a basin (e.g. Bull, 1991; Tucker and Slingerland, 1997). In contrast, incision happens as discharge increases or sediment supply from hillslopes decreases. Greater effective discharge may occur as a result of increased quantity or intensity of precipitation, increased effective runoff from catchments, or from glacial meltwater. Decreased sediment supply may be a result of increased storage in catchments or lower sediment production on hillslopes. These climatically controlled variations in discharge and sediment supply in response to glacial-interglacial oscillations are transferred downstream and become integrated throughout the entire length of large fluvial systems (e.g. Hancock and Anderson, 2002).

Sea-level fluctuations resulting from varying volumes of global ice introduce the potential for aggradation and degradation through the lower reaches of fluvial systems that reach the sea (Blum and Törnqvist, 2000). Rising sea-level results in decreased gradients that cause localized aggradation or increased sinuosity as lower reaches are flooded (Leopold and Bull, 1979; Schumm, 1993; Merritts et al., 1994). In contrast, lowered sea-level results in steeper gradients that effectively increase stream power and allow rivers to re-establish graded profiles through vertical incision (Mackin, 1948). A rapid sea-level fall can result in the formation of a knickpoint that can migrate tens of kilometers upstream creating a transient pulse of incision through the system (e.g. Crosby and Whipple, in press).

Baselevel changes in response to tectonic movement or drainage integration further influence the pattern of aggradation and degradation occurring throughout large fluvial systems, particularly in localized areas where these forces are active (e.g. Pederson et al., 2002). Tectonic movement due to faulting or epeirogeny has the potential to create or reduce accommodation space for deposition and influences the base level for specific reaches along the length of a river. Drainage integration has been a significant factor in the evolution of rivers in the interior western U.S (e.g. Longwell, 1946; Hunt, 1969). Stream capture occurs as river systems evolve and become integrated through headward erosion and stream piracy or by basin-spillover (Pederson, 2001).

Oscillating fluvial processes in response to the complex interaction between climate change, sea-level fluctuation, and tectonic/drainage integration controls often result in the formation of terrace sequences along fluvial valleys. Several studies from around the world have reported detailed fluvial chronostratigraphic records in order to

help refine our understanding of how fluvial processes respond to upstream and downstream controls. The following is an overview of the literature surrounding this debate and is organized in two sections. The first section reviews the response of major continental rivers to climate change, sea-level fluctuation, and tectonic controls . The second section examines sub-catchments within the Colorado River drainage and studies relating their response to Quaternary climate change .

Responses of continental-scale rivers

Observations and studies from several large fluvial systems throughout Europe and the United States illustrate the complex and varied responses of fluvial systems to upstream, downstream, and local controls (e.g. Straffin et al., 2000; Lewis et al., 2001; Rittenour et al., 2003; Anders et al., 2005). However, given the different climatic settings and varying thresholds that may exist within these systems, it is reasonable to expect that the fluvial responses of large rivers to global climate change will have varying lag-times and will not always be in phase (Blum and Törnqvist, 2000). For example, episodic deposition of two major braided channel belt systems along the lower Mississippi River valley occurred during marine isotope stage (MIS) 6 and 2 in response to high discharge conditions from melting glaciers (Blum and Tomqvist , 2000; Rittenour et al., 2003). In contrast, deposition of fluvial gravels along the River Loire in France occurred during marine isotope stages (MIS) Se, Sb, 4-3 , and 2, contemporaneously through both warm and cold periods (Straffin et al., 2000; Colls et al., 2001).

Not only are disparities in the timing of deposition apparent when comparing the large rivers of the world, but the mechanisms responsible for aggradation-degradation

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episodes can be spatially variable along the length of a single system. The River Seine in France is of particular interest relative to the Colorado River since it contains a welldeveloped sequence of Pleistocene fluvial terraces that record the variable influence of climate and sea- level change along the distinct middle and lower portions of the system (Antoine et al., 2000). The dominant control on river evolution and incision along the lower Seine is interpreted to be the lowering of sea-level, particularly during the coldest periods observed in the record (MIS 2, 6, 12, 16, and 22). In contrast, the middle Seine is beyond the influence of sea-level fall and has responded to climatically controlled variations in discharge and sediment supplied to the system during glacial-interglacial variations (Antoine et al., 2000).

Further complexities have been documented in the Rhine-Meuse system in westcentral Netherlands, wherein fluvial responses at a single location within the system have been interpreted to be controlled by different forcing mechanisms through time (Törnqvist et al., 2000; Wallinga et al., 2004). Sedimentologic and geochronologic analysis of a \sim 20 m thick sequence of inset fluvial sediments near the river's mouth reveals two significant episodes of degradation and subsequent aggradation over the last two glacial cycles. The older phase of incision occurred at the MIS $5/4$ transition in response to a significant sea-level fall, whereas the younger phase of incision occurred at the MIS 3/2 transition and is interpreted as the result of crustal updoming along a glacial forebulge. In addition, the older depositional episode occurred during MIS 4-3 and was interpreted to be controlled by cold and dry climatic conditions wherein relative sediment supplies were increased. The younger episode of aggradation occurred after the last glacial maximum (MIS 2) in response to collapse of the glacial forebulge.

Similar temporal variations in climate responses have been reported from the River Thames in England (Maddy et al., 2001; Lewis et al., 2001). Abrupt climate changes during the transition to glacial conditions at \sim 70 ka and the rapid climate fluctuations during deglaciation from 13 to 11 ka resulted in an increase in overall sediment grain size and a change from single-channel to a braided system. Aggradation of the River Thames is constrained through luminescence and radiocarbon geochronology and is bracketed during the latter part of MIS 5 between \sim 110 and 70 ka and during the last glacial maximum (LGM) between ~18 and 10 ka (Lewis et al., 2001).

In summary , patterns of aggradation and degradation along large continental rivers vary depending upon position within the system and generally occur in response to the interplay between up stream , downstream , and local controls . Variations in the timing of deposition and incision throughout the large fluvial systems of the world are evident , especially when comparing fluvial responses within a particular glacial period. In addition, temporal and spatial variations in response to forcing mechanisms within a single drainage have been documented throughout the world, making it difficult to correlate records. Our study focuses on the aggradation-incision responses recorded along the Colorado River at Lees Ferry where variations in local sediment supply have likely become incorporated with distant hydrologic signals from headwater sources through the Quaternary (Anders et al., 2005).

Responses of Colorado River sub-catchments

The connection between glacial-interglacial scale climate change and fluvial aggradation-incision can be readily observed in the headwater drainages of the Colorado

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River. Fluvial terraces have been traced in several alpine tributaries within the Rocky Mountains directly upstream to outwash plains and moraines emplaced during the last glacial advance (e.g. Reheis et al., 1991; Chadwick et al., 1997 ; Counts and Pederson, 2005). Detailed mapping of fluvial deposits and correlation to locally dated glacial moraines along the Henry 's Fork (Counts and Pederson , 2005), Yampa River (Madole, 1991), and within the Uinta Basin (Nelson and Osborn, 1991) indicate that the two most prominent terrace levels within these tributaries appear to have developed synchronously with the MIS 6 Bull Lake $(\sim 140 \text{ ka})$ and MIS 2 Pinedale $(\sim 20 \text{ ka})$ glacial advances. The early Wisconsin MIS 4 glacial advance is typically missing in the Rocky Mountain glacial record and often very minor in the pro-glacial tluvial record.

Although observed field relations within headwater drainages provide a reasonable correlation between glacial advances and aggradation-incision episodes, geochrono logic data from several drainages within the region and further downstream along the Colorado River provide a more precise, yet incomplete, characterization as to the timing of these fluvial responses. Adjacent to the Colorado River headwaters , wellconstrained studies from the upper Wind River have recognized several well-preserved terrace levels near the heavily glaciated Wind River Mountains (Gosse et al., I 995; Chadwick et al., 1997; Phillips et al., 1997). Minimum ages of terrace treads have been calculated using cosmogenic and U-series methods for terrace levels WR-3 (I 25 ka or 150 ka - MIS 6), WR-2 (55 ka - MIS 4), and WR-I (21 ka - MIS 2). These ages indicate that the timing of fluvial incision (terrace abandonment) generally corresponds to reported peak glaciations or subsequent deglacial transitions of the region (Hancock et al., I 999; Sharp et al., 2003).

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Fluvial incision recorded by cosmogenic surface ages of the well preserved Fremont River terraces in southern Utah at 151 ka (MIS 6) and 60 ka (MIS 4) similarly correlate well with the timing of glacial peaks or deglaciation in the region (Marchetti and Cerling, 2005). However, an additional incision event observed at \sim 100 ka does not correlate with any reported glaciation in the region and may reflect other sediment or hydrologic controls. In addition, no deposit of MIS 2 age has been reported from the Fremont drainage.

Portions of the Colorado River system located well downstream of glaciated catchments also respond to glacial-interglacial scale climate change. Geochronology of the fill terraces along the Colorado River in eastern Grand Canyon indicates that aggradation occurred from 385 to 322 ka (M5), about 130 to 90 ka (M4), and from about 75 to 60 ka $(M3)$, with significant river incision following each event (Anders et al., 2005; unpublished data). Comparison to global and regional paleoclimate records suggests that aggradation in eastern Grand Canyon began during the latter part of glacial periods and into early interglacial periods, followed by incision that begins at some point within subsequent interglacial periods (Anders et al., 2005). Notably, the authors recognize that deposition in local catchments of the Grand Canyon does not match the timing and stratigraphy along the Colorado River and hypothesize that distinct local sediment production and hydrologic processes control the aggradation and incision of side drainages.

As noted in the studies above, new geochronologic data from deposits throughout the Colorado River drainage system are providing an emerging picture as to the timing of aggradation and incision in response to upstream climatic forcing factors. However ,

discrepancies between various drainages are apparent. Deposits from the **MJS** 2 lastglacial episode that are readily observed in headwater drainages have not been identified in either the Fremont River or along the Colorado River through the Grand Canyon. In addition, the detailed yet complex record from the mainstem Colorado River and its small tributaries demonstrates the complex interaction between upstream glacial and local climatic control. Our results from Lees Ferry provide a robust chronostratigraphy that contributes to an emerging picture of the responses of this large river to climatic forcing.

SETTING

The preservation of deposits and landforms that record the fluvial history of the Colorado River in the erosional landscape of the Colorado Plateau are understandably rare. The few river gravels and terraces that remain along the river corridor generally exist in locations where erosion of softer strata or bedrock structural controls have created wide valleys. Lees Ferry is situated in a short but relatively wide valley that marks the end of Glen Canyon and the start of Marble Canyon (Figure 3.1). The Paria River, a major tributary to the Colorado River, enters at this relatively open spot in the landscape. Lees Ferry is one of the only places in the Grand Canyon region to have been used historically as a river crossing due to the ease of river access. Native Americans used the crossing for centuries before the arrival of early western settlers (Andrews, 1990). John D. Lee , for whom the area is named, established a ferry in 1874 that was used until Marble Canyon Bridge was completed in 1929. Today, Lees Ferry serves as the launching point for all Grand Canyon river trips.

Late-Paleozoic and Mesozoic sedimentary rocks dominate the bedrock geology of the area and are well-exposed from river level to the surrounding escarpment of the Vermillion Cliffs . These deposits have remained relatively undeformed through uplift of the Colorado Plateau, except for the formation of the Echo Monocline during the early Cenozoic Laramide orogeny . This structural feature has a north-south trend through the study area and brings the easily eroded Moenkopi Formation to river level (Phoenix, 1963). Differential erosion of softer strata throughout the Lees Ferry area has created a beautiful desert landscape of alternating massive cliffs and colorful slopes rising to the surrounding high plateaus.

The climate in the region surrounding Lees Ferry varies from arid low-lying river valleys to semiarid high elevations of the Vem1illion Cliffs and Paria Plateau. Vegetation patterns similarly follow an elevational gradient, ranging from upland vegetation in the higher elevations to desert shrub communities in the lower areas (Patton and Morrison, 1991). The terraces at Lees Ferry lie in lower elevational zones and are populated by sparse desert shrubs, whereas the banks of the Colorado River are lined with riparian vegetation such as willow and exotic tamarisk.

METHODS

Surficial deposits were mapped in the Lees Ferry area in order to record their stratigraphic characteristics and correlate terrace levels. Existing surficial geologic maps focusing on Holocene deposits along the Colorado and Paria rivers have been published by Richard Hereford and others at the USGS (Hereford et al. 2000; Hereford 2004) and were used where possible. Seven cross-sectional valley profiles were surveyed using a

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total station. Ten measured sections of parts of terrace fills have been described in detail in order to interpret depositional processes during aggradation episodes. Terrace gravels at Lees Ferry exhibit complex compositional variations that result from the mixing of sediment delivered from the Colorado and Paria rivers. Nine clast counts (> 100 clasts) and eight sand petrographic point counts of grain-mount thin sections stained for the identification of feldspars (>300 counts) were performed in order to distinguish sediment from local hillslope, tributary, or mainstem Colorado River provenance.

In addition to mapping and sedimentologic observations and measurements, two dating methods have been employed to constrain the timing of aggradation and subsequent terrace abandonment at Lees Ferry: optically stimulated luminescence (OSL) and terrestrial cosmogenic nuclide 10 Be exposure (TCN). Although both dating methods have been used on fluvial deposits, ages that are calculated from OSL and TCN require different interpretations. TCN ages represent the duration of surficial exposure since the onset of incision , whereas OSL ages indicate the timing of sediment deposition and burial. A vertical sequence of precise OSL ages taken from a single deposit should record the timing and duration of aggradation within a fluvial system, whereas a single TCN age from a terrace surface indicates the timing of floodplain abandonment and subsequent incision.

OSL sand samples were collected in 5 x 20 cm aluminum tubes from recently exposed road or stream cuts. Ten OSL samples were collected from Colorado River deposits and four samples were collected from Paria River deposits. OSL sample preparation and analysis of the 90 to 150 µm fraction was performed at the University of

Nebraska Luminescence Laboratory using the single-aliquot regenerative method on a RISO TL/OSL-DA-15B/C reader with blue-green light stimulation (470 nm, Hoya U340 filter). OSL results reported up to this point in the research are preliminary, with 6 to 15 disks reported for each sample. This is only about half of the expected number of disks that will be used to calculate final ages (20 total for each sample).

Four TCN surface samples were collected from only those Colorado River terrace treads that exhibited long-tenn stability in the form of well-developed desert pavements. Surface samples composed of \sim 100-200 quartzite pebbles (1-2 cm in diameter) were crushed and analyzed in a single amalgamated sample after the methods of Repka et al. (1997). A 220 cm depth-profile consisting of six sand samples was used to account for pre-depositional inheritan ce. This profile from the most prominent deposit in the area was used for all surface ages with the assumption that the Colorado River has transported well-mixed sediment with a comparable amount of inheritance during each aggradation episode. TCN samples composed of \sim 250 kg of medium sand-sized grains (355-500 μ m) were prepared at Dalhousie University and accelerator mass spectrometer analysis was performed at Lawrence Livermore National Laboratory.

RESULTS

Deposits from the Colorado River, Paria River, and small tributary washes have been identified in the Lees Ferry area. Description of each of these deposits is presented below, including geometry and landscape position, sedimentology, provenance, and age.

Colorado River deposits

Seven distinct deposits $(M1-M7)$ and ten terrace levels $(M4y, M5y, and M5m$ are erosional fill-cut terraces) have been identified along the Colorado and Paria rivers in the Lees Ferry area (Figure 3.2). Deposits generally range in thickness from 10 to 30 m and have basal straths that can be irregular, often exhibiting false straths that are significantly higher than the true basal strath. Terrace treads range in height from 5 m (M1) to \sim 180 m (M7) above the reference river stage of 380 m³/s (\sim 13,000 cfs) and generally exhibit moderately to well-developed desert pavements (Figure 3.3). Colorado River deposits are characterized by the interfingering of five distinct facies described in Table 3.1: clastsupported gravel, cross-stratified sand, immature pebble-gravel, cobble-pebble diamicton, and boulder diamicton. Sand petrogaphic analysis indicates that mainstem Colorado River sands are $\sim 75\%$ quartz and may be slightly more feldspathic than sand from Paria River deposits (Appendix 8) .

The MI deposit is well-exposed along the Colorado River in the Lees Ferry area and occurs up to \sim 5 meters above the modern river. This deposit represents middle Holocene to modem deposition of the fluvial system. The MI has been mapped and described in detail by Hereford et al. (2000), who subdivided these Holocene deposits on the basis of dated archeological remains, tree-ring dates , content of driftwood , and comparison to similar terrace sequences from established records elsewhere within the region (Hereford et al., 1986; 1998). Mapping of these Holocene deposits and studies of historic sand bars and the sediment budget along the Colorado River have been perfonned in an attempt to understand how Glen Canyon Dam is affecting aquatic habitat and the sediment and water budgets of the Colorado River (e.g. Schmidt and Graf, 1990).

Figure 3.2. Map of Quaternary alluvium in the Lees Ferry area, including the Colorado (M l-M7) and Paria **(P l-P4)** river deposits, as well as the Johnson Wash deposit (S3).

Figure 3.3. Schematic cross-sectional valley profile of the Colorado River fill terraces at Lees Ferry, showing heights, geometries, and ages of deposits. Sec Table 3.2 for error calculations on individual ages.

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Table 3.1. Sedimentary facies, descriptions, and interpretations for deposits at Lees Ferry, Arizona.

Data Point Description	Height (m) ¹	Age $(ka)^2$	Dating Method	
Colorado River				
Holocene flood deposits (M1)	$0 - 6$		Various	
Surveyed M2 terrace tread	14.4		$- -$	
M2 sand lens in fill	10.7	21 ± 1	OSL^5	
M3y sand lens in fill	12.8	39 ± 3	OSL^5	
Surveyed M3 strath	15.3		-1	
Surveyed M3 terrace tread	25.6	$- -$	--	
M3 desert pavement	25.6	38 ± 3	TCN ⁴	
M3 sand lens in fill	19.2	72 ± 5	OSL^5	
M3 sand lens in fill	22.2	71 ± 5	OSL^5	
S3 sandy unit in fill	51.9	70 ± 6	OSL^5	
S3 sandy unit in fill	60.2	40 ± 3	OSL^5	
Surveyed M4 strath (lower)	15.8		$-$	
Surveyed M4 strath (upper)	30.8		$- -$	
Surveyed M4y terrace tread	41.8		--	
M4y desert pavement	41.8	87 ± 7	TCN ⁴	
Surveyed M4o terrace tread	45.8		$- -$	
M4o desert pavement	45.8	86 ± 7	TCN ⁴	
M4 sand lens in fill	34.9	97 ± 8	OSL^5	
M4 sand lens in fill	18.3	114 ± 8	OSL^5	
Surveyed M5 strath (lower)	42.5			
Surveyed M5 strath (upper)	60.5		$- -$	
M5y sand lens in fill (lower)	45.0	121 ± 12	OSL^5	
M5 sand lens in fill (upper)	62.5	138 ± 10	OSL^5	
Surveyed M5y terrace tread	64.3		\sim $-$	
Surveyed M5m terrace tread	66.7			
M5m desert pavement	66.7	129 ± 10	TCN ⁴	
Surveyed M5o terrace tread	78.7	$- -$		
Paria River				
Holocene flood deposits (P1)	$0 - 5$		Various	
Surveyed P2 terrace tread	14.6			
Surveyed P3 strath	8.7			
Surveyed P3 terrace tread	23.1			
P3 sand lens in fill	12.2	74 ± 7	OSL^5	
Surveyed P4 strath (lower)	15.8		--	
Surveyed P4 strath (upper)	33.5			
Surveyed P4y terrace tread	38.6			
Surveyed P40 terrace tread	46.5			
P4 sand lens in fill	17.6	214 ± 23	OSL^5	
P4 sand lens in fill	23.6	96 ± 11	OSL^5	
P4 sand lens in fill	33.3	102 ± 9	OSL^5	

Table 3.2. Survey and geochronologic data for Lees Ferry deposits

¹ Referenced to a local river stage of 380 m³/s (13,000 cfs)
² 28 errors reported for all ages
³ Fine-grained Holocene overbank sediment of the Colorado River (Hereford et al., 2000)

⁴ Terrestrial cosmogenic-nuclide date of surface sample corrected for inheritance with depth-profile, minimum surface age (see

Appendix E for complete data)
⁵ OSL samples represent ages of deposition (see Appendix F for complete data)

The M2 at Lees Ferry is preserved in only two places just upstream of the confluence with the Paria River (Figure 3.2). It has a minimum thickness of 14 m and the base lies below the modem Colorado River. Sedimentology of the lower portion of the deposit (at the southern-most outcrop) is dominated by local debris flow and interfingered side canyon fluvial facies. The upper 1.5 to 5 m of the M2 is characterized by very fine to medium sand with ripple-cross stratification interpreted to have been deposited in mainstem overbank environments. This is typically interfingered with hillslope deposits. OSL geochronology from sandy units indicates that aggradation of the M2 began prior to 21 ± 1 ka and continued until sometime after this (Table 3.2). A deposit first interpreted as an M2 near the Lees Ferry boat ramp may turn out instead to be a mainstem channel facies remnant from a time between the major M3 and M2 fill depositions, based on results of an OSL sample. An age of 39 ± 3 ka from this deposit indicates that a pause in the overall incision may have occurred after the M3 episode and before the onset of M2 aggradation ~20 ka. Scdimentology of this unit indicates deposition of predominantly far-traveled cobbles (~83% of clasts) in the mainstem Colorado River channel (Table 3.1; Figure 3.7; Appendix A). Far-traveled clasts include rocks identified as quartzites and volcanic porphyries.

The M3 deposit is preserved in only one location in the Lees Ferry area. The total preserved thickness of the local M3 is at least 11 m and its base is 14 m above the reference river stage. This deposit exhibits complex interfingering of mainstem gravels, tluvial side-canyon gravels from Johnson Wash, and mainstem overbank deposits. A clast count from the middle of the deposit indicates that 68% of the clasts are of local lithologies (sandstones, limestone, and rocks of the Claron formation), whereas 32% are

far-traveled quartzites and volcanic porphyries (Figure 3.7; Appendix A). Aggradation of the M3 began by \sim 70 ka, based on two OSL ages of 72 \pm 5 ka and 71 \pm 5 ka obtained from the middle portion of the deposit, respectively (Table 3.2; Figure 3.4). A TCN date obtained from the well-preserved desert pavement on the terrace tread suggests that incision occurred prior to 38 ± 3 ka. Based on the thickness, extent of preservation, and timing of degradation of the M3 deposit in eastern Grand Canyon (Anders et al., 2005), it is likely that the full thickness of the M3 is not preserved at Lees Ferry .

The M4 is by far the most extensively preserved deposit in the Lees Ferry area. The thickness of the M4 ranges from 10 to 30 m due to topography of the basal strath. A prominent higher false strath lies 3 1 m above the reference river stage, whereas the lower planar bedrock contact lies 15 m above the reference river stage (Figure 3.3). The lower half of the M4 deposit has interfingering mainstem gravel and overbank sand facies. In contrast, mainstem gravel and sandy facies in the upper half of the deposit are interrupted by \sim 3 m thick debris flow units dominated by locally derived clasts (Figure 3.6). Clast counts within three units (units 2, 5, and 10) indicate that far-traveled lithologies $(\sim 50\%)$ are sub-equally mixed with local clast types (-50%) (Figure 3.7; Appendix A). The M4 aggradation episode began prior to 114 ± 8 ka, and continued until at least 97 ± 8 ka, based on OSL dates obtained near the base and from the upper portion of the deposit (Table 3.2 ; Figure 3.5). Degradation of the M4 was accomplished in two phases. Initial incision occurred by 86 ± 7 ka and resulted in the abandonment of the M4o surface. Incision halted for a brief time (less than the resolution of TCN methods) prior to ultimate abandonment and formation of the M4y surface by 87 ± 7 ka (Table 3.2).

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Figure 3.4. Photo of the M3 deposit showing location and ages of OSL (fill) and TCN (tread) samples. The M4 deposit is in background, including the M4y TCN age obtained from the terrace tread. For uncertainty of dates and other geochronologic information , see Table 3.2 and Appendix E and F.

Figure 3.5. Part of the M4 deposit showing location and age of OSL samples collected near the base and middle of the deposit, as well as TCN age collected from the M4o terrace tread. Photo taken looking upstream (north) at confluence. For geochronologic data and uncertainties, sec Table 3.2 and Appendix E and F.

Figure 3.6. Type sedimentary section of the M4 deposit. Visual proportions of exotic and local clasts were made in order to estimate the dominant source throughout deposition . Sedimentary facies interpretations are also included. Sedimentary descriptions can be found in Appendix C.

Figure 3.7 . Histograms of clast-count data collected from Colorado and Paria river deposits. Total counts were of 100 or more clasts. Raw data are found in Appendix A.

The *MS* deposit is preserved in three locations in the southwestern part of the Lees Ferry area. Local thickness of the *MS* ranges from 10 to 30 m due to topography of the basal strath. A significant higher strath lies 61 m above the reference river stage, whereas the lower bedrock contact lies 43 m above the reference river stage (Figure 3.3). The sedimentology of the M5 is obscured due to poor exposure; however, tributary debris flow units dominate the top of the deposit and appear to be interfingered with mainstem sand and gravel facies near the middle and base. Preliminary OSL analysis from a sand lens in the middle of the deposit suggests that aggradation was ongoing at 138 ± 10 . Subsequent degradation occurred in at least two phases, the first resulting in the formation of the M5o surface. A new floodplain $(M5m)$ was established \sim 4 m below this level some time before 129 ± 10 ka, as indicated by a TCN date. In addition, a subsequent pause in overall incision may have occurred after the *MS* episode and before the onset of M4 aggradation, as indicated by an OSL age of 121 ± 12 ka from a slightly lower landform .

The M6 and M7 deposits are erosional remnants perched high above the modem Colorado River. The top of the M6 lies 72 m above the reference river stage and is called Lees Lookout on USGS topographic maps. Remnants of this deposit are also likely preserved downstream on the high benches of Marble Canyon. The M6 has not been dated due to lack of exposure of datable material.

The M7 lies ~180 m above the present day Colorado River on Johnson Mesa in three isolated patches. The M7 has stage V soil carbonate development, and was estimated to have an age of $~500$ ka by Lucchitta et al. (2000). U-series methods were used to date the well-developed carbonate rinds of the M7 (data provided by Warren

Sharp and Richard Phillips at Berkeley Geochronology), and initial results indicate that deposition occurred >350 ka (secular equilibrium). However, a younger stage of carbonate growth occurred ~ 60 ka based on U-series analyses of carbonate coatings that envelope older rinds (Warren Sharp and Richard Phillips, personal communication). This suggests that a climate episode at ~ 60 ka (during M3 deposition) resulted in renewed soilcarbonate fonnation in this high, old soil profile.

Paria River deposits

Four distinct deposits (P1-P4) and five terrace levels (P4y is an erosional fill-cut terrace) have been identified near the mouth of the Paria River (Figure 3.2). Deposits range in thickness from 9 to 30 m and have irregular basal straths similar to those observed in the Colorado River terraces. Terrace treads have moderately-developed desert pavements and range in height from 4 m (P1) to 46 m (P4o) above the modern Paria River (Figure 3.8). Paria River deposits are sedimentologically similar to Colorado River deposits , displaying intcrfingering of the five facies described in Table 3.1. Results from sand petrology indicate that sands from the Paria River near the confluence are \sim 80% quartz, perhaps slightly more quartz rich than sand of the mainstem Colorado River (Appendix B).

Figure 3.8. Schematic cross-sectional profile of the Paria River fill terraces near Lees Ferry, showing heights, geometries, and ages of the deposits. Details of geochronology can be found in Table 3.2 and Appendix F.

The P1 is composed of several inset Holocene sandy deposits that have been the subject of several studies over the past 20 years (e.g. Hereford, 1986; Graf et al., 1991; Hereford, 2002). The P1 has an exposed thickness of 4 m and the base lies below the modem Paria River (Figure 3 .8). The suite of PI deposits have been correlated with incision due to frequent large floods during episodes of frequent positive El Nino-Southern Oscillations, and aggradation during times of low flood magnitudes (Hereford, 1986; Graf et al., 1991).

The P2 deposit is preserved along the lower Paria River valley near the confluence with the Colorado River (Figure 3.2). It has a maximum thickness of \sim 14 m, and its base lies below the modern Paria River (Figure 3.8). The P2 is characterized by the interfingering of overbank sand and hillslope overland flow facies (Table 3.1). OSL dating of this deposit has not yet been completed; however, physical correlation to the M2 just downstream suggests aggradation was ongoing \sim 20 ka.

The P3 deposit is found in three locations along the lower Paria River valley. This deposit is 9 m thick and its base lies 12 m above the modern Paria River (Figure 3.8). Although the P3 deposit is not well-exposed, fluvial gravels appear to be interfingered with overbank sand facies. An OSL date from a laterally extensive sand lens in the middle of the P3 indicates that aggradation was occurring at 74 ± 7 ka (Table 3.2).

The P4 is the most extensively preserved deposit in the Paria River valley. It has a thickn ess of 24 m and its base exhibits two distinct planar levels. The upper strath lies 31 m above the modern Paria River , whereas the lowest strath is 16 m above the Paria River. The P4 has two distinct depositional units, each overlain by a bouldery debris

flow package that followed pulses of Paria aggradation. Paria River deposits are characterized by fluvial gravels derived from the Paria catchment (including \sim 5% of the distinctive limestone clasts of the Claron Formation) interfingered with significant overbank sand lenses (Figure 3.7; Table 3.1; Appendix A). OSL geochronology from the P4 suggests that the lower depositional unit may be a buried older deposit $(214 \pm 23 \text{ ka})$, but results show a good deal of uncertainty. Aggradation of the upper depositional unit occurred at \sim 110 to 90 ka, based on two OSL ages of 96 \pm 11 ka and 102 \pm 9 ka from the upper 20 m of the deposit, which are the same age within error (Figure 3.8; Table 3.2). Although the timing of terrace abandonment has not been numerically constrained, initial degradation resulted in the formation of the P4o surface and the establishment of a new floodplain \sim 20 m below this surface (P4y surface).

Tributary side canyon deposit

Previous research at Lees Ferry performed by Kaufman et al. (2002) described a fossiliferous deposit of interfingered silty-sand, carbonate-mud, and coarse-grained sediment located within Johnson Wash (S3 in Figure 3.2). Hamblin (1994) speculated that this deposit was lacustrine in origin, having formed in an extensive lava-dam lake more than 1 Ma. However, Kaufman et al. (2002) cite sedimentologic, fossil assemblage, and stratigraphic evidence to interpret it as spring-fed carbonate deposition occurring contemporaneously with side canyon debris flows. The authors report an age of \sim 40 ka based on amino-acid dating of snails found within the deposit.

Detailed sedimentologic and geochronologic data from our study expands on the findings of Kaufman et al. (2002). The S3 is generally \sim 10 m thick, although the

irregular basal strath may increase or decrease local thicknesses. Landscape position of the S3 indicates that it is graded to and can be correlated to the M3. The lower portion of the deposit is characterized by carbonate-rich silt to medium sand deposited in a paludal environment. In contrast, carbonate-rich sands in the upper deposit are interfingered with medium beds of debris flow diamictons (Table 3.1; Appendix C), indicating significant influx of tributary sediment near the end of deposition. Two new OSL dates from the S3 indicate deposition began prior to 70 ± 6 ka and was nearly complete by 40 ± 3 ka (Table 3.2), correlating well to the M3 on the Colorado River.

DISCUSSION

Climatically induced changes in sediment supply and hydrology have resulted in the formation of thick fill terraces at Lees Ferry that are superimposed on overall incision of the region. The observed spatial extent and thickness of these deposits gives us some idea about the magnitude of climate responses during different aggradation events in middle-late Pleistocene time, although this is also influenced by valley geometry (Figure 3.2). The M4 is of particular interest since it is the most extensive deposit in the Lees Ferry area, and it is tempting to interpret the M4 as responding to a very large climate event. However, significant vertical and lateral erosion by the Colorado River at Lees Ferry just prior to deposition of the M4 was potentially facilitated by the erosional bench created by the recessive Moenkopi Formation and underlying resistant Kaibab Limestone. This contact is brought to river level right where the M4 is largest, immediately downstream of the Paria River confluence. Aggradation of the M4 occurred along this wide valley bottom and the deposit may have been protected from subsequent

erosion because of its position above the resistant Kaibab Limestone ledge. It is interesting to note that correlative deposits \sim 130 km downstream in eastern Grand Canyon are only partially preserved in isolated locations (Anders, 2003). The M3 is not well preserved at Lees Ferry, perhaps because it lies confined riverward and below the resistant Kaibab Limestone. In contrast, the M3 dominates the stratigraphy in eastern Grand Canyon since it is the youngest major aggradation event preserved within the relatively narrow canyon (Anders, 2003).

Some broad patterns can be observed in the provenance data of the Lees Ferry deposits. First, clast counts generally indicate that local sandstones are sub-equally mixed with far-traveled quartzites and volcanic porphyries in the mainstem channel facies (Figure 3.7). Second, some pink limestone clasts from the Claron Formation of the upper Paria River catchment are observed in several Colorado River clast counts below its confluence with the Paria River. This indicates that sediment from the Colorado and Paria Rivers are mixed, although accurate proportions of mixing cannot be deciphered from clast count data due to the similar clast types in each catchment. Third, debris flow deposits are typically observed in the upper portions of depositional packages, perhaps indicating that local hillslope and tributary activity is greatest near the end of aggradational episodes. In addition, sedimento logic and stratigraphic evidence indicates that the relative abundance of channel and overbank facies in the Colorado and Paria river deposits has not significantly changed through the preserved record. In particular, the presence of imbricated cobble channel facies and overbank sands with lower preservation potential throughout each deposit suggests that the depositional style of the paleo-Colorado and Paria Rivers have been relatively similar during each aggradation

episode. It is difficult, however, to directly compare plan-form proportions of modern river depositional facies to those preserved in vertical outcrops of fill terraces emplaced during aggradational modes of the river.

Regional comparison of jluvial records

Regional comparison between the chronostratigraphy at Lees Ferry and fluvial records throughout the middle-upper Colorado River catchment highlight both similarities and differences. The M2 deposit at Lees Ferry area has an OSL depositional age of 21 ± 1 ka, equivalent to the 21 ± 5 ka surface age reported from a terrace along the Wind River (Sharp et al., 2003). These deposits appear to be associated with the last glacial maximum episode reported from the headwater drainages of the Colorado River. A deposit of this age has not yet been reported from the Colorado River on the Colorado Plateau, but was hypothesized to exist under the modern river in eastern Grand Canyon (Anders et al., 2005). The presence of the M2 deposit in the Lees Ferry area suggests that the Colorado River responded lo climate change that occurred during the last glacial maximum by aggrading; however, the total thickness and full timing of deposition remains unknown.

Aggradation of the M3 is generally consistent with regional fluvial records. In particular, the onset of deposition at Lees Ferry at 72 ± 5 ka is synchronous with the timing of aggradation reported downstream in eastern Grand Canyon (Anders et al., 2005) . However , degradation of the M3 and S3 at Lees Ferry at - 40 ka appears to be later than the onset of incision $(-50-60 \text{ ka})$ reported from eastern Grand Canyon (Anders et al., 2005) , the Wind River (Hancock et al., 1999; Sharp et al., 2003), and the Fremont

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River (Repka et al., 1997). There are two possible explanations for this discrepancy. First, the M3 TCN date may be on a lower erosional terrace level rather than the true top of the original deposit, which may not be preserved at Lees Ferry. Second, incision of the M3 at Lees Ferry may have been delayed due to high influxes of sediment from local hillslope and tributary systems. Comparison to the nearby fluvial record reported from eastern Grand Canyon wherein significant side canyon aggradation was occurring from 50-30 ka supports this second explanation.

Aggradation of the M4 deposit at Lees Ferry beginning prior to 114 ± 8 ka is broadly consistent with or slightly younger than deposition reported in eastern Grand Canyon starting by 124 ± 1 ka and continuing after 118 ± 3 ka. However, subsequent degradation of the M4 at Lees Ferry at 87 ± 4 ka significantly post-dates the timing of incision reported for the stratigraphically equivalent WR-3 in the Wind River (150 ± 8 or 125 ± 37 ka). There are at least two possible reasons for this discrepancy. First, geochronologic comparison between the M4 deposit at Lees Ferry and the WR-3 in the Wind River may indicate a potential depositional lag time of \sim 20 to 40 ka associated with the downstream transfer of sedimentary signals from glaciated headwaters. Second, the influence of local climatic events in the central Colorado Plateau region may have overshadowed any distant climatic control in the headwaters. Although the M4 is the most extensive deposit in the Lees Ferry area, the timing of aggradation and degradation appear to be consistent with ages reported from the tributary S4 in eastern Grand Canyon of \sim 110 to 90 ka (Anders et al., 2005; unpublished data). Further, the incision of the Fremont River recorded by a cosmogenic surface age at 102 ± 16 ka (Repka et al., 1997) lies within this time frame.
Aggradation of the M5 before 138 ± 10 ka does not appear to correlate to any preserved deposit in eastern Grand Canyon, although deposits of this age may have subsequently been eroded in the steep canyon landscape. Abandonment of the MS at \sim 130 ka at Lees Ferry appears to significantly post-date incision reported along the Fremont River (151 \pm 24 ka) and near the headwaters along the Wind River (WR4 = 167 \pm 6 ka). Considerable scatter and a lack of data from deposits of these ages preclude any definitive correlations between the timing of incision in these regional records; however, it appears that events at Lees Ferry are somewhat younger than those occurring upstream in headwater catchments.

Although a robust dataset from the Paria River is not yet available, comparison between the fluvial records obtained from the Paria River, Johnson Wash, and the Colorado River at Lees Ferry are generally consistent. The P3 and M3 have equivalent depositional ages of \sim 70 ka, and the P4 and M4 were both aggrading \sim 100 ka. Similarly, aggradation of the Johnson Wash paludal deposit has comparable ages with the Colorado River M3 deposit (70 \pm 6 to 40 \pm 3 ka). These similarities suggest that deposition along Colorado River, tributary Paria River, and side canyon environments at Lees Ferry occurred at or nearly in-step with each other despite issuing from catchments of varying size and climatic setting. At this close proximity to their confluences, local aggradation and incision of the tributary systems may have ultimately been controlled by the base level of the Colorado River. Additional detailed geochronologic analysis of deposits located farther up the Paria drainage would allow testing of this observation.

Response of the Colorado River to climate change

Geochronology of the Colorado River deposits at Lees Ferry indicates that aggradation and degradation of this continental-scale river generally occur in response to upstream climatic controls from headwater drainages and are not affected by sea-level controls occurring \sim 1,200 km downstream at the mouth. However, the timing of fluvial responses of the Colorado River at Lees Ferry is not always perfectly consistent with fluvial records reported within the region or with independent climate records. Comparison between the chronostratigraphy at Lees Ferry and the regional Devils Hole and global SPECMAP climate records indicates that aggradation at Lees Ferry occurred during regionally cold and wet periods, whereas incision occurred during relatively warm and dry periods (Figure 3.9). However, no simple or direct relation between aggradationdegradation events at Lees Ferry and Marine Isotope Stages (MIS) is observed . Aggradation of the M2 , M3 , and M6 appears to have consistently begun during the full glacial conditions of MIS 2, 4, and 6, respectively. In contrast, incision of the Lees Ferry deposits does not appear to have occurred at an equivalent point within marine isotope stages. Incision of the M6 and incision of the M2 appears to have begun during interglacial conditions after glaciers in headwater drainages melted. In contrast, degradation of the M3 occurred well into the colder MIS 3.

Interestingly, the most prominent and best dated deposit in the Lees Ferry area (M4) was deposited during MIS 5d-b, a time in which no glaciations have been reported in the headwater drainages or in global climate records. Although the M4 does not appear to be associated with a specific glacial episode, the regional Devils Hole and global SPECMAP climate records indicate that aggradation occurred when regional

temperatures were relatively low and global ice volume was relatively high \sim 115 to 90 ka (Figure 3.9). In addition, these climate records indicate that incision of the M4 occurred during an apparent warming period in MIS 5b-a - 90 to 80 ka. Emplacement and abandonment of the M4 during relatively small global climate fluctuations within MlS 5 may indicate that fluvial responses at Lees Ferry are sensitive to relatively small changes. Local sediment production and hydrology may be significantly affected by apparently minor shifts in temperature and ice volume, ultimately resulting in the formation of terraces at Lees Ferry during times in which fluvial responses were not pronounced in other areas throughout the region or world.

CONCLUSIONS

The data from this study suggest that deposits at Lees Ferry are younger than those in equiva lent landscape positions in headwater catchments and perhaps somewhat younger than correlative deposits in eastern Grand Canyon. In addition, the presence of a predominant \sim 115 to 90 ka M4 deposit at a time without known glaciation in headwater catchments supports the conclusion that local sediment production and hydrologic factors have significantly influenced the timing and nature of aggradation and incision during certain times within the Quaternary record.

Figure 3.9. Curve representing the height of the Colorado River channel bed through time. Stippled pattern represents fill deposits, whereas the gray represents bedrock. Alluvial aggradation cycles are superimposed on overall downcutting at Lees Ferry. Solid bold lines are regressions approximating minimum and maximum incision rates at Lees Ferry. The regional Devils Hole record (modified from Winograd et al., 1992) and global SPECMAP record (modified from Martinson et al., 1987) are shown below. Shaded bars represent glacial periods. See Table 3.2 for error calculations on individual ages.

The timing of aggradation and degradation reported from the deposits at Lees Ferry illustrates the complex nature of large fluvial systems and their potential responses to local and distant climatic forcing. Anders et al. (2005) propose a revised conceptual model for dryland environments in which increased sediment production on local hillslopes in the Grand Canyon region during glacial conditions is initially stored in thick colluvial mantles. A subsequent change to warmer and wetter climate regimes initiates the transport of that stored material through incremental events down tributary channels, ultimately reaching the mainstem river. Climatically controlled aggradation-degradation cycles at Lees Ferry may be integrated with localized controls on sediment supply such as those envisioned by Anders et al. (2005). Far-traveled clasts were deposited simultaneously with local hillslope and Paria River sediment in fluvial environments, suggesting that aggradation episodes were not simply controlled by variations in discharge and sediment supply from distant headwaters. Rather, the influx of sediment from tributaries and local hillslopes has likely influenced the timing and magnitude of aggradation and degradation at Lees Ferry.

In summary, aggradation and degradation episodes at Lees Feny are influenced by hydrologic driving forces that are likely tied to climate conditions and the advance and retreat of glaciers in headwaters. However, changes in sediment supply are likely driven mostly by local sediment production from hillslopes and tributaries. Large drainages, such as the Colorado River, may exhibit patterns of deposition and incision that are controlled by this complex mixture of upstream-driven hydrology and local sediment production.

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CHAPTER 4

CONCLUSIONS

As our knowledge of the evolution of the Colorado River system improves, a complex picture of overall river incision and fluvial responses is beginning to emerge. Chronostratigraphic analysis of the well-preserved Pleistocene fluvial terraces along the Colorado River at Lees Ferry indicate a middle-late Pleistocene bedrock incision rate of 290 to 470 m/my. These incision rates are approximately two to three times higher than others reported downstream in Grand Canyon and upstream in the upper Colorado River basin. In contrast, our incision rates at Lees Ferry are similar to the recently reported high incision rates near Glen Canyon and along the Fremont River, suggesting there is a reach of faster incision along the Colorado River in the vicinity of Lees Ferry and Glen Canyon. These faster mid-late Pleistocene incision rates of the central Colorado Plateau region may be caused by: I) localized epeirogenic uplift due to tectonics and potentially enhanced by isostatic rebound; or 2) transient knickzones resulting from drainage integration $~6$ Ma, moderated by variations in bedrock resistance.

Superimposed on this long-term regional incision are several aggradation and degradation episodes that illustrate the potentially complex nature of a large river's responses to local and distant climatic forcing. Comparison between the chronostratigraphy at Lees Ferry to both fluvial records within the region and independent climate records indicates that the timing and magnitude of fluvial responses may be spatially and temporally variable. Although aggradation appears to occur during relatively cold and wet periods and incision during relatively warm and dry periods, no

simple relation between the magnitude of fluvial responses at Lees Ferry and marine isotope stages is observed. Specifically, the timing of aggradation and incision within the Lees Ferry record potentially occurs somewhat later than that which has been reported in the glaciated headwaters, as well as downstream in the Grand Canyon. In contrast, the timing of aggradation and incision often matches the timing of Grand Canyon tributaries studied in other research. In conclusion, the results from Lees Ferry should support and encourage further study into how large drainages, such as the Colorado River, may exhibit patterns of deposition and incision controlled by the interaction between upstream hydrologic forcing factors and local sediment production.

APPENDICES

Appendix A. CLAST COUNT DATA

 \sim

Figure A.1. Map of Quaternary deposits in the Lees Ferry area, showing approximate locations where clast counts were performed.

Unit: Location:			M2 (sandy flood deposit) 446827 E, 4079927 N							
Notes:							Count performed in unit 1 of M2 sandy sed description			
	Red SS	W/Y Qtz	Ylw Qtz	V Porph	B1 Chert	Lt Chert	Rd Chrt	Limestne	Claron	Rd Silt
	1.5 3.5 \mathbf{I} 1.5 1.5 3.5 10 0.5 $\sqrt{3}$ $\sqrt{2}$ 1.5 1.5 1.5 0.5	1.5 $\overline{3}$ 2.5 $\overline{2}$	9 0.5 $2.5\,$ 1.5 3.5 3 $\,$ $\,$ $\sqrt{2}$ 0.5 3.5 3.5 2.5 2.5 2.5 6 3 1.5 $\overline{3}$ \overline{c} 1.5 $\overline{2}$ 0.5 0.5 5 $\overline{2}$	5 9.5 $\frac{2}{3}$ $\overline{4}$ 2.5 1.5	\overline{c} 1 $\overline{\mathbf{c}}$ \overline{c} $\mathbf{1}$ $\overline{4}$ 2.5 2.5 1.5 1 1	1 1 0.5 1.5 0.5 3	1.5 0.5	0.5 $\overline{4}$ 0.5 $\overline{2}$ $\overline{4}$ 1.5 $\mathbf{1}$ 0.5 1 $\sqrt{2}$ $\mathbf{1}$ 1.5	2.5 1.5 $\mathbf{1}$ $\overline{2}$ 0.5 1.5	$\overline{2}$ 3.5 0.5 $\mathbf{1}$ 2.5 $\frac{2}{1}$
Avg Size $- cm$	2.1	2.3	2.8	3.9	1.9	1.3	1.0	1.6	1.5	1.8
Percent Total	17.5	4.1	26.8	7.2	11.3	6.2	2.1	12.4	6.2	7.2

Table A.1. Clast count data for the M2 deposit.

Table A.2. Clast count data for the M3y deposit.

Unit:	M ₃							
Location:		446036 E, 4078918 N		M3 deposit in culvert near Paria beach cobble unit above M3 OSL.				
Notes:								
	Red SS	W/Y SS	Qrtzite	V Porph	Bl Chert	Chert	Limestone	Shinarump
12 0.5 0.5 2.5 Average	$\overline{4}$ $\!1$ $\sqrt{6}$ 1.5 \mathfrak{Z} 0.5 \overline{c} $\boldsymbol{7}$ $\sqrt{2}$ 11 12 $10\,$ 5 $\ensuremath{\mathfrak{Z}}$ 1.5 $8\,$ \mathfrak{Z} 6 $\overline{9}$ 1.5 1.5 $\mathbf{1}$ 6 6 11 \overline{c} 11 $\sqrt{2}$ $\overline{3}$ $\sqrt{5}$ 0.5 1.5 6 \overline{c} $\,$ 5 $\,$ 3.5 $\sqrt{2}$ $\overline{\tau}$ 2.5 6 1 2.5 2.5 \mathbf{I} $\overline{4}$ $\boldsymbol{9}$ $\overline{4}$ \mathfrak{Z} $\sqrt{3}$ $\sqrt{5}$ 14 4.8	$\overline{21}$ $\overline{4}$ $\overline{\tau}$ $\overline{\mathcal{I}}$ 3 9 3 13 $\overline{4}$ $\overline{4}$ $0.5\,$ 1.5 $\sqrt{5}$ $\mathbf{1}$ 8 6.1	$\overline{5}$ 1.5 3.5 $\sqrt{2}$ $\sqrt{2}$ 2.5 $\mathbf{1}$ 1.5 $\mathbf{1}$ $\overline{2}$ 2.5 $\sqrt{2}$ 1.5 $\mathbf{1}$ 1 $\mathbf{1}$ 0.5 $\mathbf{1}$ 4.5 $\frac{2}{3}$ $\overline{2}$ 2.0	$\sqrt{6}$ $\sqrt{2}$ 11 $\sqrt{6}$ $\sqrt{2}$ $\overline{4}$ $\overline{4}$ $\frac{3}{3}$ 4.6	0.0	$\sqrt{2}$ $\sqrt{2}$ 0.5 1 $\mathbf{1}$ $\begin{array}{c} 2 \\ 4 \end{array}$ 1.5 \mathfrak{Z} $2.5\,$ $\sqrt{2}$ 1.5 $\overline{4}$ 2.1	6 $\overline{3}$ 6 5 $\mathbf{1}$ 14 $\mathfrak s$ $\ensuremath{\mathfrak{Z}}$ 0.5 $\mathbf{1}$ 5 τ 3.5 $\begin{array}{c}\n2 \\ 2 \\ 5 \\ 7\n\end{array}$ 4.5	$\overline{2}$ $\frac{3}{8}$ 4.3
Size - cm						9.6		2.2
Percent Total	41.5	$11.1\,$	16.3	6.7	$0.0\,$		12.6	

Table A.3. Clast count data for the M3 deposit.

 $\overline{\text{Unit:}}$

Notes:				M4 in gully cut north of housing terrace								
	Red $\rm SS$ $\overline{2}$ 2.5 $\overline{4}$ 15 13 6.5 $\overline{4}$ 4.5 $\sqrt{5}$ 8 3.5 15 5 $\overline{4}$ 2.5 7 4 $\overline{4}$ 2.5 Q	\mathbf{W}/\mathbf{Y} SS 1.5 $0.5\,$ $\frac{2}{2}$ 1.5 4.5 7.5 4.5 1.5 1.5 3.5 3.5 8	Yllw Qtz 1.5 0.5 4 $\sqrt{2}$ $\overline{9}$ 9.5 3 $\overline{4}$ $\overline{4}$ 2.5 $1.5\,$ 0.5 3.5 3.5 \mathbf{I} $\overline{3}$ $\,$ $\frac{2}{5}$	Wht Qtz 2.5 1.5 5.5 3.5 3.5	Dark Qtz $\overline{5}$ $\overline{4}$ 4 3.5 1 1.5 8 $\overline{4}$ 1.5 2.5 3.5 $\overline{4}$	P/Gy Qtz 10 1.5 $\overline{9}$ 22 8	Red Qtz 6 2 3.5 $\overline{4}$	$\overline{\mathbf{V}}$ Porph 3.5 2.5 $\boldsymbol{7}$ 5 5 $\overline{3}$ $\overline{3}$ \overline{c}	Blk Chert 3.5 $1.5\,$ $\sqrt{2}$ 2.5 $4.5\,$ $\frac{2}{3}$	Red Chert	Lime- stone 3.5 2.5 $\begin{array}{c} 2 \\ 5 \end{array}$ 1.5 $\frac{2}{3}$	Drk $\mathbf{S}\mathbf{S}$ 4.5
Avg. Size - cm	6.1	3.2	3.3	3.3	3.5	10.1	3.9	3.9	2.7	1.0	2.8	4.5
Percent Total	19.2	12.5	20.2	4.8	11.5	4.8	3.9	7.7	6.7	1.0	6.7	1.0

Table A.5 . Clast count data for the M4 middle deposit.

Unit: M4 (middle)

Location:

Location: Notes:		$(11 - 7)$	M4 in gully cut north of housing terrace				
	Red SS	W/Y SS	Quartzite	V Porph	Blk Chert	Chert	Limestone
	$\overline{5}$ 29 12 2.5 14 $\overline{4}$ 6 5 $\overline{4}$ 15 $28\,$ 12.5 12 21 $\sqrt{3}$ \mathfrak{Z} 11 20 9 $\overline{4}$ 14 20 25 5 $\boldsymbol{6}$ 6 3.5 12 11 39 $\begin{array}{c} 3 \\ 2 \\ 2 \\ 4 \end{array}$ 6	1 6 \mathfrak{Z} 12 $\overline{\mathcal{I}}$ 5 $\frac{2}{8}$ $\overline{4}$ $\overline{3}$ \mathbf{I} $\overline{\mathbf{3}}$ $\mathfrak z$ $\overline{3}$ 5	3 $\overline{4}$ $\overline{4}$ \mathfrak{Z} 9 $\ensuremath{\mathfrak{Z}}$ \overline{c} $\overline{}$ 3.5 \overline{c} 5 $\overline{4}$ 6 \mathbf{I} $\overline{\mathcal{I}}$ 6 \mathfrak{Z} $\sqrt{3}$ $\sqrt{2}$ 1.5 $\!1$ 6 3 2.5 $\sqrt{2}$ 2.5 $\overline{4}$ 0.5 5 $\overline{3}$ $\overline{3}$ \overline{c} $\sqrt{2}$ $\mathbf{1}$ \overline{c} $\overline{3}$ $\overline{3}$ 3.5 $7\overline{ }$	$\overline{9}$ $\overline{4}$ 3 $\overline{4}$ $\frac{2}{9}$ 5 4 $\,$ $\,$ $\mathbf{1}$ $\sqrt{3}$ $\overline{4}$ 3.5 $\overline{3}$ 2.5 $\overline{4}$ 12 $\overline{3}$	5.5 $\overline{\mathcal{I}}$ \overline{c} $\mathbf{1}$ $\mathfrak z$ $\overline{3}$ $\overline{7}$ \overline{c} $\sqrt{3}$ $\overline{4}$ $\overline{4}$ \bf{l} $\sqrt{2}$ 2.5 1.5 $\sqrt{5}$	1 \mathbf{I} \mathfrak{Z} 2.5 $\frac{2}{2}$ $\overline{\mathcal{I}}$ \overline{c} $\overline{4}$ $\sqrt{2}$ $\frac{2}{2}$	10 13 $\begin{array}{c} 13 \\ 5 \\ 3 \\ 8 \end{array}$ 31 $\sqrt{6}$ 12 $\overline{4}$
Avg. Size - cm	10.6	4.4	3.7	4.7	3.3	2.5	10.5
Percent Total	24.3	10.2	27.2	12.2	10.9	8.2	6.8

Table A.6. Clast count data for the M4 upper deposit.

Unit: M4 (upper)

	Quartzite	Red SS	W/Y SS	V Porph	Blk Chert	Chert	\label{Limes} Limestone	Claron
	$\overline{15}$	36	$\sqrt{6}$	$\overline{3}$		$\overline{1}$	$\overline{5}$	1.5
	$\overline{9}$	$1\,8$	$\,$ 5 $\,$	4.5		1.5	10	1.5
	12	16	$\sqrt{3}$	11		$\mathbf{1}$	10	4.5
	2.5	6	6.5	11		0.5	12.5	3.5
	5	9.5	$\overline{4}$	$\overline{4}$		$\mathbf{1}$	\mathbf{I}	\mathbf{I}
	6	2.5	3	3.5		\mathbf{I}	0.5	\mathfrak{Z}
	$\,$ 8 $\,$	$\sqrt{2}$		6 1.5		1.5	9.5 1.5	1.5 $\mathbf{1}$
	8 5.5	13		2.2		$\begin{array}{c} 2 \\ 2 \end{array}$	$\overline{2}$	2.5
	$\ensuremath{\mathfrak{Z}}$	0.5 $\sqrt{5}$		4.5		0.5		3.5
	\mathbf{I}	$\,$ $\,$		$\sqrt{5}$		\mathbf{I}		
	9	5		$\overline{3}$		3.5		$\begin{array}{c} 3 \\ 3 \\ 2 \end{array}$
	$\mathfrak z$	5.5		$\overline{4}$		$\mathbf{1}$		
	\mathfrak{Z}	14		16		1.5		
	\mathbf{I}	\mathfrak{Z}		1.5		\mathbf{I}		
	$\mathbf{1}$	16		\mathfrak{Z}		2.5		
	16	16		$\ensuremath{\mathfrak{Z}}$		0.5		
	13	\mathfrak{Z}		$\overline{4}$		\mathbf{I}		
	6	2.5		$\boldsymbol{7}$		\mathbf{I}		
	2.5	5		$\sqrt{3}$		$\sqrt{2}$		
	$\overline{4}$	$15\,$		7.5				
	$\,$ 5	33		3.5				
	$\sqrt{2}$			$\overline{5}$				
	16							
	7.5							
2.5	$\sqrt{3}$							
$\mathbf{1}$	0.5							
1.5	5.5							
$\overline{4}$	0.5							
3.5	13							
\bf{l} 12	10 0.5							
$\sqrt{2}$	$\overline{4}$							
$2.5\,$	1.5							
5.5	13.5							
\overline{c}	2.5							
2.5	$\mathbf{1}$							
1.5	2.5							
6	2.5							
$\overline{4}$	$\mathbf{1}$							
1.5	2.5							
9	4.5							
0.5	$\,$ 5 $\,$							
6	$\sqrt{3}$							
9	9							
5 2.5	$\overline{4}$							
	$\mathbf{1}$							
Avg. Size -	5.5	10.7	4.6	5.1	$0.0\,$	1.4	5.8	2.4
cm				14.3		12.4	5.6	8.1
Percent	28.6	13.7	3.7		0.0			

Table A.7. Clast count data for the S4 lower deposit.

Unit: S4 (lower)

Unit:	S ₄ (upper)	
Location:	445842 E, 4081450 N	
Notes:	Big Momma S4 deposit, second-most upper cemented gravel bed.	

Table A.8. Clast count data for the S4 upper deposit.

Table A.9. Clast count data for the M7 deposit.

Unit: M7 (Johnson Mesa) Location:

Table A.10. Summary clast count data for deposits at Lees Ferry, showing percent of clast type at each site . Clast count data was grouped into the most commonly observed clast types.

	Yellow Qrtzite	Other Qrtzite	\mathbf{V} Porph	Other Chert	Red SS	Other SS	Lime- stone	Black Chert	Claron	Other Local Clasts
M ₂ upper site 1	27	$\mathbf{0}$	τ	8	18	$\overline{4}$	12	11	6	$\overline{7}$
M3y Site2	47	$\sqrt{0}$	36	6	5	$\boldsymbol{0}$	6	θ	$\boldsymbol{0}$	$\boldsymbol{0}$
M ₃ Site 2	16	$\boldsymbol{0}$	$\overline{7}$	10	41	11	13	$\overline{0}$	$\overline{0}$	$\sqrt{2}$
M ₄ Lower Site 4	15	18	5	6	25	12	8	6	1	\mathfrak{Z}
M ₄ Middle Site 5	20	25	8	$\mathbf{1}$	19	13	$\overline{7}$	$\overline{7}$	$\boldsymbol{0}$	$\boldsymbol{0}$
M ₄ Upper Site 6	27	$\boldsymbol{0}$	12	$\,$ 8 $\,$	24	10	$\overline{\mathcal{I}}$	11	$\boldsymbol{0}$	$\boldsymbol{0}$
P4 Lower Site 7	42	$\boldsymbol{0}$	14	12	14	$\overline{4}$	6	$\overline{0}$	8	θ
P4 Upper Site 8	11	20	5	6	26	19	6	8	$\boldsymbol{0}$	\sqrt{a}
M7 Site 9	26	$\boldsymbol{0}$	14	20	$\overline{4}$	18	$\overline{4}$	$\mathbf{0}$	9	6

Figure A.3. Histogram of clast-count data collected from the M3y deposit. Total counts were of 111 clasts.

Figure A.5. Histogram of clast-count data collected from the M4 lower deposit. Total counts were of 110 clasts.

Figure A.7. Histogram of clast-count data collected from the M4 upper deposit. Total counts were of 147 clasts.

Figure A.9. Histogram of clast-count data collected from the P4 upper deposit. Total counts were of 122 clasts.

Appendix B. SAND PETROLOGY DATA

Figure B.1. Map of Quaternary deposits in the Lees Ferry area, showing approximate locations where sand petrology samples were collected.

	Location Description	Notes
$GC-04-LF-S1$	Sample location and depth at GC-04-LF- OSL1. Sample taken from depth of ~138 cm from surface.	Interpreted as mainstream sediment
$GC-04-LF-S2$	\sim 1 m above basal contact of M4 deposit; same location as sed description #1, north of stop sign	Interpreted as mainstem sediment. Well- cemented unit.
$GC-04-LF-S3$	$~\sim$ 6 m above basal contact of M4 deposit; same location as sed description #1, north of stop sign.	Interpreted as a combination of mainstream and Paria sediment.
$GC-04-LF-S4$	~1 m from surface of M4 deposit; same location as sed description #1, north of stop sign.	Interpreted as a combination of mainstream and Paria sediment.
$GC-04-LF-SS$	Just east of Paria trailhead parking area; sample taken from modern flood deposits of Paria River	Sampled during high monsoon flows of Paria.
$GC-04-LF-S6$	~100 m N of Lees Ferry boat launch; sample taken from modern flood deposits of the Colorado River.	Sampled from '83 Sands.
$GC-04-LF-S7$	Soil pit, ~0.5 m depth, taken from soil horizon #4. M4 deposit, younger surface.	Interpreted as a combination of mainstream and Paria sediment.
$GC-04-LF-S8$	Soil pit, ~1.5 m depth, taken from soil horizon #6. M4 deposit, younger surface.	Interpreted as a combination of mainstream and Paria sediment.

Table B.1. Description of sample location and predicted sediment source.

	$\mathbf n$	Ortz	Kspar	Plag	Bio	Musc	Lith	Ply	Carb	Grnt	Amph	Vol	Mud	Othr
								Xst		Hvy				Obscd
$GC-04-$	336	267	28		5	$\overline{3}$	11	$\overline{2}$	$\overline{4}$	5	$\overline{4}$	$\overline{0}$	$\overline{0}$	6
$LF-S1$														
$GC-04-$	356	256	23	$\overline{2}$	13	$\overline{0}$	11	9	21	3	1	$\overline{0}$	Ω	17
$LF-S2$														
$GC-04-$	355	257	23	$\overline{0}$	11	\mathbf{I}	20	9	$\overline{7}$	7	\overline{Q}	$\overline{2}$	Ω	9
$LF-S3$														
$GC-04-$	363	280	29	$\overline{2}$	14	$\overline{0}$	5	$\overline{7}$	8	$\,$ 8 $\,$	6	$\overline{0}$	3	
LF-S4														
$GC-04-$	396	333	14	3	9	$\mathbf{1}$	8	5	13	$\overline{4}$	5	$\overline{0}$	$\overline{0}$	1
$LF-S5$														
$GC-04-$	383	287	46	6	7	$\overline{4}$	$\overline{4}$	5	14	$\overline{7}$	$\overline{3}$	$\overline{0}$	$\overline{0}$	$\overline{0}$
$LF-S6$														
$GC-04-$	366	283	30	$\overline{2}$	14	θ	13	$\overline{4}$	8	5	5	$\overline{0}$	$\mathbf{1}$	$\mathbf{1}$
LF-S7														
$GC-04-$	342	263	30	6	11	$\overline{2}$	9	$\mathbf{1}$	$\overline{7}$	9	3	$\overline{0}$	1	$\overline{0}$
$LF-S8$														

Table B.2. Raw sand petrographic data collected from deposits in the Lees Ferry area.

	$\mathbf n$	Ortz	Kspar	Plag	Bio	Musc	Lith	Ply	Carb	Grnt	Amph	Vol	Mud	Othr
								Xst		Hvy				Obsed
$GC-04-$	336	79.5	8.3	0.3	1.5	0.9	3.3	0.6	1.2	1.5	1.2	0.0	0.0	1.8
$LF-S1$														
$GC-04-$	356	71.9	6.5	0.6	3.7	0.0	3.1	2.5	5.9	0.8	0.3	0.0	0.0	4.8
$LF-S2$														
$GC-04-$	355	72.4	6.5	0.0	3.1	0.3	5.6	2.5	2.0	2.0	2.5	0.6	0.0	2.5
$LF-S3$														
$GC-04-$	363	77.1	8.0	0.6	3.9	0.0	1.4	1.9	2.2	2.2	1.7	0.0	0.8	0.3
$LF-S4$														
$GC-04-$	396	84.1	3.5	0.8	2.3	0.3	2.0	1.3	3.3	1.0	1.3	0.0	0.0	0.3
$LF-S5$														
$GC-04-$	383	74.9	12.0	1.6	1.8	1.0	1.0	1.3	3.7	1.8	0.8	0.0	0.0	0.0
$LF-S6$														
$GC-04-$	366	77.3	8.2	0.6	3.8	0.0	3.6	1.1	2.2	1.4	1.4	0.0	0.3	0.3
$LF-S7$														
$GC-04-$	342	76.9	8.8	1.8	3.2	0.6	2.6	0.3	2.0	2.6	0.9	0.0	0.3	0.0
$LF-S8$														

Table B.3. Percent composition calculated from sand petrographic data for deposits in the Lees Ferry area.

Appendix C. SEDIMENTARY COLUMNS AND DESCRIPTIONS

Figure C. I. Map of Quaternary deposits in the Lees Ferry area, showing approximate locations where sedimentary descriptions were made.

Geographic location (UTM): 4079927 N, 446830 E Stop sign, Paria bridge, sandy outcrop Outcrop description: Road cut, exposure includes a strath, but is likely not the lowest part, catching side channel... Total outcrop thickness $~1.5 \text{ m}$ Bedding ranges from IO cm to 50 cm Bedding is not generally tabular or continuous; rather it is lenticular, irregular, with notable "pockets" of coarser gravel Outcrop is likely capped (upper unit) by slopewash originating in terraces above $(M4)$ **Unit I:** Thickness: 23 cm Contact: abrupt, bedrock below Extent: partially laterally extensive downstream - pinches off in lenses Sedimentary structures: crude imbrication, rip-up clasts Texture: Matrix $-Max: c.$ sand Min: silt Grain size - Max: 17 cm Min: 2 mm Roundness - subrounded - angular Sorting: poorly sorted Color: tan-reddish buff Composition: Matrix: mostly quartz Average: vf sand Average: 1-2 cm Clasts: quartzite, porphyry, sandstone, limestone - see clast count for proportions... Cement: n/a General rock name: clast to matrix supported cobble pebble gravel Secondary features: n/a Final notes: pockets of clast supported, even amount of matrix supported; lower contact is abrupt, but not smooth. **Unit 2:** Thickness: \sim 3 m Contact: abrupt decrease in grain size, wavy Extent: extensive, generally tabular top Sedimentary structures: fine portion does not appear to have any; mostly massive; clast portion has crude imbrication Texture: Matrix $-Max: vf. sand Min: silt
Grain size - Max: -
Min: -$ Grain size $-Max: -$ Roundness - too small to see well Sorting: sand is well sorted Color: light tan grayish red Composition: Matrix: mostly quartz, very few dark minerals Clasts: -- Cement: n/a General rock name: very fine sand Secondary features: n/a Average: vf sand Average: -- Final notes: fine portion contains pockets/lenses of coarse grained material; represents multiple events instead of single flood. Coarse texture: Matrix - Max: c. sand Min: vf. snad Average: m. sand Grain size - Max: 12 cm Min: 2 mm Average: 1-2 cm $Roundness - subrounded - subangular - clast supported$ Sorting: poorly sorted Color: light brownish red with speckles of other colors Composition:

Matrix: mostly quartz

Clasts: black chert, red sandstone, quartzite, porphyry, mudstone (Moenkopi), limestone – similar proportions to clast count

Cement: n/a

General rock name: clast supported pebble gravel

Secondary features: n/a

Final notes: pockets/lenses range from 20 cm to 1 m long, from 7 cm to 20 cm thick; some are lenticular, come irregular shape.

Unit 3:

Thickness: 20 cm Contact: abrupt, blankets unit 2 Extent: laterally extensive, tabular Sedimentary structures: n/a Texture: Matrix $-Max$: c. sand Min: silt
Grain size $-Max$: $\frac{1}{2}$ Min: $\frac{1}{2}$ Grain size $-Max: -$ Roundness $-$ c. sand is subangular to angular Sorting: moderately well sorted Color: light gray buff Composition: Matrix: mostly quartz C. sand: red sandstone, mudstones Cement: well cemented carbonate (?) General rock name: silt Secondary features: n/a Average: silt Average: --

Final notes: sporadic floating pebbles and cobbles (largest is 24 cm long red sandstone flagstone)

Unit 4:

Thickness: 60 cm Contact: abrupt, erosional Extent: not laterally extensive - confined to paleo channel - channel fill Sedimentary structures: crude imbrication Tex ture: Same as coarse pockets in unit 2, but redder

Final notes: more complex than initially thought; complex filling of paleo channel (?); third dimension may solve the problem.

Unit 5:

Thickness: 72 cm Contact: relatively planar, clear, marked color difference Extent: laterally extensive, interbedded fine beds and coarse lenses 15-20 cm thick Sedimentary structures: coarse lag over unit ¾; local crude imbrication Texture: Similar to unit 2; coarser portion cobble dominated; ratio of coarse pockets to fine beds is greater and closer

to 1:1; more equal dispersion; more red color

Final notes: cobbles range from subrounded to angular; majority of larger cobbles are red sandstones.

Outcrop cap: pseudo-pavement with cactus...

Figure C.2. Sedimentary units of the M2 deposit. Visual proportions of exotic and local clasts were made in order to estimate the dominant source throughout deposition.

Geographic location (UTM): Outcrop description: Stream cut exposes outcrop Outcrop-7 m tall Lower portion - quartzite clasts >1 m thick Middle portion dominated by thick clast-supported angular side canyon red clasts Top is capped with quartzites in desert pavement Flat top, undulating (highly) erosive; irregular basal strath Cosmo surface sample GC-04-LF-408 collected from top **Unit I:** Thickness: 0-110 cm Contact: obscured, lowest contact dives underground on either side of terrace Extent: laterally extensive where exposed Sedimentary structures: crude imbrication Texture: Matrix $-Max: -V$ Min: $-Vax: -V$ Grain size $-Max$: 18 cm Min: m. sand Roundness - subrounded - subangular Sorting: poorly sorted Average: -- Average: 2-4 cm Color: multi-colored, many different colored clasts; overall light brown/gray reddish tan Composition: Matrix: --Clasts: see clast count Cement: n/a General rock name: clast supported pebble cobble gravel Secondary features: calcite/gypsum ppt. under clasts Final notes: pockets mainstem fluvial clasts. **Unit 2:** Thickness: 110-700 cm Contact: abrupt contact; change in clast type and roundness; uneven wavy contact - sharp Extent: laterally extensive, filling mainstem carved channel, overriding lower fluvial gravels and reworking some of them up into debris flow deposit Sedimentary structures: n/a Texture: $Matrix - Max: --$ Min: --Grain size - Max: 30-50 cm Min: f. sand Roundness - angular - subangular Sorting: poorly sorted Color: red sandstone / sand throughout Composition: Matrix: red sand $-$ quartz (local) Average: --Average: 5-7 cm Clasts: red ss 90% ; quartzites 1 %; white ss/limestone ?% Cement: n/a General rock name: clast supported angular pebble cobble gravel Secondary features: n/a Final notes: appears to be debris flow from side canyon, one package filled in and overrode underlying fluvial sediments; additional flows followed and were separated by episodes of more regular (organized) side stream deposition.

Unit 3:

Thickness: interfingers throughout unit 2 Contact: not always clear/distinct, but grain size change marks separation Extent: laterally extensive where exposed

Sedimentary structures: crude imbrication Texture:

Matrix $-Max: -c$ Min: --
Grain size $-Max: 12$ cm Min: f. sand Grain size $-Max: 12$ cm Roundness - angular - subangular Sorting: poorly sorted Color: red Composition: Matrix: --Clasts: sandstone $-$ red $-$ side drainage Cement: n/a Average: -- Average: 1 cm General rock name: matrix supported angular pebble gravel (some cobbles) Secondary features: n/a

Final notes: interfingers, separating possible episodic debris flow activity from side drainages; this unit and previous appear to fill channel cut by main river.

Unit 4:

Thickness: 700-800 (?) cm Contact: can't see, we dug a hole in the surface of the terrace \sim 75 cm deep Extent: --Sedimentary structures: n/a Texture: Matrix $-Max: -$ Min: --
Grain size $-Max: 5$ cm Min: f. sand Grain size $-$ Max: 5 cm Roundness - rounded - subangular Sorting: poorly sorted Color: red/tan sand Composition: Average: m. sand Average: 1-2 cm

> Matrix: mostly quartz Clasts: red ss 25%; quartzites 40 %; volcanics 10%

Cement: n/a

General rock name: matrix supported pebble gravel

Secondary features: n/a

Final notes: not very similar to unit 1; surface is littered with abundant quartzite clasts; quartzite clasts spill over terrace surface and appear to overlie debris flow unit; surface or upper layer of extensive quartzite clasts would have to have been present at one time.

Figure C.3. Sedimentary units of the M3y deposit. Visual proportions of exotic and local clasts were made in order to estimate the dominant source throughout deposition.

Geographic location (UTM): 4078918 N, 446036 E

 $M3$ deposit – culvert cut and nearby hillslopes

Outcrop description:

Lower 2.38 m from basal strath to $1st$ identifiable undisturbed unit is obscured and will be described by hillslope digging

Middle portion of outcrop exposed nicely in stream/culvert cut

Upper portions begin to be disturbed by road and culvert cut and will be described across the drainage on main M3 deposit - cosmo surface sample

Total outcrop thickness \sim 9 m

Beds are typically 0.5 to 1 m thick

Hillslope deposits blanket unconformably over truncated terrace gravels

Although this is a culvert cut, middle portion is nicely exposed, allowing good characterization of the M3

Unit I:

Thickness: 0-238 cm

Contact: obscured, covered by poor exposure

Extent: laterally extensive where exposed

Sedimentary structures: digging to find; cobbles unearthed, red sand, large red ss clasts, small quartzites; start to have better exposure at ~200 cm; crude imbrication of large red ss clasts; clast supported

Sorting: poorly sorted

Color: overall light brownish to red

Composition:

Matrix: quartz - likely locally sourced (red)

Clasts: ss 68% ; quartzite 12%; others in small amounts

Cement: n/a

General rock name: clast supported pebble cobble gravel

Secondary features: calcite/gypsum ppt.

Final notes: lower contact obscured, visible thickness of unit -0.5 m, dominated by red sand and red local clasts

Unit 2:

Thickness: 238-280 cm

Contact: distinct color contrast, although unit separated out due to apparent relative abundance of quartzite clasts Extent: laterally extensive over 2 m of exposure Sedimentary structures: crude imbrication

Texture:

Sorting: poorly sorted

Color: lighter brown, tan, some red clast (but less than previous unit) Composition:

Matrix: quartz

Clasts: red ss 20%; white ss 25%; quartzite 35&; carbonate 10%; few volcanics Cement: n/a

General rock name: clast supported pebble cobble gravel

Secondary features: calcite ppt.

Final notes: unit is not as red as lower unit; noticeable amounts of quartzites present.

Unit 3:

Thickness: 280-340 cm Contact: abrupt change to sand Extent: laterally extensive, appears on both sides of gully wall Sedimentary structures: plane beds, low angle cross strata, small scale laminations (less abundant) Texture:

Matrix $-Max: --$ Min: --Grain size $-Max$: m. sand Min: silt Roundness - subrounded - subangular Sorting: well sorted Color: tan buff Composition: Average: --
Average: f. sand --Some 2 cm pebbles--Matrix: quartz sand, iron stained grains concentrated in plane beds Clasts: -- Cement: n/a General rock name: fine sand Secondary features: iron staining Final notes: No separate bedding features present; good sand, good sed structures: OSL sample GC-05-LF-OSL10 taken from this unit. **Unit 4:** Thickness: 340-630 cm Complex unit.. Pocket of coarser cobbles observed on south end of unit - grades/interfingers with smaller pebble concentrated packages to the north Distinct pockets of channelized sand observed sporadically up throughout the unit, starting \sim 1 m above lower contact Coarse grained component: Abrupt change in grain size compared to unit 3 Laterally extensive contact over planar sand unit Sedimentary structures: moderate imbrication, subtle grading (getting finer) to the north (upstream) and up deposit Texture: Grain size - Max: 25 cm Min: f. sand Roundness - rounded - subangular Sorting: poorly sorted Color: reddish tan Composition: Matrix: -- Average: 2-3 cm Clasts: see clast count ... red ss may increase upward through deposit, although size may decrease Cement: n/a General rock name: clast supported pebble cobble gravel Secondary features: calcite/gypsum ppt. under clasts Final notes: local outsized clasts observed throughout; large clast clusters on south end - including increase in quartzite clasts. Sand component: Channelized pockets 1-2 m in length / 20-50 cm in height Scattered throughout / interfingering with coarse grained component Sedimentary structures: low angle cross strata , ripple cross laminations dominate Texture: Matrix - Max: m.-c. sand Roundness - rounded - subangular Min: silt Average: f. sand

Sorting: well sorted

Color: tan buff Composition: Matrix: quartz

Cement: n/a

General rock name: fine sand Secondary features: iron staining

Final notes: GC-05-LF-OSL11 taken from upper sand pocket

-- IT'S RAINING HARD-

Unit 5:

Thickness: 630-950 cm

Covered slope - loose description made in gully along main M3 deposit where cosmo surface sample collected

1st Hole Dug: At 640 cm - ~50 cm deep / with 30 cm exposed of undisturbed sediment No contacts visible Sed structures not apparent Texture: Matrix $-Max: -Max: -Min: -S
Grain size - Max: 6 cm Min: m. sand$ Grain size $-Max: 6 cm$ Roundness - subrounded - subangular Sorting: poorly sorted, but not extremely Color: tannish light brown, red/gray components Composition: Matrix: quartz sand Clasts: red ss 60% ; quartzites 10% Cement: n/a General rock name: matrix supported pebble gravel Secondary features: n/a Final notes: poorly exposed, but good enough... 2nd Hole Dug: At 790 cm Texture: $Matrix - Max: --$ Min: --Grain size - Max: 12 cm Min: f. sand Roundness - subrounded - subangular Sorting: poorly sorted Color: tan buff Composition: Matrix: quartz sand Average: -- Average: c. sand $-2-5$ mm Average: -- Average: 2 cm Clasts: red ss 10%; quartzites 60%; volcanics 5%; chert 2% Cement: n/a General rock name: clast supported pebble cobble gravel Secondary features: n/a

Final notes: has appearance of good mainstem sand, lots of quartzites

Figure C.4. Sedimentary units of the M3 deposit. Visual proportions of exotic and local clasts were made in order to estimate the dominant source throughout deposition .

Geographic location (UTM): 407097 N, 445507 E S3 paludal deposit of Kaufman et al... Outcrop description: Stream/gully cut Well exposed lower portion on south side of road Can trace bedding near top of lower portion to bedding at lower portion of top section across road (north) Deposit is \sim 10 m thick Encompasses carbonate spring deposits and local debris flows Bedding ranges from 20 to 1 m **Unit** I: Thickness: 62 cm Contact: sharp with bedrock Extent: apparently continuous but obscured by large slump Sedimentary structures: low angle cross strata Texture: Matrix - Max: c. sand Min: f. snad Grain size - Max: 13 cm Min: 2 cm Roundness - subrounded - angular Sorting: poorly sorted - matrix supported Color: yellowish red/tan Composition: Matrix: grains fizz, some quartz Clasts: -- Cement: n/a General rock name: --Secondary features: iron staining, gypsiferous nodules Final notes: clasts floating in matrix **Unit 2:** Thickness: 25 cm Contact: clear contact; color change; loss of clasts Extent: extensive, relatively tabular/wavy Average: m. sand Average: 5 cm Sedimentary structures: planar bedding, accentuated by iron staining Texture: Matrix - Max: m. sand Min: silt Grain size $-Max$: $--$ Min: $--$ Roundness ---Sorting: moderately well sorted Color: yellowish buff Composition: Matrix: mostly quartz, some grains fizz Clasts: -- Cement: n/a General rock name: -- Secondary features: iron staining along bedding planes Final notes: -- **Unit 3:** Average: vf. sand Average: --

Thickness: 52 cm Contact: unclear contact; gradual color change Extent: same as unit 2 Sedimentary structures: wavy bedding near top Texture: Matrix - Max: c. sand Min: silt

Average: vf. sand

Grain size $-Max$: $--$ Min: $--$ Roundness ---Sorting: fairly poorly sorted Color: light grayish with light rust streaks Composition: Matrix: mostly quartz, some grains fizz Clasts: -- Cement: n/a General rock name: -- Secondary features: iron staining Final notes: wavy beds, possible grain size reduction at bed contacts

Unit 4:

Thickness: 118 cm **Same as unit 3 - differences noted** Contact: abrupt; color change to yellow Color: Dijon mustard

Texture: mostly fine sand with significant silt and finely disseminated (FDC) carbonate; few coarse grains Secondary features: some vertical iron staining and gypsiferous crystals, iron staining more significant (similar wavy beds as unit 3).

Unit 5:

opposite side of gully for convenience Thickness: 75 cm Contact: abrupt and wavy Extent: laterally extensive Sedimentary structures: faint wavy bedding Texture: Matrix - Max: c. sand Min: silt
Grain size - Max: 4 cm Min: 0.5 cm Grain size $-Max: 4 cm$ Roundness - well rounded to angular Sorting: moderately: well sorted Average: f. sand Average: 2.5 cm Color: weathered color - reddish buff; fresh color - white grayish dijon mustard Composition: Matrix: mostly quartz, FDC Clasts: quartzite, chert, sandstone, siltstone (all from local rocks) Cement: n/a General rock name: Secondary features: gypsum crystals, FDC Final notes: -- -- floating in matrix--

Unit 6:

Thickness: \sim 300 cm Contact: gradual, not well defined Extent: laterally extensive Sedimentary structures: same as unit 5 Texture: Matrix - Max: m. sand Min: vf. sand Grain size $-$ same as unit 5 Roundness --Sorting: moderately well sorted Color: medium reddish brown Composition: Matrix: same as unit 5 Average: f. sand Clasts: same as unit 5; clasts appear to cluster along pseudo-bedding planes Cementy punky stuff: silt to tine sand; well sorted; carbonate cemented; distributed in discontinuous planes parallel to bedding

Cement: n/a

General rock name: -- Secondary features: carbonate cement in discontinuous layers; cemented layers seem to increase towards top - become more continuous Final notes: jump across road...

Unit 7:

Thickness: 40 cm Contact: obscured/covered - best guess across road Extent: laterally extensive, but probably pinches out Sedimentary structures: none, maybe some crude imbrication Texture: Matrix - Max: vc. sand Min: vf. snad Grain size - Max: 25 cm Min: 2 cm Roundness - angular to subangular Sorting: poorly sorted Color: reddish brown Composition: Matrix: mostly quartz Clasts: -- array of local clasts Cement: compacted General rock name: -- Secondary features: n/a Final notes: likely debris flow Average: f. sand Average: 5 cm

Unit 8:

Thickness: 45 cm Contact: clear contact Extent: laterally extensive, planar in an interfingering way Sedimentary structures: faint bedding planes Texture: Matrix – Max: vc. sand Min: vf. Sand Grain size $-Max: Roundness - -$ Sorting: moderately well sorted Color: medium reddish brown Composition: Min:-- Average: m. sand Average: -- Matrix: mostly quartz, small grains of local siltstones Clasts: -- Cement: compacted General rock name: --Secondary features: n/a Final notes: occasional floating clasts 0.5 to I cm

Unit 9:

Thickness: 83 cm Contact: clear contact; noted by grain size change Extent: tabular, laterally extensive Sedimentary structures: good imbrication in pebbles; crude imbrication in cobbles; coarsens upward at base Texture: Matrix $-$ same as unit 8 Grain size - Max: 17 cm Min: 2 mm Roundness - angular to subangular Sorting: poorly sorted - clast supported Color: varies Composition: Average: 5 cm

Matrix: same as unit 8 Clasts: array of local clasts, primarily red ss Cement: n/a

General rock name: -- Secondary features: n/a Final notes: good candidate for debris flow

Unit 10:

Thickness: 55 cm Contact: change in grain size, semi-abrupt Extent: laterally extensive, tabular Sedimentary structures: moderate imbrication in pebbles; planar bedding at bottom Texture: Matrix $-$ same as unit 8 Grain size $-$ Max: 6 cm Min: 2 mm Roundness - angular Sorting: poorly sorted Color: medium reddish brown Composition: Matrix: same as unit 8 Average: 2 cm Clasts: local clasts, especially siltstones (Moenkopi) Cement: compacted (for all compacted units there may be minor carbonate cementing) General rock name: -- Secondary features: n/a Final notes: clasts concentrated along bedding planes, 1 to 3 cm thick beds with clasts; clast beds spaced 2 to 10 cm near bottom half, upper half mostly massive.

Unit 11:

Thickness: 74 cm Contact: clear contact; change in grain size Extent: lenticular, channel shaped -30 m in lateral extent Sedimentary structures: crude imbrication in large boulders Tex ture: Matrix $-$ same as unit 7 Grain size $-Max: 80 \text{ cm}$ Min: 0.5 cm Roundness - angular Sorting: very poorly sorted Color: medium reddish brown Composition: Matrix: same as unit 7 Average: 18 cm Clasts: array of local clasts including Navajo Sandstone Cement: compacted General rock name: -- Secondary features: n/a Final notes: big boulders clast supported with finer matrix in between; big boulders disappear down paleoslope.

Unit 12:

Thickness: 85 cm **Same as unit $10**$

Unit 13:

 $*$ *moved ~20 m to NW, 10 m E of GC-04-LF-OSL16 sampling site** Thickness: 92 cm Contact: sharp, noted in OSL site description; based on color change Extent: thickness varies laterally Sedimentary structures: n/a Texture: Matrix – Max: f. sand Min: silt Grain size $-Max: -$ Roundness --- $Min: -$ Average: vf. sand Average: --

Sorting: fairly well sorted Color: light yellowish buff Composition: Matrix: mostly quartz, some FDC Clasts: -- Cement: n/a General rock name: -- Secondary features: some iron staining in upper parts; irregular shaped carbonate pieces (cm scale); wavy shaped to

jagged

Final notes: bedding appears draped in mound shape towards top.

Unit 14:

Thickness: 14 cm to only a few cm Contact: sharp - wavy draping over underlying unit Extent: wavy, capping several mounds in area Sedimentary structures: n/a Texture: Matrix $-Max: f.$ sand Min: silt
Grain size $-Max: -$ Min: --Grain size $-Max: -$ Roundness ---Sorting: moderately well sorted Color: light reddish brown yellow Composition: Matrix: mostly carbonate with some quartz grains Clasts: -- Cement: carbonate General rock name: --Secondary features: iron staining Final notes: punky texture... from rain erosion (?) Average: f. sand Average: --

Unit 15:

Thickness: 55 cm, may vary throughout deposit Contact: obscured, not clear Extent: only found on highest points -3 or 4 mounds Sedimentary structures: n/a Texture: Matrix - Max: vf. sand Min: silt Grain size $-Max: -Min: -$ Roundness --Sorting: fairly well sorted Color: light whitish gray Composition: Matrix: quartz sand and FDC Clasts: -- Cement: n/a General rock name: --Secondary features: maybe iron staining near top Average: silt Average: -- Final notes: unit of Kaufman et al. snails; capped by unit similar to unit 14 in a draping manner...

Figure C.5. Sedimentary units of the S3 deposit. Visual proportions of exotic and local clasts were made in order to estimate the dominant source throughout deposition.

Geographic location (UTM):

Thick M4 deposit near the stop sign at Paria River bridge

Outcrop description:

Bottom -4 m forms cliff (well cemented); upper -6 m is less consolidated and more recessive Bed thickness is -0.5 to 1 m, with some apparently thicker beds (-2 m) towards top Bedding appears laterally extensive, with interspersed channel form (sandy) within tabular units

Unit I:

Thickness: varies from 0.5 to 1.1 m Contact: abrupt lower bedrock strath contact Extent: laterally extensive, but varying thickness Sedimentary structures: very slight imbrication, rip-up clasts (Moenkopi) Texture: Grain size - Max: 90 cm Min: m. sand Roundness - subrounded to subangular Sorting: poorly sorted Color: 7.5 YR 7/4; varying clast colors Composition: Matrix: quartz grains Average: 3-4 cm Clasts: quartzite 15%; red and white ss 55%; carbonate 10%; Moenkopi 5% Cement: carbonate General rock name: clast supported bouldery cobble gravel Secondary features: carbonate cement, carbonate pendants on underside of clasts

Final notes: appears to be a mixture of local and far-traveled clasts; very large clasts (bou lders are unique to this lower unit - compared to next few units).

Unit 2:

Thickness: 55 cm, although varies up to 75 cm on bottom Contact: sharp - irregular Extent: laterally extensive with varying thicknesses on bottom Sedimentary structures: slight imbrication Texture: Grain size - Max: 20 cm Min: m. sand Roundness - rounded to subangular Sorting: --Color: 5 YR 7/3; varying clast colors Composition: Matrix: quartz grains Average: 0.5-1 cm Clasts: ss 35%; quartzite 30%; volcanics 5%; carbonate 13 %; Moenkopi 3% (more worked than unit 1) Cement: carbonate

General rock name: clast supported cobble gravel

Secondary features: calcite cement

Final notes: larger clasts are sandstones; less matrix than unit 1; channel cut of sandy gravel with cross strata.

Unit 3:

Thickness: ~85 cm Contact: abrupt and laterally planar Extent: laterally extensive Sedimentary structures: imbrication (especially in smaller clasts) Texture: Grain size - Max: 10 cm Min: m. sand Average: 0.5 cm Roundness - rounded to subangular Sorting: poorly sorted, but clasts are generally unimodal Color: 7.5 YR 7/3; varying clast colors, but white seems to dominate Composition:

Matrix: quartz grains Clasts: ss 5%; quartzite 40%; volcanics 1%; carbonate/chert 42% (chert coming out of Kaibab ?) Cement: carbonate General rock name: clast supported pebble gravel Secondary features: calcite cement (well developed) Final notes: finer grained than previously defined units; dominantly white (chert/carbonates??); lighter colored matrix.

Unit 4:

Thickness: 90 cm Contact: gradational Extent: laterally extensive, tabular unit Sedimentary structures: massive Texture: Grain size - Max: 40 cm Min: m. sand Roundness - rounded to subangular Sorting: -- Color: 5 YR 7/3 Composition: Matrix: quartz grains Average: 0.5-1 cm Clasts: ss 53%; quartzite 30%; volcanics 4%; Moenkopi and local clasts 8% Cement: carbonate General rock name: clast supported cobble gravel Secondary features: calcite cement Final notes: similar in appearance to unit 2, but sandstone clasts are larger.

Unit 5:

Thickness: 190 cm Contact: gradational, but tabular and continuous Extent: laterally extensive Sedimentary structures: massive Texture: Grain size - Max: 15 cm Min: m. sand Roundness - rounded to subangular Sorting: poorly sorted Color: matrix -7.5 YR $6/4$: clasts $-$ look at unit 1 Composition: Matrix: quartz grains Average: I cm Clasts: ss 23%; quartzite 37%; volcanics 6%; chert 2%; weathered quartzite(?) 13% Cement: carbonate General rock name: clast supported pebble cobble gravel Secondary features: calcite cement, not as well developed

Final notes: more recessive weathering profile; in between units 5 and 6 there appears to be a small, \sim 6 cm thick, channel fill, but will be included in next unit.

Unit 6:

Thickness: 40 cm Contact: marked by the abrupt appearance of cobbles Extent: laterally extensive, appears tabular except for channel form at the base Sedimentary structures: slight imbrication, some channel forms at the base Texture: Grain size - Max: 24 cm Min: m. sand Average: 0.25-1.5 cm Roundness - subrounded to subangular Sorting: poorly sorted, with zones of slightly better poorly sorted Color: matrix - 10 YR 7/2; clasts - look at unit 1 (some yellow quartzites) Composition: Matrix: quartz grains

Clasts: ss 25%; quartzite 35%; volcanics 17%; chert 3%; weathered quartzite $(?)\%$; limestone 10% Cement: carbonate (not as well developed)

General rock name: clast supported cobble pebble gravel Secondary features: slight calcite cement Final notes: has channel forms and is cut by the upper unit.

Unit 7:

Thickness: 105 cm Contact: abrupt appearance of boulders Extent: laterally extensive Sedimentary structures: n/a Texture: Grain size $-Max: 83 cm$ Min: f. sand Roundness - rounded to subangular Sorting: extremely poorly sorted Color: matrix -5 YR $6/6$; clasts - reddish sandstones Composition: Matrix: quartz grains Clasts: red ss 61% ; quartzite 27% Cement: n/a General rock name: clast supported cobble boulder gravel Secondary features: n/a

Final notes: big boulders.

Unit 8:

Thickness: 165 cm Contact: obscured Extent: laterally extensive, appears tabular Sedimentary structures: n/a Tex ture:

Grain size - Max: 42 cm Min: f. sand Average: 1.5-2 cm Roundness - subrounded to subangular

Average: 36 cm

Sorting: poorly sorted, with lenses of coarser cobbles and boulders between more gravelly portions Color: matrix -2.5 YR $5/8$; clasts $-$ same, more red ss Composition:

Matrix: quartz grains Clasts: ss 56%; quartzite 28%; volcanics 10% Cement: n/a General rock name: clast supported boulder cobble gravel Secondary features: n/a Final notes: getting sandier, vegetation increasing, less carbonates, still steep slope.

Unit 9:

Thickness: 250 cm Contact: gradual, marked by angular appearance Extent: laterally extensive Sedimentary structures: n/a Texture: Grain size $-Max: 76 cm$ Min: f. sand Roundness - subrounded to angular Sorting: poorly sorted Color: matrix -2.5 YR $5/6$; clasts $-$ same as all Composition: Matrix: quartz grains Clasts: ss 68%; quartzite 23%; volcanics 3% Cement: n/a General rock name: clast supported boulder cobble gravel Secondary features: n/a Average: 13 cm

Final notes: angular appearance, appears to have colluvial component, but channels with locally rounded gravels are present and may be reworking locally delivered cobbles and boulders of sandstone.

Unit 10:

Thickness: 158 cm Similar to units 6 and 8 But max grain size is 12 cm, min in fine sand, average is 0.5-1 cm Color is redder, no vegetation, laterally extensive

Unit 11:

Thickness: 160 cm Contact: planar, abrupt appearance of predominantly cobbles Extent: laterally extensive, but slightly lenticular Sedimentary structures: faint / slight imbrication Texture: Grain size $-Max: 42 \text{ cm}$ Min: f. sand Roundness - subrounded to subangular Sorting: poorly sorted Color: matrix - red; clasts - same as all Composition: Matrix: quartz grains Clasts: ss 52%; quartzite 35%; volcanics 5% Cement: n/a General rock name: clast supported cobble gravel Secondary features: n/a Average: 2 cm

Final notes: still has red appearance, more sandstones, less quartzites than unit 10, smaller quartzites but sandstones are red and subangular to subrounded.

Unit 12:

Thickness: 124 cm Contact: gradational with the marked appearance of more abundant gravel sized quartzites Extent: laterally extensive, sitting on top of boulders Sedimentary structures: slight imbrication, nothing definitive, hinted at in structure Texture: Grain size - Max: 13 cm Min: f. sand Average: 0.5-1 cm Roundness - subrounded Sorting: poorly sorted Color: matrix $-$ reddish, but lighter than units 8 through 10; clasts $-$ same as all Composition: Matrix: quartz grains Clasts: -- ss 68% ; quartzite 23%; volcanics 3% -- $(?)$ Cement: n/a General rock name: clast supported cobble gravel

Secondary features: n/a

Final notes: got dark, working too hard... time to sleep for the day.

Figure C.6. Sedimentary units of the M4 deposit - M4-1 description. Visual proportions of exotic and local clasts were made in order to estimate the dominant source throughout deposition.

Geographic location (UTM):

Just south of stop sign, below housing terrace in south side of gully leading directly up to houses M4 deposit, including the lowest portion of the deposit, near lowest surveyed strath

Outcrop description:

Fairly decent gully cut, exposing \sim 3/4 of the deposit very well

Total thickness is \sim 20 m

Bedding ranges from 10 cm to > 1 m

Bedding is not always tabular or continuous, several lenticular/channel shaped beds included

Unit I:

Thickness: 70 cm Contact: abrupt contact with bedrock strath Extent: apparently extensive, although dives quickly underground Sedimentary structures: crude imbrication Texture: Matrix - Max: vc. sand Min: vf. sand Grain size $-Max: 19 cm$ Min: 2 mm Roundness - subrounded to rounded Average: m. sand Average: 1.5 cm Sorting: moderately to fairly poorly sorted – generally clast supported

Color: matrix - light reddish brown gray; clasts - assorted

Composition:

Matrix: mostly quartz

Clasts: red ss 22%; quartzite 12%; porphyry 3%; black chert 7%; other ss 12%

Cement: carbonate in some places General rock name: --

Secondary features: calcite cement

Final notes: unit is more obscured than most, although some limited good exposure exists; matrix is very sandy; has a colorful salt and pepper look; there are actually many ≤ 0.5 cm clasts throughout matrix.

Unit 2:

Thickness: 110 cm Contact: difficult to see, slight decrease in grain size, specifically a loss of large clasts Extent: apparently extensive Sedimentary structures: crude imbrication in larger clasts Texture: Matrix - Max: vc. sand Min: vf. sand Grain size - Max: 12 cm Min: 2 mm Roundness - subrounded to rounded Average: m. sand Average: 1-2 cm Sorting: overall poorly sorted $-$ clast and matrix supported areas throughout

Color: matrix – light grayish brown; clasts – colorful salt and pepper appearance Composition:

Matrix: mostly quartz

Clasts: red ss 8%; quartzite 20%; porphyry 5%; black chert 8%; other ss 12%; limestone 6% Cement: n/a

General rock name: --

Secondary features: n/a

Final notes: very sandy unit with pebbles throughout; quartzites are present but do not visibly dominate.

Unit 3:

Thickness: 28 cm Contact: fairly abrupt, better cemented, sandy lens separating Extent: not continuous, thickens down paleoslope Sedimentary structures: planar cross strata, pebble lags on bedding planes, crude imbrication Texture:

Matrix - Max: m. sand Min: vf. sand Average: f. sand

Sorting: zones or beds of fairly well (sand) and fairly poorly (pebbles) sorted; matrix and clast supported zones Color: matrix $-$ light reddish brown; clasts $-$ speckled appearance Composition:

Matrix: mostly quartz

Clasts: red ss 9%; quartzite 18%; porphyry 3%; black chert 6%; other ss 19%; limestone 3% Cement: carbonate

General rock name: --

Secondary features: calcite cement

Final notes: lowest ~10 cm is composed of well sorted sand lens that pinches out; upper portion contains pebbles that are generally sorted on bedding planes, or form planar bedding planes throughout.

Unit 4:

Thickness: 26 cm Contact: fairly sharp Extent: lenticular, not extensive, but observed in most of outcrop $(\sim 3m \arccos)$ Sedimentary structures: crude imbrication in larger clasts Texture: Matrix $-Max$: vc. sand Min: vf. sand Grain size $-Max: 1.5 cm$ Min: 2 mm Roundness - subrounded to subangular Sorting: overall moderately sorted - matrix supported Color: matrix $-$ light grayish brown; clasts $-$ --Composition: Average: m. sand Average: 0.4 cm

Matrix: mostly quartz

Clasts: red ss, quartzite, black chert, limestone $-$ all less than 5% of deposit Cement: n/a

General rock name: --

Secondary features: some gypsum coatings Final notes: very sandy unit, almost no clasts.

Unit 5:

Thickness: 153 cm Contact: fairly abrupt, appearance of large clasts Extent: generally tabular, extensive Sedimentary structures: good imbrication Texture: Matrix - Max: vc. sand Min: vf. sand Grain size $-Max$: 48 cm Min: 2 mm Roundness - subroundcd to rounded Average: m. sand Average: $3-4$ cm (red ss -9 cm) Sorting: overall poorly sorted - clast and matrix supported areas throughout Color: matrix $-$ light grayish brown; clasts $-$ Composition: Matrix: mostly quartz Clasts: see clast count Cement: some carbonate General rock name: -- Secondary features: calcite cement

Final notes: interfingering of definite mainstem river gravels with clusters of large red sandstone cobbles and boulders; few lenses of generally small pebbles and sand.

Note: continually moving semi-laterally up gully

Unit 6:

Thickness: 45 cm Contact: fairly abrupt, marked by dark organic and iron staining beds Extent: continuous in outcrop

Sedimentary structures: low angle beds, planar beds Texture:

> Matrix - Max: c. sand Min: f. sand
Grain size - Max: 12 cm Min: 2 mm Roundness - subrounded to subangular

Average: m. sand Average: bimodal $- 1$) subang. 0.5 cm; 2) subround. 4.5 cm

Sorting: overall fairly poorly sorted – clast supported (mostly small clasts)

Color: matrix - gravish buff with dark gray and rust

Composition:

Matrix: mostly quartz

Clasts: red ss 8% ; quartzite 20%; porphyry 5%; black chert 8% ; other ss 12%; limestone 6% Cement: n/a General rock name: --

Secondary features: paleosol (?)

Final notes: if paleosol, would have been part of previous unit.

Unit 7:

Thickness: 19 cm (variable due to lens shape) Contact: faintish, marked by decrease in grain size Extent: lenticular, pinches out Sedimentary structures: fining upward, plane beds at top where sand dominates Texture: Matrix - Max: vc. sand Min: f. sand
Grain size - Max: -- Min: --Grain size $-Max: -$ Roundness - subrounded to subangular Average: m. sand Average: -- Sorting: fairly poorly sorted, becomes fairly well sorted at top where sand dominates $(\sim 7 \text{cm})$ Color: matrix - light grayish buff Composition: Matrix: mostly quartz Clasts: red ss 8%; quartzite 20%; porphyry 5%; black chert 8%; other ss 12%; limestone 6% Cement: compacted General rock name: --Secondary features: n/a Final notes: there are some few pebbles, less than 1 cm near bottom of unit.

Unit 8:

Thickness: 378 cm Contact: abrupt change in grain size Extent: apparently tabular and extensive Sedimentary structures: imbrication, planar cross strata in sand lenses Sand lens: vf. to c. sand, average: m. sand; subrounded, moderately well sorted; mostly quartz; occasional small pebbles < I cm Majority of deposit: Texture: Matrix - Max: vc. sand Min: vf. sand Average: m. sand Grain size - Max: 36 cm Min: 2 mm Average: 2-3 cm (red sandstone - 13 cm)

Roundness - subrounded to rounded Sorting: overall poorly sorted - clast supported Color: matrix - light grayish buff; clasts - speckled with light gray buff reddish

Composition:

Matrix: mostly quartz

Clasts: see clast count. ..

Cement: carbonate in some places, especially near top

General rock name: --

Secondary features: calcite cement, some gypsum

Final notes: sand lenses throughout, although generally spaced every 1.2 m or so - some as close as 20 cm; there are some areas of slightly lower average grain size that are still clast supported, but much sandier; interfingered red ss boulders appear in a few places (spaced \sim 1.5 m) and are often clustered together on a single bed; some sand lenses are small, others are quite extensive.

Unit 9:

Thickness: 160 cm

Contact: fairly easy to see, appearances of red sandstone boulders Extent: laterally extensive, obscured on other side of ravine to north Sedimentary structures: crude imbrication, stacked boulders Texture:

> Matrix - Max: vc. sand Min: vf. sand Grain size - Max: 43 cm Min: I cm Roundness - subrounded to subangular

Average: m. sand Average: 7-8 cm

Sorting: overall poorly sorted - clast supported - big boulders touching, filled with fines and pebbles Color: sandstones give more reddish appearance

Composition:

Matrix: quartz and local clast chips

Clasts: red ss 50%; quartzite 13% ; other ss 20%; others 10%

Cement: n/a

General rock name: --

Secondary features: n/a

Final notes: not likely one event; lenses of more organized clasts that are smaller; boulders generally stacked on single planar surface.

Unit 10:

Thickness: \sim 300 cm Contact: good, clear, appearance of fluvial dominated gravels Extent: extensive, tabular Sedimentary structures: imbrication, planar cross strata in sand lenses Sand lens: vf. to c. sand, average: m. sand; subrounded, moderately well sorted; mostly quartz; occasional small pebbles < I cm Majority of deposit: Texture: Matrix - Max: vc. sand Min: vf. sand Average: m. sand
Grain size - Max: 36 cm Min: 2 mm Average: 2-3 cm (Average: $2-3$ cm (red sandstone -13 cm)

Roundness - subrounded to rounded Sorting: overall poorly sorted - clast supported

Color: matrix $-$ light grayish buff; clasts $-$ speckled with light gray buff reddish Composition:

Matrix: mostly quartz Clasts: see clast count..

Cement: carbonate in some places, especially near top

General rock name: --

Secondary features: calcite cement, some gypsum

Final notes: sand lenses throughout, although generally spaced every 1.2 m or so - some as close as 20 cm; there are some areas of slightly lower average grain size that are still clast supported, but much sandier; interfingered red ss boulders appear in a few places (spaced - 1.5 m) and are often clustered together on a single bed; some sand lenses are small, others are quite extensive.

Unit II:

Thickness: $~260$ cm **Same as unit 8**

Unit 12:

Thickness: ~350 cm **Same as unit 8 (approximately)** Covered slope, best estimate...

Unit 13:

Thickness: ~200 cm **Same as unit 9**

Unit 14: Thickness: - 130 cm **Same as unit 8**

Figure C.7. Sedimentary units of the M4 deposit, M4-2 description. Visual proportions of exotic and local clasts were made in order to estimate the dominant source throughout deposition. Sedimentary facies interpretations arc also included.

Geographic location (UTM): 4079938 N, 446458 E Approximately 300 m north of Ranger Station Prominent sandy road cut on west side of road Outcrop description: Extremely large boulder truncates sand and gravel lenses Unit is generally light tan/yellow with local reddish clasts throughout, especially in upper 0.5 m Lenticular and tabular beds present

Unit I:

Thickness: 75 cm Contact: obscured Extent: extensive, tabular Sedimentary structures: n/a Texture: Grain size - Max: 26 cm Min: silt Average: 0.5-1 cm Roundness - subrounded to angular Sorting: moderately sorted Color: tan/red/brown/gray - overall tannish red Composition: Matrix: mostly quartz Clasts: red ss 50% ; quartzite 15%; chert 15%; volcanics 2%; carbonate 5% Cement: n/a General rock name: clast supported pebble gravel with overlying cobble lag Secondary features: n/a

Final notes: cobble lag on upper boundary - red sandstone.

Unit 2:

Thickness: 25 cm Contact: undulates around boulder lag of unit I - sharp Extent: not extensive, channel scour/lenticular Sedimentary structures: low angle cross strata with some 1-2 mm clast lags on paleo-surfaces Texture: Grain size - Max: 5 cm Min: f. sand Roundness - subrounded to angular Sorting: moderately sorted – grading Color: tan to reddish brown, lighter at top Composition: Matrix: mostly quartz Clasts: red ss 30%; quartzite 15%; chert 30% Cement: n/a General rock name: coarse sand Secondary features: n/a Average: 0.2 cm

Final notes: unit coarsens upward in three apparently distinct packages.

Unit 3:

Thickness: 40 cm Contact: gradual, sharp in some places Extent: fairly extensive, tabular Sedimentary structures: n/a Texture: Grain size - Max: 21 cm Min: silt Roundness - subrounded to subangular Sorting: poorly sorted Color: tannish red - more red than other units Composition:

Average: 4 cm

Matrix: mostly quartz Clasts: red ss70%; quartzite 10% ; chert 10% Cement: n/a General rock name: clast supported pebble cobble gravel Secondary features: n/a Final notes: no real imbrication, but apparent preferred orientation.

Unit 4:

Thickness: 50 cm Contact: sharp Extent: channel shaped, scour Sedimentary structures: small scale pebble $(1-2 \text{ mm})$ overlying finer grains (no cross bedding) Texture: Grain size $-$ Max: 7 cm Min: silt Roundness - subrounded to subangular Sorting: somewhat poorly sorted Color: light tannish red Composition: Same as unit 2... Cement: n/a General rock name: matrix supported pebble gravel Secondary features: n/a Average: <1 cm Final notes: generally harder than previous units; apparent channel form.

Unit 5:

Thickness: 80 cm Contact: sharp/abrupt Extent: tabular, extensive Sedimentary structures: laminations apparent due to weathering; 20 cm scale cross strata Texture: Grain size - Max: f. sand Min: clay Roundness - subrounded to subangular Sorting: fairly well sorted Color: tan Composition: Matrix: quartz, clay Cement: n/a General rock name: matrix supported pebble gravel Secondary features: clay formation (?) Average: silt Final notes: mud crack weathering pattern - external pattern.

Unit 6:

Thickness: 95 cm Contact: abrupt change in grain size Extent: not extensive, thins to the south Sedimentary structures: n/a Texture: Grain size - Max: 30 cm Min: vf. sand Average: 4 cm Roundness - subrounded to subangular Sorting: poorly sorted Color: light reddish tan Composition: Matrix: mostly quartz Clasts: red ss70%; quartzite 10%; chert 10%; volcanics 2% Cement: n/a General rock name: clast supported pebble cobble gravel Secondary features: n/a Final notes: better rounded than previous units; local fine grained section void of bigger clasts.

Unit 7:

Thickness: 81 cm **Same as unit 5** But with a slightly coarser (0.5 cm) 20 cm thick package separating the two.

Unit 8:

Thickness: 105 cm lnterbedded unit of the description to follow with material similar to unit 5 Contact: abrupt at each boundary Extent: laterally extensive Sedimentary structures: crude imbrication Texture: Grain size - Max: 4 cm Min: f. sand Roundness - subrounded to subangular Sorting: fairly poorly sorted Color: light red Composition: Matrix: mostly quartz Clasts: same as unit 3 Cement: n/a General rock name: clast supported pebble gravel Secondary features: n/a Final notes: -- Average: 0.5 cm

Unit 9:

Thickness: 120 cm Contact: gradual Extent: laterally extensive Sedimentary structures: crude imbrication Texture: Grain size $-Max$: 20 cm Min: f. sand Roundness - subroundcd to subangular Sorting: poorly sorted Color: multi colored (yellow, black, white, red, brown) Composition: Matrix: mostly quartz Clasts: red ss 30%; quartzite 40%; chert 10% Cement: n/a General rock name: clast supported cobble pebble gravel Secondary features: n/a Average: 3 cm Final notes: better: more quartzite than observed in other units in this deposit.

Figure C.8. Sedimentary units of the upper M4 deposit, M4-3 description. Visual proportions of exotic and local clasts were made in order to estimate the dominant source throughout deposition.

Geographic location (UTM): 4081455N, 445846 E Big P4 deposit upstream along the Paria River Outcrop description: Some good exposures throughout deposit, particularly where units are cemented Sand units are present in either lenses or pockets (good for OSL samples throughout) Total outcrop thickness \sim 20 m Not sure whether this is a P3 and P4, or just a P4 with younger and older surfaces **Unit I:** Thickness: 120 cm Contact: abrupt and sharp of bedrock Extent: laterally extensive Sedimentary structures: excellent imbrication, plane beds, faint ripples Texture: Grain size - Max: 35 cm Min: f. sand Roundness - rounded to subangular Sorting: poorly sorted with pockets of well sorted sand Color: light reddish tan Composition: Matrix: mostly quartz Clasts: see clast count.. Cement: carbonate Average: 2 cm General rock name: clast supported / matrix supported cobble pebble gravel Secondary features: calcite cement Final notes: large unit, extensive, well cemented, pockets of sand. **Unit 2:** Thickness: 275 cm Contact: gradual, marked by a decrease in grain size Extent: laterally extensive Sedimentary structures: localized crude imbrication Texture: Grain size $-Max: \geq 1m$ Min: f. sand Average: pebbles -1.5 cm; boulders $-15-30$ cm Roundness - subrounded to subangular Sorting: poorly sorted Color: tannish red, but huge variation between localized clasts Composition: Matrix: mostly quartz Clasts: boulders - local lithologies; pebbles - same as unit 1 Cement: carbonate General rock name: clast supported boulder cobble pebble gravel Secondary features: calcite cement Final notes: boulders come loose easily ... smashed Ben's leg... scary; unit not as well exposed as unit 1.

Unit 3:

Thickness: 210 cm **Mixture of unit I and 2** With gradual increase in sand towards top.

Unit 4:

Thickness: 200 cm Contact: obscured, but noted by significant increase in sand Extent: not laterally extensive, but large pocket of sand Sedimentary structures: small scale cross strata, few coarsening upward sequences, massive portions Texture: Grain size $-Max$: m. sand Min: f. sand Roundness – rounded to subangular Sorting: well sorted Color: light tannish gray Composition: Matrix: mostly quartz Clasts: -- Cement: n/a General rock name: fine sand Secondary features: n/a Average: f. sand Final notes: large unit of Paria River sand, GC-05-LF-OSL3 collected from this unit; some few pebbles observed

within massive matrix.

Unit 5:

Thickness: 155 cm Contact: sharp, increase in grain size Extent: laterally extensive Sedimentary structures: good imbrication, cross bedding in sand lenses Texture: Grain size $-Max: 23 cm$ Min: f. sand Roundness - rounded to subangular Sorting: poorly sorted with pockets of well sorted sand Color: tannish light red Composition: Matrix: mostly quartz Clasts: more red sandstones than unit I Cement: carbonate General rock name: clast supported cobble pebble gravel Secondary features: calcite cement Average: 2-3 cm

Final notes: finer at base, gets coarser towards middle, and then finer at top where sand lenses appear.

Unit 6:

Thickness: 295 cm Covered slope - sand weathering unit **Same as unit 4**

Unit 7:

Thickness: 145 Contact: fairly abrupt Unit is rich in red sandstone cobbles (-15 cm) Clast supported pebble cobble gravel Crude imbrication Pebble pockets are present (similar to other units below) Matrix is cemented

Unit 8:

Thickness: 270 cm **Same as unit 5, although lacking the abundant red sandstones** Quartzites (rounded) appear to dominate and some sand pockets ore present Not as well cemented, but still has carbonate cement

Unit 9:

Thickness: 170 cm GC-05-LF-OSL2 collected from this unit Medium scale bedding (5-10 cm) cross bedding observed in places Planar beds also present

Fine to medium sand, moderately sorted, subrounded Unit seems to pinch out to the west, away from the Paria River, although pockets can be seen in other units.

Unit 10

Thickness: 90 cm **Same as unit 5** Sandstone clasts not quite as large Unit seems slightly more subangular throughout, with several local clasts present

Unit 11:

Thickness: 770 cm Contact: abrupt, change in grain size and type Extensive: laterally extensive , holding up highest terrace level with coarse grained exposure Sedimentary structures: n/a Texture: Grain size $-Max: \ge 1 \text{ m}$ Min: vf. sand Roundness - subangular to angular Sorting: very poorly sorted Color: red sandstone Composition: Matrix: mostly quartz Average: 9 cm

Clasts: almost entirely red sandstone, with low amounts of other clast types Cement: matrix hard to dislodge, compacted General rock name: clast supported angular boulder cobble gravel Secondary features: n/a

Final notes: deposit looks like it was a catastrophic event; side canyon or hillslope sediment; apparent almost open framework structure; boulders seem to be concentrated on top and bottom with more consistent cobbles making up the middle portion; walking up through the deposit, appears to maybe not be all one event - maybe multiple pockets that are distinctly separate.

Unit 12:

Thickness: 50 cm

Rounded fluvial gravels lay on surface, must have been above originally; some river gravel/sand lenses maybe present in the guts of unit 11, along the upper portion.

Appendix D. TOTAL-STATION SURVEY DATA

Figure D.1. Map of Quaternary deposits in the Lees Ferry area, showing approximate locations where survey transects were measured.

Table D.1. Raw survey data collected from the deposits along the Colorado and Paria rivers at Lees Ferry. Cross sections from each of nine transects were created using the raw easting, northing, and elevation survey data provided by the total station and were not projected in UTM coordinates due to the intercomparable nature of the study. Cross sections were generally measured perpendicular to the drainage axis; however, slight turns were sometimes introduced in order to capture terrace relations. In addition, two transects were started from one location and then completed after moving the total station. Multiple benchmark shots were included in order to connect the two survey lines. To correct for turns, all data were projected at right angles onto a line perpendicular to the drainage axis. The "distance along transect" column represents where data plotted along this line. In addition, the elevation of survey points within each transect was adjusted in relation to a high water datum (13,000 cfs).

Survey Transect through
Kaufman spring deposit

Survey Transect #4 -- Location B

Northern-Mid Survey Transect -- Near Stop Sign (cont...) / Hillslope
Profile

Lower Paria River transect -- just upstream of Paria trail parking area Survey Transect #6
Inst. Height (m)

Figure D.2. Raw Survey points from Transects 1A and 1B.

Figure D.3. Raw Survey points from Transect 2.

Figure D.4. Raw Survey points from Transect 3.

Figure *D.S.* Raw Survey points from Transect 4A.

Figure D.6. Raw Survey points from Transect 5.

Figure D.7. Raw Survey points from Transect 6.

Figure 0.8 . Raw Survey points from Transect 7.

Figure D.9. Raw Survey points from Transect 8.

Appendix E. TERRESTRIAL COSMOGENIC ¹⁰BE NUCLIDE DATA

Figure E.1. Map of Quaternary deposits in the Lees Ferry area, showing approximate locations where TCN samples were collected. Gray text indicates samples that were collected but are no longer being processed .

Table E.1. Site, chemical, and calculated data for terrestrial cosmogenic ¹⁰Be nuclide ages.

GC-04-LF-401

June 12, 2004

local ion:

UTM-- 445932 E 4078639 N, Zone 12.

 \sim 2 km south of Housing terrace, just above (west) of Paria gully ppt. gauge.

 \sim Elevation: 3230 feet (taken off topo map).

Surface Characteristic:

Possible M4 terrace deposit, M4y terrace surface. Dimensions: 1 km length, 100 m width (max), tilting slightly to southwest away from river. No rubification of clasts. Most clasts are not fractured on surface. Av horizon \sim 1 cm thick. Fragments of pedogenic carbonates on surface. Max clast size: \sim 20 cm. Majority of clasts are very well rounded. 20-30% clasts are sub-angular. Less than 1 plant/sq. meter (prickly pear and sage grass). Lithology: quartzite and chert dominate.

I surface sample of mostly 2 cm diameter clasts collected. Greater than 100 pebbles collected within 100 m of road ending - Future site or soil profile. Should give min. age of deposit. Recent erosion, possibly indicated by poor pavement, but no apparent gullying. Generally a flat surface. A voided sampling near plants. Sample collected $~-4$ m from edge of terrace.

GC-04-LF-404 ...

October 17, 2004 Series of Depth-Profile Samples # I Soil Pit dug at location of surface sample GC-04-LF-40I. 3 bulk-density samples collected through the vertical sequence.

locmion:

UTM-- 445932 E 4078639 N, Zone 12.

 \sim 2 km south of Housing terrace, just above (west) of Paria gully ppt. gauge.

 \sim Elevation: 3230 feet (taken off topo map).

Site Characteristics:

Possible M4 terrace deposit, M4y terrace surface.

Bioturbation extended to \sim 10 cm depth, therefore sampling began at 30cm rather than 50 cm as outlined in Gosse sampling strategy ... Out of zone of mixing.

Sampling for all depths performed around level line within ± 2.5 cm due to the presence of large clasts and pebbles.

Sampling for all depths performed around all three walls in order to increase the sampling swath. Samples from all depths were collected of all material in pit and then sieved in order to collect both the sand and pebble fraction.

GC-04-LF-404.30 S & **P**

Sample collected at 30 cm depth as specified above.

GC-04-LF-404.60 S & **P**

Sample collected at 60 cm depth as specified above.

GC-04-LF-404.100 S & P

Sample collected at 100 cm depth as specified above.

GC-04-LF-404.140 S & **P**

Sample collected at 140 cm depth as specified above.

GC-04-LF-404.180 S & P

Sample collected at 180 cm depth as specified above.

GC-04-LF-404.220 S & P

Sample collected at 220 cm depth as specified above.

GC-04-LF-405 ...

October 18, 2004

Series of Depth-Profile Samples #2

3 bulk-density samples collected through the vertical sequence .

Location:

UTM: 446458 E 4079938 N, Zone 12.

 \sim 300 m north of Lees Ferry Ranger Station.

Prominent sandy road cut on west side of road.

Possible M4 terrace deposit.

Elevation: \sim 3260 feet.

Outcrop Characteristics:

Deposit is tabular and lenticular bedded, with inter-fingering fine sands and silts (overbank) and lenses of pebble-cobble-gravels (side stem dominated).

Gravels display imbrication in a southward direction.

Outcrop exists due to ν 4 m high road cut.

Sand lenses could represent overbank Colorado River (light-tan color of fine sands).

Sampling for all depths performed around level line within ± 2.5 cm.

Samples from all depths were collected of all material in outcrop and then sieved in order to collect both the sand and pebble fraction.

Horizon shielding may need to be corrected for...

Compass Bearing = $C.B.$ Horizon Angle = H.A.

GC-04-LF-405.78 S & P

Sample collected at 78 cm depth as specified above.

GC-04-LF-405.148 S & P

Sample collected at 148 cm depth as specified above.

GC-04-LF-405.198 S & P

Sample collected at 198 cm depth as specified above.

GC-04-LF-405.300 S & **P**

Sample collected at 300 cm depth as specified above.

GC-04-LF-405.431 S & **P**

Sample collected at 431 cm depth as specified above.

GC-04-LF-405.563 S & **P**

Sample collected at 563 cm depth as specified above.

GC-04-LF-406

October 18, 2004

location:

UTM-- 446456 E 4079811 N, Zone 12.

 \sim 4 m north of 1st power line east of the main road.

On the north-western flanks of the housing terrace.

~Elevation: 3260 feet.

Surface Characteristic:

Possible M4 terrace deposit, M4y terrace surface.

Surface is \sim 2 m lower than the highest Housing terrace surface.

Dimensions: 12 m in N/S direction, and $5-8 \text{ m}$ in E/W direction.

Located on the flanks of the Housing terrace .

No rnbification of clasts.

Av horizon \sim 4 cm thick (maybe thicker because of enhanced slopewash).

Good desert varnish, although many of the darkest clasts have likely been previously varnished in the Shinarump deposit.

Many clasts come from the Shinarump, good quartzites, many of them sampled in depth-profile #2. Many clasts are fractured at the surface.

Mainstem (yellow) quartzites are not as abundant as at sample location for GC-04-LF-401.

Majority of clasts are not well-rounded, although many well rounded clasts are fractured.

Few pedogenic carbonates at the surface (1% of clasts).

Less than 1 plant / sq. meter.

Moderately-well to well-developed pavement on surface.

Max clast size: 20 cm.

A vcrage clast size: 2 cm.

I surface sample of mostly 2 cm diameter clasts collected. Greater than 100 pebbles collected $~4$ m from edge of terrace. Should give min. age of deposit. Generally a flat surface, although slopes slightly towards ravine to the north. Avoided sampling near plants.

Inheritance will be calculated in depth-profile samples.

GC-04-LF-407

October 18, 2004 *location:* UTM-- 446485 E 4079472 N, Zone 12. The Housing Terrace (south end). \sim 5 m south of the fence corner of the southern-most house. \sim Elevation: 3265 feet. *Surface Characteristic:* Possible M4 terrace deposit, M4o terrace surface. Bigger clasts than M4y sample GC-04-LF-406. Dimensions: \sim 1 km (N/S) by 200-300 m (E/W).

This is THE Housing terrace.

Some volcanic clasts (not many) present at surface.

No rubification of clasts. Very few pedogenic carbonates observed on the surface. Av horizon $~-4$ cm thick. Moderately developed desert varnish. Well-developed desert pavement. Most clasts are rounded, although some sub-angular and fractured clasts are present. Mainstem (yellow) quartzites are more abundant than at $GC-04-LF-406$, and perhaps equal to $GC-04$ -LF-401. Less than 1 plant / sq. meter. Max clast size: 20 cm.

Average clast size: 3-4 cm.

1 surface sample of mostly 2 cm diameter quartzite clasts collected.

Greater than 100 pebbles collected \sim 3 m from edge of terrace.

Should give min. age of deposit/terrace surface.

Generally a flat surface, although surface slopes slightly to the south (same direction as modern river). Avoided sampling near plants.

Inheritance will be calculated in depth-profile sample $#1$.

GC-04-LF-408

Oc tober 19, 2004

Location:

UTM-- 447685 E 4080059 N, Zone 12.

Possible M2 terrace deposit.

Located \sim 70 m west of bathroom at the Lees Ferry boat ramp.

Sample collected \sim 3 m NW of telephone pole.

 \sim Elevation: 3150 feet.

Surface Characteristic:

Dimensions: \sim 25 m (E/W) by 6-7 m (N/S).

More volcanic clasts than any other surface samples previously.

No rubification of clasts.

Very few pedogenic carbonates observed on the surface.

Av horizon \sim 3 cm thick.

Moderately well-developed pavement.

Desert varnish is moderately well-developed, although not as good as other sample surfaces.

Most clasts are rounded, very few fractured clasts are present.

Mainstem (yellow) quartzites and volcanics dominate the surface clasts composition.

Less than 1 plant / sq. meter.

Max clast size: 15 cm.

Average clast size: 1-2 cm.

1 surface sample of mostly 2 cm diameter quartzite clasts collected.

Greater than 100 pebbles collected \sim 3 m from edge of terrace.

Should give min. age of deposit/terrace surface.

Sample collected in NW portion of the terrace \sim 3-4 m from edge.

Generally a flat surface, although slopes gently to the east towards the Colorado River.

Avoided sampling near plants and telephone pole.

Inheritance will be calculated from depth-profile samples.

GC-04-LF-409 October 19, 2004 Location:

UTM-- 446031 E 4078859 N, Zone 12. Possible M3 terrace deposit. Located just above Paria beach parking lot to the SW. Sample collected in center of deposit, \sim 3-4 m from edges. ~Elevation: 3200 feet. *Surface Characteristic:* Fewer volcanics than at GC-04-LF-408 sample site. Dimensions: -8 m (E/W) by 8 m (N/S). No rubification of clasts. Few pedogenic carbonates observed on the surface, although slightly more than at GC-04-LF-408. Av horizon \sim 2.5 cm thick. Well-developed pavement. Well developed desert varnish, more than GC-04-LF-408. Most clasts are rounded to sub-rounded. Few fractured clasts. Less than 1 plant / sq. meter. Max clast size: 22 cm. Average clast size: I cm.

1 surface sample of mostly 2 cm diameter quartzite clasts collected. Greater than 100 pebbles collected \sim 3-4 m from edge of terrace. Should give min. age of deposit/terrace surface. Generally a flat surface, although slopes gently to the east towards the Colorado River. Avoided sampling near plants. Inheritance will be calculated from depth-profile samples.

CC-04-LF-410

October 19, 2004 *Location:* UTM-- 445585 E 4078458 N, Zone 12. Possible M6 terrace deposit, M6y terrace surface. Located above M4y where soil pit dug, on the southern most finger of the bird-shaped mapped unit. \sim 1 km south and west of Paria beach parking lot. ~Elevation: 3350 feet. *Surface Characteristic:* Dimensions: \sim 500 m (E/W) by 400 m (N/S). Clasts in pavement have algae rinds on them. No rubification of clasts. Well-developed pavement. Moderately to well developed desert varnish. Moderate amount of pedogenic carbonates observed on the surface. Av horizon \sim 3.5-4 cm thick. Most clasts are rounded to sub-rounded.

Some clasts are fractured.

Quartzites dominate the surface-clast composition.

More broken chert than other deposits.

More limestone (<5%) than other surfaces sampled.

Less than 1 plant / sq. meter.

Max clast size: 25 cm.

Average clast size: 3 cm.

I surface sample of mostly 2 cm diameter quartzite clasts collected. Greater than 100 pebbles collected $~6$ m from edge of terrace. Sample collected on the southern and eastern most portion of the terrace surface. Should give min. age of deposit/terrace surface. Avoided sampling near plants. Inheritance will be calculated from depth-profile samples.

GC-0S-LF-501

May 26, 2005 *Location:* UTM-- Possible M4 deposit, M4y terrace surface. Located southeast of ranger station $~100$ m, just above main road. Part of survey transect 3... surface is approximately 2 m lower than the main housing terrace, just to the west of staff houses. ~ Elevation: 3350 feet. *Swface Characteristic:* Dimensions: \sim 50 x 30 m (N/S). Well-developed pavement. Moderately to well developed desert varnish. Av horizon \sim 2-3 cm thick. Most clasts are rounded to sub-rounded. Quartzites dominate the surface-clast composition. Less than 1 plant / sq. meter. Max clast size: 25 cm. Average clast size: 3 cm.

1 surface sample of mostly 2 cm diameter quartzite clasts collected . Greater than 100 pebbles collected \sim 6 m from edge of terrace. Should give min. age of deposit/terrace surface. A voided sampling near plants. Inheritance will be calculated from depth-profile samples.

Appendix F. OPTICALLY STIMULATED LUMINSCENCE DATA Ages and errors shown in this appendix reflect data completed as of December 31, 2006, and may differ slightly from those reported in the thesis body due to the completion of more recent work.

Figure F.1. Map of Quaternary deposits in the Lees Ferry area, showing approximate locations where OSL samples were collected. Gray text indicates samples that were collected but are no longer being processed.

Sample #	UNL Sample #	Location Description	UTM Coordinates	Depth Below Surface (m)
$GC-04-$ LF- OSL1	UNL-1127	M4 terrace deposit, roadcut near Ranger Station	446458 E, 4079938 N	1.2
$GC-04-$ LF- OSL ₂	UNL-1139	P4(?) terrace on thick Paria River deposit - upper sample	445838 E, 4081493 N	13.5
$GC-04-$ LF- OSL ₃	UNL-1145	P4(?) terrace on thick Paria River deposit - middle sample	445842 E, 4081450 N	10.0
$GC-04-$ LF- OSL ₄	UNL-1138	P4(?) terrace on thick Paria River deposit - lower sample	445876 E, 4081459 N	30.0
$GC-04-$ LF- OSL ₅	UNL-1141	P3 terrace downstream of thick Paria deposit	445996 E, 4081085 N	6.2
$GC-04-$ LF- OSL ₆	UNL-1146	P2/flood deposit (?) closest to Lonely Dell Parking	446690 E, 4080594 N	8.0
$GC-04-$ LF- OSL7	UNL-1147	P2/flood deposit (?) closest to Lonely Dell Parking	446690 E, 4080594 N	3.0
$GC-04-$ LF- OSL8	UNL-1134	M2 near Paria stop sign, super sandy flood	446827 E, 4079827 N	3.2
$GC-04-$ $LF-$ OSL9	UNL-1128	M4 terrace deposit, up small drainage near stop sign - lower sample	446807 E, 4079863 N	22.0
$GC-04-$ $LF-$ OSL10	UNL-1133	M3 terrace deposit in culvert/gully cut near Paria beach - lower sample	446036 E, 4078918 N	6.5
$GC-04-$ LF- OSL11	UNL-1142	M3 terrace deposit in culvert/gully cut near Paria beach - upper sample	446036 E, 4078918 N	3.8
$GC-04-$ LF- OSL12	UNL-1149	M4 terrace deposit in separatede island of M4 on west side of road	446255 E, 4079606 N	14.0
$GC-04-$ LF- OSL13	UNL-1140	M4 terrace deposit on east side of road up to ranger station - middle	446292 E, 4079546 N	9.7
$GC-04-$ LF- OSL14	UNL-1148	M4 terrace deposit on east side of road up to ranger station - upper	446288 E, 4079636 N	3.0
$GC-04-$ LF- OSL15	UNL-1143	Spring deposit of Kaufman - lower sample – S3.	445504 E, 4078111 N	9.2
$GC-04-$ $LF-$ OSL ₁₆	UNL-1144	Spring deposit of Kaufman - upper sample - S3.	445504 E, 4078111 N	1.2
$GC-04-$ LF- OSL17	UNL-1353	M5 deposit taken near cosmo surface sample GC-04- $LF-410$	445585 E, 4078458 N	2.0
$GC-04-$ LF- OSL18	UNL-1352	M5y deposit taken from terrace west of campground - middle of deposit	445797 E, 4079287 N	5.0
$GC-04-$ LF- OSL19	UNL-1351	M2 deposit taken from terrace west of bathroom at ramp - top of deposit	447685 E, 4080059 N	1.5

Table F.1. OSL sample numbers and location descriptions for samples collected at Lees Ferry.

De (Gy) Error Age (ka)

GC-05-LF-OSL2 P4, Lees Ferry UNL-1139

Sample location:

Sample descript: medium-scale cross bedding, planar lamina, f to m sand UTM 12 4081493 **N,** 445838

166

±

GC-05-LF-OSL3 P4, Lees Ferry **IINI -1145**

Sample descript: f to m sand, small-scale cross strata, some pebbles present Sample location: UTM 12 4081450 N, 445842 E

167

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GC-05-LF-OSL4 P4, Lees Ferry UNL-1138

168

De (Gy) Error Age (ka) ±

GC-05-LF-OSL5 P3, Lees Ferry **UNI-1141**

Sample descript:

low angle and planar Sample location: UTM 12 4081085 N, 445996 E 169

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GC-05-LF-OSL8 M2, Lees Ferry UNL-1134

Sample descript: silt to f sand, no regular bedding **Sample location:** UTM 12 4079827 **N,** 446827 E

170

±

Age (ka)

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GC-05-LF-OSL9 M4, Lees Ferry **UNL-1128**

Sample descript: The misand, faint small-scale critical model ocation: UTM 12 4079863 N, 446807 E

Age (ka)

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GC-05-LF-OSL10 M3, Lees Ferry
UNL-1133

Sample location:

Sample descript: silt to f sand, planar and low angle cross strata (small-scale) UTM 12 4078918 N, 446036 E

Age (ka)

 \pm

GC-05-LF-OSL11 S3, Lees Ferry

Sample descript. Sit to 1 sand, massive
Sample location: UTM 4079606 N, 446255 E

De (Gy) Error Age (ka)

GC-05-LF-OSL13 M4 , Lees Ferry UNL-1140

±

Age (ka)

±

GC-05-LF-OSL 15 53, Lees Ferry UNL-1143

Sample location: UTM 12 4078111 N, 445504 E

GC-05-LF-OSL16 S3, Lees Ferry
UNL-1144

Sample location: UTM 12 4078111 N, 445504 E

176

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GC-06-LF-OSL17 M5, Lees Ferry **UNL-1353**

CODDIGS hennies, Sample descript: UTM 12 4078458 N, 445585 E 177

Age (ka)

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De (Gy) Error

GC-06-LF-OSL 18 MS, Lees Ferry UNL-1352

Sample location: UTM 12 4079287 N, 445797 E

GC-06-LF-OSL19 M3y, Lees Fe UNL-1351

De (Gy) 96.24

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wt Mean=

Sample descript: vf to m sand, fairly well sorted, small pebble lens above sampled unit **Sample location:** UTM 12 4080059 N, 447685 E