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S. L. Edburg
J. A. Hicke
P. D. Brooks
E. G. Pendall
B. E. Ewars
U. Norton

See next page for additional authors

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Cascading impacts of bark beetle-caused tree mortality on coupled biogeophysical and biogeochemical processes

Steven L Edburg1, Jeffrey A Hicke2, Paul D Brooks3, Elise G Pendall4, Brent E Ewers4, Urszula Norton5, David Gochis6, Ethan D Gutmann6, and Arjan JH Meddens7

Recent, large-scale outbreaks of bark beetle infestations have affected millions of hectares of forest in western North America, covering an area similar in size to that impacted by fire. Bark beetles kill host trees in affected areas, thereby altering water supply, carbon storage, and nutrient cycling in forests; for example, the timing and amount of snow melt may be substantially modified following bark beetle infestation, which impacts water resources for many western US states. The quality of water from infested forests may also be diminished as a result of increased nutrient export. Understanding the impacts of bark beetle outbreaks on forest ecosystems is therefore important for resource management. Here, we develop a conceptual framework of the impacts on coupled biogeophysical and biogeochemical processes following a mountain pine beetle (Dendroctonus ponderosae) outbreak in lodgepole pine (Pinus contorta Douglas var latifolia) forests in the weeks to decades after an infestation, and highlight future research needs and management implications of this widespread disturbance event.

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In a nutshell:

- Bark beetle disturbances differ from fire or logging disturbances because bark beetles alter forest functioning without immediately affecting forest leaf area or stem density
- Understanding the cascading ecosystem impacts of bark beetle outbreaks requires knowledge of the impacts on coupled biogeophysical and biogeochemical processes
- Critical biogeophysical impacts include changes in water and energy cycling, and alterations to the ecosystem services associated with water supply
- Critical biogeochemical impacts include changes in carbon and nitrogen cycling, and alterations to the ecosystem services associated with carbon sequestration

1Laboratory for Atmospheric Research, Department of Civil and Environmental Engineering, Washington State University, Pullman, WA (sedburg@wsu.edu); 2Department of Geography, University of Idaho, Moscow, ID; 3Department of Hydrology and Water Resources, University of Arizona, Tucson, AZ; 4Department of Botany and Program in Ecology, University of Wyoming, Laramie, WY; 5Department of Plant Sciences and Program in Ecology, University of Wyoming, Laramie, WY; 6Research Applications Laboratory, National Center for Atmospheric Research, Boulder, CO; 7Environmental Science Program, University of Idaho, Moscow, ID

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Likewise, critical biogeochemical impacts include reductions in plant C uptake, increases in decomposition, and potential loss of nutrients. An example of “coupled” biogeophysical and biogeochemical processes is the influence of canopy structure (leaf area and stem density) on the amount of precipitation captured by the foliage (and therefore on soil moisture), the effects of soil moisture on soil decomposition and plant growth, and the interaction between soil nutrients, decomposition, and plant growth (Figure 2).

Biogeophysical and biogeochemical impacts following bark beetle infestation have the potential to severely affect both natural resources and economic values. For example, snow from mountain ecosystems is the major source of water for more than 60 million people in the western US and Canada (Bales et al. 2006); changes in forest structure following bark beetle epidemics alter the amount, timing, and partitioning of this resource (Rex and Dubé 2006; Pugh and Small 2012). Post-insect-infestation tree mortality also affects C and N cycling in forests. Although most of these forests are net C sinks (e.g., Schimel et al. 2002), insect-related disturbances may cause them to release C to the atmosphere (Kurz et al. 2008). Nutrient cycling within affected forest ecosystems will also be modified, with reduced plant uptake increasing water and nutrient export. As a result, the aggregate impact of insect outbreaks may have consequences for regional and global weather and climate systems as well as for water supply and C storage.

Here, we present a chronological model of ecosystem impacts to help inform future management decisions and to identify future research areas that will improve understanding of insect-related disturbances. Our model focuses on the characteristic time scales of a mountain pine beetle (Dendroctonus ponderosae) outbreak in lodgepole pine (Pinus contorta Douglas var latifolia) forests (Figure 2), beginning in the initial days and weeks after infestation (Stage 1; Figure 3a), proceeding through a phase in which needles turn red in the months to years following the outbreak (Stage 2; Figure 3b), to the gray phase that occurs as needles fall off dead trees within 3–5 years following attack (Stage 3; Figure 3c), and finally to tree regeneration and snagfall in the decades following the outbreak (Stage 4; Figure 3d). Pine stands that are affected by mountain pine beetle infestations are typically dominated by lodgepole pines (>80% of the stem density), although spruce (Picea spp) and fir (Abies spp) are also found therein. Understory vegetation may be extensive in some stands (e.g., Brown et al. 2010). We anticipate that our conceptual model will provide a framework for future investigations of the impacts of bark beetles on forest ecosystems.

**Coupled biogeophysical and biogeochemical impacts**

**Stage 1: green attack (days to weeks)**

Mountain pine beetles preferentially infest and kill larger diameter host trees (pines), leaving smaller diameter trees and understory vegetation unaffected (Shore and Safranyik 1992). Beetles introduce blue stain fungi (Grosmannia clavigera) into tree xylem, which decrease and eventually prevent water transport (Paine et al. 1997). Sap flux studies have shown that a drop in transpiration occurs within one month of infestation and that the rate of change is dependent on fungal virulence (Yamaoka et al. 1990). This finding contrasts with experimentally girdled trees (simulating beetles feeding on phloem) that took up to five growing seasons to die, whereas those inoculated with blue stain fungi died in one growing season (Knight et al. 1991).

The initial impact of mountain pine beetle-induced mortality on lodgepole pine trees is hypothesized to occur in three stages. First, water transport in the stem shuts down, which results in the same response of stomatal closure as tree response to drought. Stomatal conductance and plant hydraulics are closely coordinated (Ewers et al. 2007). Second, a drop in stomatal conductance leads to a

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Figure 1. Areas affected by bark beetles from 1997–2010 (in the western US) and 2001–2010 (in British Columbia, Canada). Color of each grid cell represents the last year that bark beetle damage was observed by aerial surveys.
drop in photosynthesis because of CO₂ limitation (Flexas et al. 2004). Third, attacked trees die of C starvation due to lack of CO₂ as a result of reduced stomatal conductance (Adams et al. 2009; McDowell 2011), hydraulic failure (Sala and Hotch 2009; Sala et al. 2010), or a combination of the two (McDowell et al. 2011). Additional research is needed to determine the specific mortality mechanism (McDowell et al. 2011).

The decrease in water uptake described above leads to several mountain pine beetle/blue stain fungus impacts on lodgepole pine during the epidemic’s tree mortality phase. Lower hydraulic conductance within affected trees causes a reduction in N uptake, and reduced N availability further decreases hydraulic conductance, causing a feedback loop (Ewers et al. 2000). Tree hydraulic conductance prompts reductions in stomatal conductance during the weeks immediately after successful bark beetle colonization of a stand (Figures 2 and 4). In severe outbreaks, the reduced stomatal conductance of affected trees cascades from the tree to the canopy to the ecosystem.

Stage 2: fading trees and red attack (months)

During this stage, stand structure (leaf area and stem density) remain unmodified (Figure 5); however, cascading ecosystem impacts begin. As stomatal conductance declines in infested trees and as more trees are killed within a stand, the magnitude of impacts expands to the ecosystem scale (Figure 2). We hypothesize that evapotranspiration rates measured by eddy covariance will decline approximately in proportion to the amount of mortality that occurs within the footprint of the eddy covariance tower during the months following the outbreak, initially driven by a decline in transpiration (Figure 4). The degree to which tree-scale impacts will cascade to ecosystem-scale processes will depend on the number of trees killed, understory characteristics, and pre- and post-disturbance stand structure.

We also hypothesize that the rapid decline in tree transpiration will leave more water in the soil within the first few months post-outbreak (Figure 4; Clow et al. 2011).
Lodgepole pine forests are typically water-limited during the middle and late growing season (Knight et al. 1985; Fahey and Knight 1986), but in the first summer of attack we anticipate soil moisture to remain elevated (Morehouse et al. 2008). Growing-season soil moisture is hypothesized to be highest during Stage 2 as transpiration is reduced, while the now-red canopy shades the ground and minimizes atmospheric turbulence at ground level in the same fashion as an uninfested stand, reducing soil evaporation. Higher soil water contents are expected to reduce soil and litter temperatures and moderate the daily and seasonal variations of soil and litter moisture (Griffin et al. 2011). We believe that the increased soil moisture content and lower temperatures play key roles in the coupling between biogeophysical and biogeochemical processes (e.g. plant growth and decomposition; see coupling in Figure 2), which initiates ecosystem-scale impacts.

Altered soil moisture and temperature, together with reduced uptake of nutrients by trees, initial root death, and inputs of relatively N-rich litter, are expected to increase the amounts of mineral N available in soils within just a few months of the outbreak (Figure 6; Clow et al. 2011; Griffin et al. 2011). Extractable nitrate and ammonium concentrations will increase until microbial immobilization and plant uptake rates increase. These elevated concentrations will likely reduce plant hydraulic conductance as a result of lower root-to-leaf area ratios and higher resistance to cavitation (Ewers et al. 2000), both of which will limit how rapidly plant transpiration increases with succession. High resin-ammonium accumulation and/or soil nitrate concentrations have been reported by Morehouse et al. (2008), Clow et al. (2011),

Figure 3. Photographs of different stages of bark beetle outbreaks. (a) Stage 1, green attack; (b) stage 2, red attack; (c) stage 3, gray phase; (d) stage 4, snagfall and re-growth.
In the months following beetle-caused tree mortality, net C uptake for several years following beetle-induced mortality will fall; the magnitude of this change will depend on stand and outbreak characteristics, similar to stand-level evapotranspiration rates. Although modeling studies report large net C losses (Kurz et al. 1978) and demonstrate a potential for loss of nutrients from the system in either gaseous or dissolved form.

Forest–atmosphere C exchange will continue to change in the months following beetle-caused tree mortality. Plant C uptake will fall; the magnitude of this change will depend on stand and outbreak characteristics, similar to stand-level evapotranspiration rates. Although modeling studies report large net C losses (Kurz et al. 2008), eddy covariance measurements indicate that surviving vegetation can maintain net C uptake for several years following beetle-induced mortality (Brown et al. 2010). Root respiration will decrease within months following an infestation, as belowground C allocation is initially low during and after disturbance (Goulden et al. 2011). Model simulations suggest the reduction of root respiration will lower total soil respiration (root plus decomposition) for months to years following the initial infestation, despite enhanced decomposition of fine roots as a result of the increased soil moisture content and high N availability (Edburg et al. 2011). However, Morehouse et al. (2008) measured a constant soil respiration within the first several months following attack, suggesting that the decreases in root respiration were compensated for by increases in decomposition fluxes.

**Stage 3: gray attack (years)**

Dead trees typically enter the gray phase within 3–5 years of the initial infestation, as needles fall from killed trees (Amman et al. 1990). Needlefall causes a decrease in total leaf area (Figure 5), which has cascading impacts on coupled biogeochemical and biophysical processes (Figure 2). That is, a drop in leaf area reduces the aerodynamic drag of the canopy, thereby reducing the amount of momentum (ie momentum of air) absorbed by the canopy and altering surface exchange fluxes (Edburg et al. 2010). A reduction in leaf area also decreases the amount of solar radiation (Musselman et al. 2008) and precipitation (Zhang et al. 2004; Pugh and Small 2012) caught by the foliage, resulting in increased sunlight and water available at the soil surface. We argue that these changes will lead to decreased sublimation of snow from the canopy and increased surface sublimation and evaporation as more radiation and momentum of air penetrate to the snow surface and ground surface (Musselman et al. 2008; Royer et al. 2011). Furthermore, snow redistribution will be substantially modified because it is a function of wind and turbulence within and above the forest (Stottlemeyer and Troendle 1999). Peak soil moisture should occur earlier in the spring as the timing of snow melt changes (Volotchi et al. 2009), whereas lower transpiration will allow soil moisture to remain elevated later into the growing season (starting in Stage 1; Figures 2 and 4). Surface soil moisture may become considerably more spatially variable as the increasingly patchy canopy results in locally enhanced solar input and turbulence, which, in turn, amplifies evaporation from the land surface. These increases will also partially offset reductions in transpiration (Gustafson et al. 2010; Royer et al. 2011).

Carbon and N cycling during the gray attack period is altered as a result of increased needle decomposition and continued root decomposition from Stage 2 (Edburg et al. 2011). Because decomposition and microbial uptake of N are dependent on soil temperature and moisture, the C and N cycles are strongly dependent on the aforementioned biogeophysical changes to canopy structure and water cycling during this stage (as shown in Figure 2). We further hypothesize that increased substrate availability from needle and root inputs to soil, which are greater than background rates (Edburg et al. 2011), will lead to increased decomposition (Figure 6). However, total soil respiration rates will continue to be suppressed unless decomposition rates increase enough to compensate for the loss of root respiration during this stage.

Given the potential compensating effects of changes in root respiration and decomposition, the response of surviving trees and understory vegetation is critical in determining the net ecosystem C balance during this stage (Romme et al. 1986; Brown et al. 2010; Pfeifer et al. 2011). As decomposition increases, N immobilized in microbial biomass may be retained in the ecosystem and potentially transferred to vegetation as microbial biomass turns over (Brooks et al. 1998). Alternatively, both denitrification (Brooks et al. 1997) and hydrologic N export (Brooks and Griffin et al. (2011). This initial response has been observed in many forested ecosystems following disturbance (eg Likens et al. 1978) and demonstrates a potential for loss of nutrients from the system in either gaseous or dissolved form.

**Figure 4.** Time series of changes to stomatal conductance (Gs), transpiration (T), evaporation (E), and soil moisture (Sm) during and following mountain pine beetle-caused mortality in a lodgepole pine forest. Morehouse et al. (2008) reported increased soil moisture after a bark beetle outbreak in a ponderosa pine (Pinus ponderosa) forest in the months to years following attack.
Williams 1999) may reduce plant-available N and limit the growth of regenerating vegetation. Increases in soil moisture, nutrients, and sunlight available to the understory are hypothesized to promote plant growth (Romme et al. 1986; Brown et al. 2010; Pfeifer et al. 2011; Royer et al. 2011). Because lodgepole pines can be limited by both water and N (Fahey and Knight 1986), we expect the net effect on productivity of increased water and sunlight and the variations in N availability to be site dependent because of differences in soil properties (water-holding capacity) and pre- and post-outbreak species composition (eg rates of transpiration; Fahey and Knight 1986). In some locations, surviving trees and understory vegetation respond favorably to increased water and N in the gray stage, whereas in other N-limited locations, further N losses may outweigh increases in sunlight or water and favor symbiotic N fixation by the understory, potentially postponing forest regeneration.

**Stage 4: snagfall and re-growth (decades)**

Snags begin to fall and the understory and surviving trees increase growth rates (productivity) in Stage 4, as the decrease in canopy cover that began in Stage 3 continues (Figures 2 and 5). The rate of snagfall is a function of soil type, soil temperature and moisture, and tree species (Lewis and Hartley 2005). Canopy openings created by snagfall during this period accumulate more snow than surrounding forests because less snow is intercepted (caught) by the canopy (Golding and Swanson 1978) and wind re-deposition of snow increases (Stottlemeyer and Troendle 1999). As new seedlings and surviving trees grow during the following decades, the opening is filled in and interception of snow increases until the canopy is closed (Figure 5).

Additions of C from snags and coarse woody debris combined with labile C in fresh litter and root exudates will stimulate decomposition and thus immobilization of N (Figure 2; Harmon et al. 1986; Edburg et al. 2011). We expect the magnitude of the increases in decomposition and immobilization to be smaller than the increases in Stage 3 but they will be longer lasting because C and N from the snags will enter the soil at a slower rate and because the substrate quality is lower than that of roots and needles (Figure 6). Large C inputs to the soil from snagfall and high rates of N immobilization may lead to large belowground stores of C and N, which could be susceptible to groundwater N export as in Stage 3.

The availability of sunlight, water, and N will play a key role in the growth of understory vegetation, surviving trees, and tree seedlings (Fahey and Knight 1986; Romme et al. 1986). The rate of plant growth in this phase is also linked to snagfall (Figure 2). After snagfall, we expect increased sunlight to reach the surface and elevated soil moisture, leading to enhanced plant growth over the decades in Stage 4. However, we also expect available N to play an important role in plant growth during this phase. Additional C inputs to the soil from snags will decrease C limitations on microbes and thus increase the immobilization of N, which will reduce plant-available N, thereby restricting plant growth (Figure 6). This reduction in growth is hypothesized to last until the snags decompose (Figure 5). At this point, immobilized N from snag decomposition may become available to plants, leading to increases in plant growth, with perhaps higher productivity and leaf area as compared with pre-disturbed productivity (Edburg et al. 2011).

Recovery of net ecosystem productivity (NEP = 0) to steady-state conditions was hypothesized by Odum (1969), whereas Goulden et al. (2011) demonstrated that near-steady-state conditions were achieved about 150 years after fire in the boreal forests of northern Canada. We anticipate faster recovery rates following bark beetle disturbances for several reasons: the amount of tree mortality within a given location is often lower than for stand-replacing fires, in which the majority of trees are killed; in addition, bark beetles prefer larger host trees, leaving non-host trees, smaller diameter trees, or shrubbery or herbaceous vegetation intact (Shore and Safranyik 1992). Thus, post-outbreak stands may include a considerable quantity of surviving vegetation that is able to take advantage of increased nutrients, light, and water, and so undergo rapidly increasing growth rates (Romme et al. 1986; Veblen et al. 1991; Brown et al. 2010; Pfeifer et al. 2011). Successional changes after bark beetle outbreaks (Veblen et al. 1991) may also play a key role in recovery.
Conclusions

We face many challenges in testing and refining this conceptual framework of cascading ecosystem impacts of bark beetle-caused tree mortality. Evaluation of our framework requires datasets spanning the time frame of the cascading impacts (weeks to decades). We currently have only limited observational datasets that describe the location, severity, and duration of tree mortality. These dynamic processes must also be integrated into ecosystem models, which can then be evaluated against coordinated field observations. Long-term observational studies that span pre- and post-outbreak periods are the most useful because they capture the temporal dynamics needed for model evaluation. Amiro et al. (2010), for instance, synthesized a chronosequence of eddy covariance measurements for several disturbances; however, bark beetle outbreaks were poorly represented in this synthesis because of the difficulty in maintaining long-term studies, locating eddy covariance towers within potential outbreak areas, and using such towers in topographically complex and geographically remote areas.

Quantifying how changes in stand-scale processes alter biogeochemical and biogeochemical cycling at regional scales remains an open research question. At the small catchment scale, where affected host species comprise a large fraction of the catchment area, a reduction in transpiration can be expected to increase stream flow. However, at the regional scale, if additional unaffected stands are included, it is unclear whether canopy-scale impacts will have any notable effects. Stand-level productivity can be substantially reduced following a severe bark beetle outbreak, suggesting that regional impacts occur as a result of severe outbreaks (e.g. the mountain pine beetle attack in British Columbia; Kurz et al. 2008). However, other regional occurrences are not as severe as that in British Columbia (Pfeifer et al. 2011), and the consequences of regional-scale events in North America may depend on the duration, size, and severity of attacks. Furthermore, it is unknown how impacts on biogeochemistry and biogeochemistry from severe bark beetle outbreaks compare to other forcing signals, such as those associated with annual climate variability. Several key research themes that should be addressed in the future are:

- resolve tree mortality mechanisms caused by bark beetles and associated blue stain fungi;
• characterize how changes in soil moisture affect the coupling between biogeophysical and biogeochemical processes, including vertical and lateral moisture distribution;
• characterize and quantify site dependency of soil microbial feedbacks and the response of surviving trees and understory vegetation, including the potential export of nutrients;
• quantify the spatial extent and severity of outbreaks at watershed to regional scales;
• quantify changes in biogeophysical cycling across a wide range of spatial and temporal scales at sites with different successional trajectories; and
• isolate (or at least more clearly attribute) beetle-driven changes in ecosystem-scale water, energy, and biogeochemical budgets from those driven by natural climate variability and long-term climate and ecosystem trends.

In our opinion, addressing the aforementioned research themes will increase the understanding of the coupled biogeophysical and biogeochemical impacts of bark beetle outbreaks and lead to improved pre- and post-disturbance management decisions.

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